# Density Currents in Shear Flows: Effects of Rigid Lid and Cold-pool Internal Circulation, and Application to Squall Line Dynamics

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## **Summary**

In this paper, the idealized density current models developed in previous studies that contain a rigid lid are evaluated. It is shown, through numerical experiments, that a realistic stable layer placed at the level of the rigid lid acts very much like a rigid lid, in confining the environmental flow to the levels below and in controlling the behavior of density currents that propagate in the flow. When this stable layer is replaced by a neutrally stable layer, the behavior of the density current is very different. These results support the hypothesis that in the atmosphere, the tropopause or a tropospheric inversion layer acts much like a rigid lid in the idealized density current models, and the solutions of these models are applicable to organized quasi-two-dimensional convective systems such as squall lines in the atmosphere.

The effects of cold pool internal circulations on density current behavior are also examined. When a single circulation is present initially inside the cold pool, the direction of circulation that gives rise to a smaller shear across the density interface leads to a smoother interface and a much more steady density current head. Large-amplitude eddies develop along the interface when the direction of circulation is reversed. When two layers of shear are initially present inside the cold pool, the flow pattern inside the cold pool tends to be dominated by the circulation of the lower layer. In this case, the behavior of the density interface and density current head depends on more than just the initial cross-interface shear.

The overall flow pattern in the density current in which the cold pool circulation contains rearward flows at the ground level bears a close similarity with that found in mature squall line systems, as shown by comparisons with a simulated squall line. When the cold pool in a squall line is defined in terms of the equivalent potential temperature, the solutions of idealized density currents in sheared flows become more readily applicable to squall lines. Keywords: Density current, Squall line dynamics

#### **1. Introduction**

It is generally recognized that the interaction between the environmental shear and the cold pool of a thunderstorm outflow (density current) plays an important role in supporting longlived squall lines (Moncrieff 1978; Moncrieff 1992; Thorpe et al. 1982 TMM82 hereafter; Rotunno et al. 1988 RKW88 hereafter; Fovell and Ogura 1988; and Xue 1990). Density currents also occur in the forms of sea-breeze fronts (e.g., Simpson et al. 1977) and the gust fronts associated with cold frontal rainbands (e.g., Carbone 1982). Recent detailed observational studies further revealed the importance of density current type sea-breeze and cold outflow fronts in the initiation and organization of convection (e.g., Atkins et al. 1995; Weckworth and Wakimoto 1992).

To improve our understanding of the interaction between density currents and their environment, nonlinear, two-fluid steady-state models of increasing degree of realism have been developed in recent years, by Xu (1992, X92 hereafter), Xu and Moncrieff (1994, XM94 hereafter), Liu and Moncrieff (1996a), Xue et al (1997, XXD97 hereafter), and Xue (2000, X2000 hereafter). An idealized three-fluid model was also considered in Moncrieff and So (1989). Compared to the classic inviscid density current model of Benjamin (1968), these models considered additional elements: (i) constant and non-constant environmental shear, (ii) internal cold pool circulation, (iii) negative vorticity generation associated with energy loss along the interfacial layer between the density current and its environment, and (iv) density stratification and latent heating. These idealized models show, through mathematical analyses, how the depth, propagation speed and shape of the density current are controlled or influenced by the environmental shear and cold pool strength. The results from these rather idealized theoretical models have been validated, to the extent possible, by carefully designed high-resolution numerical experiments (e.g., Xu et al. 1996; Xue et al. 1997; Xue 2000; Liu and Moncrieff 1996b).

As pointed out in XXD97 and discussed further in X2000, the density currents in most of the earlier studies are subject to a potentially important limitation: the flow is restricted by a rigid upper boundary. The rigid lid assumption makes the theoretical treatment viable because the top boundary of the problem is a streamline that can be pre-determined from environmental conditions. Typically, the presence of this rigid lid forces the flow above the density current to run through a narrow channel, thereby reducing the pressure (through the Bernoulli effect) in the region. The reduced pressure often contributes significantly to the support of deep density currents. XXD97 and X2000 suggested that in the real atmosphere a strong inversion layer or the tropopause may act as a rigid lid in channeling the flow underneath therefore the results from the idealized density current models with a rigid lid is applicable to the atmosphere. When this rigid lid is replaced by a stable layer, internal gravity waves are forced in the stable layer by disturbances from below. The gravity wave solutions have to be coupled with the density current solution below, making theoretical treatment difficult. In this paper, we resort to a numerical model for the answer. A set of numerical experiments are conducted in which a stable layer replaces the rid lid of a previously studied density current in a sheared flow, and it is found that the density current solution below is only quantitatively different from those under a rigid lid. The solution is very different when this stable layer is also removed.

When we do believe that the tropopause acts like a rigid lid, the atmospheric density current in the form of thunderstorm outflow, defined in terms of the potential temperature deviation from the environment, seldom reaches more than a third of the tropospheric depth, however. This is related to the typical thermodynamic structure in the vertical hence the limit in cold air supply. As a result, the real atmospheric density currents tend to be highly turbulent because their height is far below the steady state solution. To examine this issue, XXD97 conducted an experiment (experiment SLSA in that paper), where the environmental shear is strong enough to support a deep current having a steep frontal slope while the cold air supply is insufficient for it to reach the steady-state depth. It was found that even though the flow was highly turbulent and the cold pool was shallow, the numerical solution still exhibited a branch of deep jump-type flow, with the envelope of cold air and frontal slope more or less agreeing with the steady-state solution (see Fig.12 of XXD97). The results suggest that even in the cases where steady state solution cannot be reached due to the lack of cold air supply, the theoretical solution of the density current depth and speed still offer a very useful guidance on the overall flow behavior.

The previously studied density current models can be made more applicable to quasi-twodimensional atmospheric convective systems such as squall lines if we generalize the two-fluid models to include any flows with fluids of different densities. In a well-established squall line system (see, for example the conceptual model of TMM82 in Fig.1), we can consider that the entire colder (defined in terms of equivalent potential temperature because the convective region is saturated) descending branch rear-to-front inflow together with the rotor in the head region constitutes the density current. *The density difference and the dynamic (pressure-gradient) force balance across the flow interface* (shown as thick dashed line in Fig.1), *together with Bernoulli energy conservation along the streamlines, determine the overall (quasi-steady) flow structure as they do the density current*. In this sense, the results of our idealized two-fluid density-current models can be extended to explain the squall line dynamics, at least qualitatively.

In mature squall line systems, the circulation associated with the downdraft branch and the rotor circulation in the head region of the density current / cold pool are common features. In most previous studies, the effect of the circulation inside density current is neglected. Xu and Moncrieff (1994) included a cold pool circulation with constant vorticity in their idealized (inviscid two-fluid) steady-state model of a density current propagating into a flow with uniform vertical shear. They found that, in comparison with the environmental shear, the strength of the cold pool internal circulation has only a small quantitative effect on the solution of the current (see their Fig. 2) while the direction of the circulation was found to have no bearing on the solution. A counterclockwise internal circulation associated with positive vorticity has exactly the same effect on the solution as a clockwise circulation. The original idealized density current model for uniform shear environment has since been generalized for non-constant environmental shears and verified by numerical experiments (XXD96, XXD97, and X2000). Since the solutions of the latter are obtained from the same set of physical principles, we expected that the funding of XM94 remains qualitatively valid for the idealized models with non-constant shear. XM94's results, however, clearly depend on the assumption that the flow system is inviscid. The idealized models neglect the fictional effect between the two fluids even though shear there is usually large. In a real world as well as in numerical experiments, the development of Kelvin-Helmholtz (KH) eddies at the interface hence the resultant turbulence mixing clearly depends on the crossinterface shear. For this reason, the direction of the internal circulation is expected to affect the behavior of the density current, in perhaps a significant way. Understanding such effects through numerical experimentations is another purpose of this paper.

In the following, we will first summarize the basic principles used by and the key and relevant results from the previous theoretical models. Section 3 discusses the design of numerical experiments for density currents and section 4 will present and discuss results from these experiments. To further illustrate the applicability of the density current model results to squall

line systems, we present in section 5 results from a realistic squall line simulation and compare the key flow features in the simulated squall line with those in density currents. Finally, summary and discussions are given in section 6.

### 2. Idealized density currents in a two-layer shear flow

In this section, we briefly discuss the theoretical models of idealized two-dimensional density currents. Xue (2000) considers an idealized steady-state density current in an environmental flow with two constant shear layers under a rigid lid. The model, with an additional cold pool circulation with constant vorticity  $\gamma$ , is illustrated in Fig. 2. The variables in this model are non-dimensionalized using the following scaling:

$$(x,z) \leftarrow (x,z)/H, \quad (u,w) \leftarrow (u,w)/U, \quad p' \leftarrow p'/(\rho_0 U^2),$$

where the variables on the left hand side of the arrow are non-dimensional.

In Fig. 2, the shaded area indicates the cold pool of an idealized density current. The flow is assumed to be in a steady state in a reference frame moving with the current. The x and z are the horizontal and vertical coordinates; u and w are the horizontal and vertical velocities, respectively; H is the depth of the domain bounded by two rigid boundaries;  $U \equiv (gH\Delta\rho/\rho_0)^{1/2}$ is the velocity scale; g is the acceleration of gravity;  $\Delta\rho \equiv \rho_1 - \rho_0$  is the density difference between the denser fluid inside the cold pool ( $\rho_1$ ) and the lighter fluid outside ( $\rho_0$ ); and  $p' \equiv P - P_0$  is the perturbation pressure, which is the difference between total pressure P and reference pressure  $P_0$ . The reference pressure  $P_0$  is the unperturbed pressure in the upstream inflow associated with constant density  $\rho_0$ , where  $\rho_0$  satisfies the hydrostatic relation  $P_0 = g\rho_0(H-z)$ . Viewed in a framework moving with the density current front, the low-level environmental flow moves towards the head from the right (the upstream). The flow outside the cold pool is separated into two layers, each with a constant vorticity. This constant vorticity dictates that the horizontal flow at the far downstream has the same shear as that at the far upstream in both the upper and lower layer. The circulation inside the cold pool, considered in this paper, is assumed to have a constant vorticity,  $\gamma$ .

In Fig. 2,  $\alpha$  and  $\beta$  are, respectively, the constant lower and upper-level shears of the environmental inflow.  $d_0$  ( $0 \le d_0 \le 1$ ) is the depth of the lower layer in the far upstream and  $c_0(>0)$  the constant inflow speed (absolute value) at the lower boundary.  $d_1$  is the depth of the lower shear layer above the cold pool in the far downstream, and  $c_1(>0)$  is the speed of outflow atop the cold pool. Finally, h (0 < h < 1) is the depth of the density current head. Without considering the cold pool internal circulation (as in X2000), the flow is fully described by seven non-dimensional control parameters,  $\alpha$ ,  $\beta$ ,  $d_0$ ,  $c_0$ ,  $d_1$ ,  $c_1$ , and h.

Following the discussion of Benjamin (1968), the remote system-relative environmental inflow and outflow are constrained by mass continuity, vorticity conservation, flow-force balance and the conservation of Bernoulli energy along streamlines. It can be shown that only three of the seven parameters are independent, and physical solutions can be found given any three of these parameters. Xue (2000) discusses the cases where the shears of inflow ( $\alpha$  and  $\beta$ ) and the height of the inflow layer interface ( $d_0$ ) are prescribed. Once the inflow shears ( $\alpha$  and  $\beta$ ) and layer interface height ( $d_0$ ) are given, the ground-level inflow speed ( $c_0$ ), the depth of density current

(*h*), the outflow speed  $(c_1)$  and layer depth  $(d_1)$  are determined by the model. A numerical procedure is used to find the solutions for *h*,  $c_0$ ,  $c_1$  and  $d_1$  and the results are plotted as curves. When the cold pool circulation is considered, the shear (vorticity) of this circulation becomes an additional external parameter. Since these eight parameters describe the flow away from density current head region where the flow is horizontal, the results from the aforementioned procedure describe the far-field solutions.

According to X2000, the density current becomes deeper and propagates faster relative to the surface environmental flow as the shear in the low-level inflow increases from negative through positive values. The effect of the upper-level shear is similar – increased upper-level shear supports deeper density current. When the shears in both layers are equal, the solution degenerates to that of X90 for constant shear inflow and to that of Benjamin (1968) when the shear in both layers is zero. In general, when the overall shear in the inflow is positive, the density current depth is more than half of the vertical channel depth.

The effects of cold pool internal circulation on the density current in a constant environmental shear were examined in XM94. It was found that the internal circulation of constant vorticity has only a relatively small (typically less than ten percent) quantitative effect and the effect is independent of the direction of circulation.

Apart from the far-field solutions, the flow structure near the front is also important. In the case of thunderstorm outflow, the shape of the front has a direct influence on the vertical orientation of low-level updraft. An upshear-tilted updraft allows condensed water to be downloaded underneath the updraft without interrupting the low-level inflow, thus long-lived squall line type convection can be maintained (e.g., TMM82; RKW88).

von Karman (1940) and Benjamin (1968) showed analytically that the angle at the stagnation point between the density front and the horizontal is  $60^{\circ}$  for an idealized inviscid density current in a uniform (non-sheared) environmental flow with a free-slip lower boundary condition. It was further shown in X92 and XM94 that this  $60^{\circ}$  angle is independent of the inflow shear and cold-pool circulation. The shape of the frontal interface away from the lower boundary is determined by the continuity condition of pressure across the interface and the solution can be obtained numerically (Xu et al. 1992). According to X92 and XM94, when the positive inflow shear is very strong, the frontal interface slope can become steeper than  $60^{\circ}$  at the mid-depth (z=h/2) location. It is expected that these findings can be extended to the model with two layers of inflow shear.

In the following two sections, we will discuss the design of and results from densitycurrent numerical experiments.

# 3. The design of numerical experiments for density current

As in XXD97 and X2000, a modified version of the Advanced Regional Prediction System (ARPS, Xue et al. 1995; Xue et al. 2000) is used for our density current simulations. The ARPS is used in its simplest dry two-dimensional mode with no physics. To facilitate direct comparison of numerical results with the theoretical solutions, the Boussinesq approximation is assumed. The base-state density  $\rho_0$  and potential temperature  $\theta_0$ , are set to a constant. For the experiments with rigid top lid, rigid free-slip conditions are applied at both top and bottom boundaries. For those with a stable upper-layer, wave-radiating upper-boundary condition is used (Xue et al. 2000). Open conditions are also used at the upstream and downstream boundaries. More details on the model setup for rigid lid cases can be found in XXD97 and X2000. As in earlier studies, we do not include sub-grid scale turbulence parameterization, but rely on the high spatial resolution to resolve the turbulent eddies often found to the rear of the current head. We apply a small amount of horizontal computational mixing/diffusion on the velocity fields only. We avoid applying mixing in the vertical for fear of affecting the vertical wind profile. The monotonic flux-corrected transport (FCT) scheme of Zalesak (1979) is used to advect potential temperature field.

The equations actually solved by the model are those for horizontal velocity u, vertical velocity w, perturbation pressure p' and perturbation buoyancy b in non-dimensional form, as given in XXD97 as Eq.(3.1). The same scaling parameters as in XXD97 are used here for non-dimensionalization of the model equations, except for the definition of H.

Length scale:	<i>H</i> (the height of the rigid lid or the bottom			
	of the upper stable layer)			
Velocity scale:	$U \equiv (gH\Delta\rho/\rho_0)^{1/2}$			
Pressure scale:	$P \equiv \rho_0 U^2 = g H \Delta \rho$			
Time scale:	$T \equiv H/U$			
Buoyancy scale:	$B \equiv g\Delta \rho / \rho_0 = g\Delta \theta / \theta_0$			

In the dimensional ARPS model, we choose for convenience H=1 km,  $\Delta\theta=3^{\circ}$ K,  $\theta_{0}=300^{\circ}$ K, g=10 ms<sup>-2</sup>. The other derived scaling parameters are U=10 m/s, P=120 Pa and T=100 s. Since our results will be presented and discussed in non-dimensional space, the results are not limited to the applications of shallow domain.

The density current in the numerical model is generated by placing an initially static block of cold air (b = -1) in the middle portion of an elongated computational domain of size  $40 \times D$  non-dimensional units consisting of  $767 \times (80 \times D+1)$  grid points. Here D is either 1 or 3. Therefore, the grid spacings are  $\Delta x = 0.05$  and  $\Delta z = 0.0125$  in the horizontal and vertical directions, respectively. The initial cold block is specified to be more than 8 non-dimensional units in width, and it is, in most cases, able to supply sufficient cold air for achieving and maintaining a quasi-steady density current. The initial shape of the cold pool is specified in the same way as in XXD97 [c.f. Eq.(3.3) there]. The interface has a slope of  $60^{\circ}$  at the frontal nose.

To examine the effect of rigid top lid versus a stable layer, we conducted a set of three experiments, listed as LID0, LID1 and LID2 in Table 1. Experiments LID1 and LID2 are based on control experiment LID0, which is essentially experiment SLS in XXD97 (see Table 1 in XXD97). In this experiment, a rigid lid is placed at non-dimensional height D=1, and the shear in the inflow is confined to the low level depth of  $(d_0=)$  0.2 and the shear magnitude is a relatively large value of  $(\alpha=)$  3. No cold pool internal circulation is considered in this set of experiments, i.e.,  $\gamma=0$ . The theoretical model predicts a deep density current of depth (h=) 0.77 (see Fig. 2 of XXD97). The initial flow for the experiment is specified in the same way as in XXD97, i.e., the upstream and downstream flow is determined from the specified parameters and the corresponding theoretical model solutions, and the interior flow is obtained by solving a steady-state vorticity equation expressed in terms of the streamfunction. The top boundary condition for the streamfunction is constant following the upstream lower boundary and along the density current interface. This initial setup closely resembles the steady state solution of the theoretical model.

Experiments LID1 and LID2 are variations of LID0. In LID1, the rigid lid in LID0 is removed and a stable layer of depth 2 and a static stability (N) of  $2 \times 10^{-2}$  s<sup>-1</sup> (roughly corresponding to the stability of an isothermal layer) is placed above the original channel of depth 1. The purpose here is to examine how much this stable layer acts like a rigid lid, and if the density current solution is significantly altered by this change. This experiment employs a waveradiating top boundary condition at height 3, allowing gravity waves to exit the boundary freely. In experiment LID2, the stable layer in LID1 is replaced by a neutral layer of the same depth and a rigid lid is placed at height 3. In both experiments, the initial cold block is the same as in LID0, and the zero-shear upper-level inflow is extended to the top of the domain. The initial flow near and above the cold block is again obtained from the vorticity equation, but this time with the constant-streamfunction upper-boundary condition specified at z=3. Significant initial flow adjustment is expected due to the presence of a stable layer.

Experiment LID2 is essentially a case of density current propagating in a neutrally stable environmental flow of depth 3. If we rescale the problem using the new domain depth of 3, then the low-level inflow shear is about 0.067 in depth (in terms of the total domain depth). According to the theoretical models of XXD97, the steady-state density current depth should be more than 0.5 of the domain depth, which, in terms of our current scaling, gives a depth of more than 1.5. The initial depth of the cold block is 0.77 in LID2 and is specified according to the solution with a domain depth of 1, therefore significant adjustment is expected. Because the initial depth is significantly lower than the expected depth therefore the cold air supply is likely to be insufficient, a steady-state solution may not be able to establish. This is a situation that is similar to experiment SLSA in XXD97.

The other set of experiments examine the effects of internal circulation inside the density current or cold pool. We introduce such circulation by specifying the far-field (downstream) flow shear. We assume that the net inflow into the density current and therefore the vertically averaged (through the depth of the current) flow speed at the far downstream is zero. From Fig. 1, we can see that both positive (du/dz > 0) and negative circulations are possible inside the generalized density current of a convective system, with the negative circulation typically associated with the rotor circulation and the positive one with the downdraft circulation to the rear. At different stages of a squall line development, one of these may dominate. When the rotor circulation, one obtains a two-layered circulation within the density current, with a layer of positive vorticity located above a layer of negative vorticity. In other cases where the rotor circulation is weak and the rear-to-front inflow jet (see, e.g., the conceptual mode of Houze 1989) is fully developed, inside the density current, one would expect a negatively sheared layer located above the positively sheared downdraft circulation with rear-to-front inflow in-between (c.f. Fig. 1).

We conceptualize the above situations by considering single and double circulations of different directions inside the cold pool and examine their effects on the behavior of the density current in a set of four numerical experiments. In experiment ICN1, a single cold-pool internal circulation with negative vorticity of magnitude 1 is considered. Experiment ICP1 is the same as ICN1 except for the direction of the internal circulation. These two experiments corresponds to the cases of  $\gamma = -1$  and +1 as illustrated by Fig. 2.

In experiments IC2N2 and IC2P2, two internal circulations are included, corresponding to two layers of shear of opposite signs at the downstream boundary of the density current (see Figs.10a and 11a). In IC2N2, the lower layer has negative and upper-layer positive vorticity of magnitude 2. IC2P2 is the same except that the internal circulations have opposite directions. The

magnitude of 2 is chosen so that the flow speeds at the ground level and at the density interface are the same as those in the corresponding single layer cases. Therefore, the cross-interface shear is the same in the single and double circulation cases when the flow direction below the interface is the same.

In this second set of (four) experiments, the initial configurations are based on the theoretical model solution for an upper-level-shear-inflow case with a rigid lid and zero internal circulation. Inflow shear below 0.2 is zero and above is 2. This reference case is considered in X2000 as experiment L0U2 (see Table 1 there) and it supports a relatively deep density current. The results from L0U2 are included in Table 1. The theoretical depth of the current (without internal circulation) is 0.72 and the model-simulated depth is 0.68.

In the numerical solutions, transient activities are expected to the rear of the current head due to KH instability. When the simulation results are compared with the idealized model, we determine the height of the simulated current at a location where the flow interface becomes more or less horizontal and KH eddies are insignificant. In the following section, we will discuss the results of these numerical experiments.

## 4. Results of numerical experiments of density current

All numerical simulations are carried out to a non-dimensional time T=24. Ensemble time averages are produced, for simulated fields over the period T=12 to T=18 at a sampling interval of 0.5. To avoid spatial smoothing due to the current propagation, the instantaneous fields are translated in x before averaging so that their frontal noses collocate with the nose at time T=12. Instantaneous fields at selected times and the time-averaged fields will be shown. All fields shown will be non-dimensional and the flow speed will be density-current-front relative.

## (a) Rigid lid versus a stable layer

The results from experiments LID0, LID1 and LID2 are discussed in this section. Fig. 3 and Fig. 4 show the flow and potential temperature fields from these experiments at non-dimensional time T=12 and 18. Fig. 5 shows the corresponding fields averaged over the period from T=12 through 18.

The simulated density current in LID0 is essentially stationary relative to the grid and the head is in a nearly steady state. The height of the current height is estimated to be between 0.75 and 0.78 from T=12 and 18, very close to the theoretical value of 0.77. The ground-level inflow relative propagation speed is estimated to be 0.79 between T=12 and 18, again very close to the theoretical value of 0.78. Since LID0 is essentially a re-run of experiment SLS in XXD97, Fig. 3a and Fig. 5a should closely match Fig.10a and Fig.10b in XXD97, respectively, with the small differences being attributed to the use of different advection schemes. Given the very deep density current, the environmental flow is channeled rapidly through a shallow depth above the current, resulting in a large pressure drop over the head due to the Bernoulli effect. This effect is clearly shown in Fig. 6a. This channeling effect and the resultant pressure drop is instrumental in creating a large upward pressure gradient that supports such a deep and quasi-steady density current.

The simulated density current in LID1 is not as stationary as that in LID0. At T=12, the front is located close to x=1.5 and at T=18 close to x=2.0. As a result, the ground-level-inflow-relative front propagation speed of front between T=12 and 18 is about 0.85 versus the theoretical value of 0.78, a difference of about 10 percent. The density current height, measured by the 299K potential temperature contour (corresponding to non-dimensional buoyancy of -0.5) in the time averaged plot (Fig. 5b) is about 0.78, very close to the theoretical value 0.77 of the

rigid lid case. The simulated head is not as steady as the rigid lid case, however, and the KH eddies develop more quickly along the interface and gain large amplitudes (Figs.3b and 4b). The obviously new features found in this run are the internal gravity waves in the stable layer forced by the flow and eddy activities below. The forcing is strong enough for wave overturning to occur at T=18 (Fig. 4b); in fact, the overturning of isentropes remains evident in the time averaged plot (Fig. 5b). What is equally important to us is that right above the current head, the 301K isentrope that was originally located near z=1 level has now been pushed up to nearly z=1.2 height by flow channeled over the current head. The raised isentropes make for more room for the neutral inflow to pass over the head, resulting less acceleration there.

The pressure drop due to the Bernoulli effect is subsequently less in LID1, as seen from Fig. 6. The minimum pressure over the head is about -0.34 as compared to -0.42 in LID0. This reduced pressure drop should lead to weaker vertical pressure gradient force hence shallower density current. The latter does not seem to be happening given that the simulated current height is close to or even slightly larger than the theoretical value. This can be explained by the fact that the effective depth of domain in which the current propagates is increased by almost 20 percent due to the raised lower boundary of the stable layer. If one re-scales the density current depth by the new increased domain depth of 1.2, the non-dimensional height of 0.78 would be equivalent to 0.65, significantly lower than the original theoretical value of 0.78. In another word, for a density current propagating into an environmental flow capped by a layer of strong static stability, the decreased channeling effect above the density current hence decreased support from vertical pressure gradient is roughly compensated by the increased scale depth due to the upward shift of the base of stable layer. The depth of the density current head (based on the original scaling) ends up to be very close to that predicted by a theoretical model with a rigid lid. This is true at least for the inflow configuration considered here.

The speed of the current in this case is found to be about 10 percent faster than the rigid lid case. This can again be understood quite well by re-scaling the velocity values. Since the velocity scale we need for non-dimensionalization (see section 2) is  $U \equiv (gH\Delta\rho/\rho_0)^{1/2}$ , a 20 percent increase in H leads to roughly 10 percent increase in the velocity scale. Using the new velocity scale that is 10 percent larger, we obtain a non-dimensional propagating speed that is almost exactly the same as the theoretical value. The exact values here should be viewed with some caution, however, because when one changes the scale depth *H*, the effective depth of the low-level shear is also changed. Still, the above arguments should be qualitatively valid. The slightly larger pressure gradient across the frontal nose at the ground level in LID1 is consistent with slightly deeper current head and less pressure drop above and therefore the faster propagation speed.

The results from experiment LID2, in which the rigid lid is in effect extended to z=3 level, are shown in Figs.3c, 4c and 5c. The behavior of the simulated density current in this case is quite different from the previous two cases. The current head is much more unsteady. Large eddies develop very quickly behind the head region and grow in size. In fact, the interface never reaches a quasi-steady horizontal state in the head region. As a result, the time-averaged field (Fig. 6c) shows significant smoothing due to eddy mixing in the potential temperature field, starting right behind the front. In time averaged field, the two-degree temperature difference spans over a depth of about 1.4, with the negatively buoyancy air reaching a height of 1.9. The current also propagates significantly faster than previous two cases, with the ground-level inflow-relative speed being about 0.98, roughly 25% over that of LID0 case. As mentioned earlier, this is a case that is similar to experiment SLSA considered in XXD97. For a channel of depth 3, the

specified initial cold pool is much lower than the steady state solution given by the theoretical model. In this case, the inflow has a positive shear of magnitude of 3 and a depth of about 7 percent of the channel depth, the expected current depth should be more than 0.5 of the channel depth, i.e., more than 1.5. Clearly, the model is trying to establish such a deep current, but for the lack of sufficient cold air supply, a quasi-steady state cannot be reached. The flow is therefore much more turbulent, but still, a tendency to establish a deeper cold pool and a deep jump-type updraft is obvious, as was the case in SLSA of XXD97 (see, e.g., Fig.10 there).

Furthermore, because of the large space above the cold pool, the channeling effect for the inflow is much weaker, the pressure drop due to Bernoulli effect is much smaller. As shown by Fig. 5c, much of the channeling effect occurs above the 1.2 level where the magnitude of pressure perturbation is less than 0.3 units (Fig. 6c), smaller than those in the previous two cases. The two centers of minimum pressure below this level in Fig. 6c are associated with two horizontal eddies seen in Fig. 5c, due to the dynamic (centrifugal) effect in vortices. The deeper though somewhat diffused cold pool and the smaller pressure drop aloft lead to higher pressure underneath the density current. The maximum surface pressure perturbation in Fig. 6c is 0.444 versus 0.291 in LID0, and the large cross-frontal pressure gradient causes the front to propagate significantly faster than the LID0 case.

In summary, the results from experiments LID0, LID1 and LID2 show that a realistic stable layer placed at the level of a rigid lid acts very much like a rigid lid, in confining the environmental flow to the levels below and in controlling the behavior of density currents that propagate in the flow. For a low-level-shear-inflow case examined here, the density current obtained has essentially the same depth as that under a rigid lid (placed at the height of the base of the stable layer), the inflow-relative propagation speed of the current is faster, but the difference is less than 10 percent. If this stable layer is replaced by a neutrally stable layer under a much higher lid, the behavior of the density current is very different. The current attempts to reach a much deeper depth consistent with the much deeper channel and the flow becomes much less steady become of the lack of sufficient cold air supply needed to reach a quasi-steady state. These behaviors can be quantitatively understood by re-scaling the parameters using the new, modified channel depth.

#### (b) Effect of density current internal circulations

In this section, we examine the results from the second set of experiments that include circulations inside the cold pool. Figures 7 and 8 show the fields at T=0 and T=18, respectively, from experiment ICP1 and ICN1, in which a single circulation of, respectively, positive and negative vorticity is specified inside the cold pool at initial time (Fig. 7). Other settings of the experiments are the same as those for experiment L0U2 in X2000, and are specified according to the solution of the corresponding theoretical model. In L0U2, the simulated density current remains essentially stationary relative to the model grid, and its depth is close to the theoretical value (0.68 versus 0.72). Figure 9 shows the time-averaged fields corresponding to those in Fig. 8. Inferring from the theoretical study of XM94 on inviscid density currents, the internal circulation should have only very small effect on the density current, and for inviscid cases considered there, the effect does not depend on the direction of internal circulation.

It is clear that the flow behaviors in these two cases are very different. The turbulent eddies along the interface in ICP1 are large and strong, developing along the frontal interface quickly. On the contrary, the interface in ICN1 is rather smooth for a significant distance behind the head region. This interface is also remarkably steady. In fact, this is the case we have seen so far that is closest to the solution of inviscid density current considered in the theoretical studies.

The interface seen from the time-averaged field (Fig. 9b) is almost as sharp as the instantaneous one in Fig. 8b, indicating very little transient activities along the interface. The interface in ICP1 is significantly diffused, however (Fig. 9a). Such large differences are obviously due to the difference in the cross-interface shear. In ICP1, the flow below the interface is in an opposite direction as the strong front-to-rear flow above, creating a large shear across the interface therefore a favorable condition for KH eddies to develop. In ICN1, the flow associated with the internal circulation is in the same direction below the interface as that above the interface, reducing the cross-interface shear (even compared with the case without internal circulation), therefore the interface is much more stable with respect to KH instability. The fast rearward advection of any vortices developed along the interface by flows both below and above the interface further reduces the disturbances found in the head region.

The height of density current, measured by the maximum height of the -0.5 buoyancy contour in the time-averaged field, is about 0.6 for ICP1 and 0.71 for ICN1. The former is significantly lower than the theoretical value of 0.72, mostly due to the diffusive eddy activities near the head. The height in ICN1 is very close to the theoretical value. In fact, it is even closer (0.70 versus 0.68) than the case without internal circulation (L0U2). In both cases, the speed of the current is very close to the theoretical value, with that in ICP1 moving slightly slower and that in ICN1 slightly faster. The difference in both cases is actually smaller than in L0U2. In a sense, by introducing an internal circulation that reduces the cross-interface shear, we obtain a density current in a numerical model that is even more close to the solution of a two-fluid inviscid theoretical density current model. These results further cross-validate our theoretical models and numerical simulations of density current. The numerical results also show that for realistic flows, the direction of the cold pool circulation does matter.

In the next two experiments, the cold pool contains at the initial time two share layers having vorticity of opposite signs. The initial flow patterns are shown in Figs.10a and 11a. In IC2P2, the lower layer in the cold pool has a positive shear / circulation and the upper layer has a negative shear. As a result, there exists a rear-to-front inflow jet at the mid-level of cold pool while the flow right under the current interface has the same direction as the flow above. This configuration resembles the flow pattern in a mature squall line system. Based on the single layer results seen earlier, one would expect relatively smooth density interface due to the relative small shear across the interface. It turns out to be untrue. As seen from the time evolution of the flow in Fig. 10, the upper negatively sheared layer in the cold pool appears to be dynamically unstable. Large KH eddies quickly develop along this shear layer and the density interface. The negativevorticity-bearing air becomes wrapped into the vortices / eddies and the less perturbed part of cold pool becomes dominated by the positive-vorticity-bearing air from the lower layer. This layer is apparently more stable, aided at least in part by the fixed lower boundary. At a result, the behavior of the density current is mainly influenced by the direction of circulation in the lower (internal) shear layer. The time-averaged fields in Fig. 10d show several interesting features. First, the buoyancy gradient originally concentrated in a single thin layer at the interface between the environmental flow and the cold-pool air is now concentrated in two distinct zones - one between the negative-vorticity-bearing air and the environmental air above, and one separating the lower-layer positive circulation and the negative one above. Positive rear-to-front flow is concentrated in the second zone. As pointed out earlier, such a flow pattern resembles that in a mature squall line. In a mature squall line, positive vorticity is generated baroclinically on the backside of the cold / heavy downdraft and is fed into the low-level rearward outflow, forming a circulation that is similar to the lower layer positive circulation. Baroclinically generated negative vorticity along the frontal interface and that generated at the rear edge of the upper-level outflow

contribute to the establishment of the upper negative circulation inside the (generalized) cold pool (Weisman 1992). Our current density current model does not attempt to fully address the issue of vorticity generation though, the main goal here is to investigate the behavior of the density current and the associated flows given an initial distribution of vorticity. Such distributions are idealizations of the more realistic situations, and are believed to be representative of the reality.

Finally, let us examine the results of experiment IC2N2, in which the initial cold pool circulations are reversed in direction (Fig. 11a). Negative vorticity is found in the lower layer and positive vorticity between this layer and the density interface. By design, the shear across the interface is the same as in ICP1, in which large eddies develop along the interface due to strong KH instability. Figure 11 shows the time evolution and the time average of the flow. Interestingly, despite the large cross-interface shear, the interface becomes rather smooth at later times (e.g., T=18), and the large amplitude eddies are only found a distance of almost 3 from the frontal zone. The flow pattern resembles that of ICN1 more than ICP1. As in the previous case, the instability of and the inability to maintain the upper positive share layer are obviously responsible for such a behavior. The lower-layer negative circulation is more stable, and with time the head region of the cold pool becomes dominated by this negative circulation, which, together with the flow above the cold pool, sweeping the eddies developing along the interface rearward and establish a more eddy-free density current interface. By T=18, significant, though large in amplitude, eddies are only found about 3 length units behind the frontal zone. Some positive vorticity is still found inside the cold pool at this time though. The interior flow at this time can perhaps be linked to fully developed rotor circulation in squall line systems. This situation is usually found in squall lines with significantly rearward-tilting updraft and rearwarddisplaced downdraft.

The results from the above two experiments show that the flow pattern inside a cold pool tends to be dominated by circulation located near the ground, which is dynamically more stable. The behavior of the density interface and density current head depends on more than just the cross-interface shear or the direction of circulation right below the interface.

# 5. Results from a high-resolution 2-D squall line simulation

To further illustrate the applicability of our density current models and the related simulation results to squall line dynamics, we present in this section the results from a high-resolution 2-D squall line simulation. A version of the ARPS (Xue et al. 1995; Xue et al. 2000; Xue et al. 2001) is used in a 2D cloud model mode with Kessler-type warm rain microphysics and 1.5-order subgrid-scale turbulence closure. A special feature used for this study is the synchronized flux-corrected transport scheme developed recently by the current author. When used to advect potential temperature, water vapor and the hydrometeor species together, the procedure of flux-correction synchronized for thermodynamic and microphysical variables results in, in addition to monotonicity (hence positive-definiteness of water variables), better conservation of quantities such as equivalent potential temperature along air parcels. Ice processes, radiation, PBL and land surface processes are neglected.

The computational domain is  $500 \times 19.2 \text{ km}^2$ . The vertically grid spacing is 100 m uniformly below 3 km then stretched gradually to 500 m at the model top. The horizontal grid spacing is 200 m, a resolution significantly higher than that used by most simulation studies of this type. The convection is triggered by an initial elliptic thermal bubble that is 20 km in horizontal extent and 2.8 km in the vertical and is centered 1.4 km above ground and 200 km

from left boundary. In the figures to be presented, the center of this initial bubble is marked as x = 0. Open lateral boundary conditions are employed. The top boundary is rigid and a Rayleigh damping layer is placed above 14 km.

The Weisman and Klemp (1982) analytical thermodynamic sounding is used, with the potential temperature and temperature set to 343K and 213K, respectively, at the tropopause at 12 km height. Potential temperature at the surface is 300K. Specific humidity is kept constant below 1.2 km at 15 g/kg.

The environmental wind profile has a constant -22.5 m/s speed below 2 km, and the speed is linearly reduced to zero at 3.285 km, and remains zero above. This profile is different from the typical low-level linear-shear profiles studied by TMM82, RKW88, Fovell (1988), and others, and was previously examined by Xue (1990). The latter showed that such 'step' inflow profiles support long-lasting squall lines that tend to produce more precipitation than low-level linearly sheared profiles with similar total shear magnitude. In this paper, we focus our discussions on flow patterns near the cold pool of simulated squall line and on the relevance of density current dynamics in a typical squall line system. First, the precipitation rate of the simulated squall line is plotted in a time-space diagram in Fig. 12. It can be seen that the overall system is quasi-steady, with new cells being continually re-generated at the leading edge of cold pool and propagating to the left (at a typical speed of 14 m/s). The overall squall line system moves to the left at a speed of about 2.6 m/s, due to the unusually strong (22.5 m/s) low-level inflow.

Similar to the density current results presented earlier, we perform an ensemble time average on the squall line fields, over a period of 1 hour and centered on 10 hours. Data were sampled at 5 min intervals and all fields were shifted in horizontal direction before averaging so that the gust fronts collocate with that at 9.5 hours. As shown by Fig. 12, this is a period that the system is rather steady (in the system-relative reference frame). The averaging removes most of the small-scale and transient signals and accentuates the broader scales of flow. The instantaneous fields at 9.5 hours are shown in Fig. 14. The velocities shown are system relative.

The shading in Fig. 13a shows regions of relatively low equivalent potential temperature  $(\theta_{a})$ . Because in most of the domain shown, the air is saturated as indicated by the simulated radar reflectivity contours, the  $\theta_e$  contours give a good representation of the air parcel trajectories, and of streamlines in this (quasi-steady, again in the system-relative reference frame) case. It is clear that the cold pool is consisted of low  $\theta_e$  air that has been brought down from the mid-levels to the ground level by the downdraft circulation. This cold pool is separated by the updraft from the mid-level low  $\theta_e$  air ahead (to the right) of the squall line, in a similar way as the cold pool in the density current model is separated by the updraft from the mid-level upstream environmental flow (see, e.g., Fig. 9). It can been seen that the air in the main updraft has similar  $\theta_e$  values as that directly over the cold pool, in a manner that is similar to the lighter, uniform-density fluid in the idealized two-fluid density current model (c.f. Fig. 2 and Figs.8-11). The main difference between the two is the presence of a layer of low  $\theta_e$  air at the mid-level ahead (to the right) of the main updraft found in the squall line system. However, since the air in this region has little vertical motion, the vertical distribution of regular buoyancy instead of  $\theta_e$  is more relevant. The perturbation buoyancy field in Fig. 13b shows near-zero values in this region, therefore the dynamics of flow in the region is not much different from that of neutrally stable flow in the idealized two-fluid density current model. Judging from the  $\theta_e$  field, the 'generalized' cold pool exhibits a raised head of more than 5 km in depth. This is, however, still lower than that predicted by a steady-state density current model, and as a result we observe a density current head with significant, transient eddy activities (Fig. 14). This situation is also similar to density current experiment SLSA in XXD97 and experiment LID2 discussed earlier. The density current head as indicated by the buoyancy field (Fig. 13b) is about 3.5 km deep, shallower than that seen from the  $\theta_e$  field.

Figure 13c shows that a zero and positive-vorticity bearing updraft is separated from the cold pool and downdraft circulation by a slanting layer of negative vorticity. Most of this negative vorticity (shaded) is generated, through horizontal buoyancy gradient (more clearly seen in Fig. 14b), along the cold pool-updraft frontal interface and at the upper levels on the left side of positively buoyant updraft. The negative vorticity thus generated is transported rearward by the front-to-rear updraft branch of flow in the form of growing eddies, much like the KH eddies found in the density current simulations (Fig. 14b). Most of the positive vorticity in the downdraft branch is generated at the back edge of the cold pool, where the horizontal buoyancy gradient reverses direction (Figs.13c and 14c). Overall, the flow pattern in the squall line system bears close similarities with those of two-fluid density currents with a near-ground cold-pool circulation that contains positive vorticity. In this regard, we can compare instantaneous vorticity and flow fields in Fig. 14c with those in Figs.10d and 8a, and the time averaged fields in Fig. 13c with those in Figs.10e and 9a. The perturbation pressure field in Fig. 13d shows a pattern similar to those of simulated density currents, such as that in Fig.6b. The main difference is in the additional layer of positive pressure perturbations in the squall line that is associated with upperlevel (above 7 km level) convective outflow. It is the result of hydrostatic response of the atmosphere to condensational heating. Finally, we point out that the updraft of simulated squall line has an orientation close to  $45^{\circ}$  at the mid-levels and the gust front is at an angle close to  $60^{\circ}$ to the horizontal direction at the surface.

#### 6. Summary and Discussion

In this paper, the effect of the imposed rigid lid in the idealized density current models developed in the previous studies is examined. It is shown, through numerical experiments, that a stable layer located at and above the level of an original rigid lid acts in a similar way as the rigid lid. The slightly decreased channeling effect above the density current due to the raised base of stable layer tends to be offset by the effect of increased channel depth. For the particular density current and inflow configuration considered, the simulated density current has an almost identical depth as that obtained under a rigid lid. The density current propagates slightly faster than the latter case. The behavior of density current is very different if the stable layer is also removed. All of these results can still be adequately explained by the theoretical model with a rigid lid through the rescaling of parameters based on the new, effective channel depth. These results support our previous suggestions that in the atmosphere, the tropopause or a tropospherical inversion layer acts in a similar way as a rigid lid found in our idealized density current models, and the results of these models are applicable to organized quasi-two-dimensional convection systems in the atmosphere.

The effects of cold pool internal circulations in simulated density currents are also examined. When a single circulation is present inside the cold pool, the direction of circulation that gives rise to a smaller shear across the density interface leads to a smoother interface and a much more steady density current head. Large-amplitude eddies develop along the interface when the circulation is reversed. When two shear layers are initially present inside the cold pool, the flow pattern inside the cold pool tends to be dominated by circulation located near the ground, which is dynamically more stable. The behavior of the density interface and density current head depends on not just the cross-interface shear or the direction of circulation right below the interface as suggested by the single circulation results. The overall flow pattern in the density current in which the cold pool circulation contains rearward flows at the ground level bears a close similarity with that found in a mature squall line system, as is shown through the comparison with a simulated squall line presented in this paper. It is argued that the density current in a squall line can be viewed in terms of the equivalent potential temperature and such a generalized density current can reach a depth of five or more kilometers in depth. The solutions of density currents in sheared flows from idealized models can help us understand the fundamental flow dynamics in squall lines, in particular, the interaction of shear inflow with the cold pool.

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Experiment	Specified Parameters			Theoretical <sup>*</sup> /Simulated Values		
	α	β	γ	D	h / simulated $h$	co/simulated co
LID0	3	0	0	1	0.77/0.76	0.78/0.79
LID1	3	0	0	3, stable upper layer	0.77/0.78	0.78/0.85
LID2	3	0	0	3, neutral throughout	0.77/1.13	0.78/0.98
L0U2	0	2	0	1	0.72/0.68	0.97/0.95
ICP1	0	2	1	1	0.72/0.60	0.97/0.96
ICN1	0	2	-1	1	0.72/0.71	0.97/0.98
IC2P2	0	2	2	1	0.72/0.51	0.97/0.95
IC2N2	0	2	-2	1	0.72/0.70	0.97/0.99

# Table 1. Initial settings and simulated parameter values for the numericalexperiments of density currents

\* Note that the theoretical values listed for ICP1 through IC2N2 are those corresponding to L0U2, which does not include any cold pool internal circulation.

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Fig. 5. Same as Fig. 3 except that the fields plotted are averaged over a period from T=12 through 18.



Fig. 6. The time-averaged non-dimensional perturbation pressure field for experiments (a) LID0, (b) LID1, and (c) LID2, corresponding to those in Fig. 5. The contour interval is 0.005.



Fig. 7. Non-dimensional velocity vector and buoyancy (contour) fields at time initial time (T=0), for (a) experiment ICP1 and (b) ICN1. A single circulation with positive (clockwise) vorticity of 1 is specified inside the cold pool at initial time in ICP1 and a single circulation of opposite direction (counterclockwise) in ICN1. Areas with negative vorticity are shaded. For ICP1, negative vorticity appears only at the flow interface at this time.



Fig. 8. As Fig. 7 but at non-dimensional time T=18.



Fig. 9. As Fig. 7, but for fields averaged over the period T=12 through 18.



Fig. 10. Non-dimensional velocity vector and buoyancy (contour) fields from experiment IC2P2 (a) at time T=0, (b), T=6, (c) T=12, (d) T=18, and (e) averaged over the period from T=12 through 18. The initial cold pool circulation can be seen in (a). Areas with negative vorticity are shaded.



Fig. 11. Same as Fig. 10 but for experiment IC2N2, in which the initial cold pool circulations [shown in (a)] have opposite directions as in IC2P2.



Fig. 12. Time-space diagram of precipitation rate from the simulated squall line. Regions with rainfall rate exceeding 150 mm/h are shaded. Contour interval is 25 mm/h.



Fig. 13. Fields from simulated squall line averaged over a one-hour period at 5 min intervals. In (a), areas with equivalent potential temperature less than 336K are shaded. Think lines are reflectivity contours at 20, 40 and 55 dBZ. Wind vectors are plotted in all panels. Other panels include, (b) the temperature-equivalent perturbation-buoyancy field (ms<sup>-2</sup>) including water loading with negatively buoyant areas lightly shaded and those colder than -2K heavily shaded; (c) vorticity (multiplied by 1000, s<sup>-1</sup>) field with negative areas lightly shaded and those less than -20 heavily shaded; and (d) perturbation pressure field (Pa). The figures are plotted to the physical scale, therefore the wind vectors show the actual direction of the flow.



Fig. 14. Same as Fig. 13, except for instantaneous fields at 9h 30 min.