

# Severe Weather Ingredients

**Material by Tom Galarneau Andrew Lyons,  
Rich Thompson and Harry Weinman**

# **Ingredients for Organized Severe Thunderstorms**

- **Instability (Lapse Rates)**
- **Lift (QG, mesoscale, convective scale)**
- **Moisture (return flow, soil moisture, evapotranspiration)**
- **Vertical wind shear**



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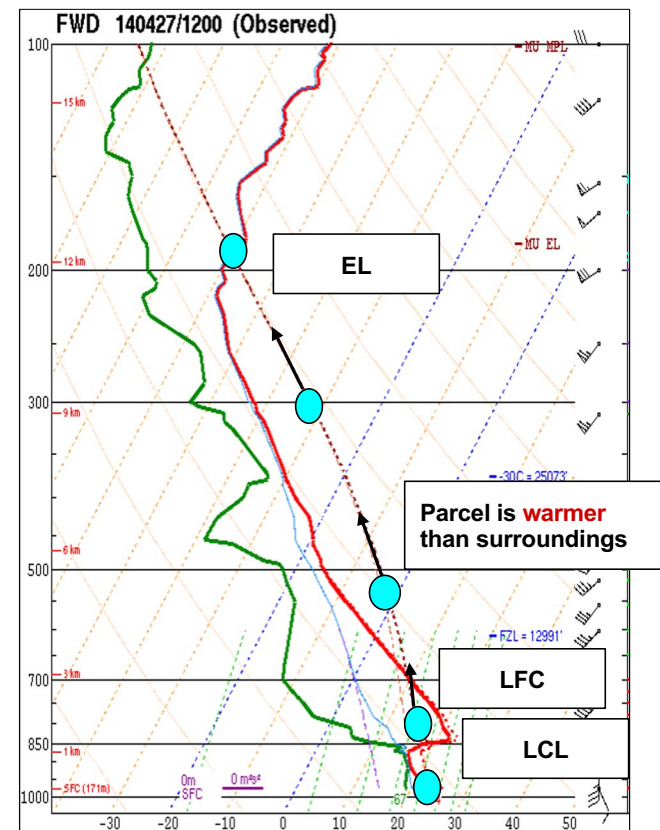
This lecture

- a separate lecture

Key ingredient for  
organized severe storms

# Instability: The Basics

- Give a parcel a nudge and see what happens!
- We determine an environments stability by its lapse rate (Change of environmental temperature with height)

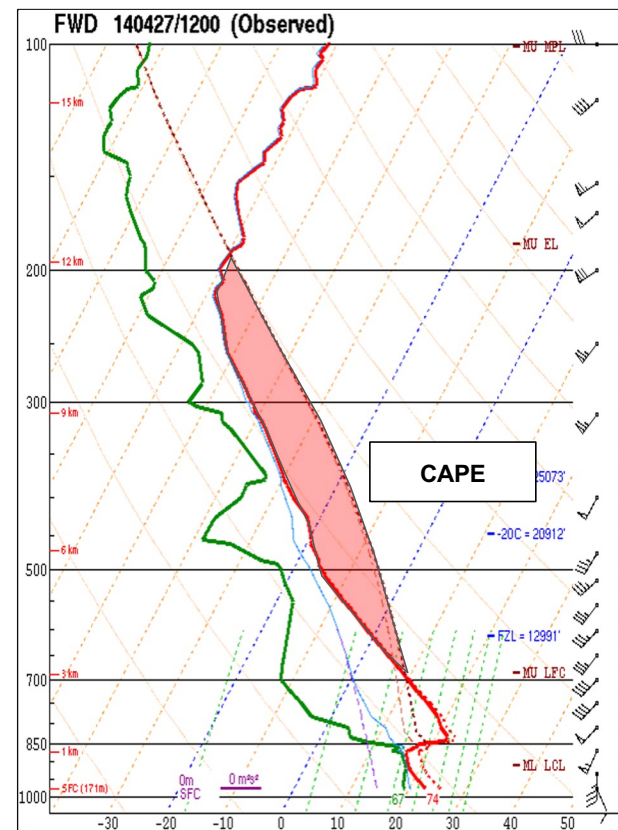


# Instability: The Basics

- Integrate the depth where parcel is warmer/less dense than its surrounding
- CAPE (Convective Available Potential Energy)

$$CAPE = \int_{LFC}^{EL} g \frac{T_{v_p} - T_{v_{env}}}{T_{v_{env}}} dz = \left[ \frac{kg \ m^2}{s^2} \frac{1}{kg} \right] = \left[ \frac{m^2}{s^2} \right]$$

$\searrow$   
 $= \left[ \frac{J}{kg} \right]$

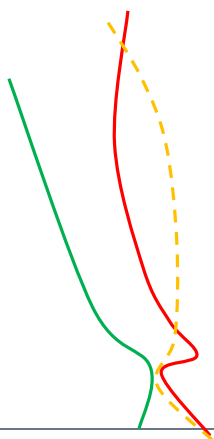


# Instability: The Basics

Take a small finite layer depth  
within CAPE profile

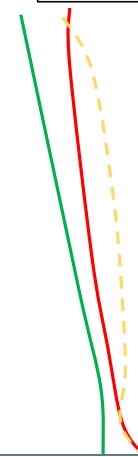
$$\frac{CAPE}{\Delta Z} = \left[ \frac{m^2}{s^2} \right] \left[ \frac{1}{m} \right] = \left[ \frac{m}{s^2} \right] \longrightarrow$$

Upward  
Acceleration!



**Short/thick CAPE:**  
Drives larger upward acceleration

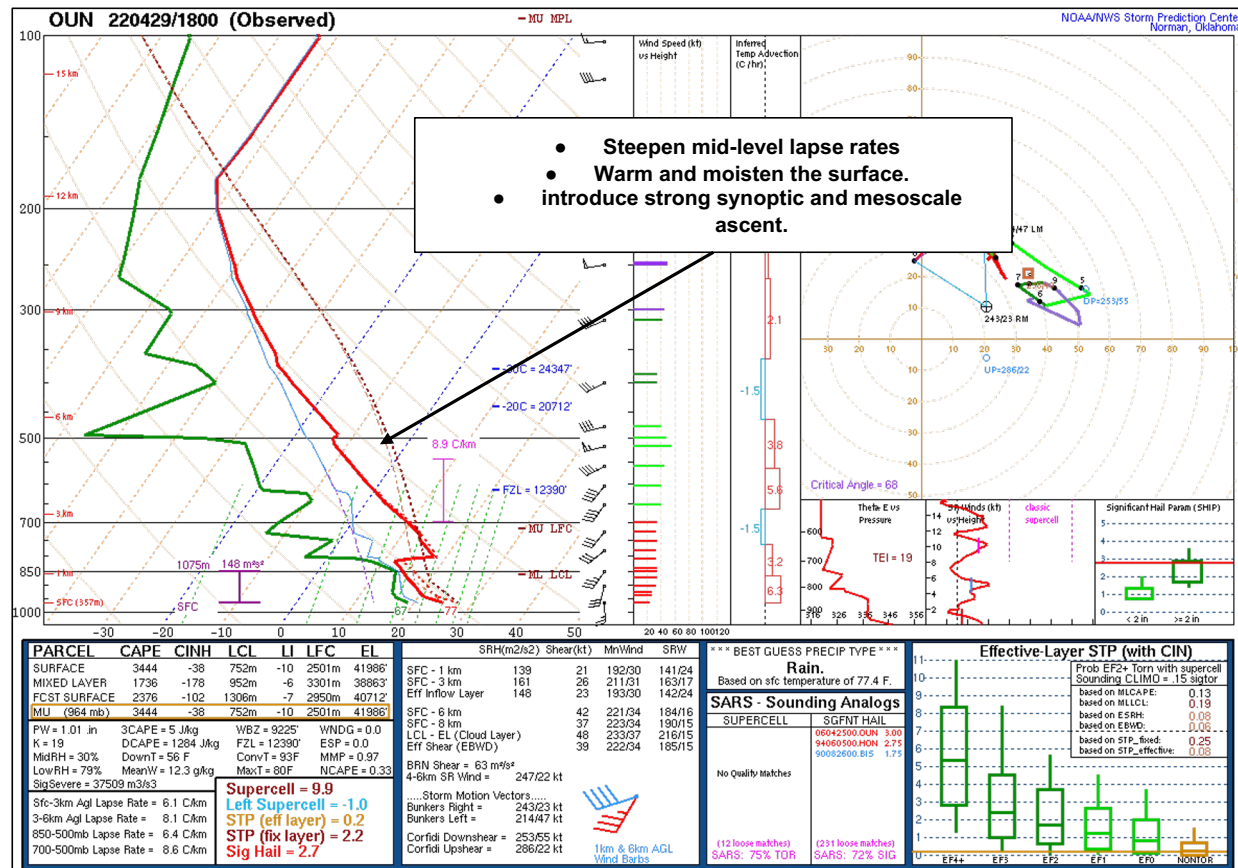
1. Offsets negative effects of dry air entrainment into updraft
2. Mitigates effects of precipitation drag on updraft



**Tall/narrow CAPE:**  
Smaller upward acceleration

Water loading and entrainment more  
substantially weaken updraft

So, how do we get wider CAPE?



# CAPE/CIN Changes Independent of $\gamma$ Tendency

- CIN can be reduced and/or CAPE increased by:

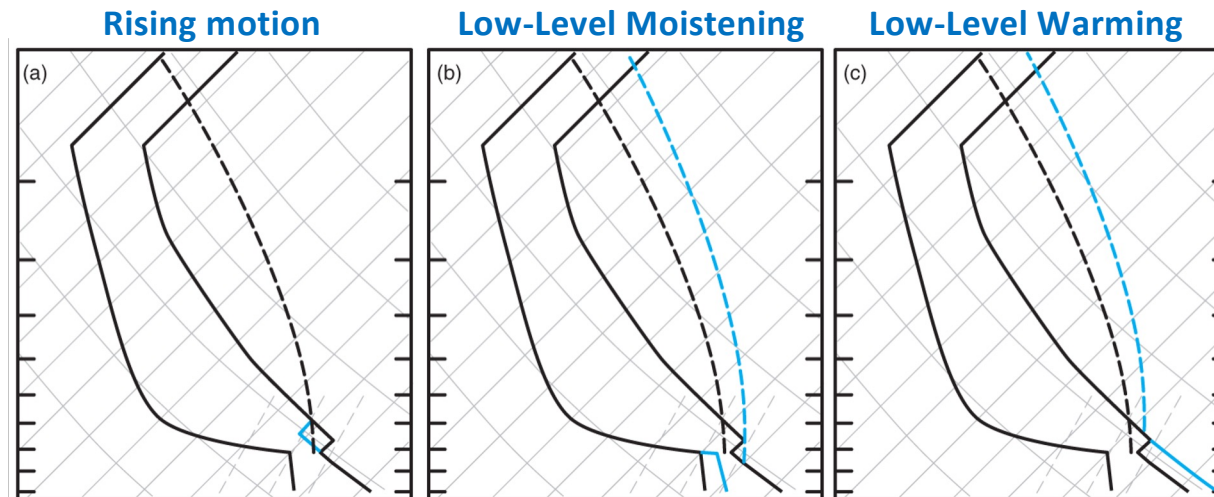


Figure 7.9

CIN can be reduced by (a) large-scale rising motion, (b) low-level moistening (e.g., moisture advection), and (c) low-level warming (e.g., insolation), despite the fact that the CIN modifications may not be accompanied by lapse rate changes, at least not over a significant depth. In (a)–(c), the isotherms and isentropes are solid gray lines, the constant mixing ratio lines are gray dashed lines, the sounding and trajectory taken by an air parcel lifted from the surface are solid and dashed black curves, respectively, and the modified sounding and parcel trajectory are blue solid and dashed curves, respectively. In (a), for clarity, only the temperature profile has been modified (the moisture profile has not been modified in accordance with the vertical motion that has been imposed in the layer of the capping inversion). Note that (b) and (c) are also accompanied by increases in CAPE. Conversely, CIN is augmented by large-scale descent, boundary layer cooling (although this would typically not occur without a concurrent stabilization of the lapse rate), and boundary layer drying (not shown).

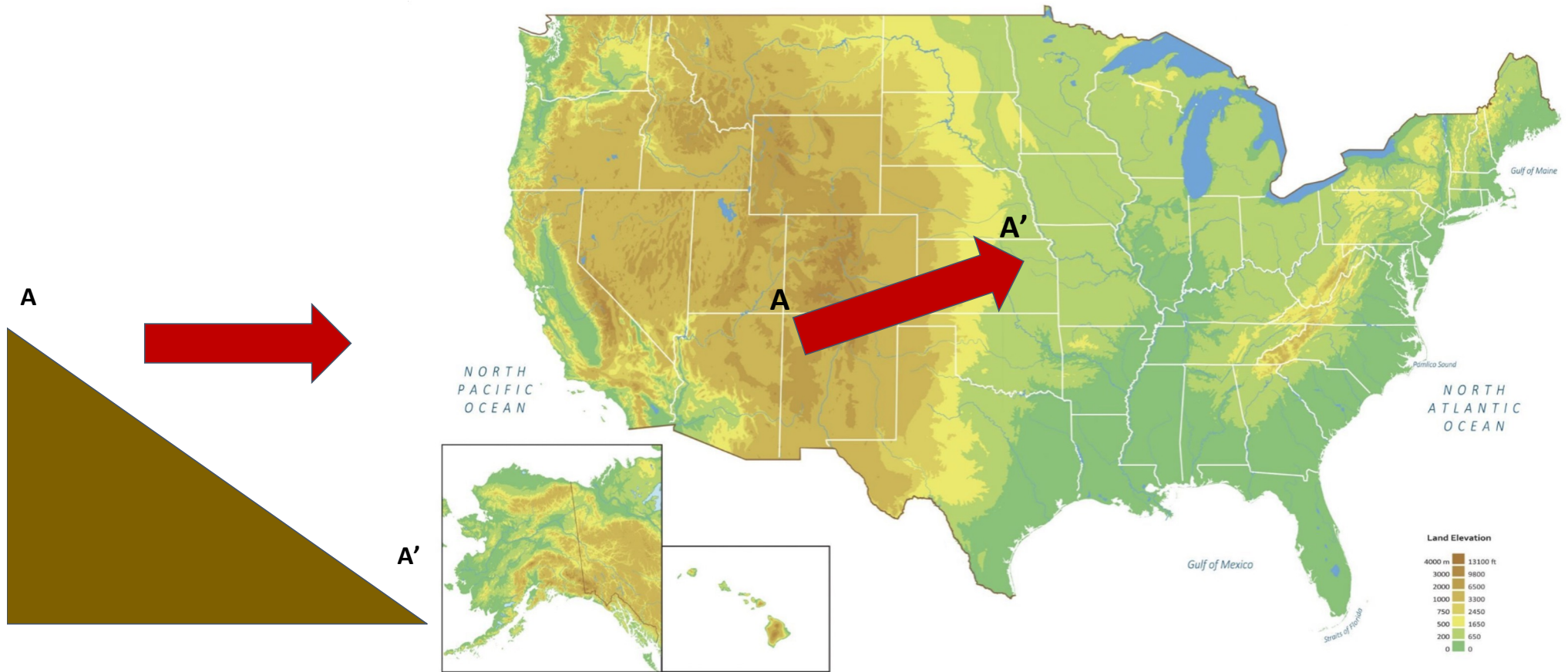
(Markowski and Richardson 2010, Fig. 7.9)

# **Instability: Elevated Mixed Layer (EML)**

- **Deep mixed layer forms over Rockies and Sierra Madre in response to surface heating**
- **Stronger heating and/or cooler temperatures aloft results in deeper mixed layer**
- **Mixed layer is advected eastward and becomes elevated east of the Rockies and Sierra Madre (really, east of the dryline)**
- **Differential advection (eastward advection of EML above northward advection of moisture from GoM) created “loaded gun” profile**

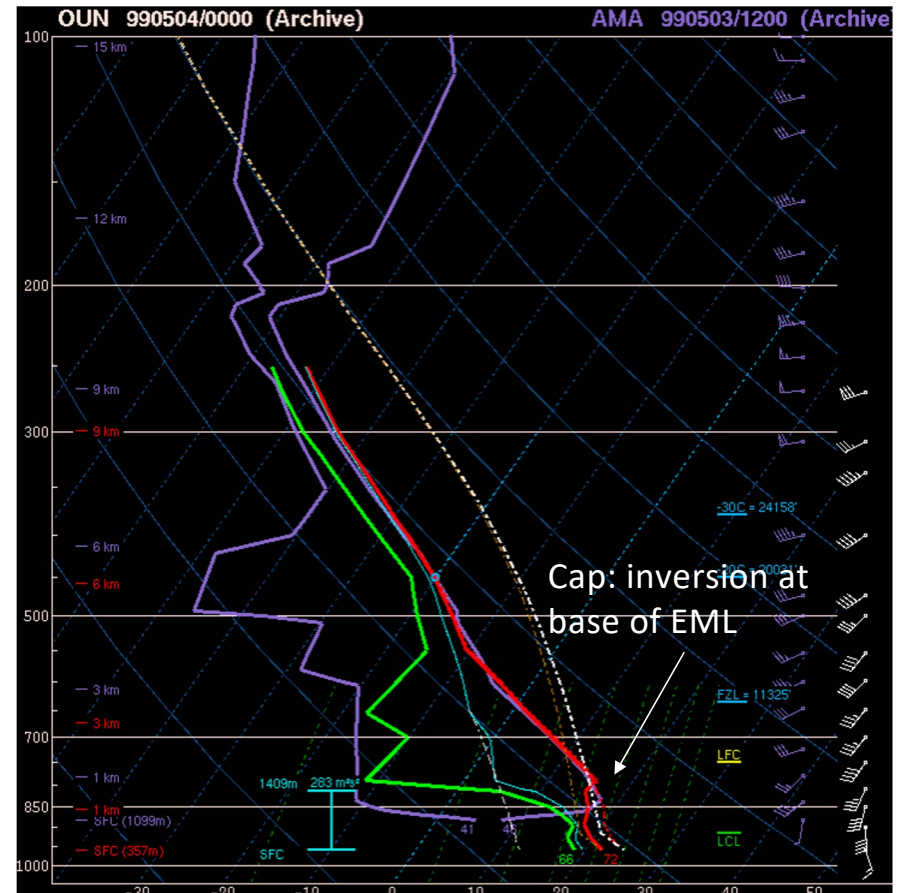
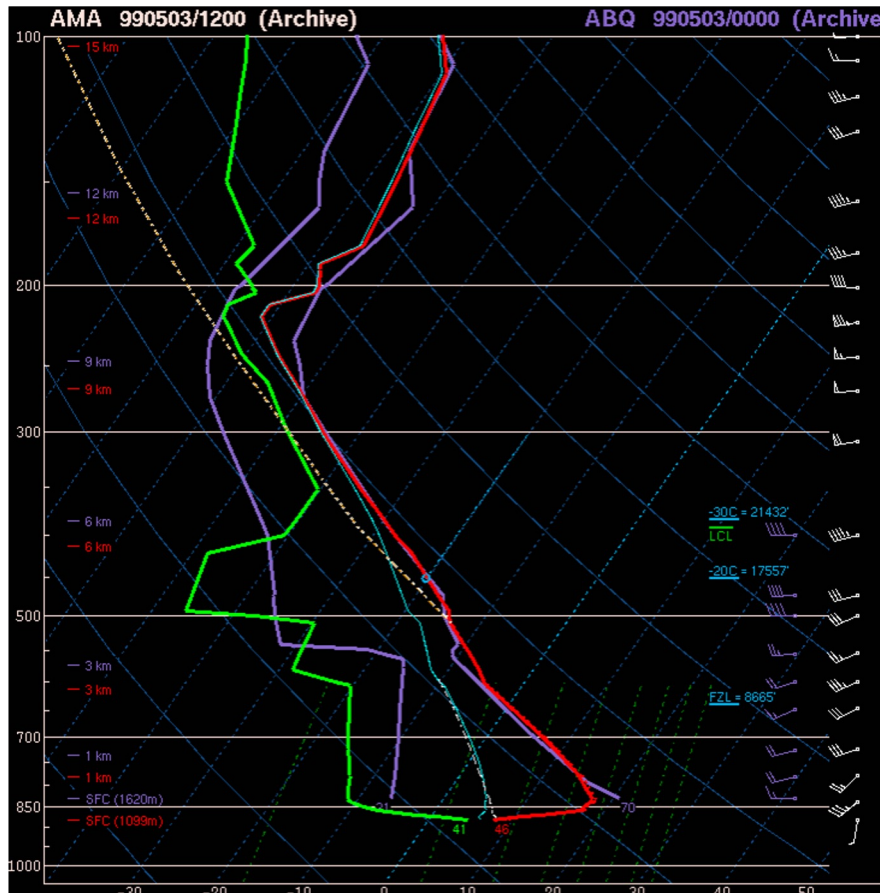


# Elevated Mixed Layer (EML) Differential Advection

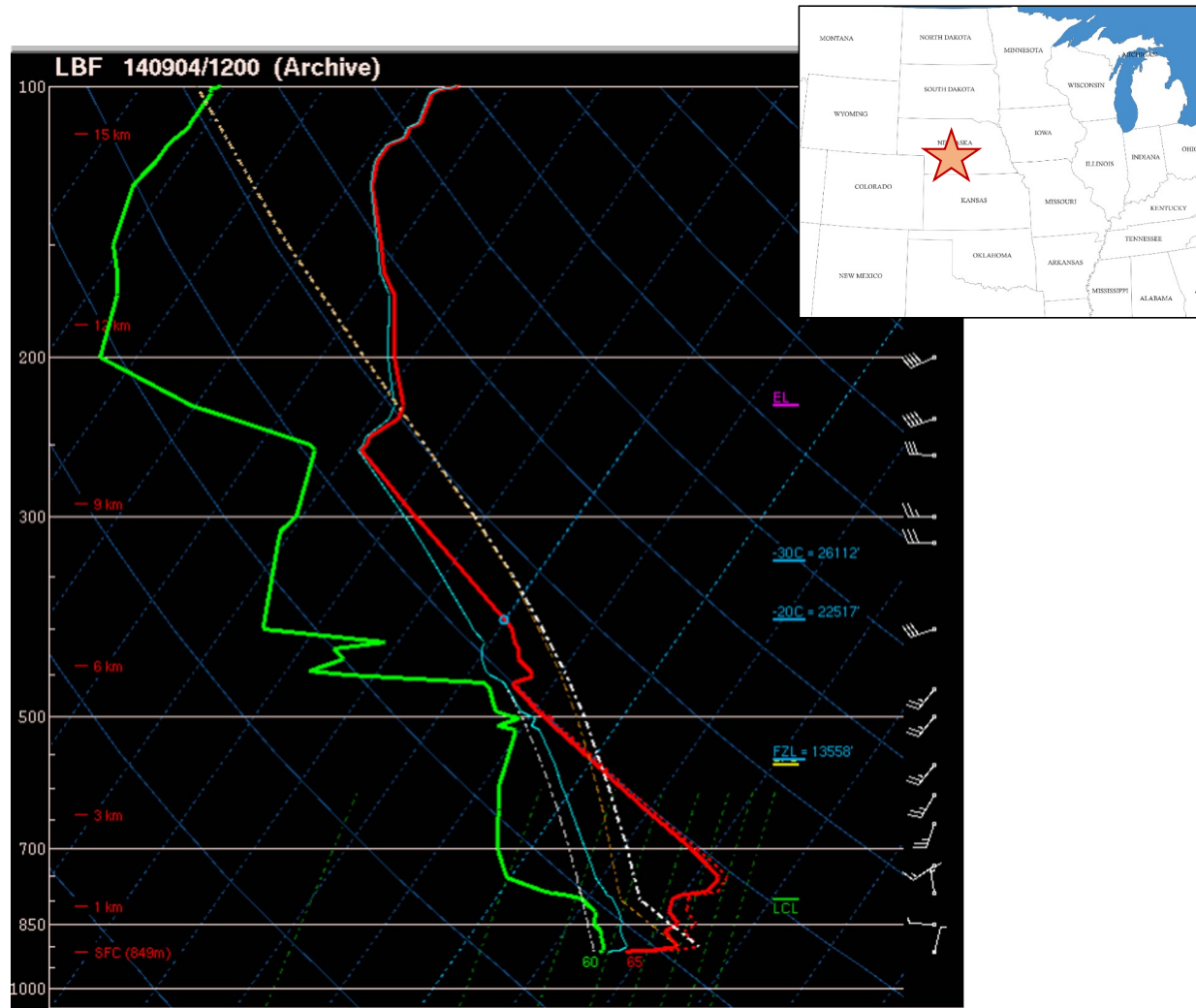




# EML at OUN (0000 UTC 4 May 1999)

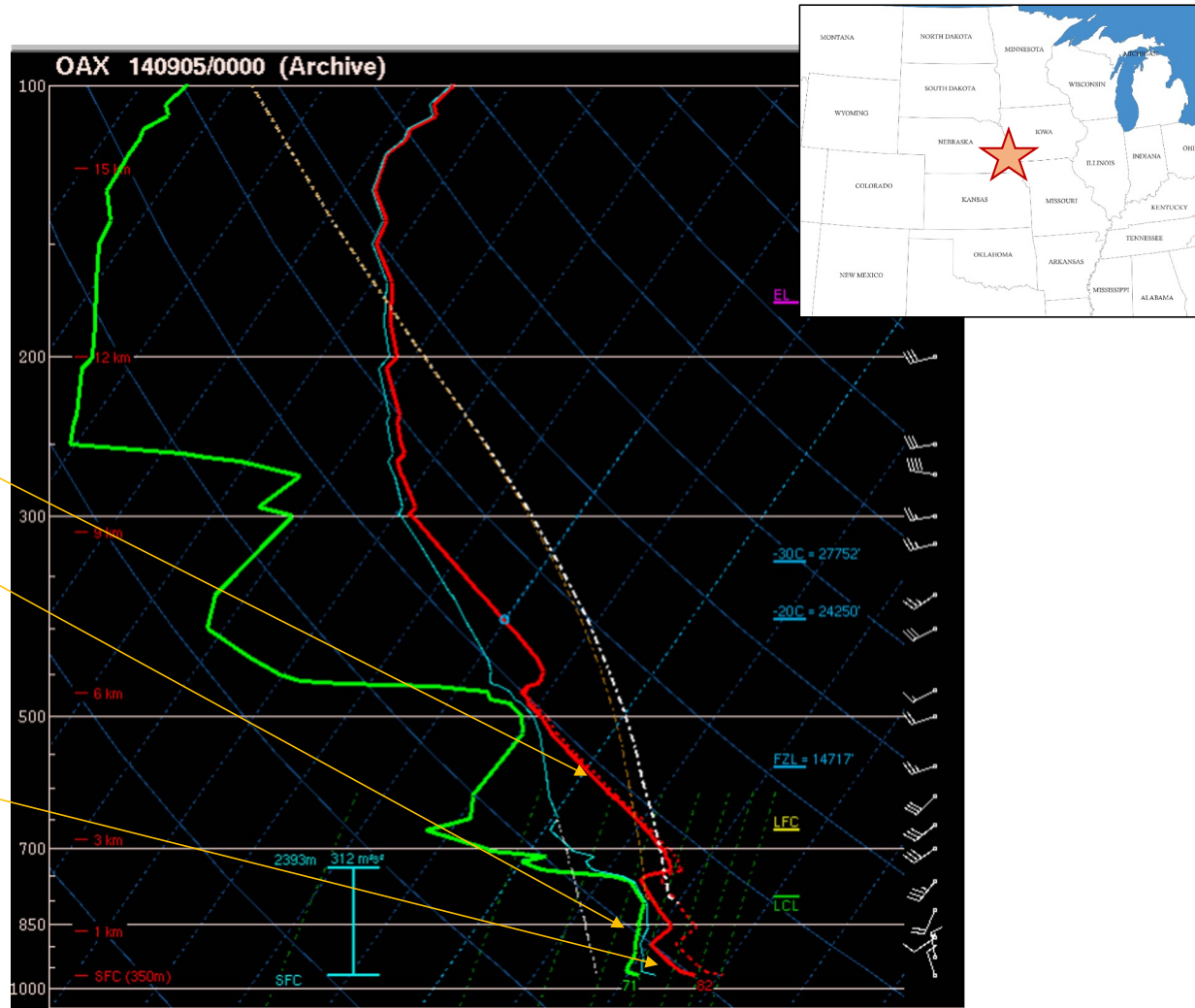






Cap strength is related to:

- EML  $\theta$
- Depth/quality of moisture
- Surface heating and mixing in moist sector



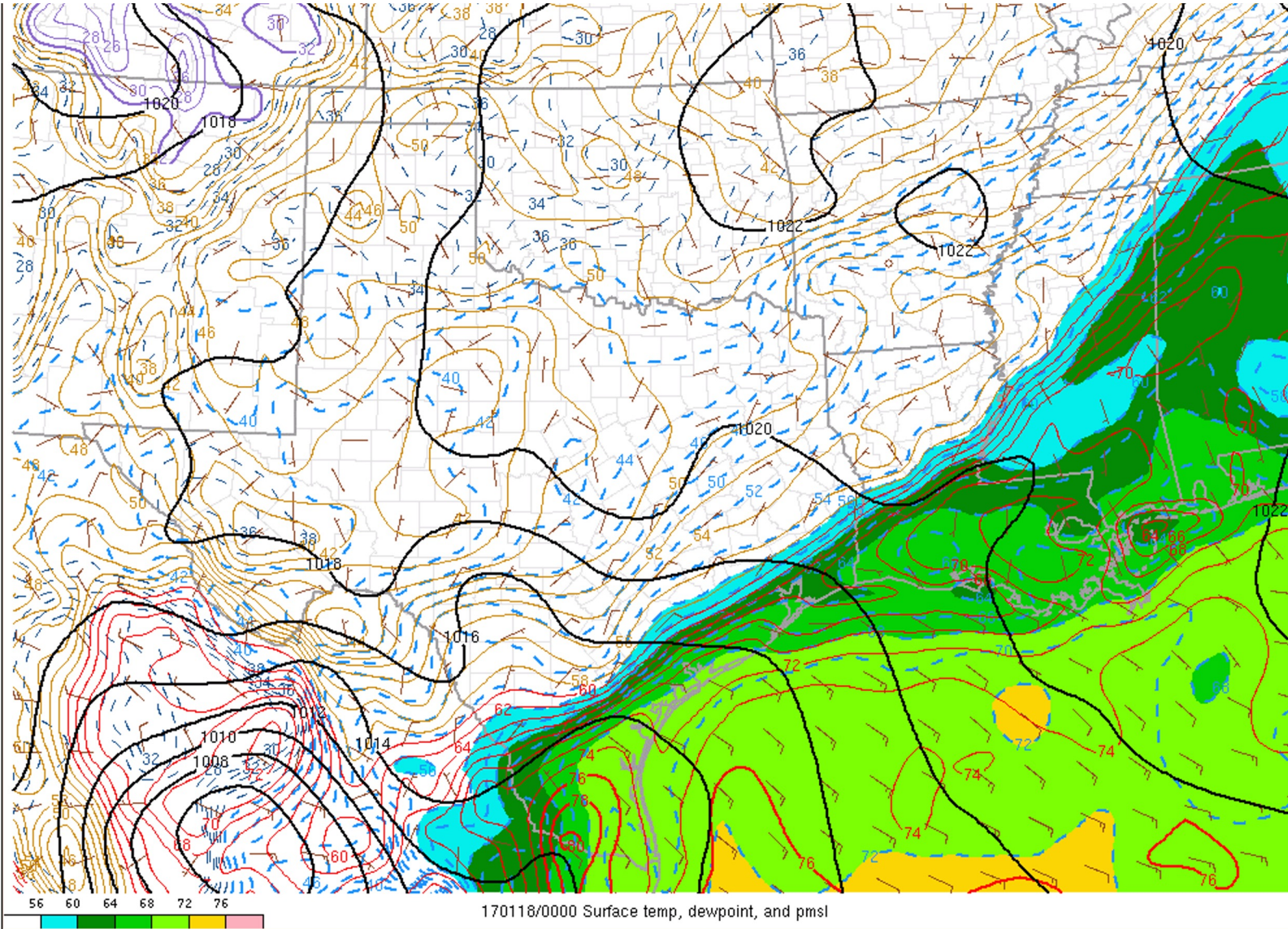
# Lift

- **Synoptic-scale lift**
  - 1-10 cm/s for many hours (6 hours ~2 km ascent (SFC-800 hPa))
  - 6 hours to reach LFC
  - Conditions environment for deep, moist convection over broad area
    - Lapse rate stretching
    - Removal of inhibition
  - Layered clouds in regions of isentropic lift; some elevated storms can form
- **Mesoscale lift**
  - 1 m/s for minutes to hours (1 hour ~3.6 km ascent (SFC-650 hPa))
  - < 1 hour to reach LFC
  - Narrow zones of ascent along boundaries and terrain features
  - Direct initiation of thunderstorms
- **Storm-scale lift**
  - 10 m/s for minutes (15 mins ~9 km of ascent (SFC-300 hPa))
  - 5 minutes to reach LFC
  - Storm maintenance and propagation (supercells, MCSs, squall lines, etc.)

# Synoptic-Scale Lift

- **QG sources**
  - Warm advection/isentropic lift and differential vorticity advection
- **Jet streaks (also QG)**
  - Straight jet: ascent in right entrance and left exit regions
  - Curved jet: ascent downstream and along jet core





- SfcOA Diag - - RAP

**Image overlays:**

- ☒ County Boundaries
- ☐ County Warning Area
- ☐ Hiways & Cities
- ☐ ARTCC Regions
- ☐ NWS Watches & Warnings
- ☐ SPC Day1 Outlook

**Image underlays:**

- ☒ None
- ☐ Radar
- ☐ Terrain
- ☐ Population
- ☐ Surface Obs

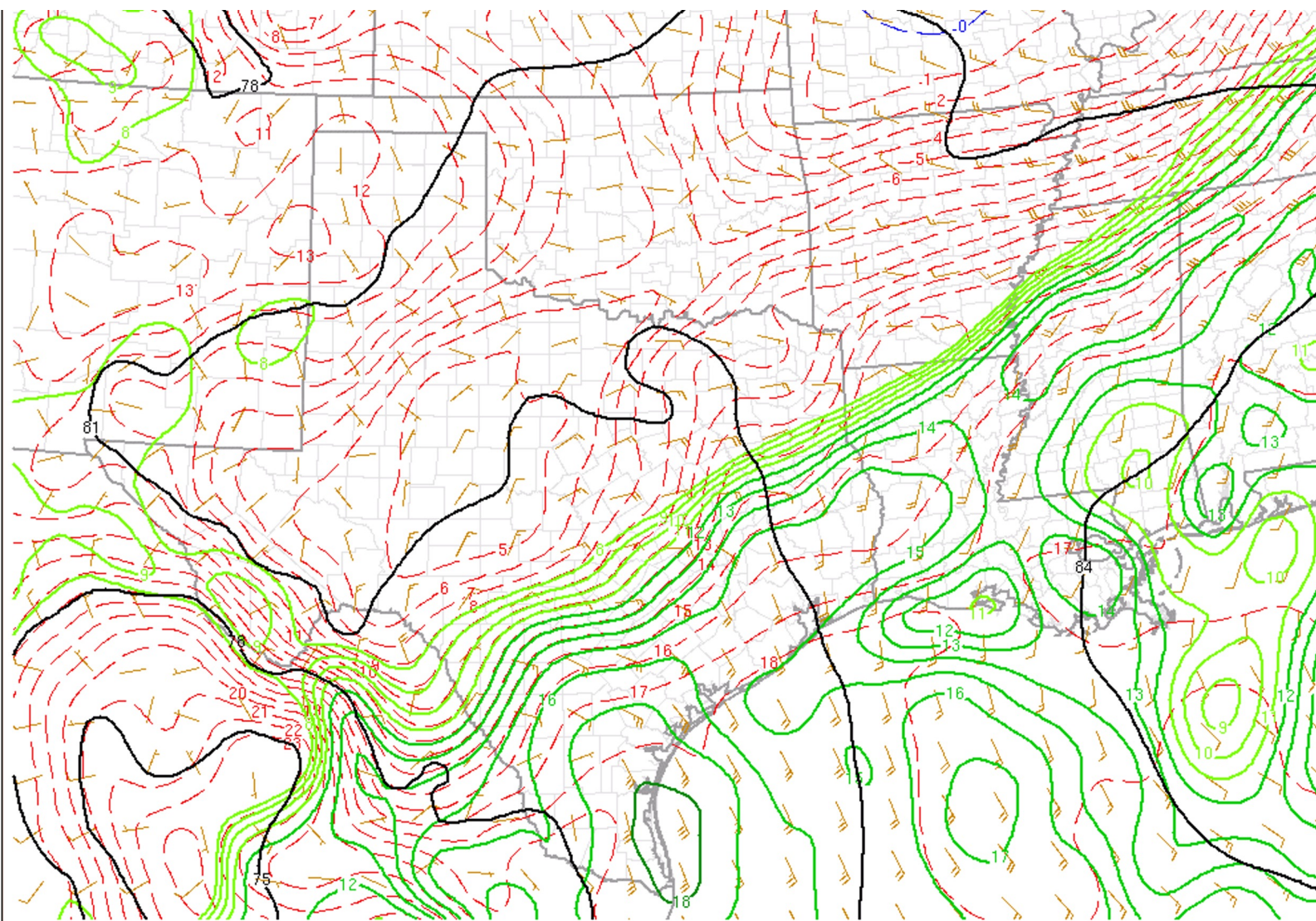
**Current SPC Product**

Show popup images? ☒

**Day1 Convective Outlook**  
Issued at 1957 UTC  
Probabilities: **H**

**Day1 National Fire Outlook**  
Issued at 1819 UTC  
This list updates automatically





170118/0000V001 925mb hght/temp/dwpt/wind

- SfcOA Diag - - RAP

#### Image overlays:

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#### Current SPC Product

Show popup images? ☒

#### Day1 Convective Outlook

Issued at 1957 UTC

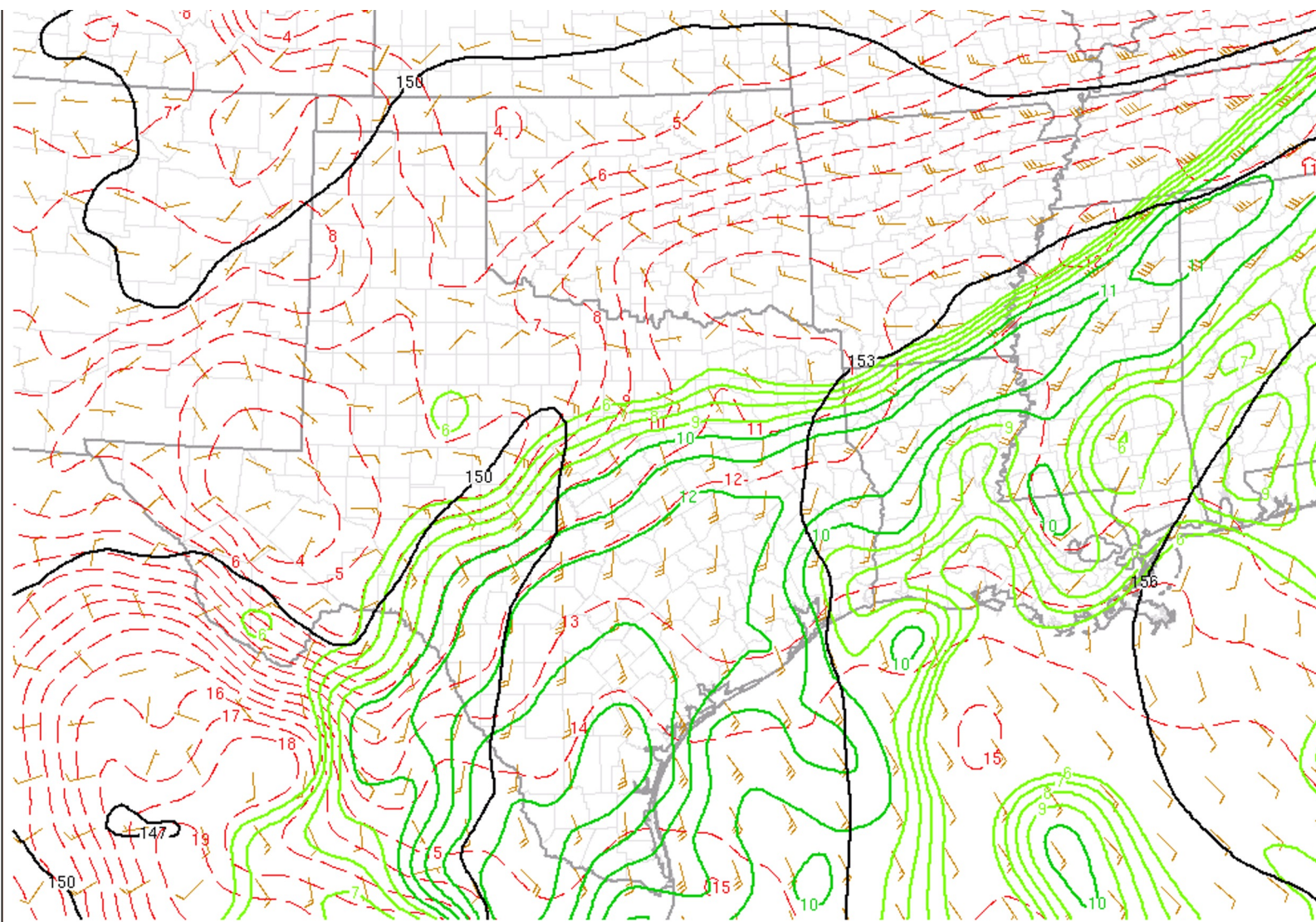
Probabilities: **H**

#### Day1 National Fire Outlook

Issued at 1819 UTC

This list updates automatically





170118/0000V001 850mb hght/temp/dwpt/wind

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#### Current SPC Product

Show popup images? ☒

#### Day1 Convective Outlook

Issued at 1957 UTC

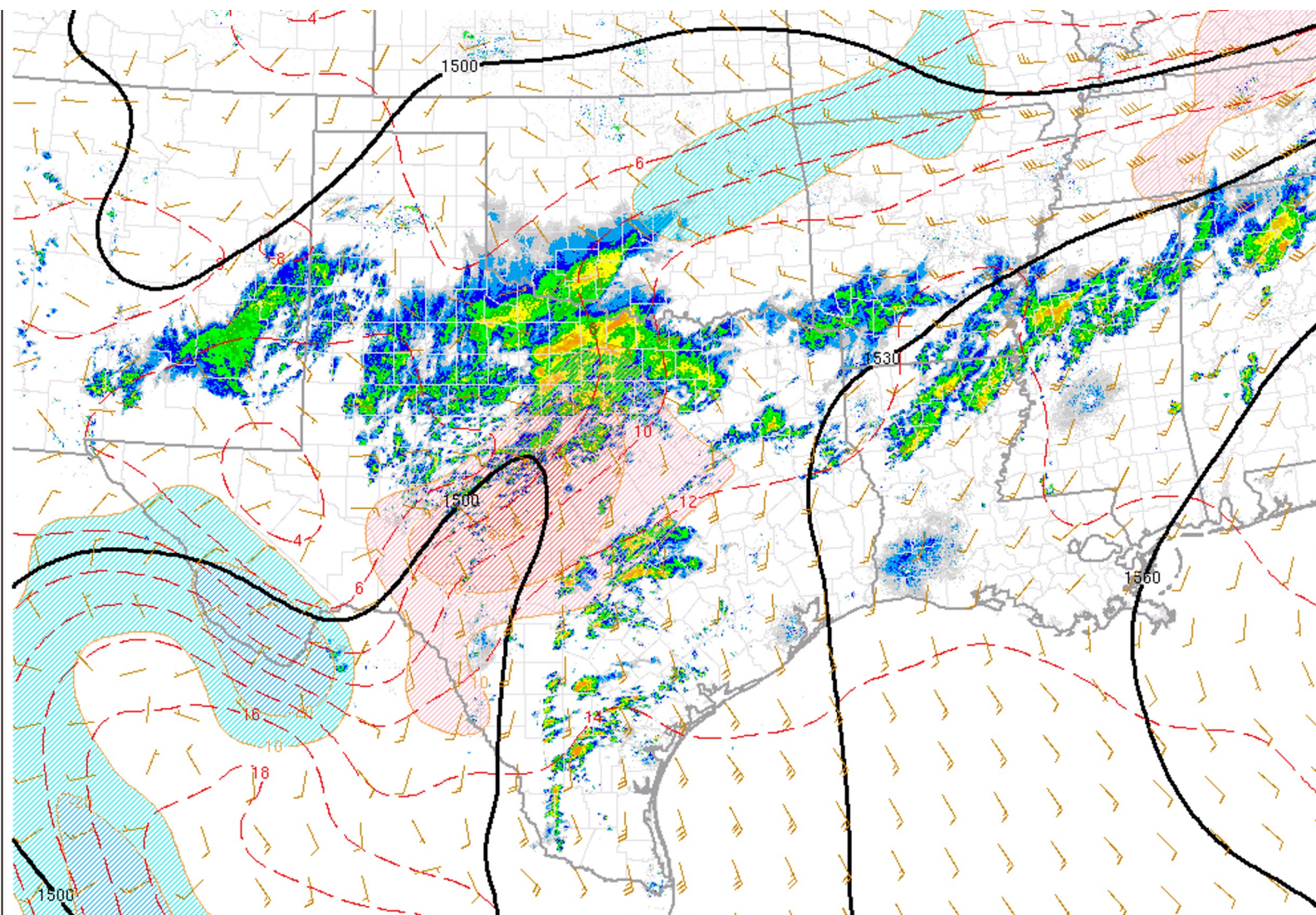
Probabilities: **H**

#### Day1 National Fire Outlook

Issued at 1819 UTC

This list updates automatically





20170118/0045 RADAR  
170118/0000V001 850 mb temperature advection (fill - 10e-5 C/s)

- SfcOA Diag - - RAP

**Image overlays:**

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- ☐ ARTCC Regions
- ☐ NWS Watches & Warr
- ☐ SPC Day1 Outlook

**Image underlays:**

- ☐ None
- ☒ Radar
- ☐ Terrain
- ☐ Population
- ☐ Surface Obs

**Current SPC Prod**

Show popup images? ☒

**Day1 Convective Outl**

Issued at 1957 UTC

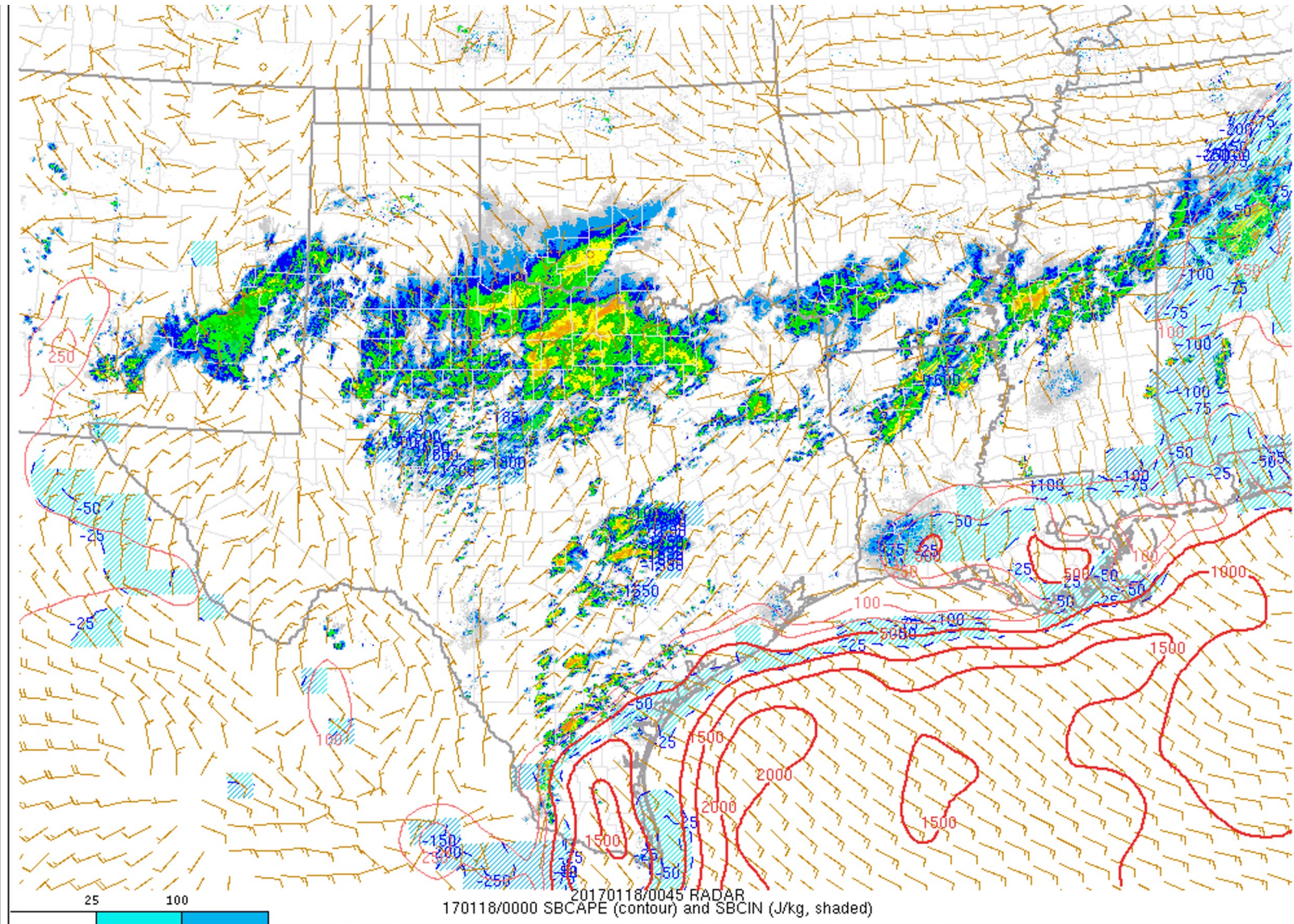
Probabilities: **T**

**Day1 National Fire Ou**

Issued at 1819 UTC

This list updates automatically





- SfcOA Diag - - RAP

#### Image overlays:

- ☒ County Boundaries
- ☐ County Warning Area
- ☐ Hiways & Cities
- ☐ ARTCC Regions
- ☐ NWS Watches & Warnings
- ☐ SPC Day1 Outlook

#### Image underlays:

- ☐ None
- ☐ Radar
- ☐ Terrain
- ☐ Population
- ☐ Surface Obs

#### Current SPC Product

Show popup images? ☒

#### Day1 Convective Outlook

Issued at 1957 UTC

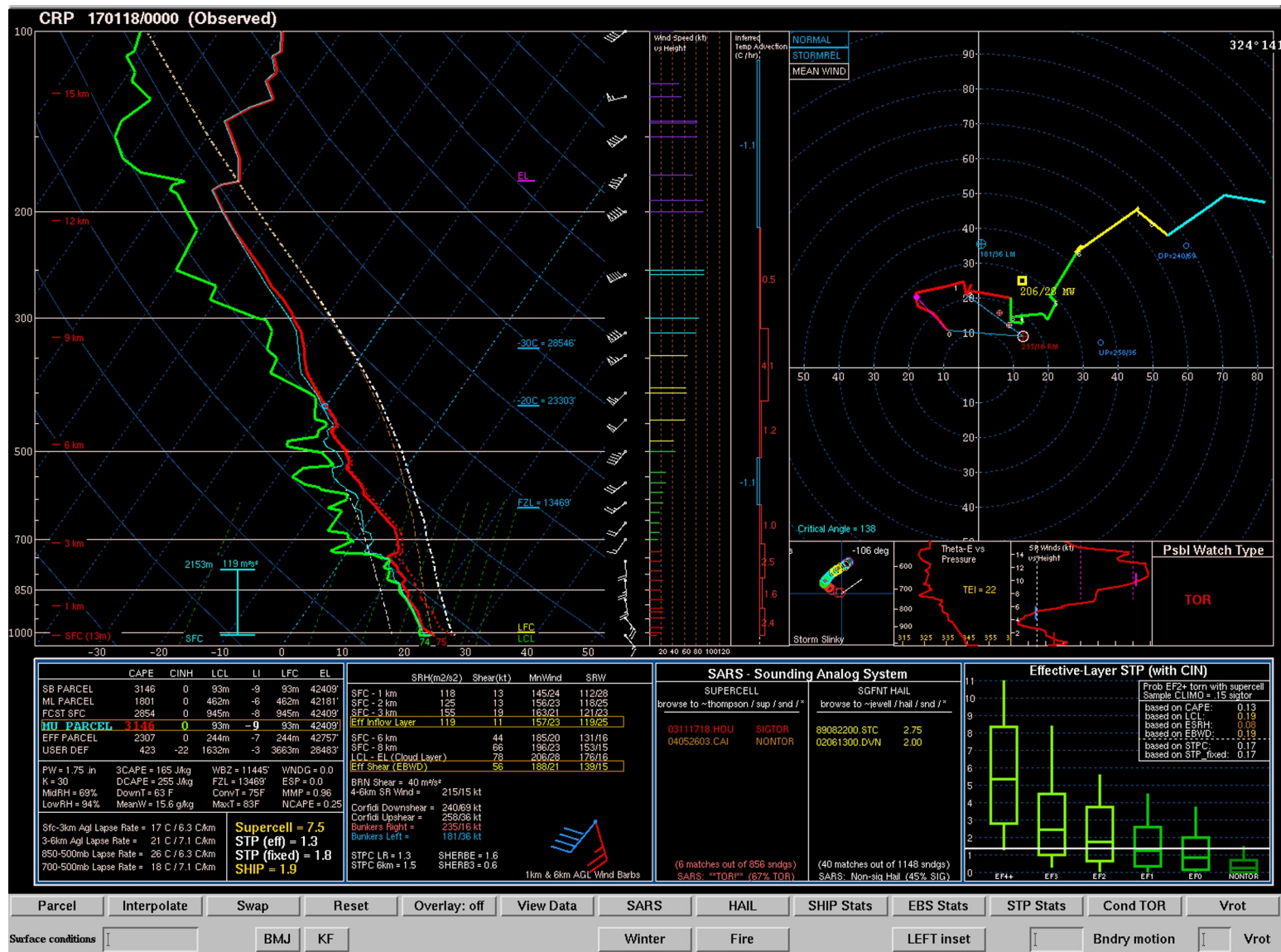
Probabilities: **H**

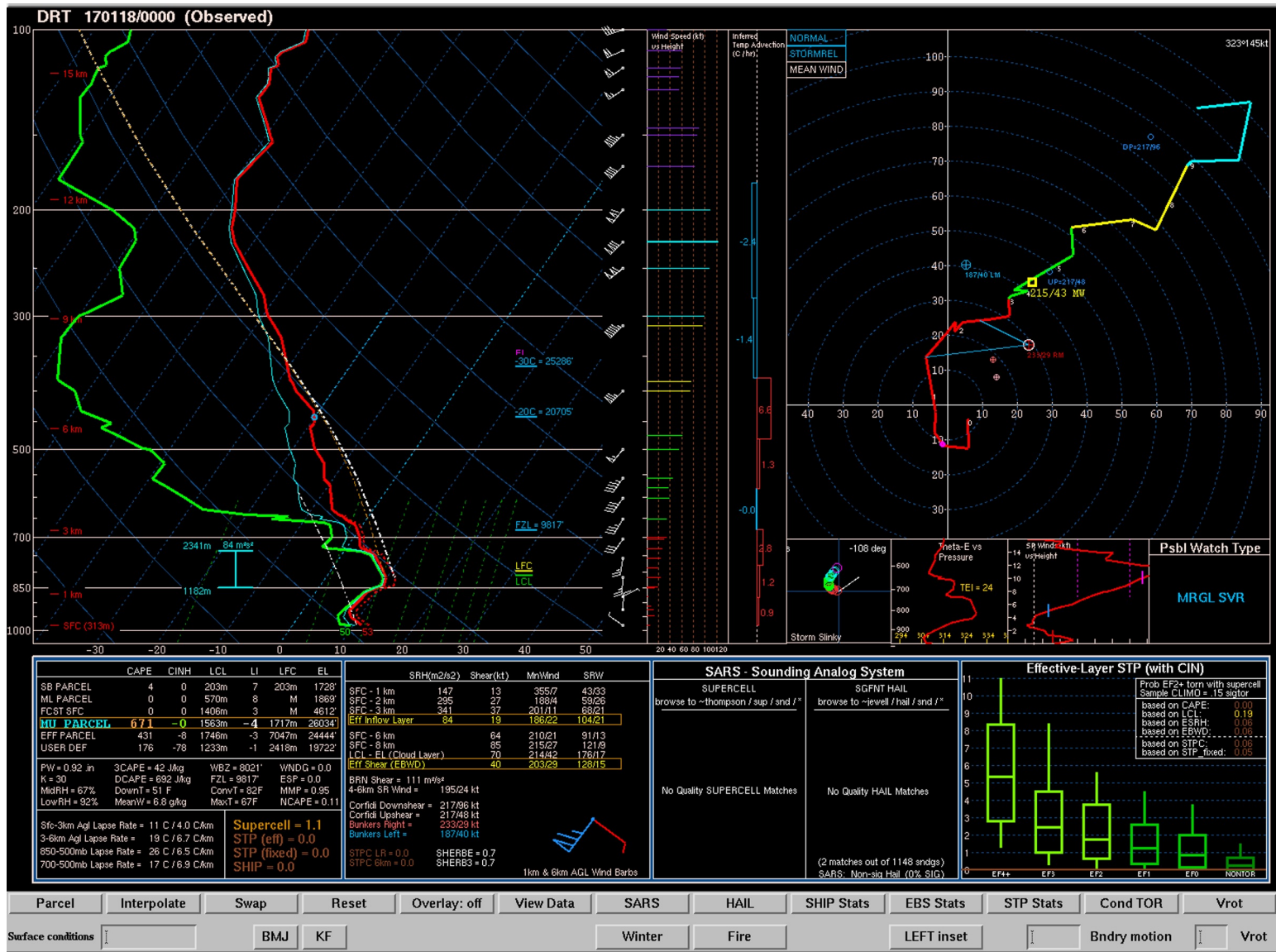
#### Day1 National Fire Outlook

Issued at 1819 UTC

This list updates automatically







# Synoptic-Scale Lift Summary

- **Gradual ascent over many hours**
  - Warm advection and differential vorticity advection are primary drivers
- **Primary role of QG ascent is to precondition environment for convection**
  - Can also initiate elevated storms in regions of warm advection
  - Most surface-based storms are triggered by mesoscale ascent

# Mesoscale Lift

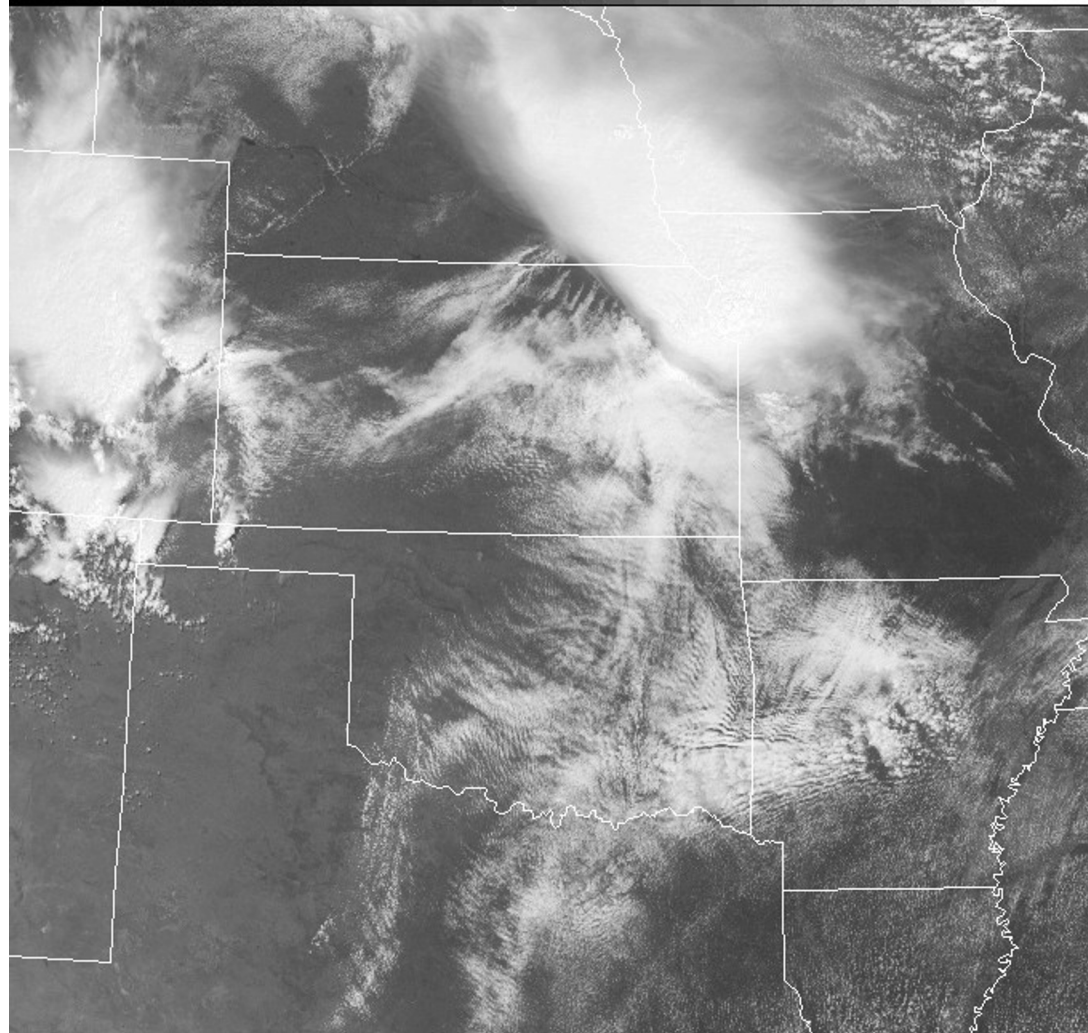
- Focused and narrow zones of lift for minutes to hours
- Fronts
  - Rising on warm side with frontogenesis
  - Isentropic ascent on sloped warm front
- Outflow boundary
  - Lift is governed by depth/strength of cold pool and low-level vertical wind shear [Rotunno-Klemp-Weisman (RKW) Theory]
- Dryline
  - Lift on dry/hot side (lower density air)
  - Lift governed by depth of mixing west of dryline and depth of moist layer east of dryline

1740 UTC Tue 24 May 2011

Visible Satellite

<http://adds.aviationweather.gov>

0 6 9 12 15 18 21 24 27 30 33 36 39 42 45 48 51 54 57 60 63 66 69 72 75 78 81 84 87 90 93 96 99



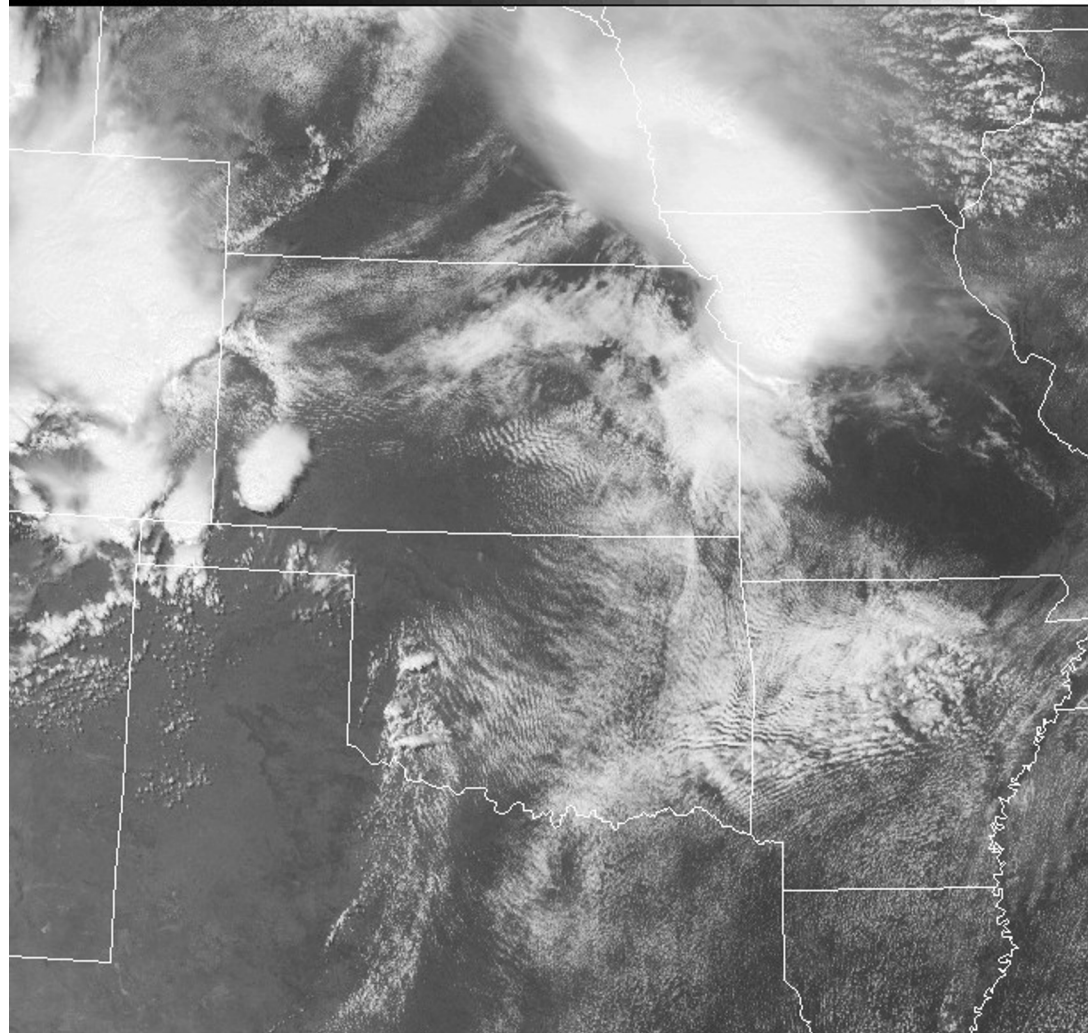


1845 UTC Tue 24 May 2011

Visible Satellite

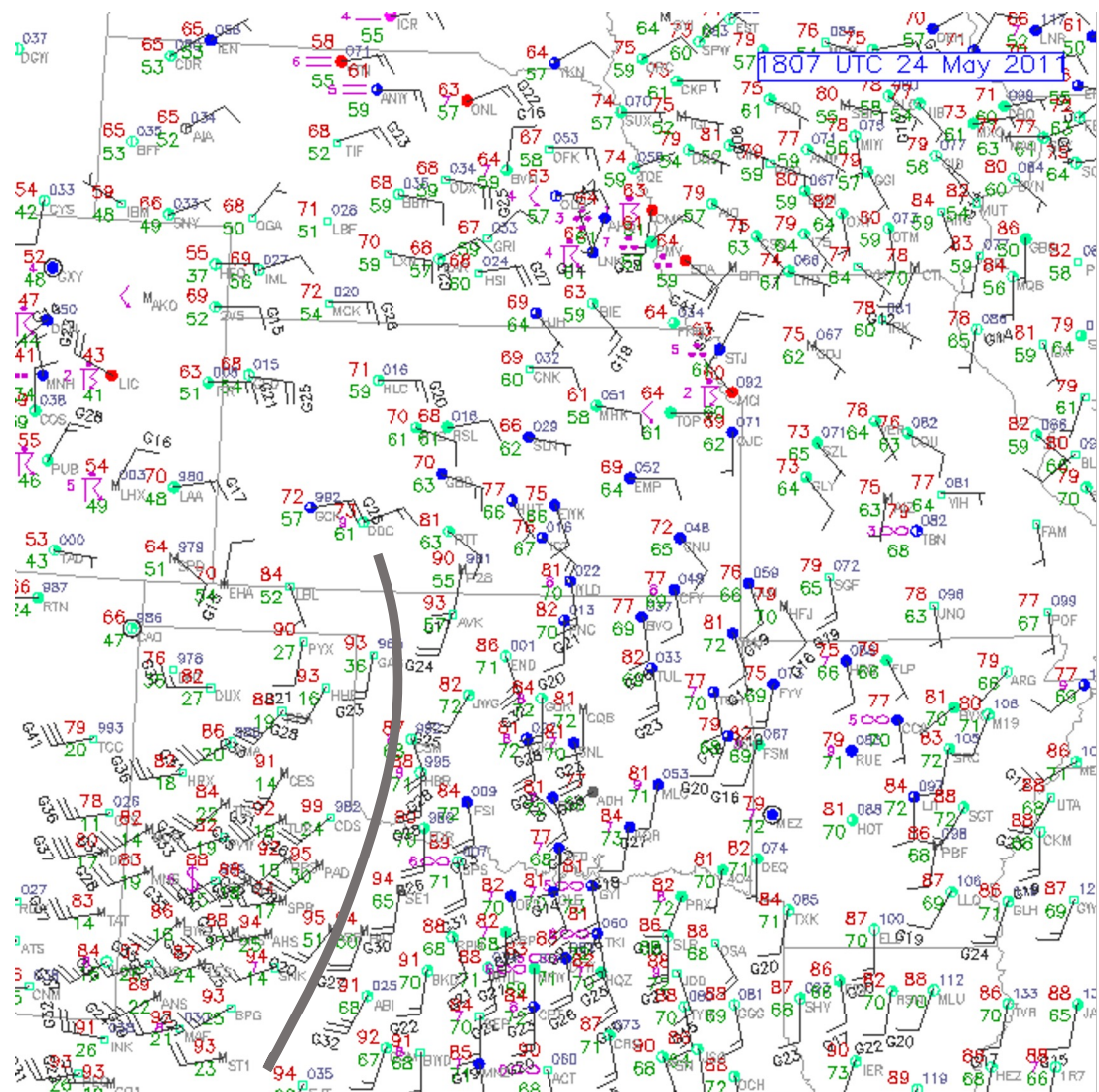
<http://adds.aviationweather.gov>

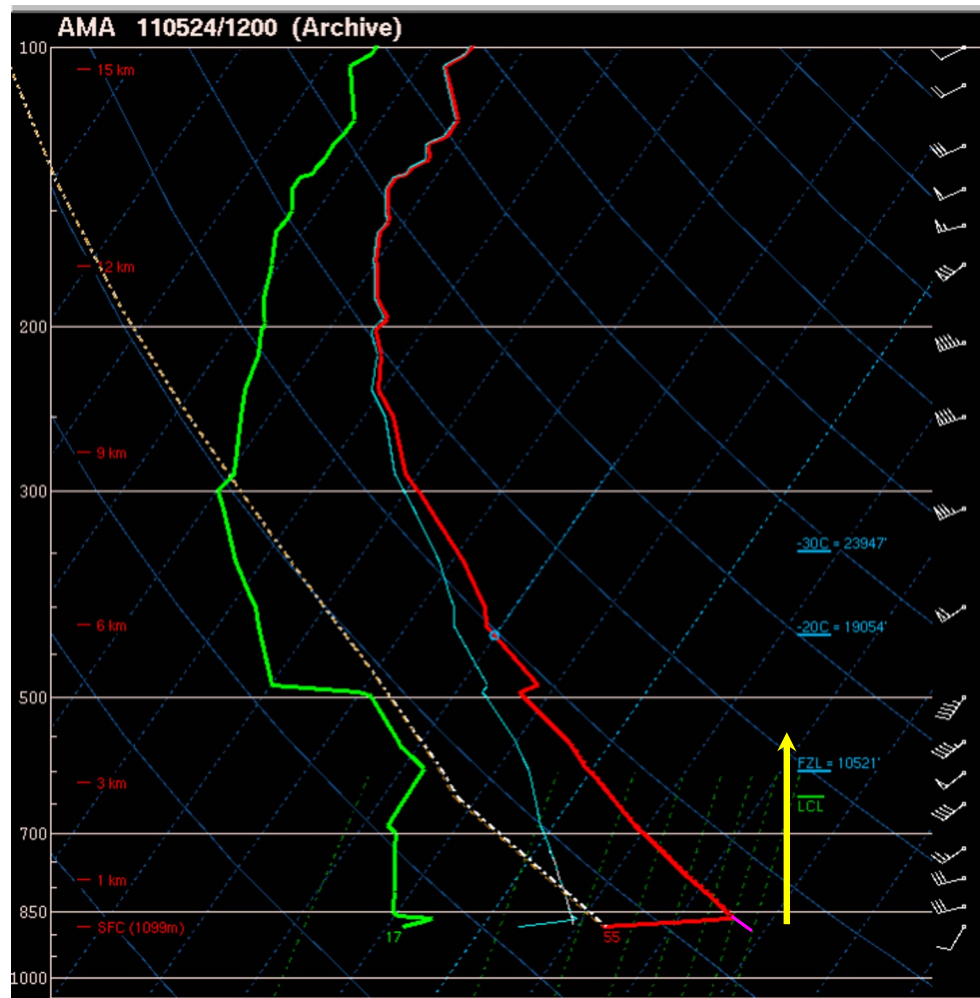
0 6 9 12 15 18 21 24 27 30 33 36 39 42 45 48 51 54 57 60 63 66 69 72 75 78 81 84 87 90 93 96 99



# Dryline

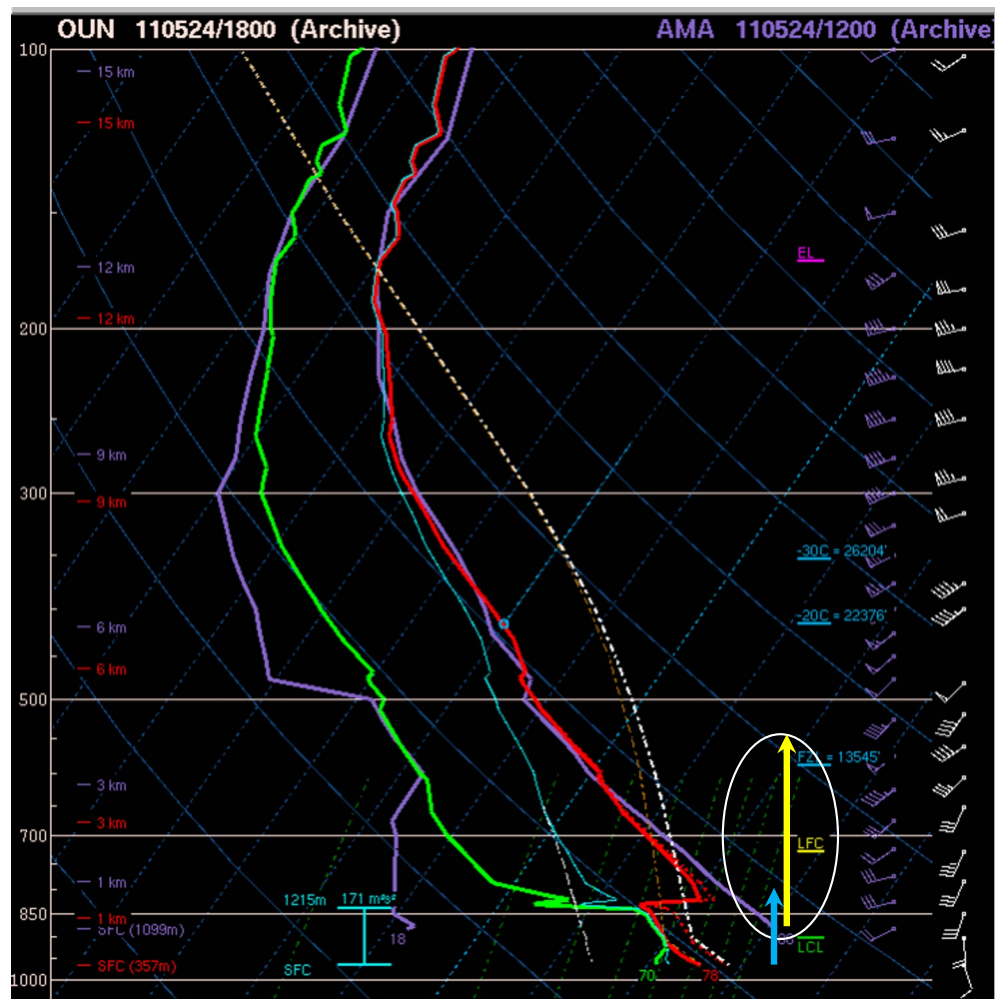
- Focused ascent along dryline can be  $\sim 1$  m/s ( $\sim 1$  hour to reach LFC)
  - Lift depends on depth of mixing west of dryline and depth of moist layer ahead of dryline
  - Maximum vertical motion along dryline scales with height of moist layer
- Convection initiation linked to the residence time of air in the zone of ascent
  - Winds above the surface parallel to dryline keeps air in zone of lift longer
- Convection will initiate in points or bands
  - Usually the mode of initiation for surface-based supercells in central/southern Plains

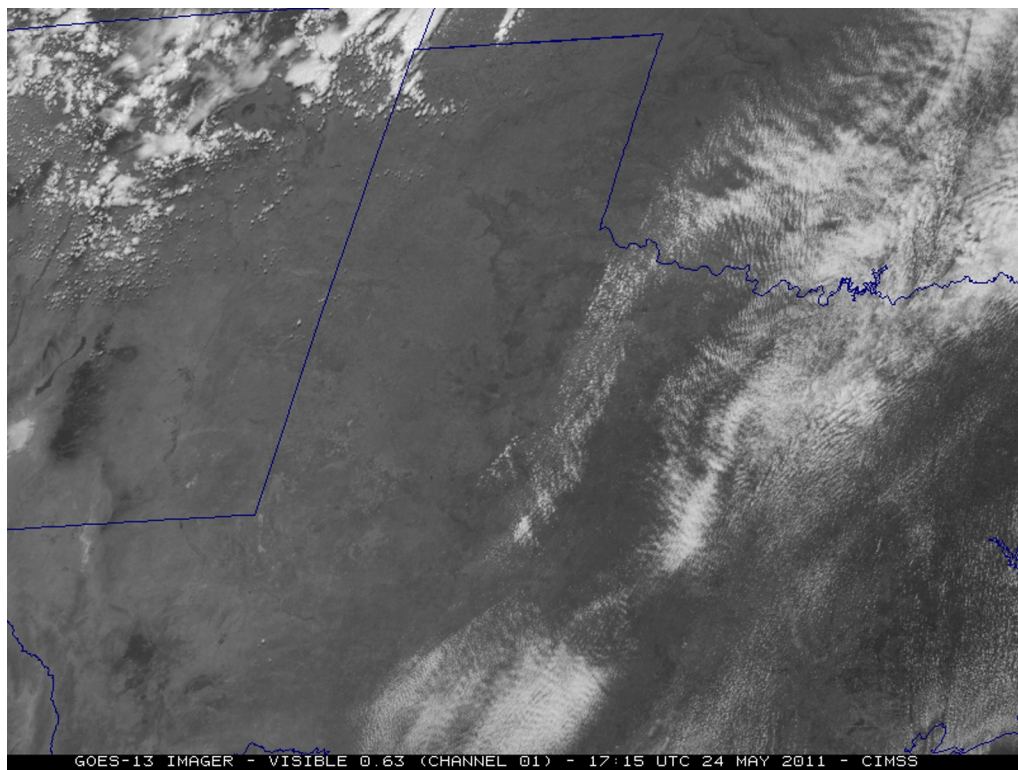












GOES-13 IMAGER - VISIBLE 0.63 (CHANNEL 01) - 17:15 UTC 24 MAY 2011 - CIMSS

# Moisture

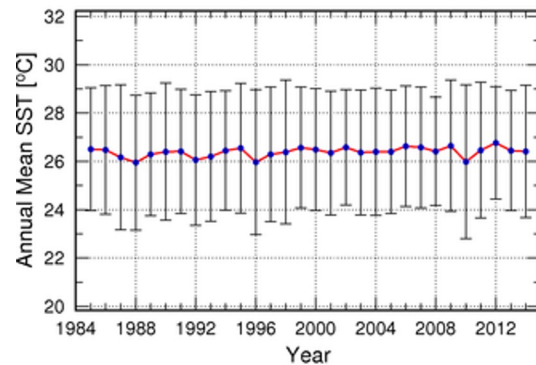
- Moisture in PBL needs to be large and deep enough for convection
- Return flow
- Inland vertical mixing/diurnal processes
- Evapotranspiration



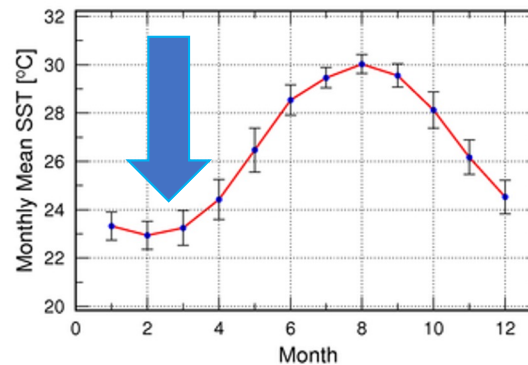
# Low-level Moisture Return Flow

- Key elements include air mass modifications over warmer water and return flow trajectories
- Air mass modification is a multi-date process
- Ask yourself the following:
  - Where is the air coming from?
  - What are the underlying ocean characteristics?
  - What is the character of the returning moist layer?

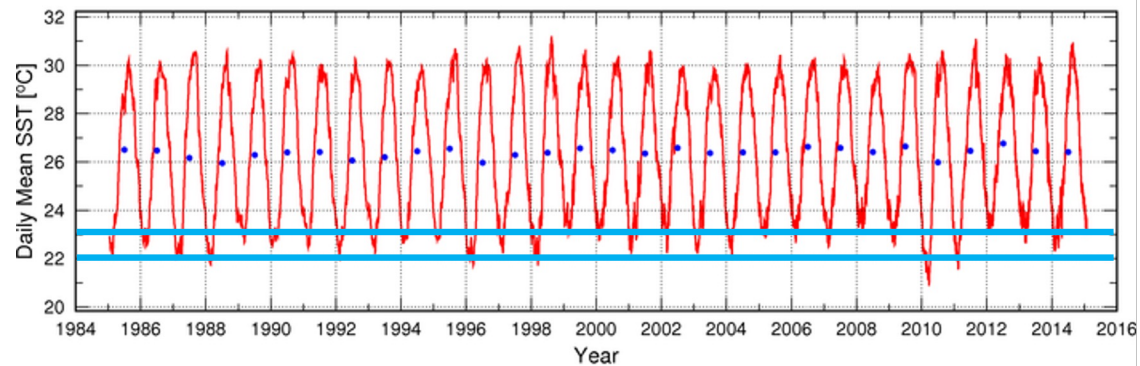
GOM Annual Mean SST



GOM Monthly Mean SST

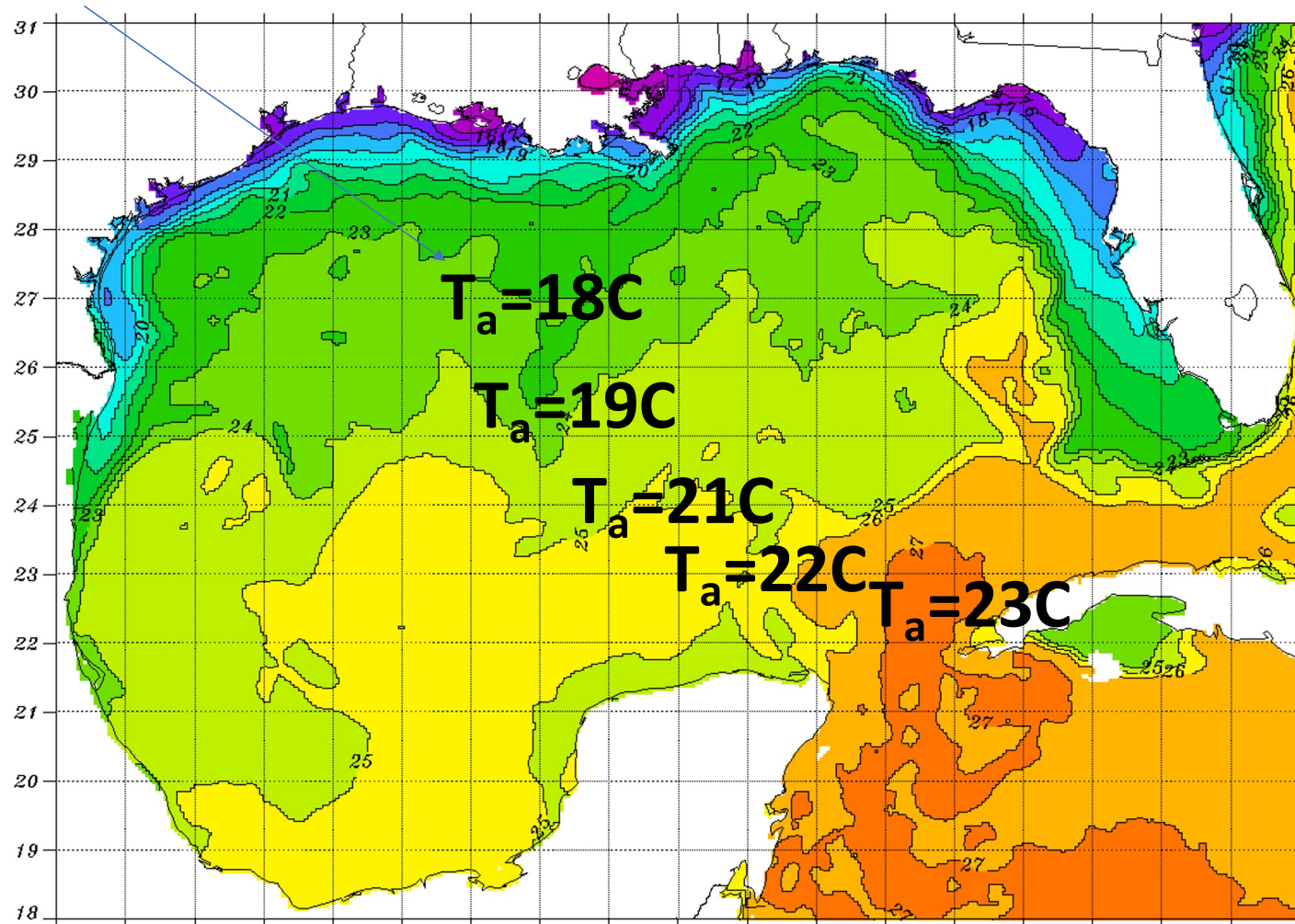


GOM Daily Mean SST

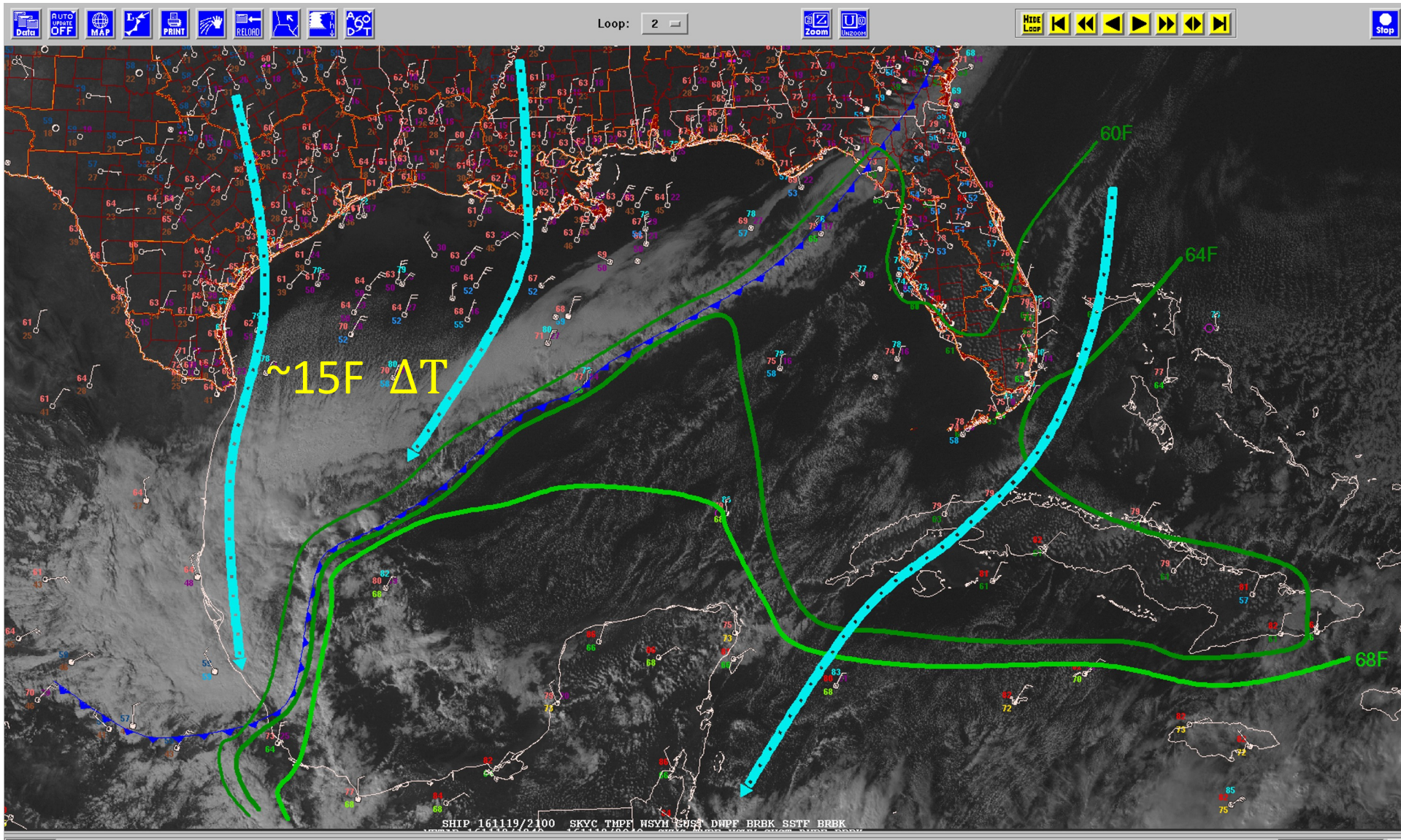


NOAA/NESDIS GEO-POLAR BLENDED 5 km SST ANALYSIS  
FOR THE GULF OF MEXICO

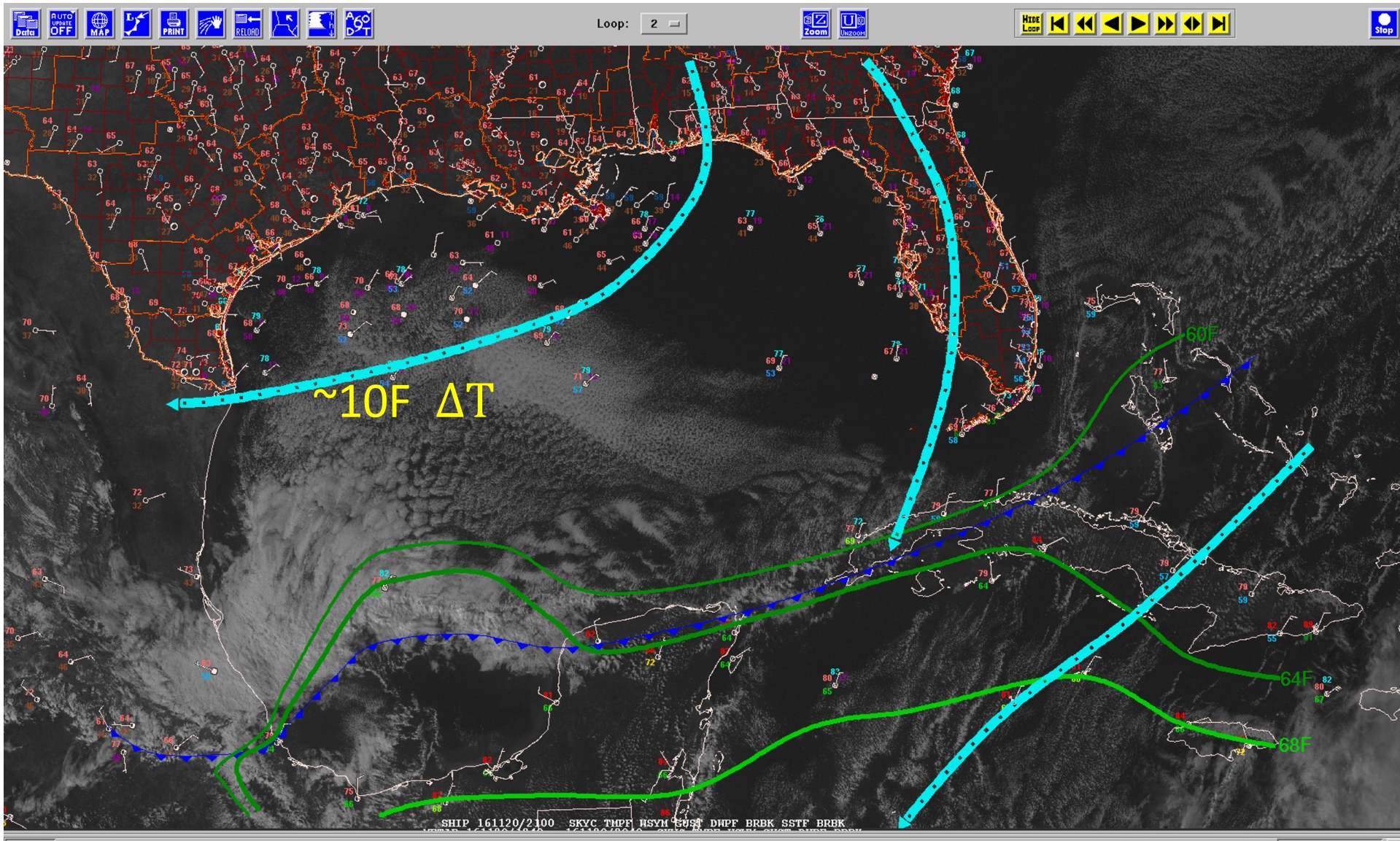
Air temperature after cold front passage



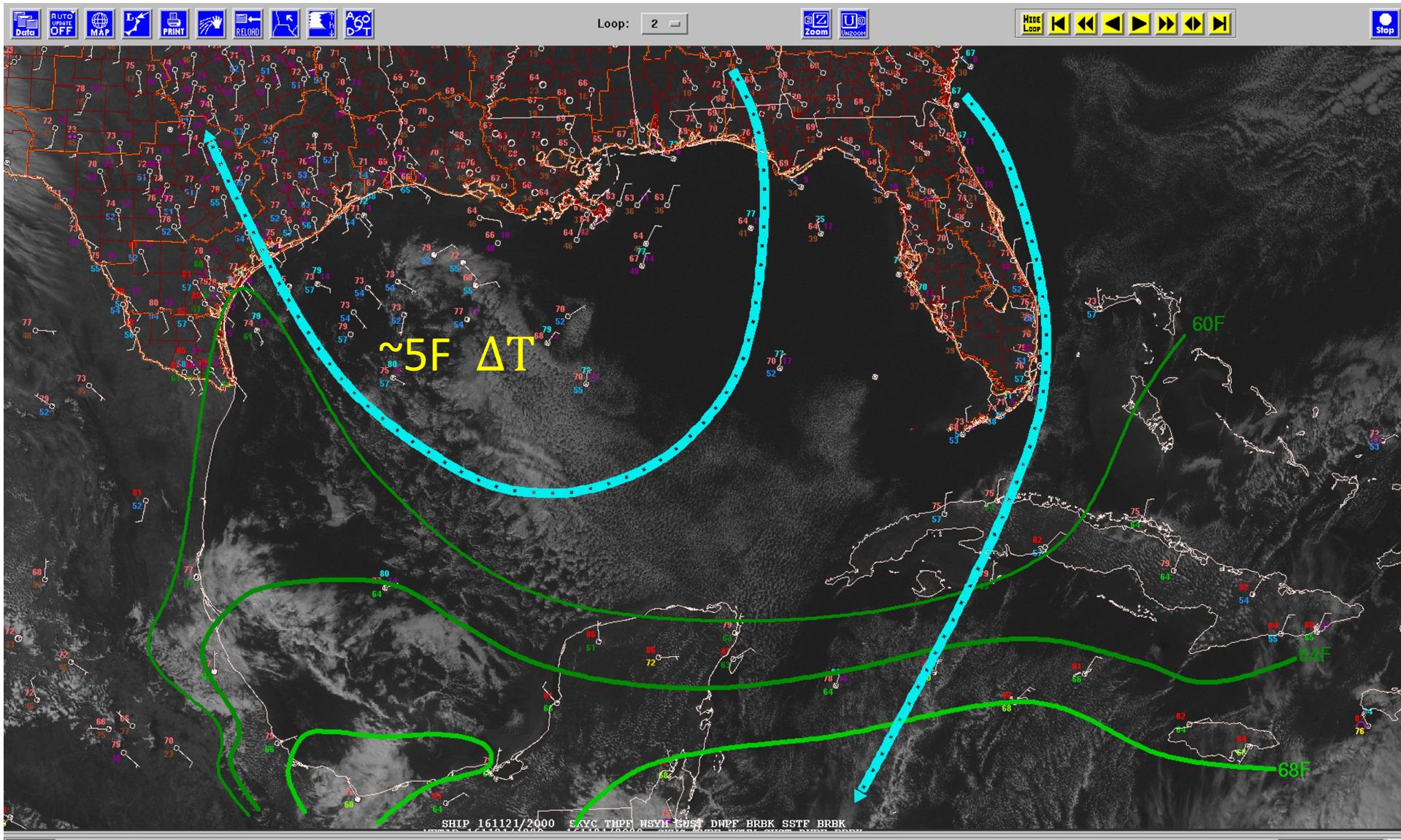


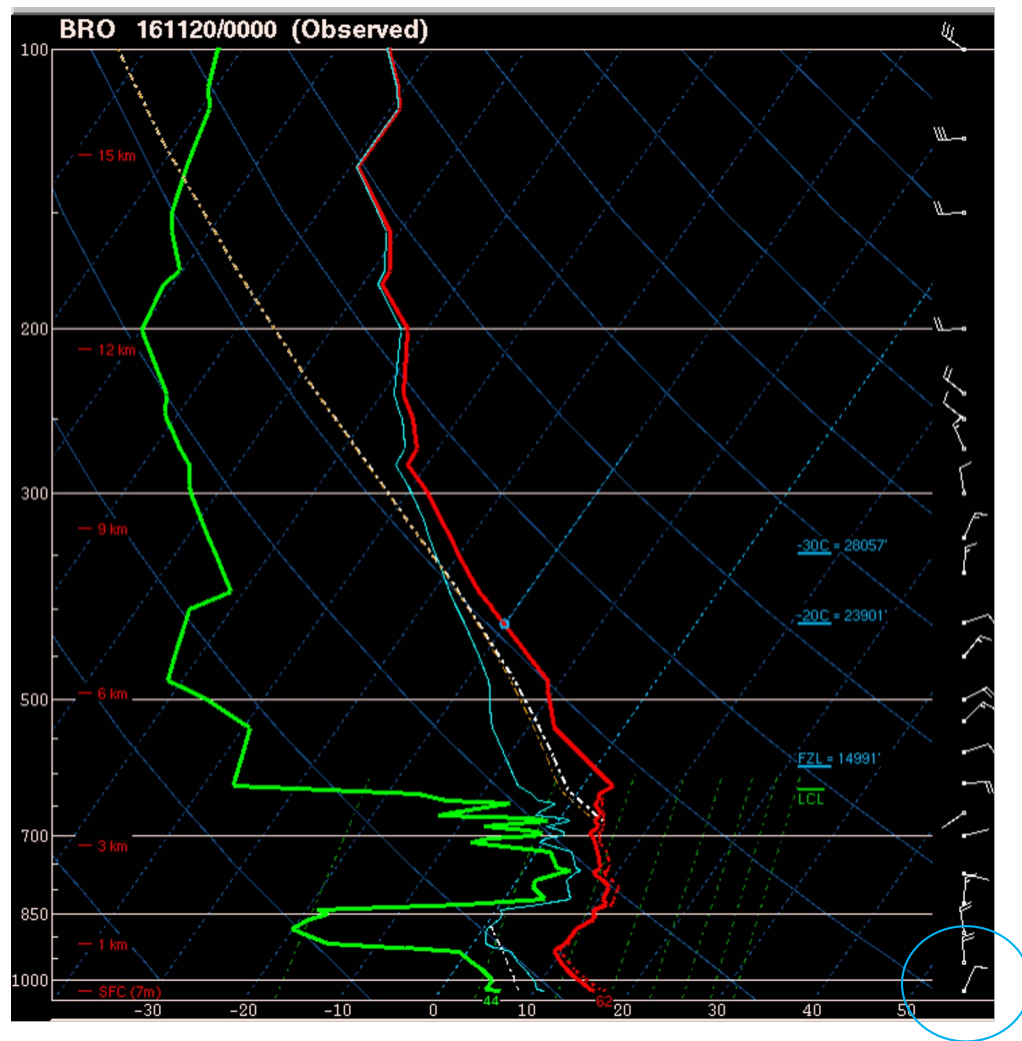




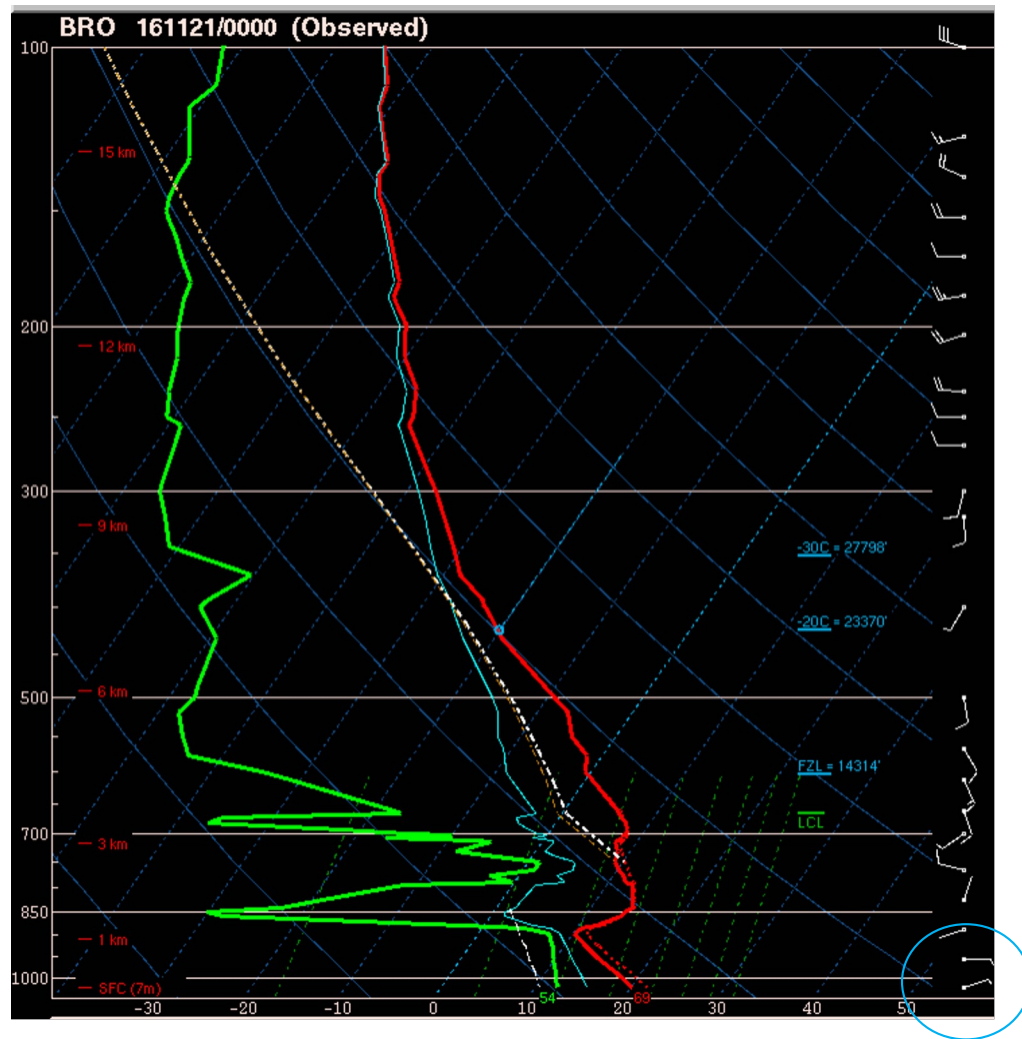


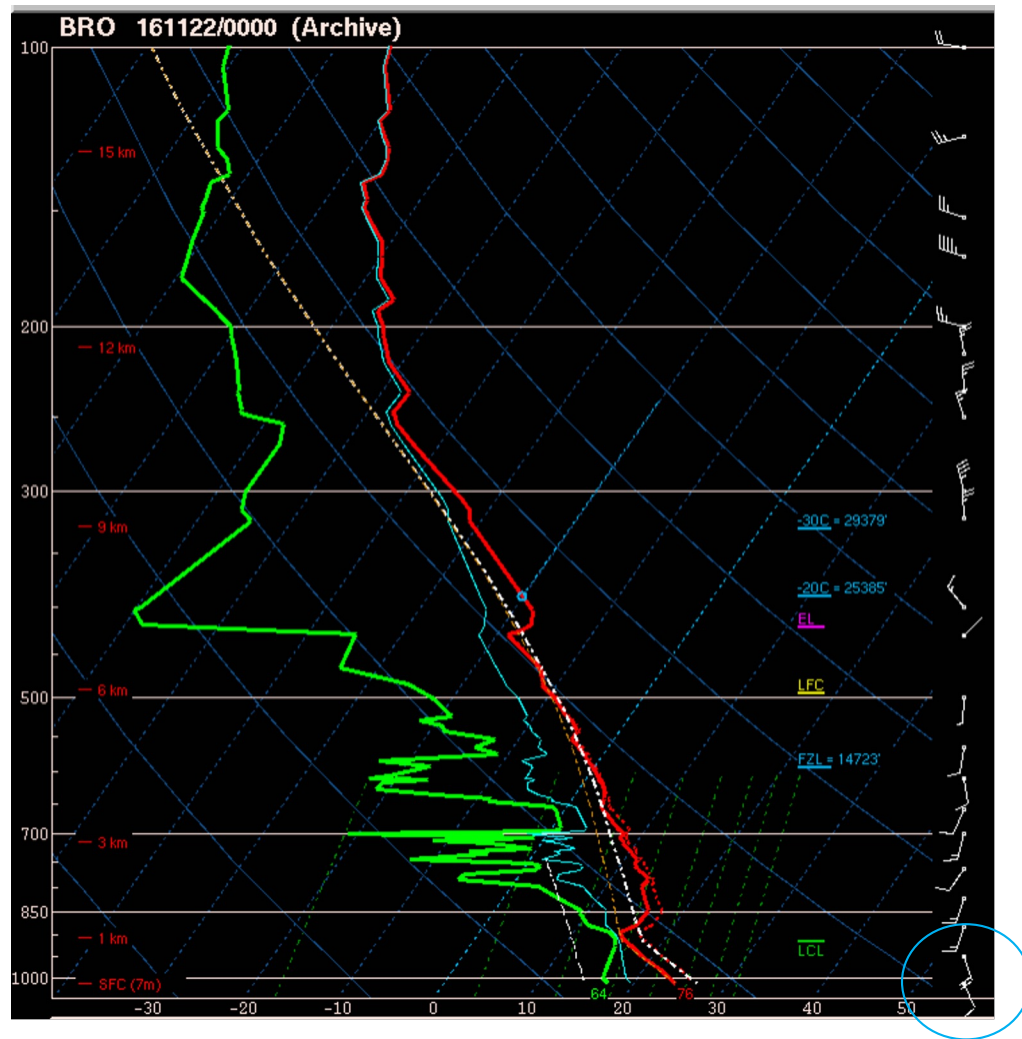


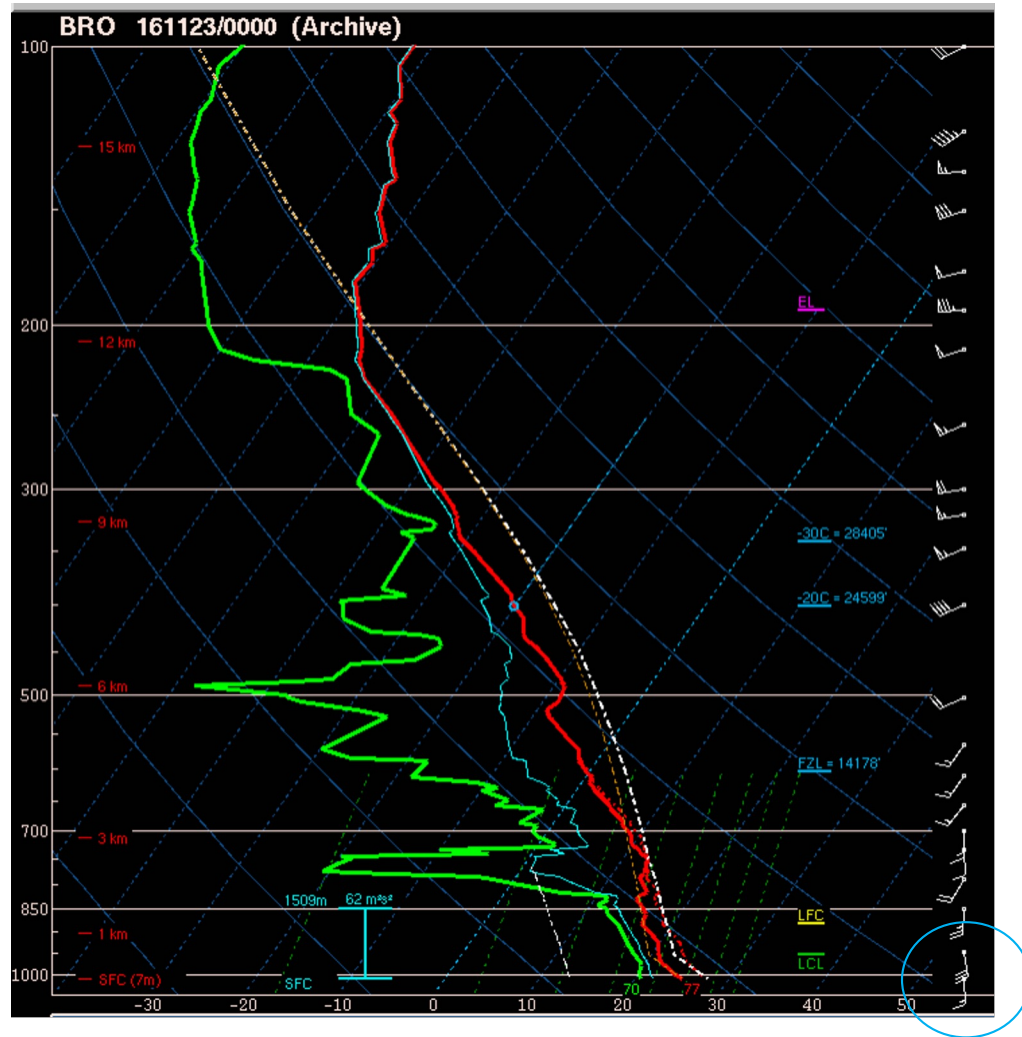






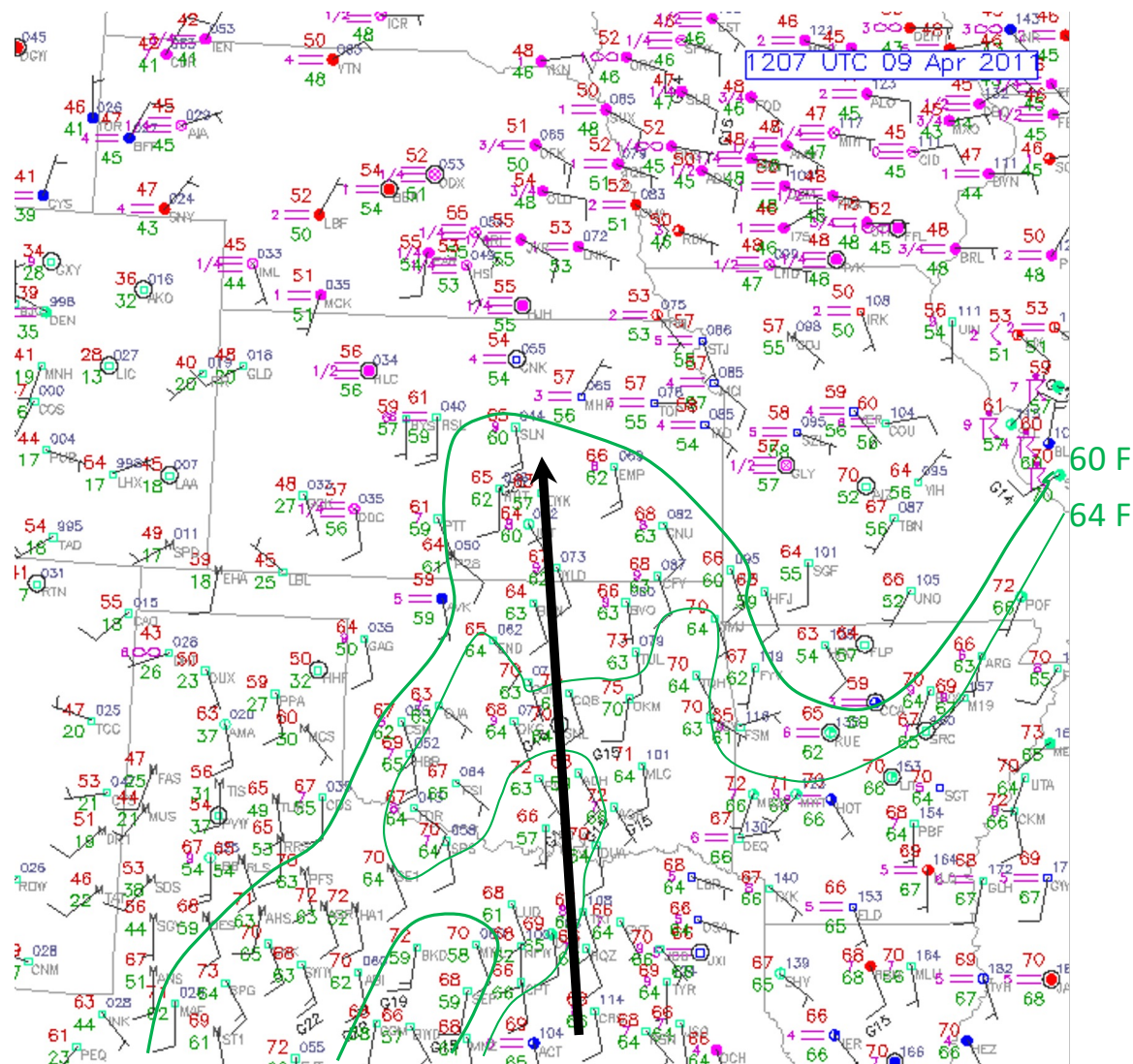




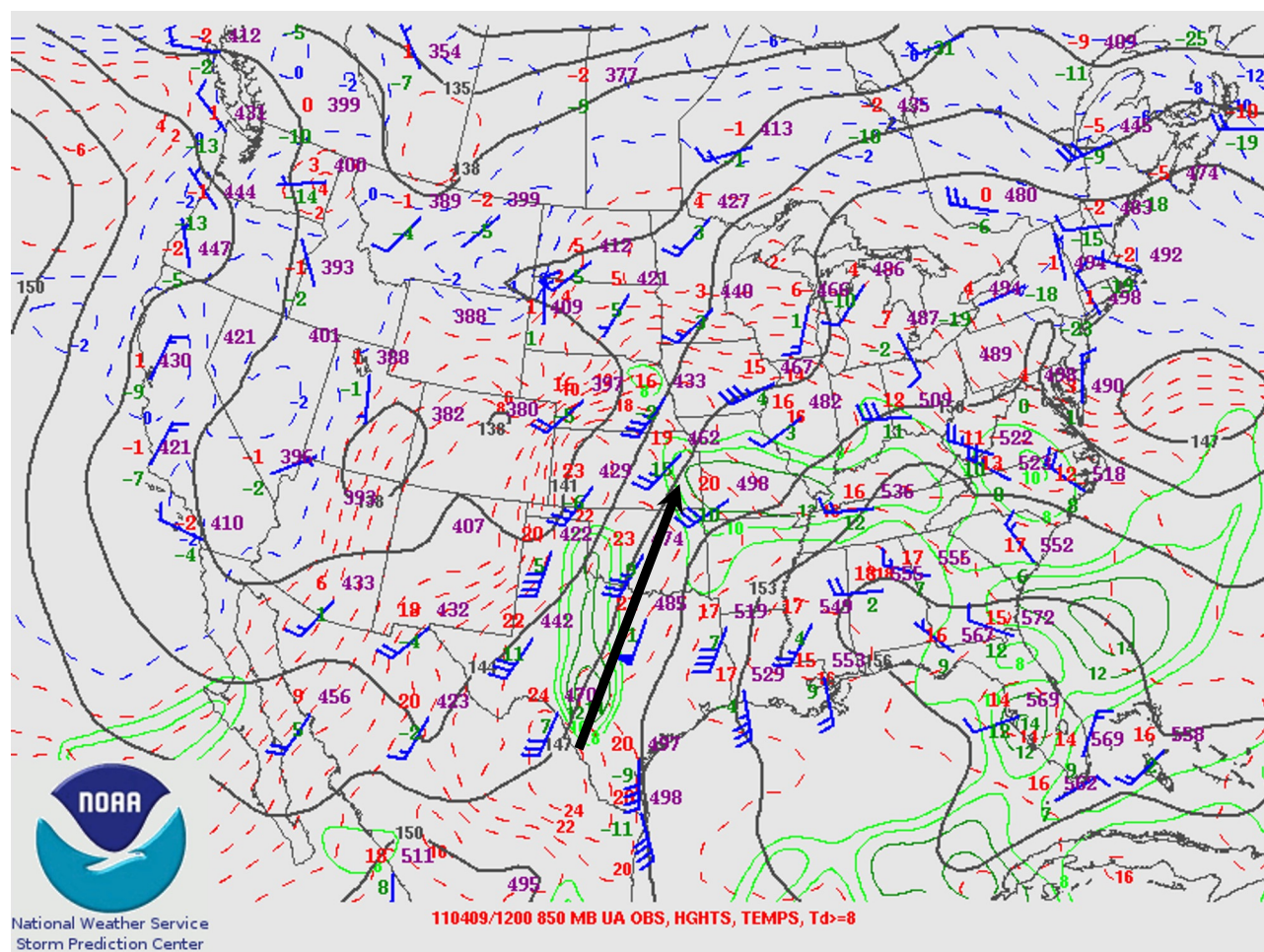


# Inland Vertical Mixing

- **Surface moisture can decrease during the daytime when:**
  - Daytime heating/mixing extends deeper than original moist layer depth
  - Moisture advection is not enough to offset mixing
  - Compounded by dry soil, little green vegetation, and ongoing drought
- **Daytime mixing is governed by:**
  - Vertical moisture structure
  - Height and strength of lid/cap
  - Upstream moisture sources
  - Local moisture sources



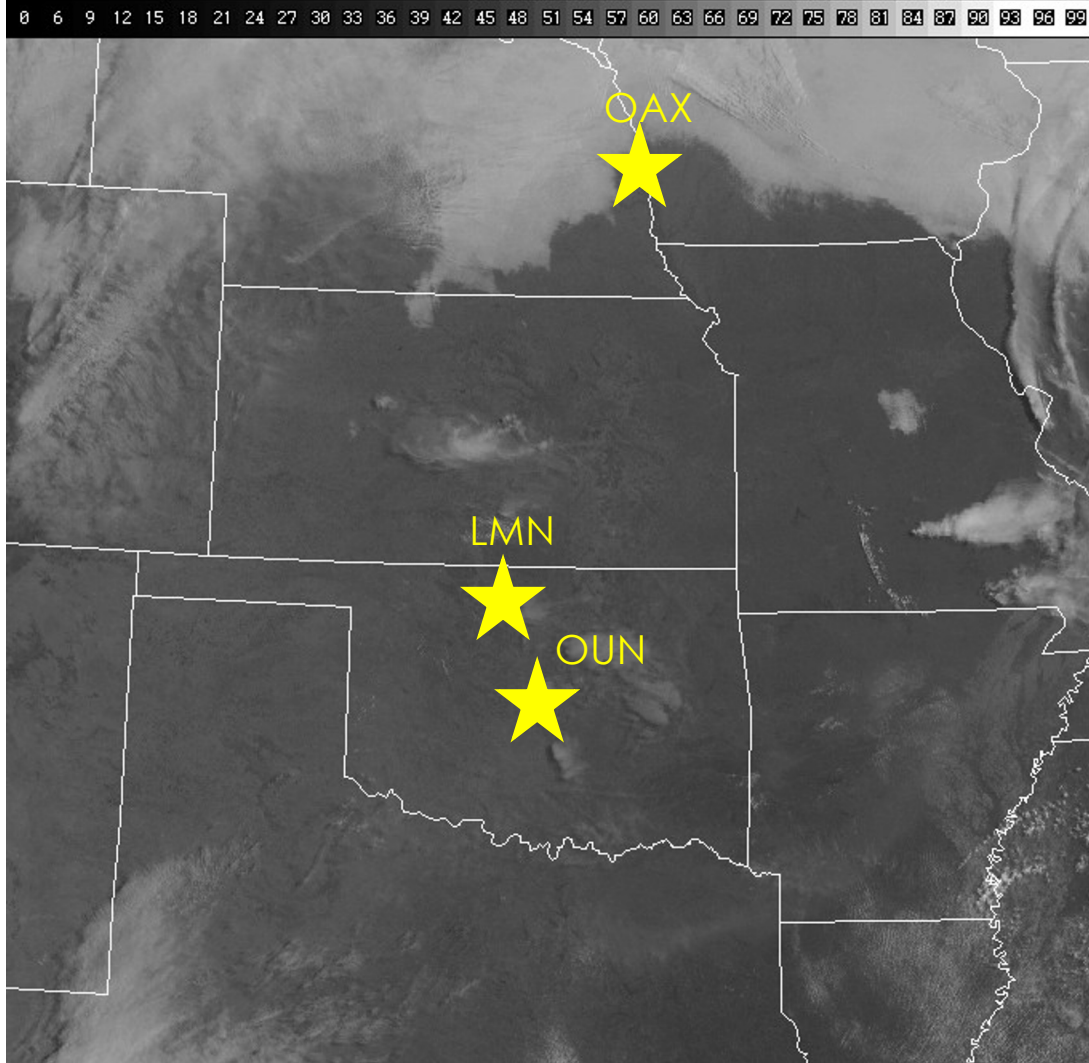




1445 UTC Sat 09 Apr 2011

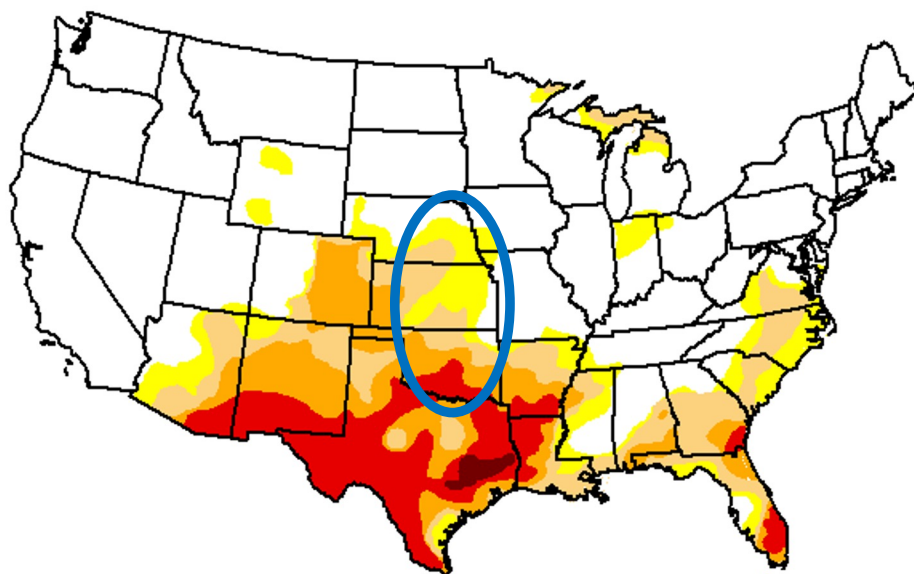
Visible Satellite

<http://adds.aviationweather.gov>





# U.S. Drought Monitor CONUS



**April 5, 2011**

(Released Thursday, Apr. 7, 2011)

Valid 7 a.m. EST

*Drought Conditions (Percent Area)*

	None	D0-D4	D1-D4	D2-D4	D3-D4	D4
<b>Current</b>	62.46	37.54	28.75	18.17	8.75	0.42
<b>Last Week</b> 3/29/2011	61.76	38.24	28.14	18.22	6.41	0.00
<b>3 Months Ago</b> 1/4/2011	60.50	39.50	21.74	8.50	2.60	0.00
<b>Start of Calendar Year</b> 1/4/2011	60.50	39.50	21.74	8.50	2.60	0.00
<b>Start of Water Year</b> 9/28/2010	60.05	39.95	13.16	3.09	0.30	0.00
<b>One Year Ago</b> 4/6/2010	71.57	28.43	9.10	2.00	0.00	0.00

## Intensity:

D0 Abnormally Dry	D3 Extreme Drought
D1 Moderate Drought	D4 Exceptional Drought
D2 Severe Drought	

*The Drought Monitor focuses on broad-scale conditions. Local conditions may vary. See accompanying text summary for forecast statements.*

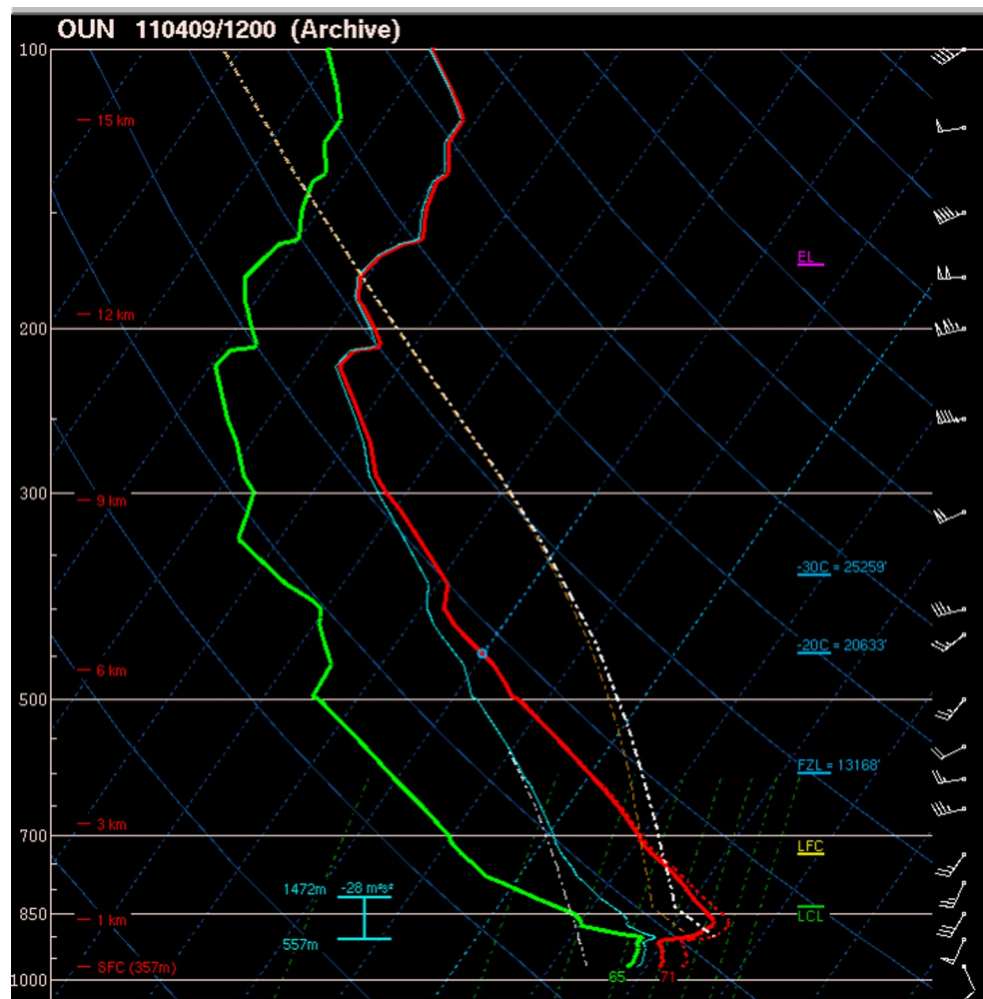
## **Author(s):**

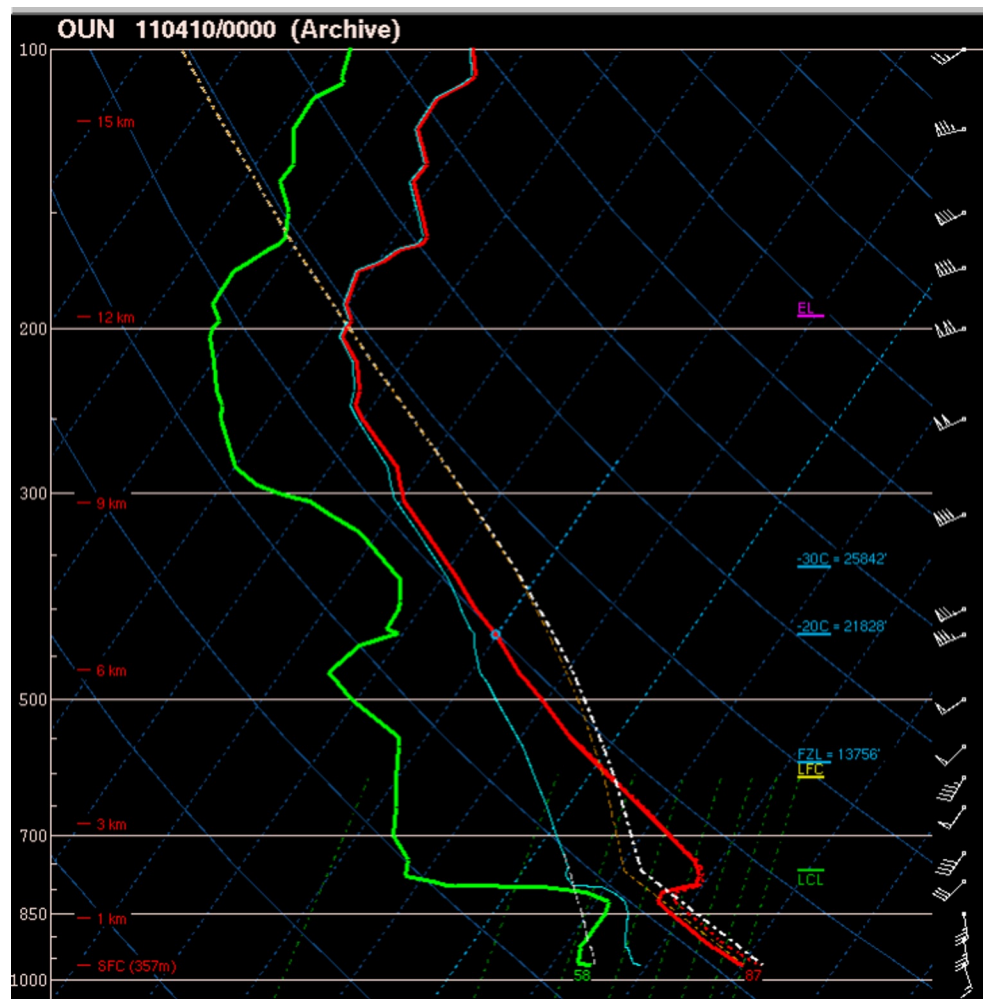
Mark Svoboda

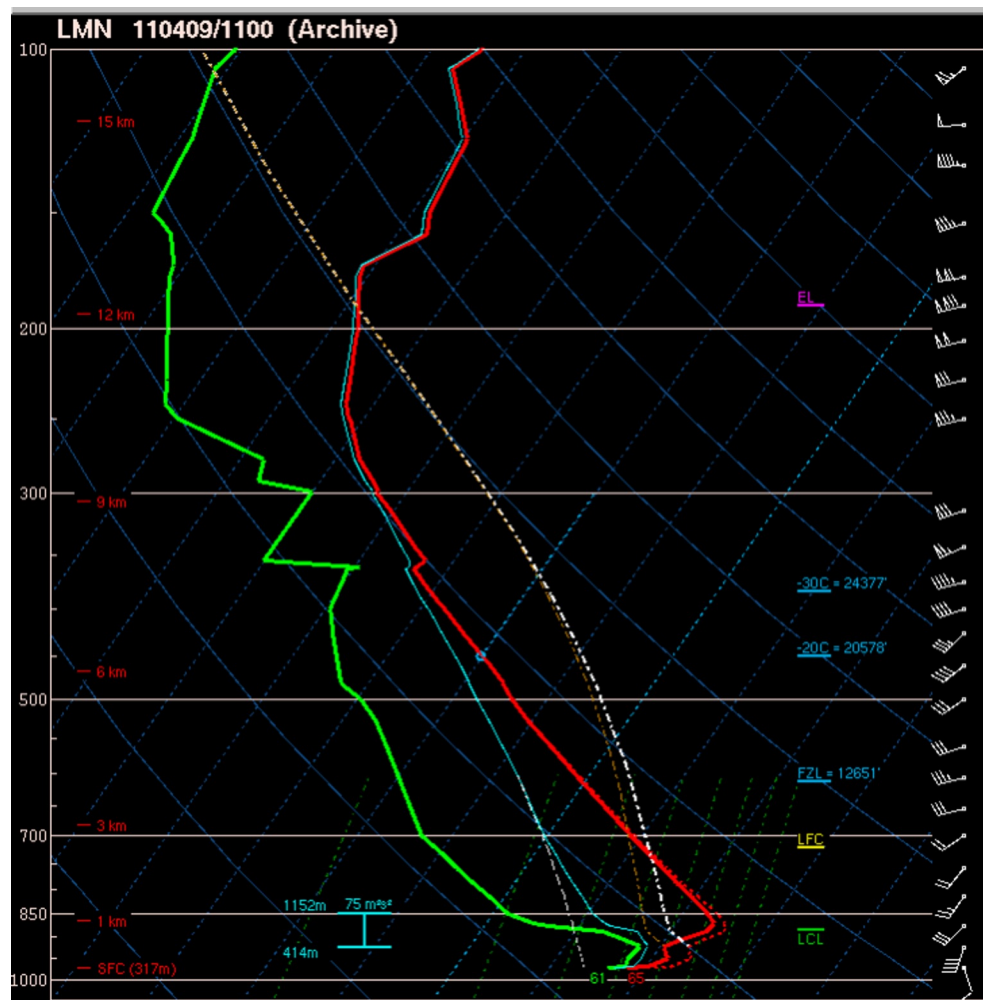
National Drought Mitigation Center

USDA

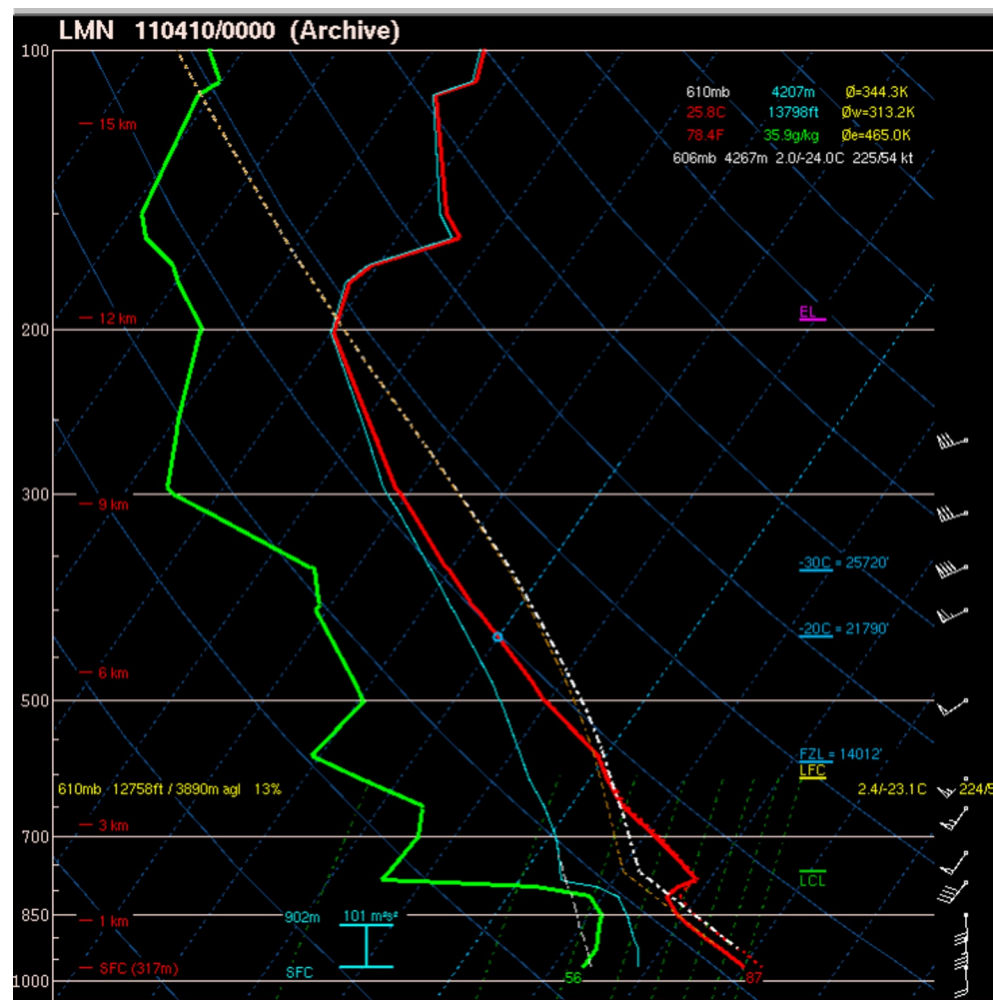


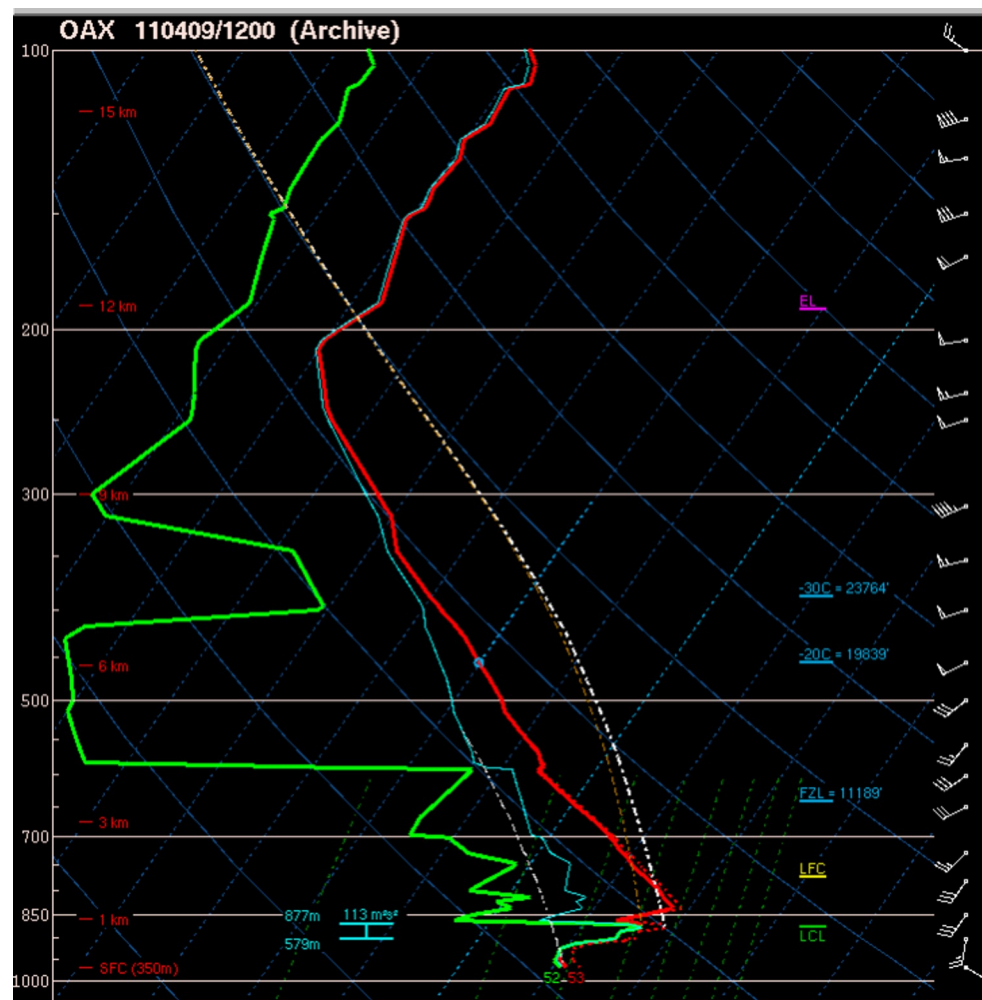


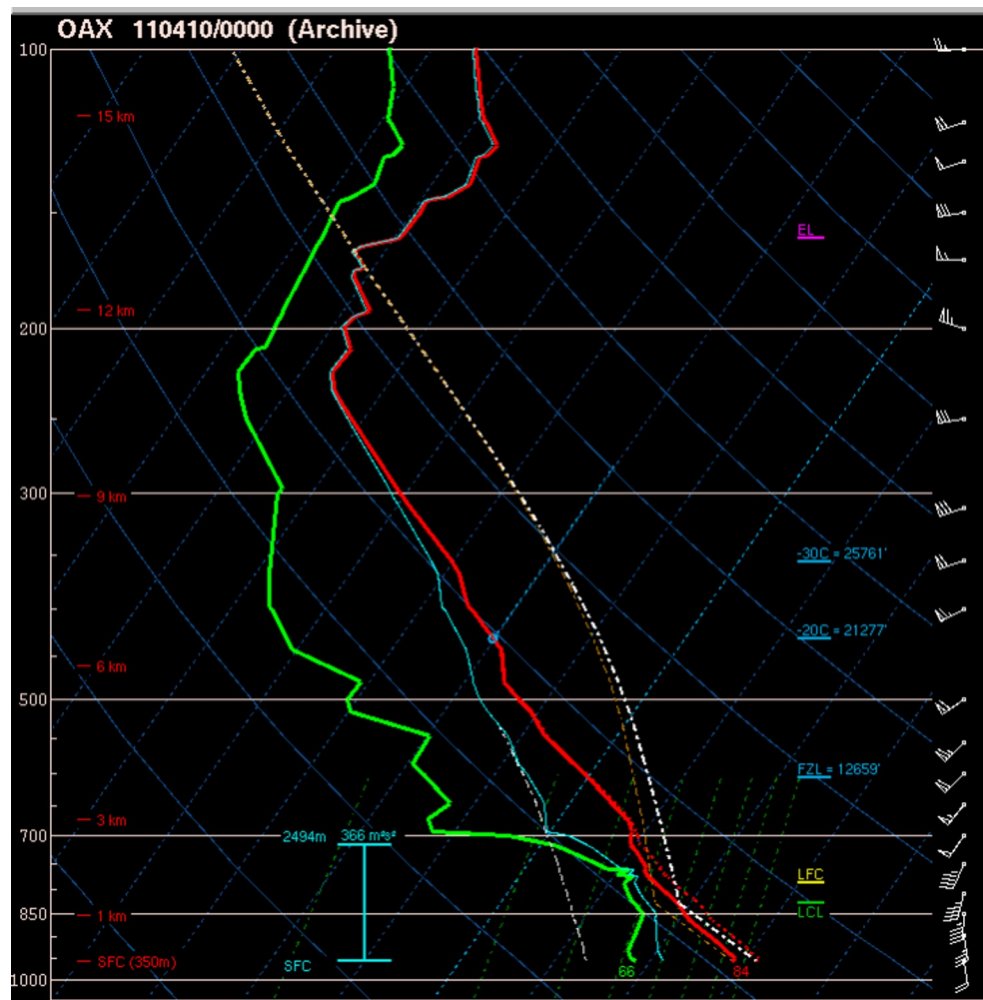






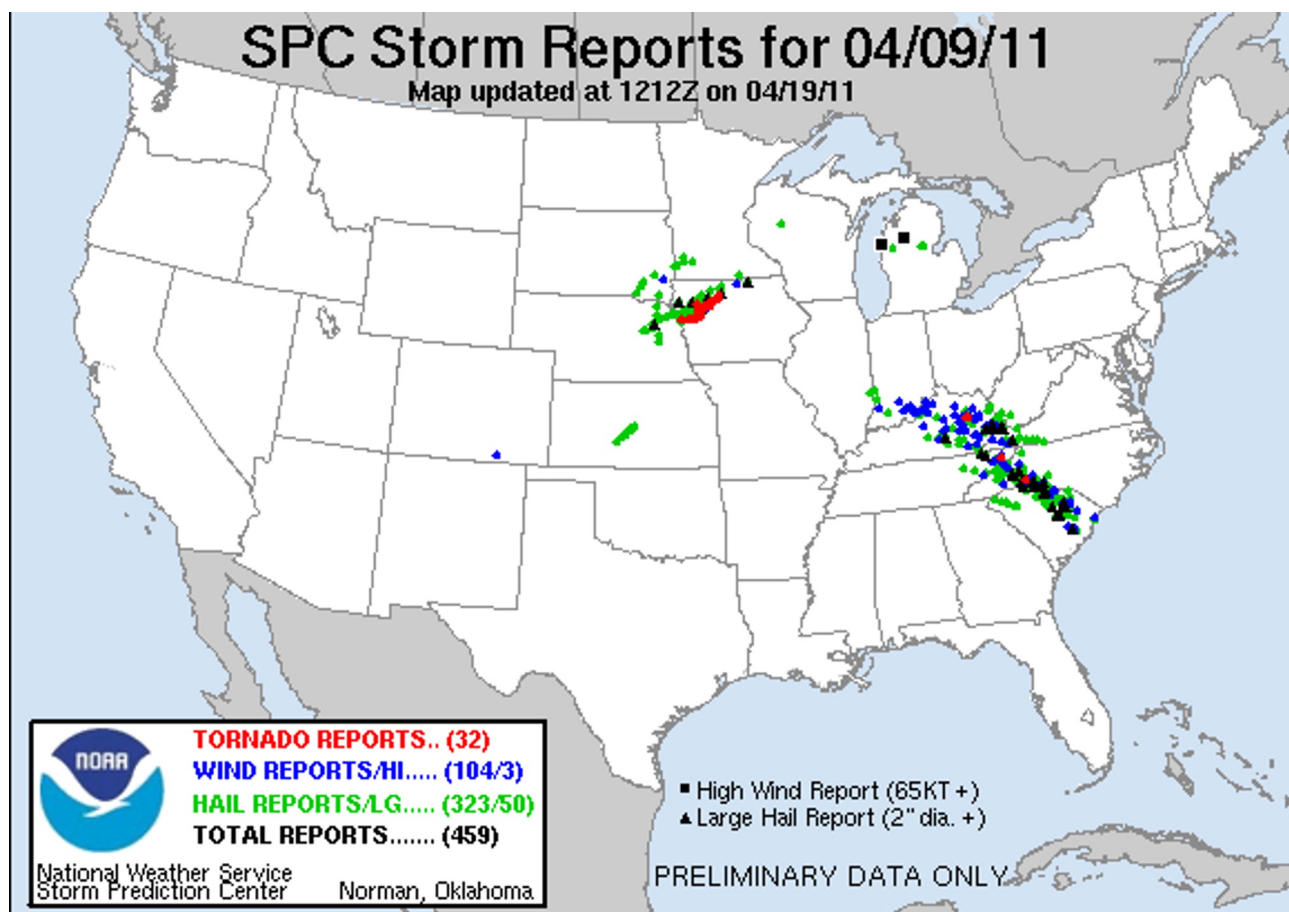






# SPC Storm Reports for 04/09/11

Map updated at 1212Z on 04/19/11

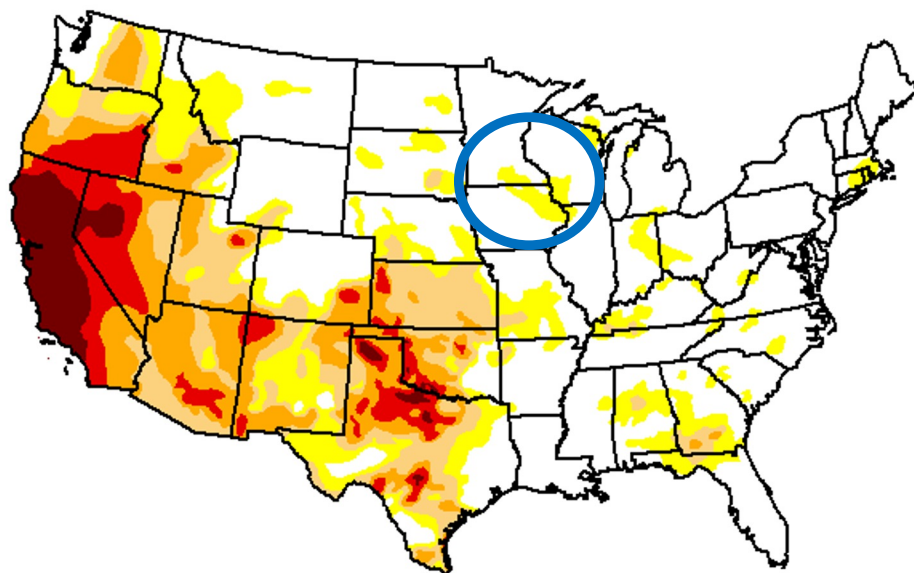




# Evapotranspiration

- Need moist soil and growing vegetation
- Plentiful rain previous 1-2 weeks
- Maturing crops (wheat, corn, or canola)
- Capped boundary layer to trap moisture; relatively weak winds
- Almost always a significant return flow contribution (in addition)

# U.S. Drought Monitor CONUS



**August 19, 2014**

(Released Thursday, Aug. 21, 2014)

Valid 8 a.m. EDT

*Drought Conditions (Percent Area)*

	None	D0-D4	D1-D4	D2-D4	D3-D4	D4
<b>Current</b>	51.93	48.07	33.56	21.62	10.12	3.80
<b>Last Week</b> 8/12/2014	52.81	47.19	33.61	22.03	10.24	3.79
<b>3 Months Ago</b> 5/20/2014	52.36	47.64	38.12	28.30	14.47	4.99
<b>Start of Calendar Year</b> 12/31/2013	48.24	51.76	30.95	16.67	3.96	0.37
<b>Start of Water Year</b> 10/1/2013	39.57	60.43	41.21	20.70	3.06	0.29
<b>One Year Ago</b> 8/20/2013	40.02	59.98	45.61	32.23	10.54	1.32

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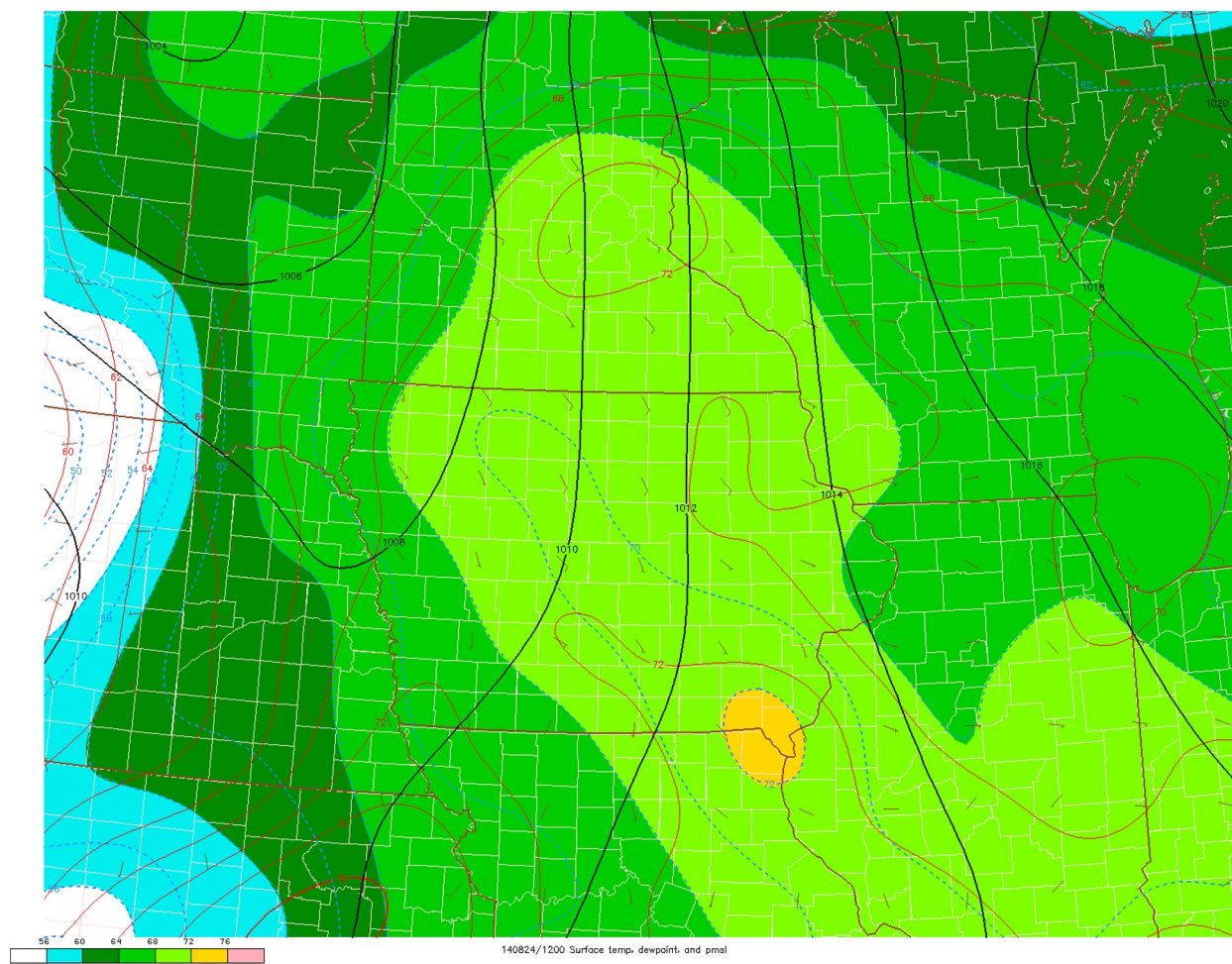
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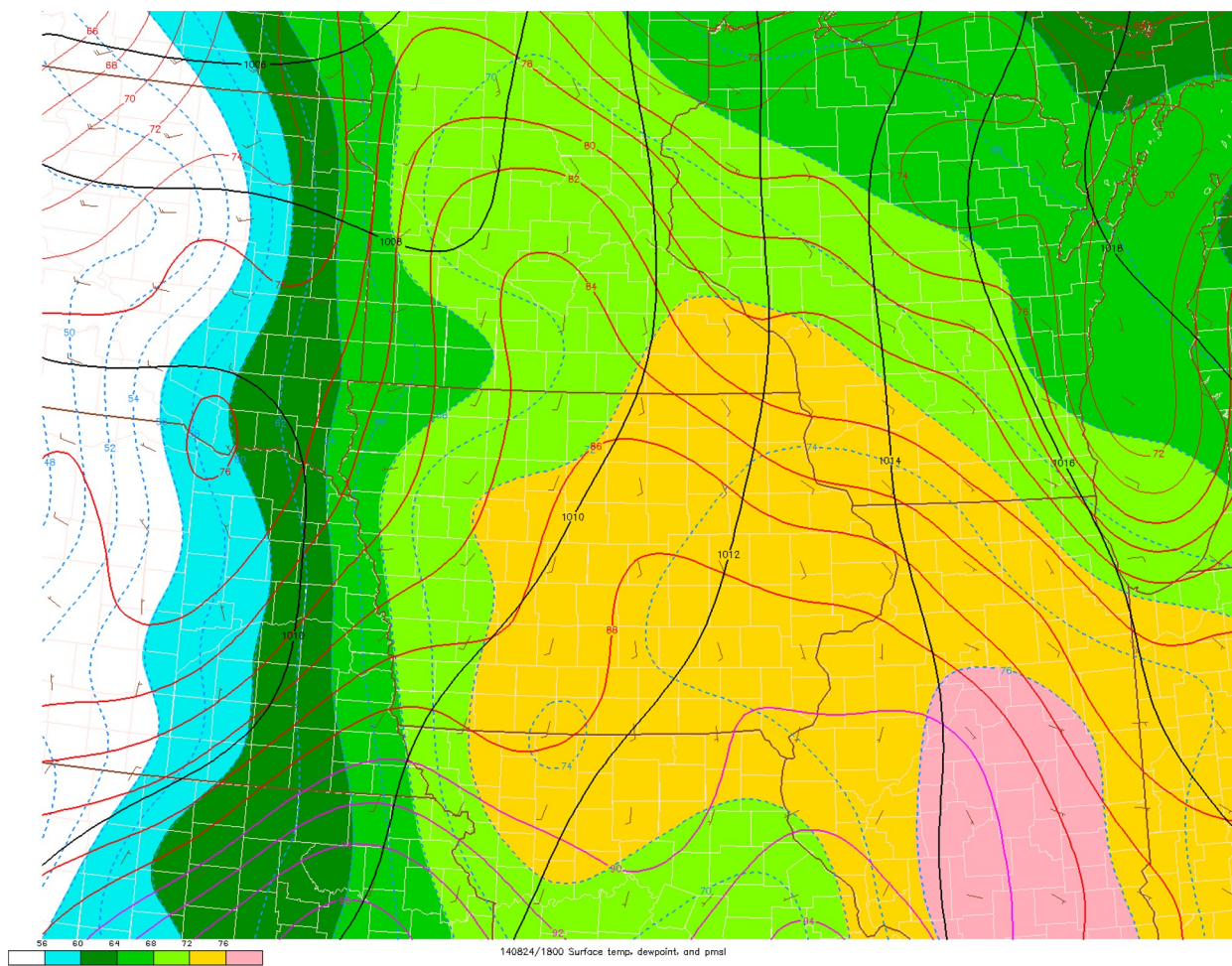
Richard Tinker

CPC/NOAA/NWS/NCEP

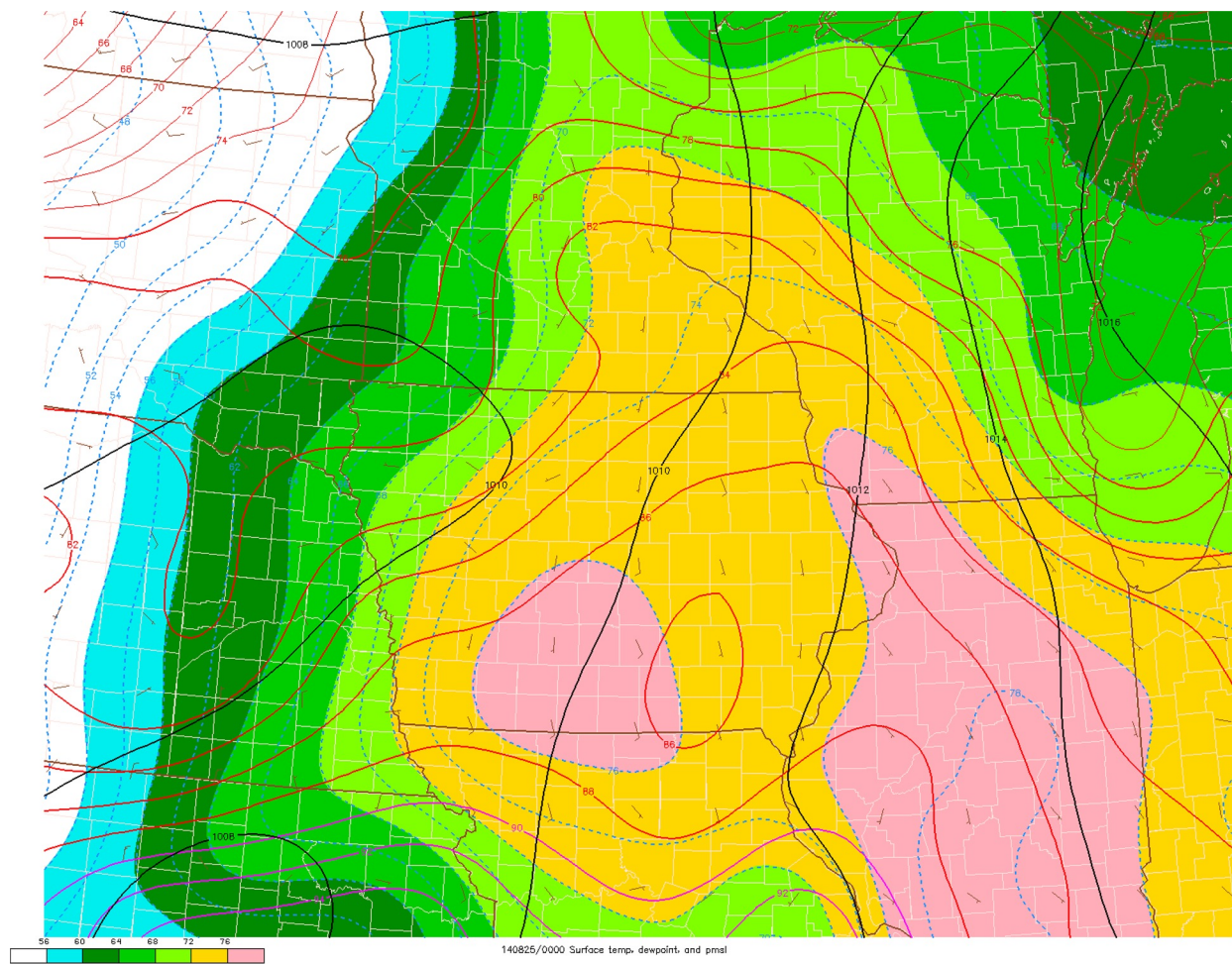
**USDA**

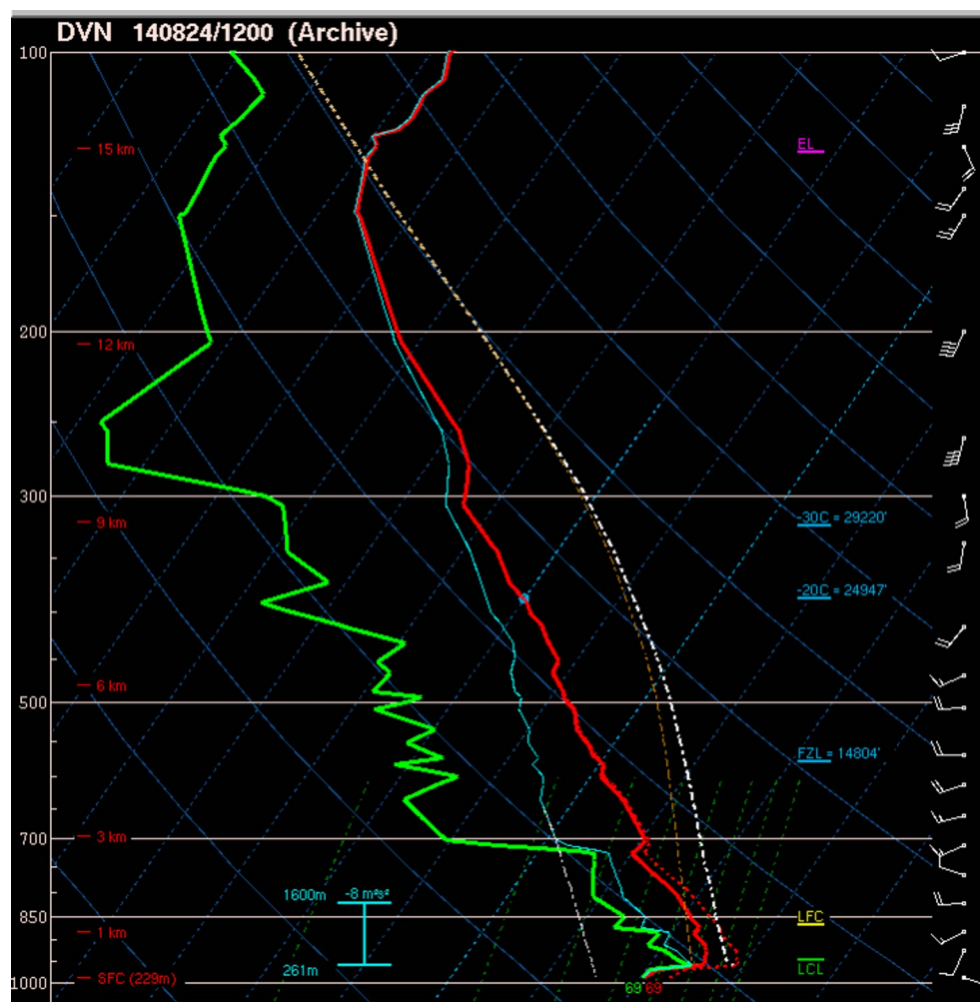


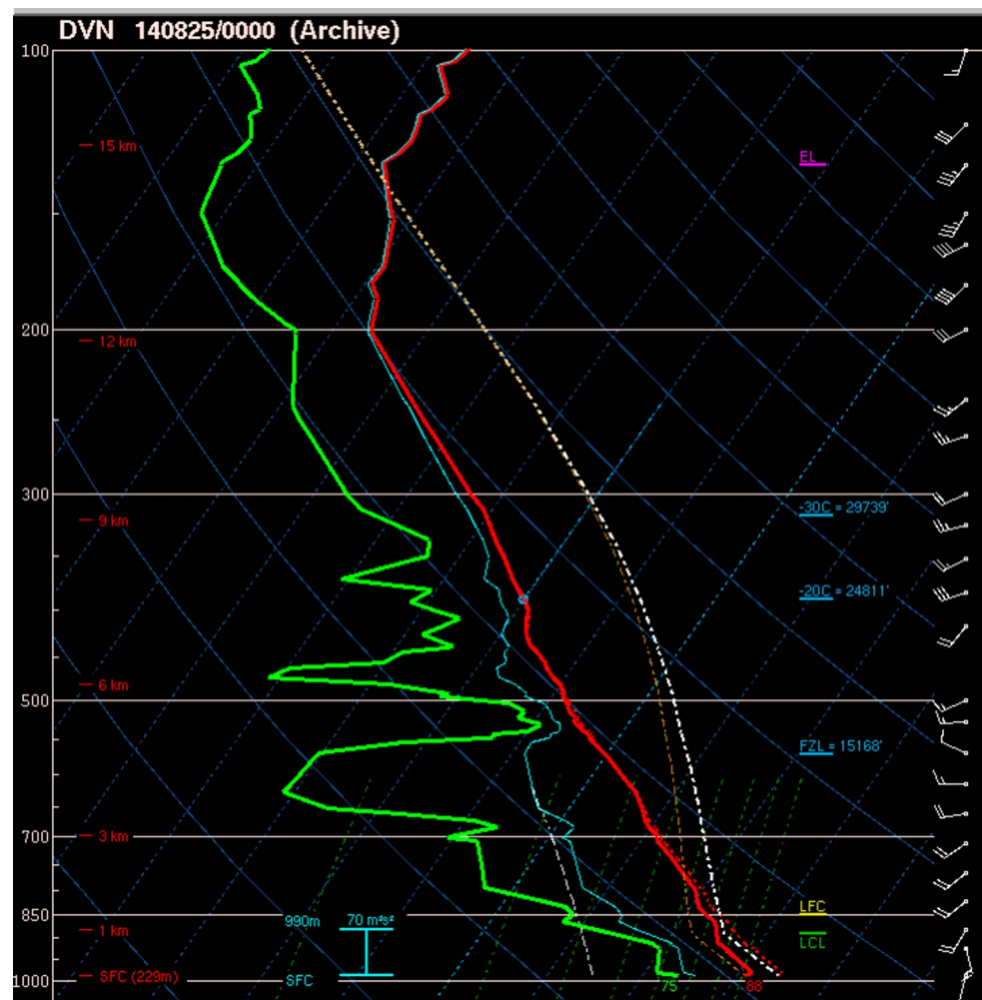


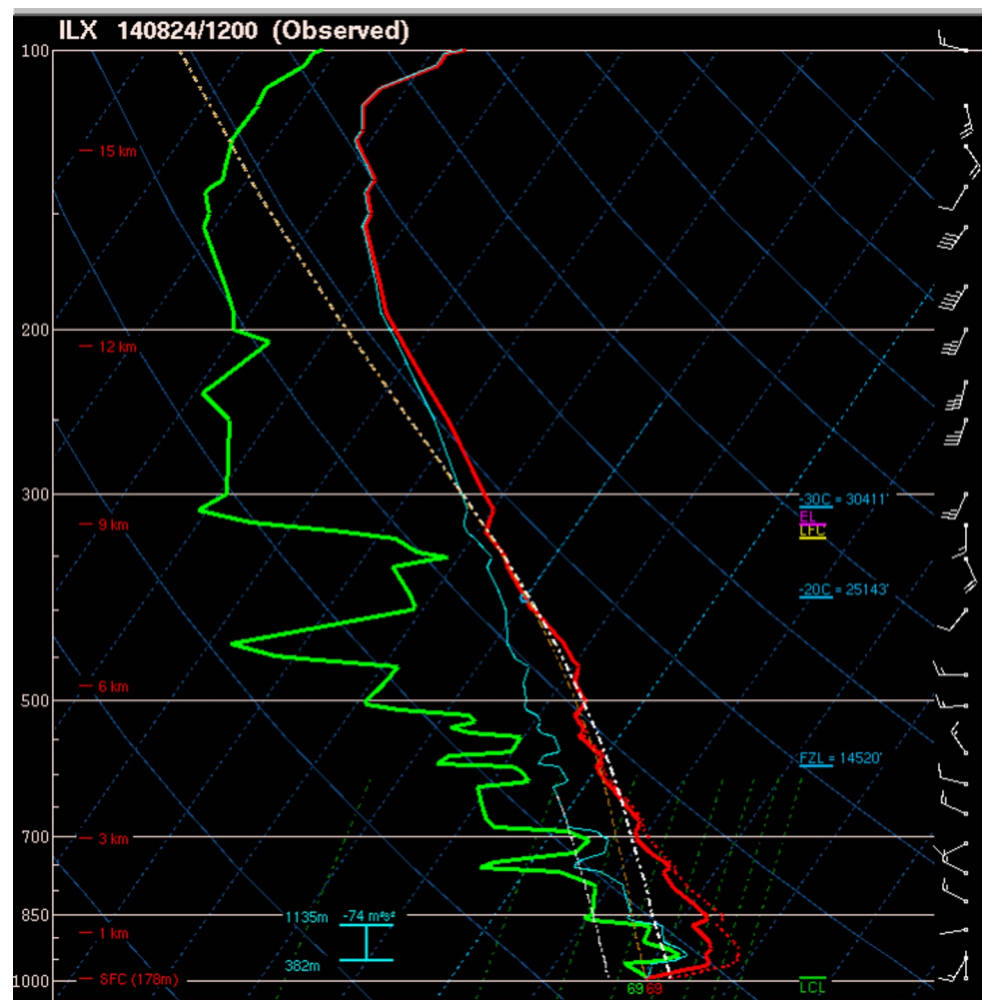




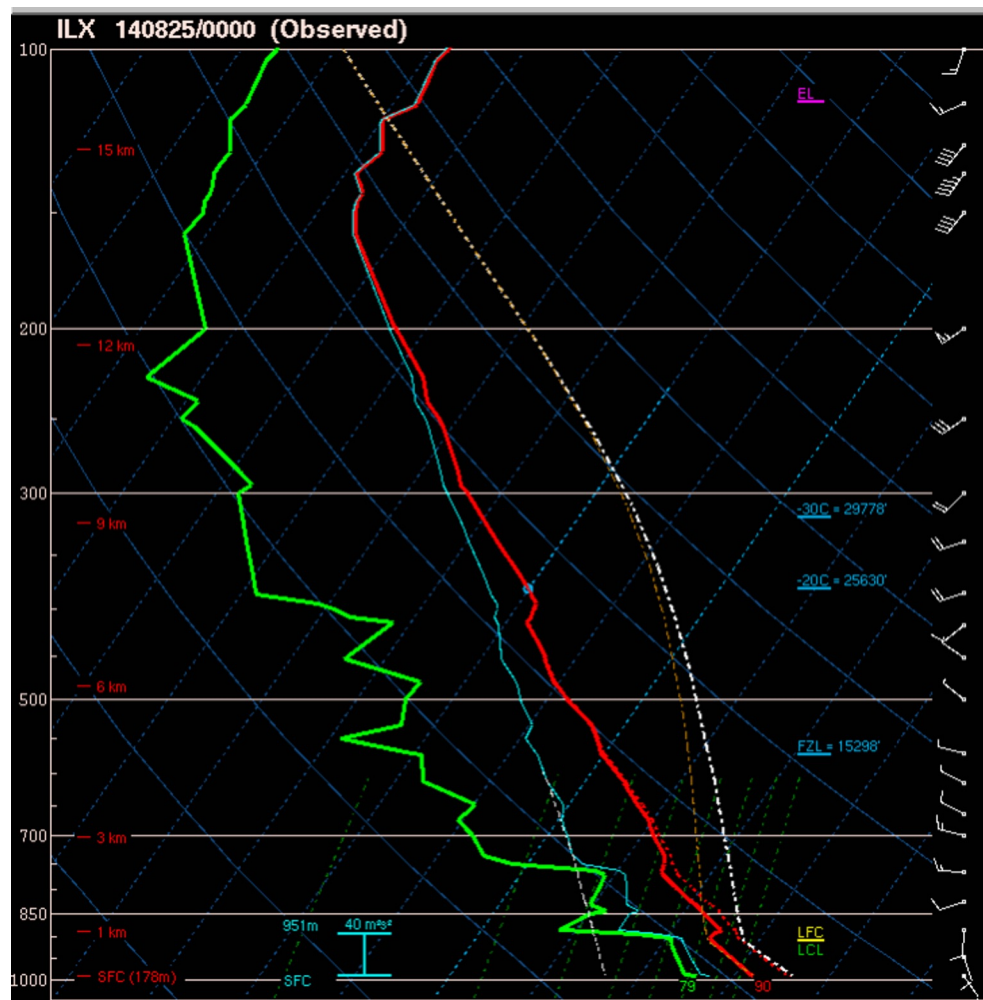






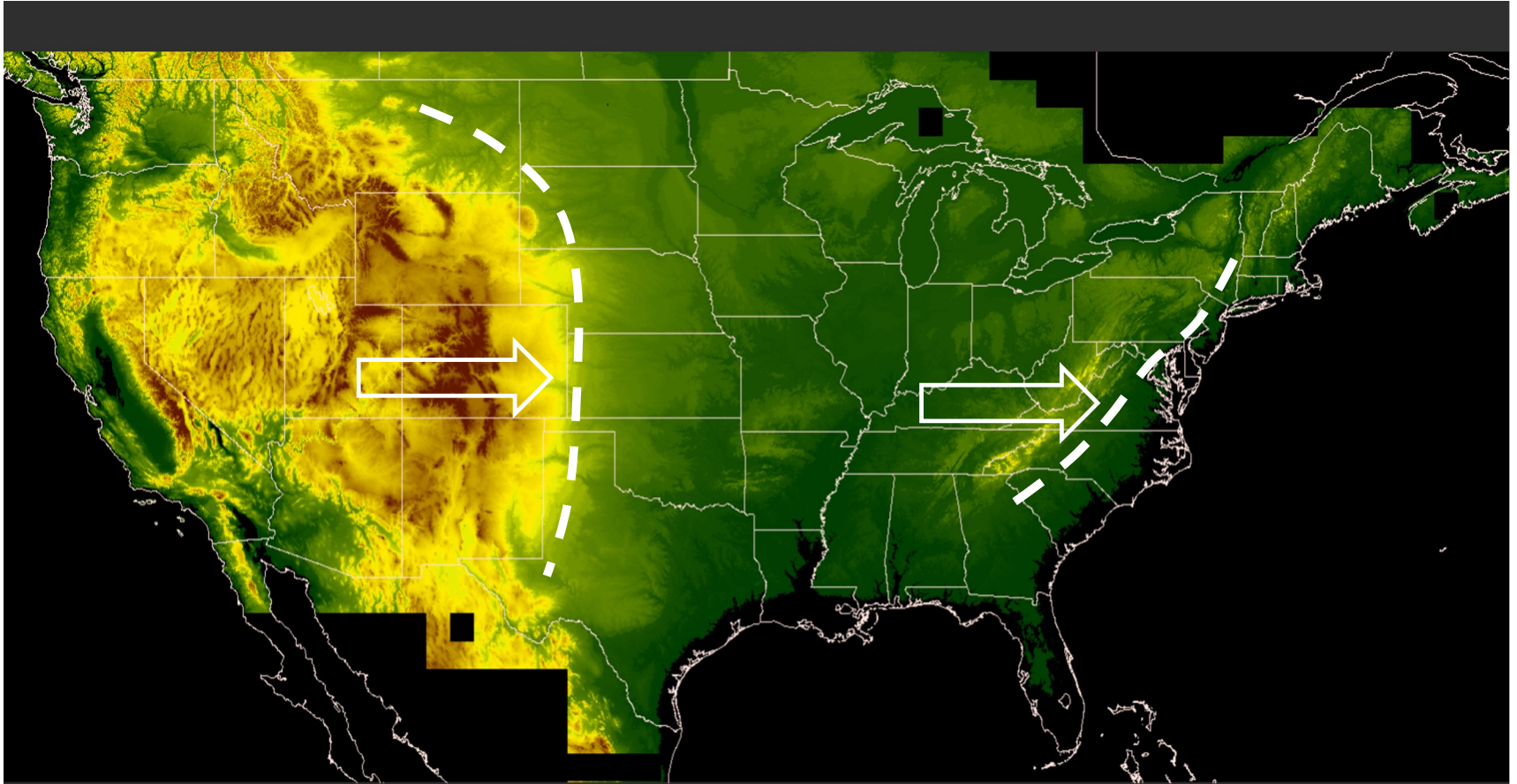






# Lee Cyclogenesis

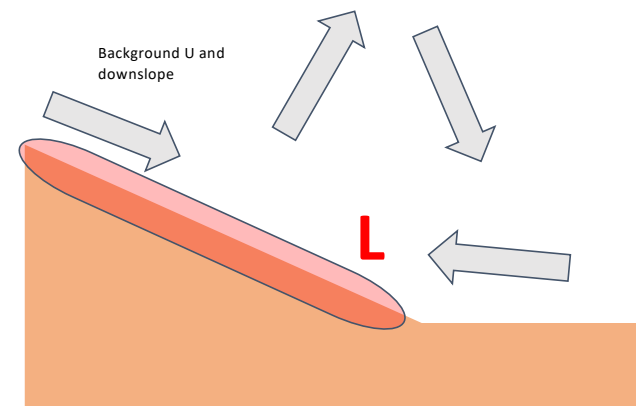
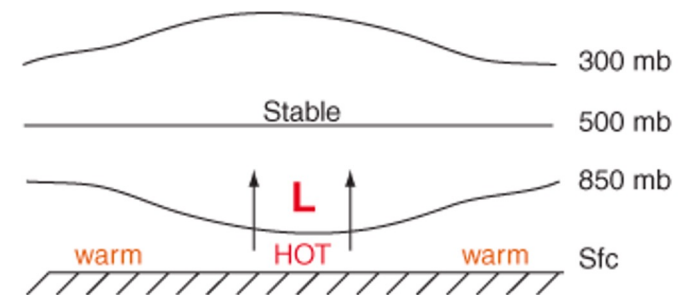
- **Midlevel flow crosses high terrain:**
  - Subsidence and warming E of mountains
  - Warming of column leads to “warm core” low formation
  - Strongest pressure falls with strongest flow crosses highest terrain
- **The lee trough/cyclone deepens before arrival of strongest Q-G forcing for ascent:**
  - Head start on differential advection, “loaded gun” sounding, and veering winds with height



951027/1200 GOES8 SMD

# Lee Cyclogenesis

- **Mid-level flow crosses high terrain:**
  - Subsidence and warming E of mountains
  - Stronger diabatic heating across slope results in rising motion from induced katabatic flow. (Frontogenesis)
  - Warming of column and rising motion leads to “warm core” low formation
  - Strongest pressure falls with strongest flow crosses highest terrain
- **The lee trough/cyclone deepens before arrival of strongest Q-G forcing for ascent: (More QG Later)**
  - Head start on differential advection, “loaded gun” sounding, and veering winds with height

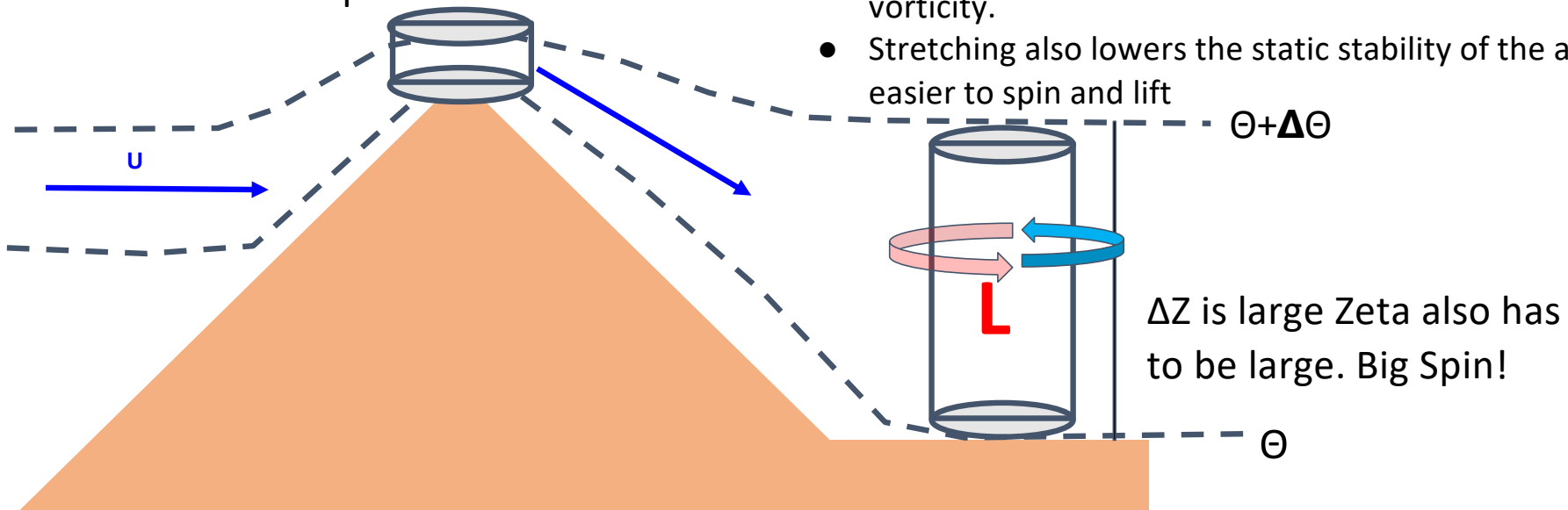




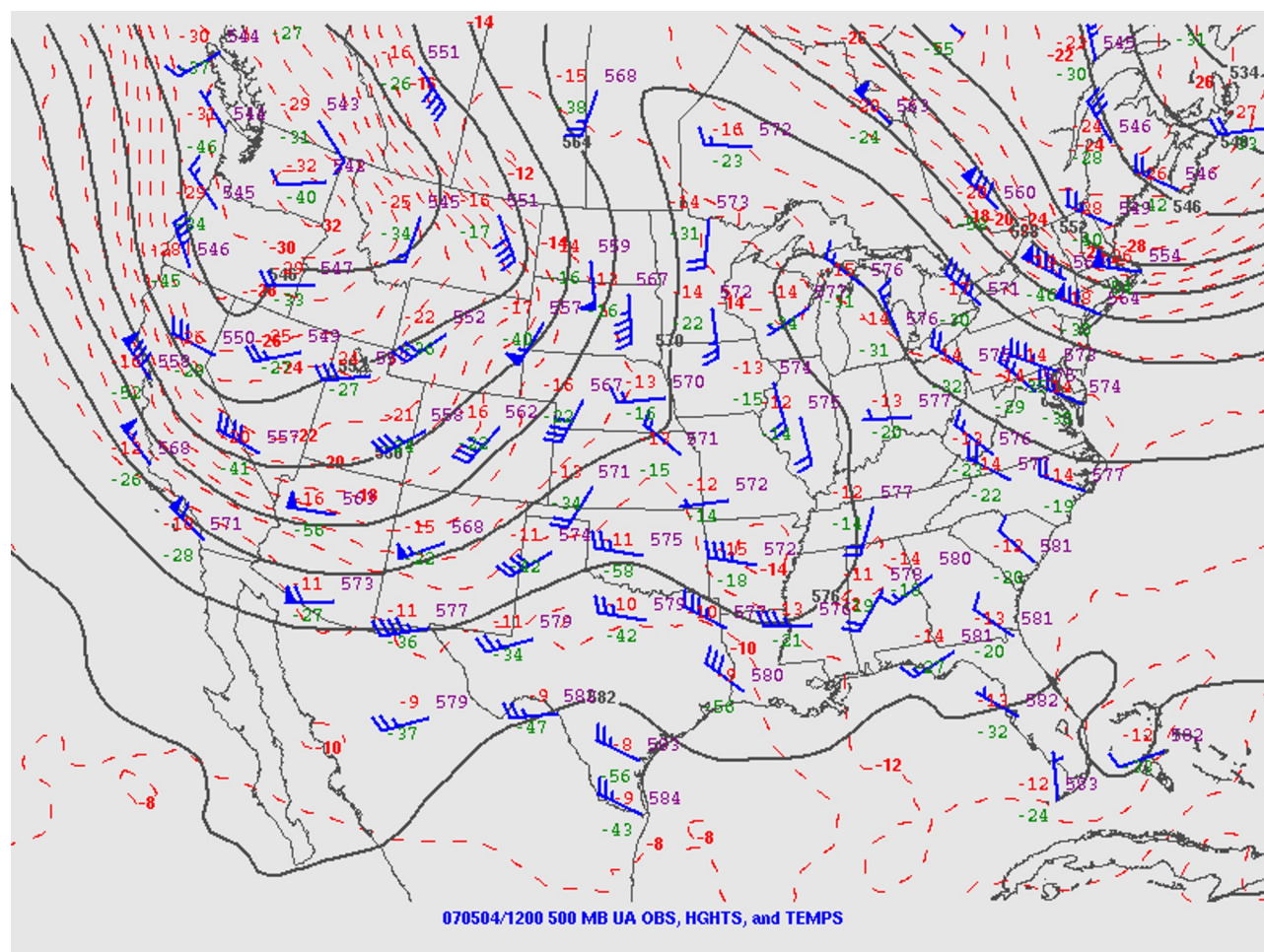
# Lee Cyclogenesis PV Framework

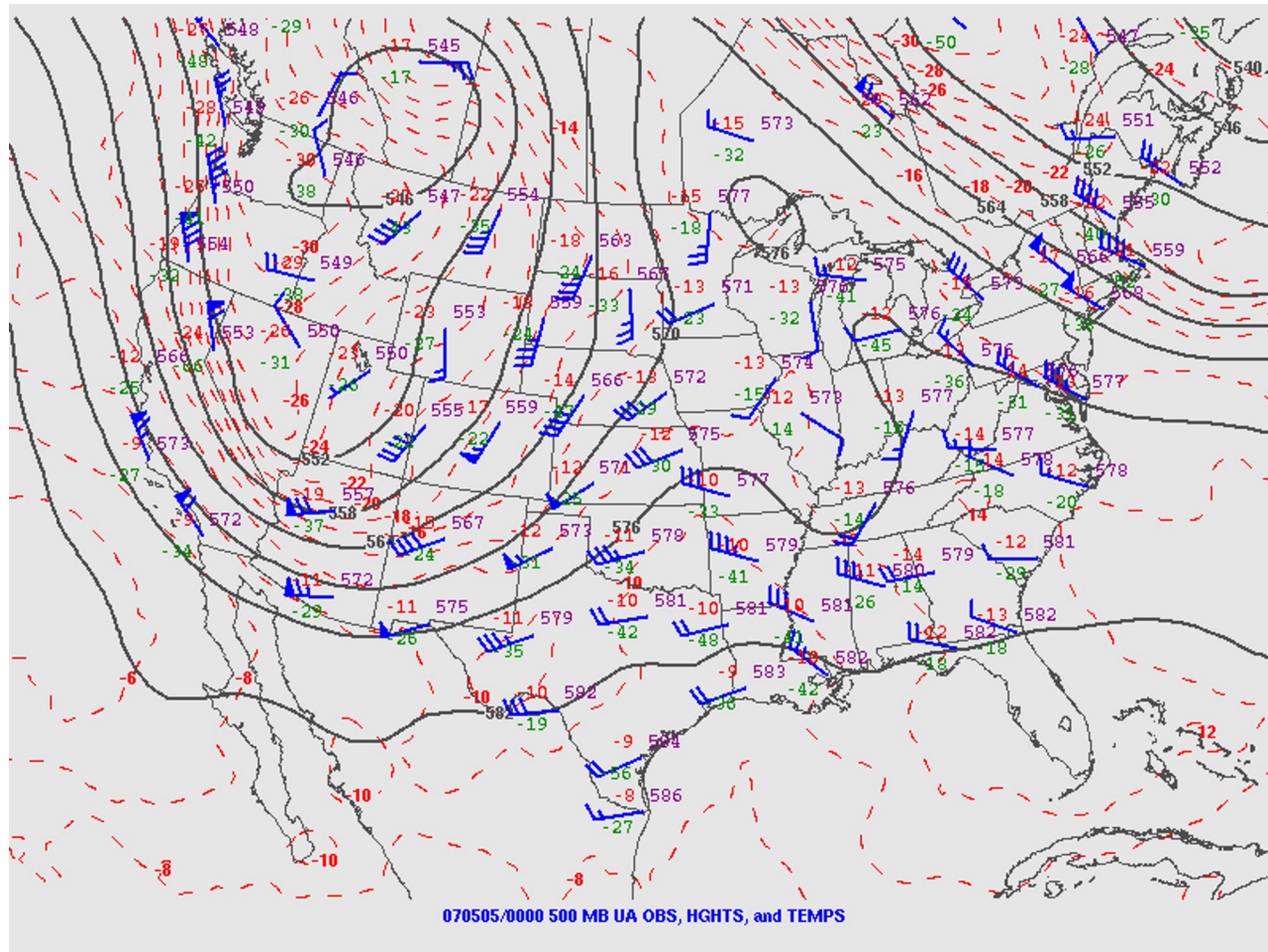
$$Pv = \frac{\zeta_{\theta} + f}{\Delta Z} = \text{const}$$

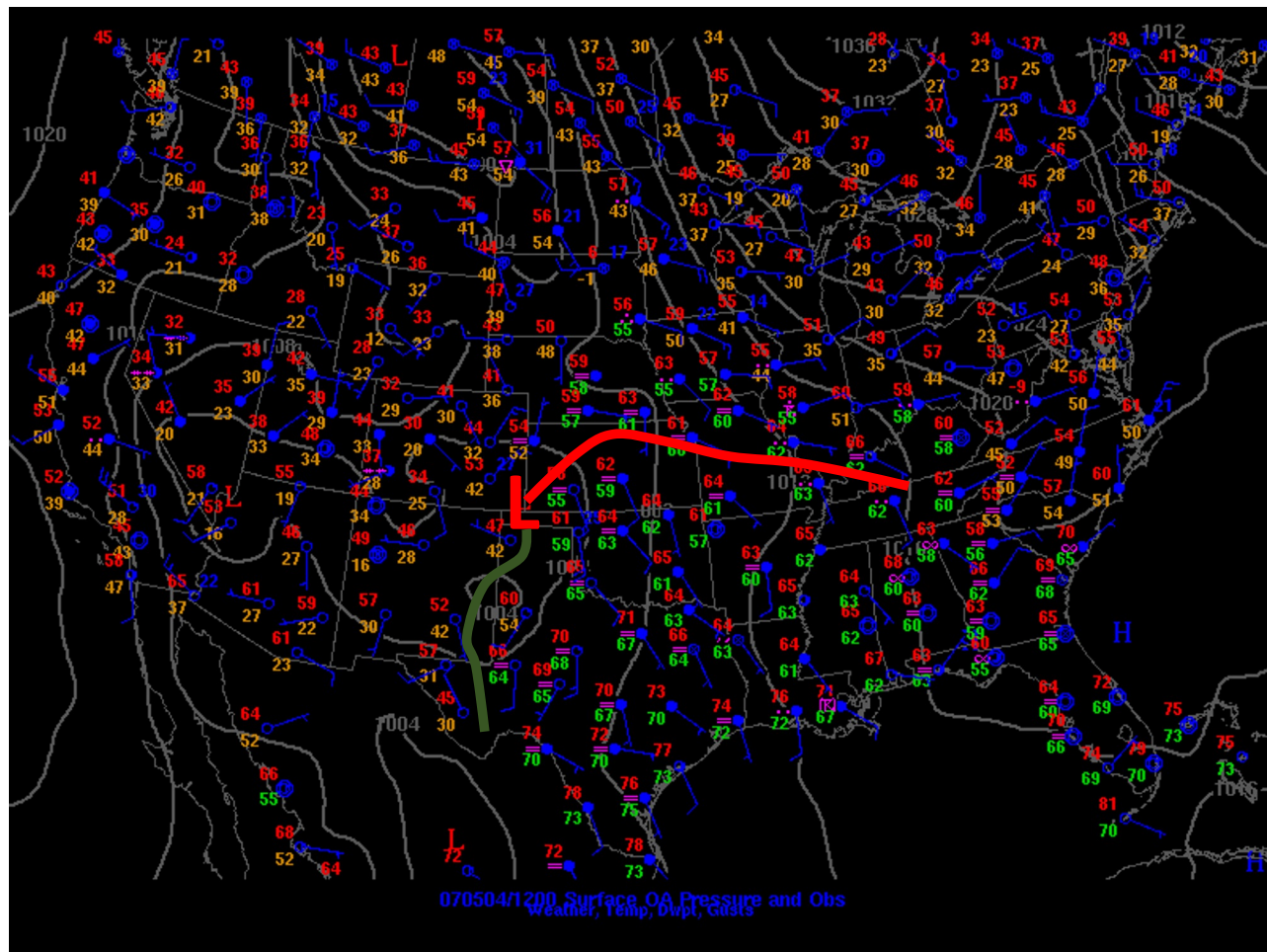
$\Delta Z$  small Zeta is also small little spin



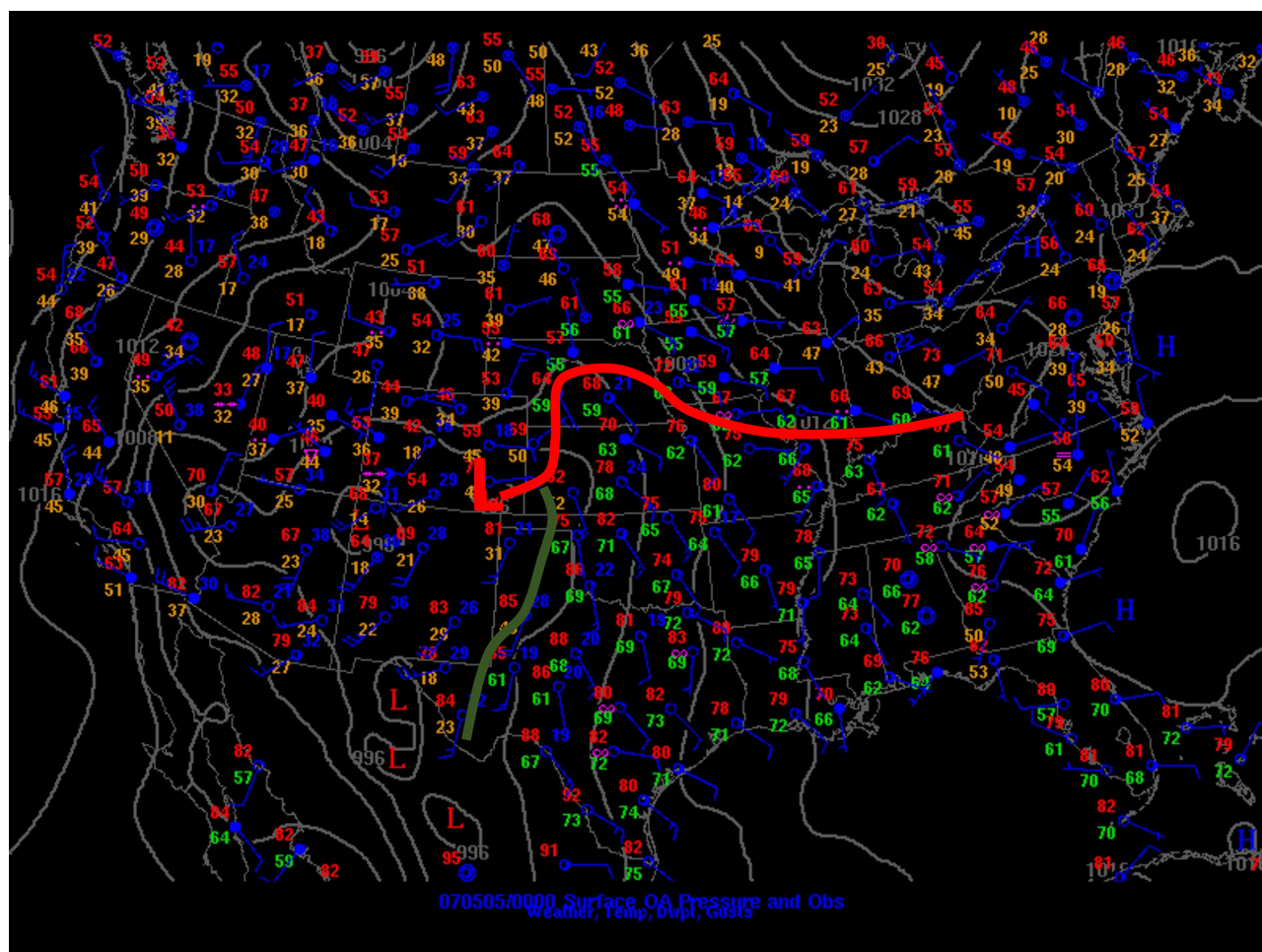
- Conservation of Potential Vorticity
  - Assumptions:
    - Constant density/stable
    - Adiabatic
    - Frictionless
- Stretching of the column vertically requires an increase in relative vorticity between the theta surfaces inducing positive vorticity.
- Stretching also lowers the static stability of the atmosphere. easier to spin and lift









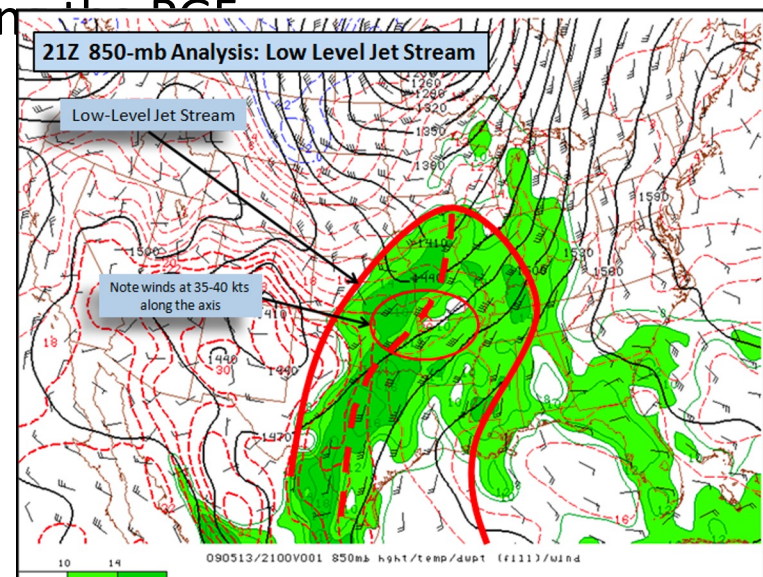
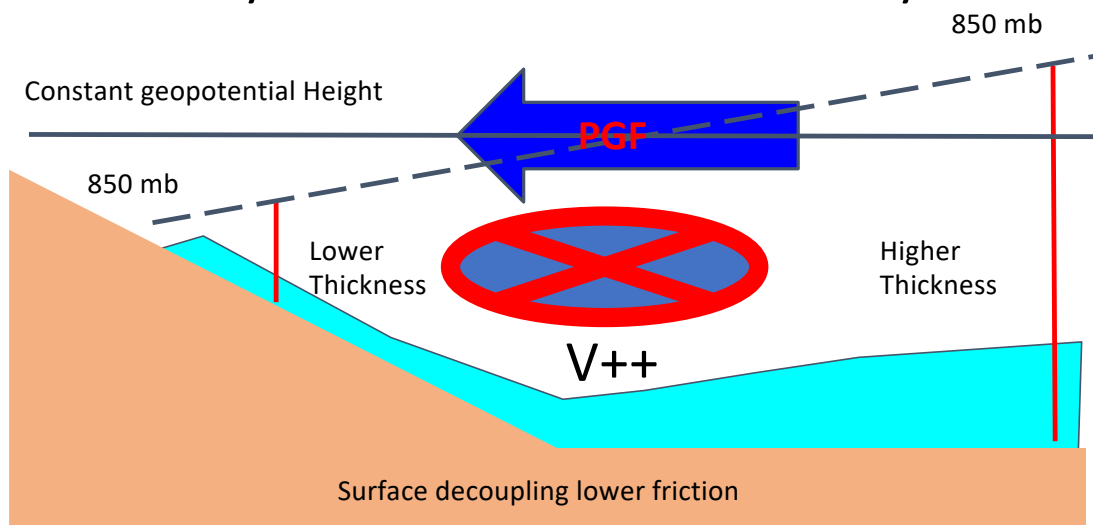


# Nocturnal Low-Level Jet (LLJ)

- Plains often see early nocturnal low-level jet (LLJ) ramp up with lee cyclogenesis
- Related to two primary factors:
  - Boundary-layer decoupling and loss of surface friction (“inertial oscillation”)
  - Diurnal temperature variations over sloped terrain (thermal wind)
- Part of the process that can favor late evening/early overnight tornadoes (with favorable moisture/CAPE)

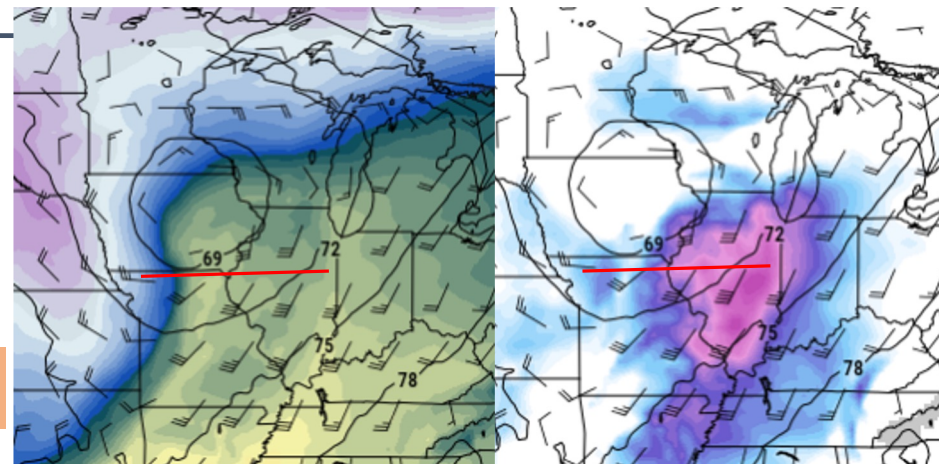
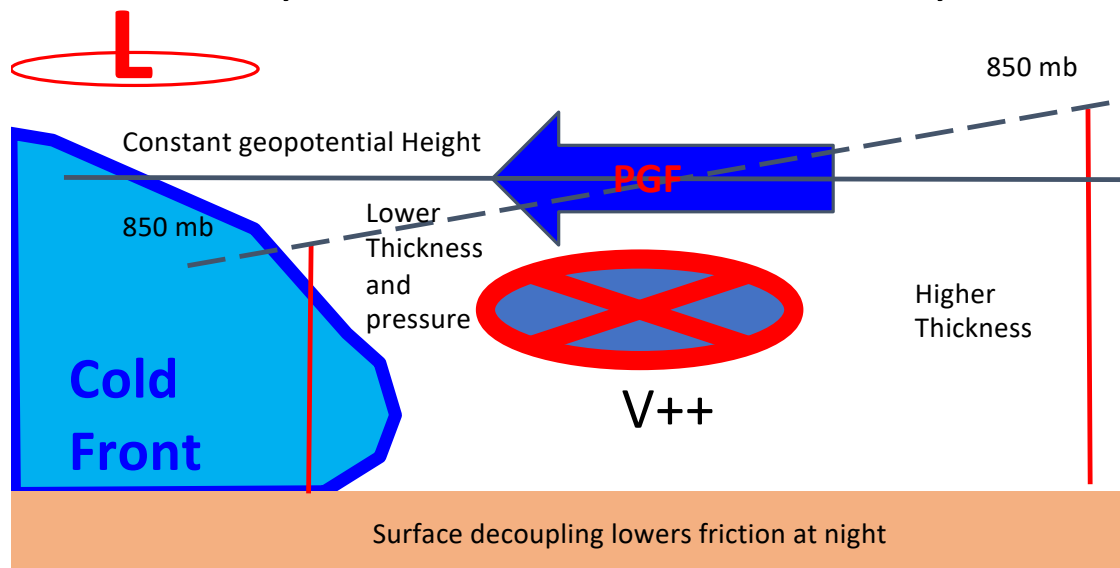
# Nocturnal Low-Level Jet (LLJ) Formation

- LLJs form in response to difference in terrain heights, heat fluxes and thickness across pressure surfaces that induces a horizontal Pressure Gradient Force.
- Coriolis forcing turns the momentum northward resulting in an enhanced southerly flow.
- Surface decoupling favors lower friction and acceleration of the jet above the surface
- Lee Cyclones can also enhance it by increasing the PGF



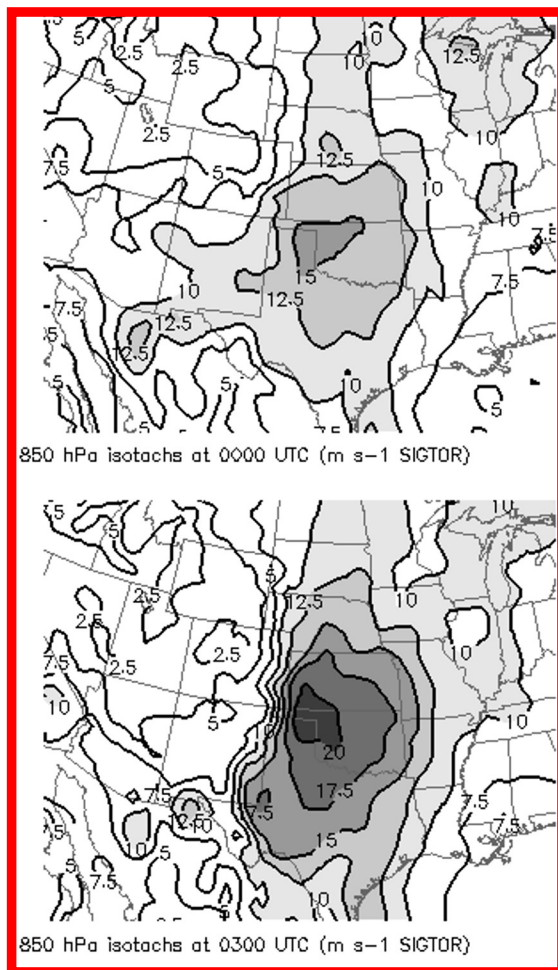
# Cyclone induced Low-Level Jet (LLJ) Formation

- LLJs form in response to difference in heat fluxes and thickness across pressure surfaces that induces a horizontal Pressure Gradient Force.
- Coriolis forcing turns the momentum northward resulting in an enhanced southerly flow.
- Surface decoupling favors lower friction and acceleration of the jet above the surface
- Lee Cyclones can also enhance it by increasing the PGF.

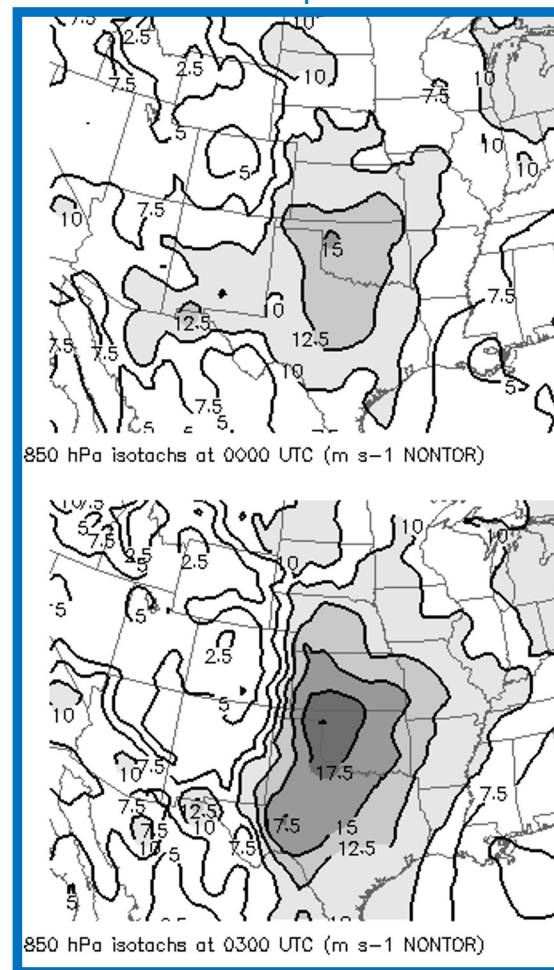




### EF2+ supercell events



### Nontornadic supercell events



# Lee Cyclogenesis Summary

- Lee cyclogenesis occurs where upper-level flow is perpendicular to terrain features (westerly flow for Rockies)
- Lee cyclogenesis more robust with lower static stability
- Lee cyclone helps drive low-level moisture return AND vertical windshear (more later)
- Nocturnal low-level jet associated with lee cyclone is driven by inertial oscillation and thermal wind → increased low-level shear and tornado threat in evening and after dark if adequate CAPE