

Severe Weather Ingredients

**Material by Tom Galarneau Andrew Lyons,
Rich Thompson and Harry Weinman**

Ingredients for Organized Severe Thunderstorms

- **Instability (Lapse Rates)**
- **Lift (QG, mesoscale, convective scale)**
- **Moisture (return flow, soil moisture, evapotranspiration)**
- **Vertical wind shear**

Ingredients for Organized Severe Thunderstorms

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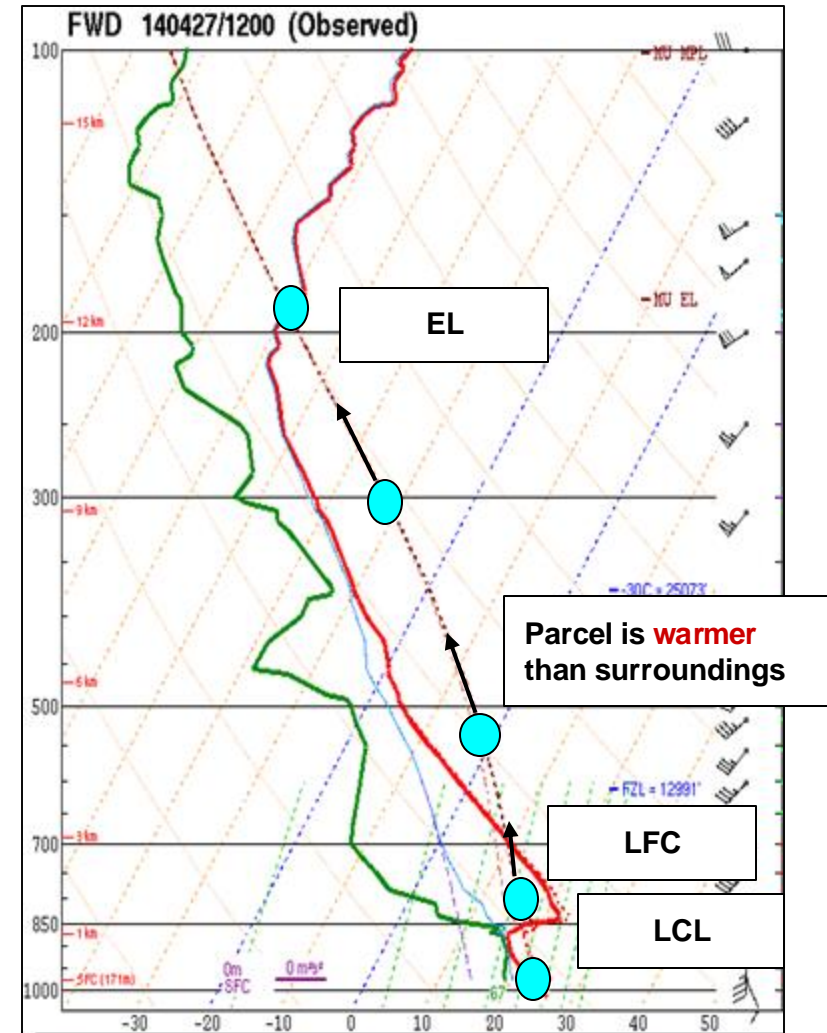
This lecture

- a separate lecture

Key ingredient for
organized severe storms

Instability: The Basics

- Give a parcel a nudge and see what happens!
- We determine an environments stability by its lapse rate (Change of environmental temperature with height)

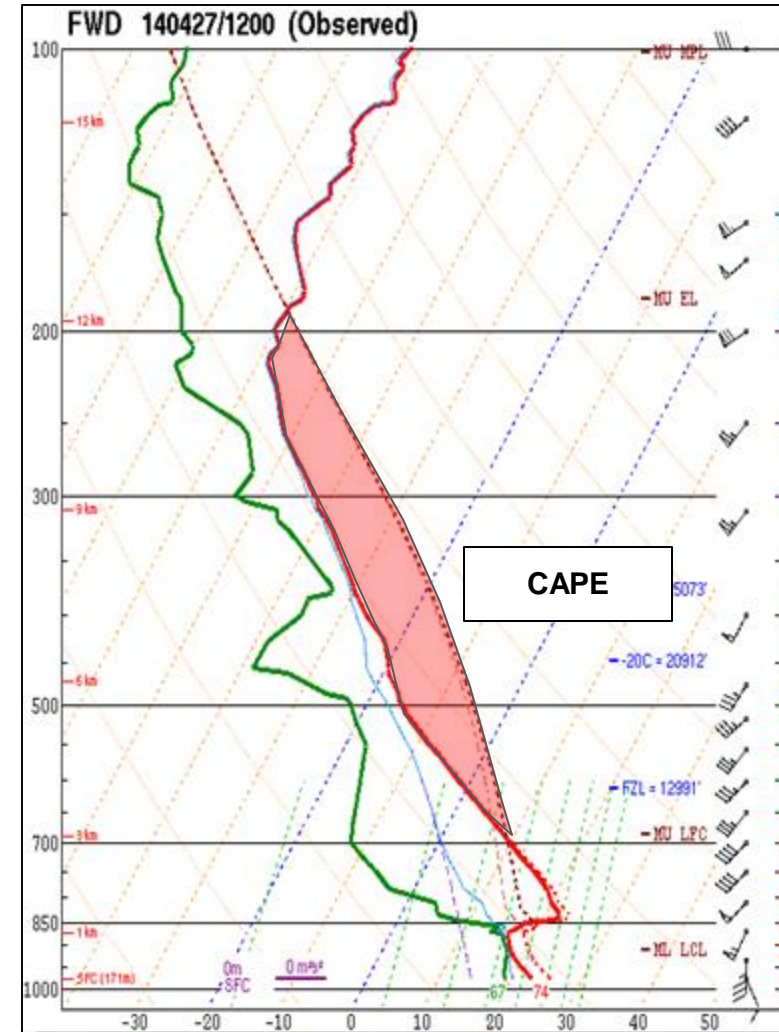


Instability: The Basics

- Integrate the depth where parcel is warmer/less dense than its surrounding
- CAPE (Convective Available Potential Energy)

$$CAPE = \int_{LFC}^{EL} g \frac{T_{v_p} - T_{v_{env}}}{T_{v_{env}}} dz = \left[\frac{kg \ m^2}{s^2} \frac{1}{kg} \right] = \left[\frac{m^2}{s^2} \right]$$

$= \left[\frac{J}{kg} \right]$

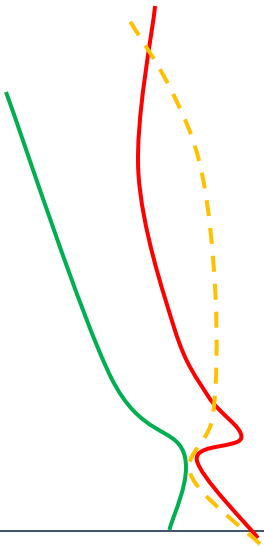


Instability: The Basics

Take a small finite layer depth
within CAPE profile

$$\frac{CAPE}{\Delta Z} = \left[\frac{m^2}{s^2} \right] \left[\frac{1}{m} \right] = \left[\frac{m}{s^2} \right] \longrightarrow$$

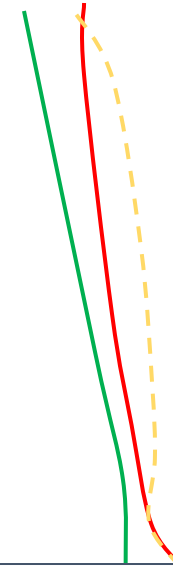
Upward
Acceleration!



Short/thick CAPE:

Drives larger upward acceleration

1. Offsets negative effects of dry air entrainment into updraft
2. Mitigates effects of precipitation drag on updraft

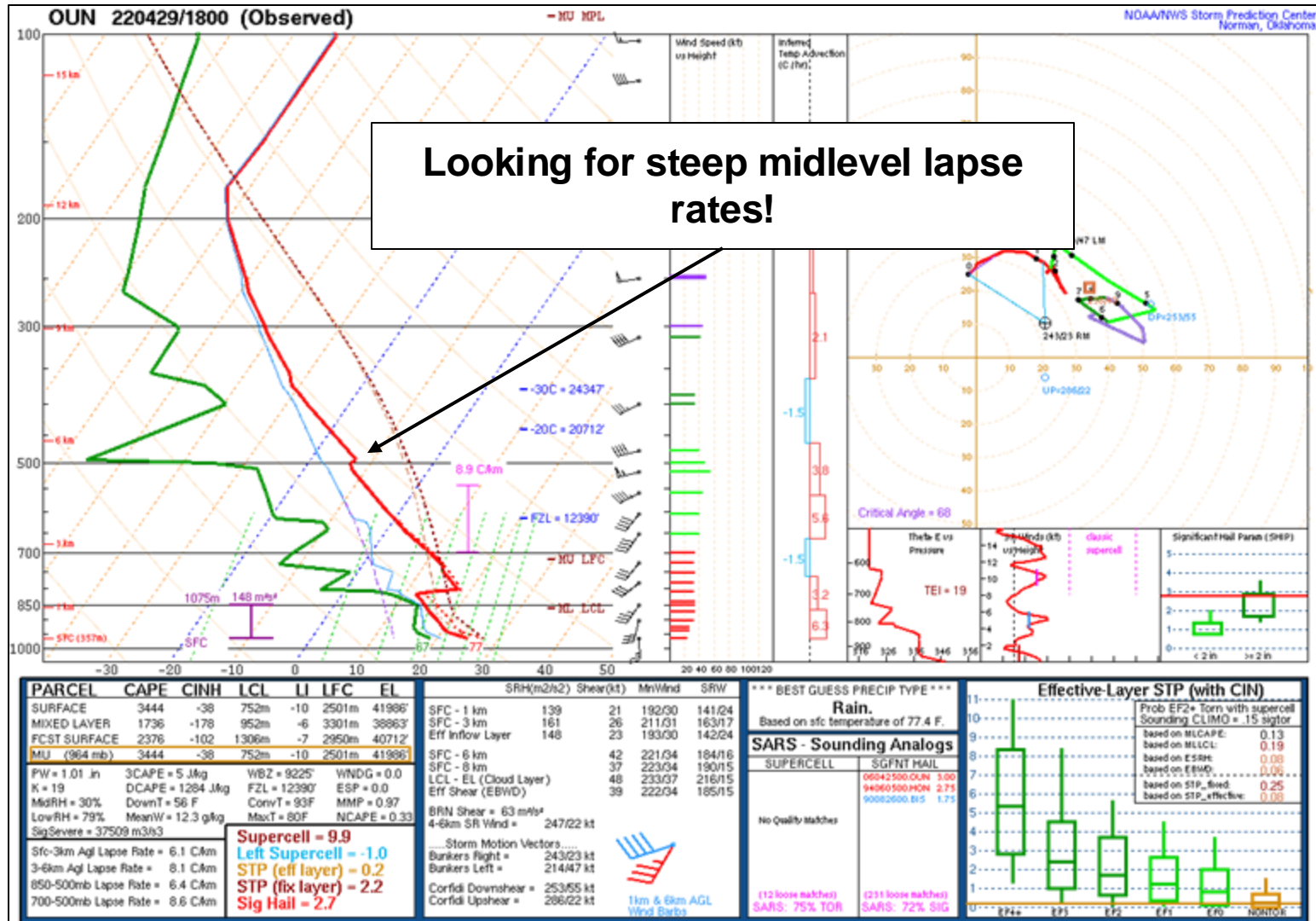


Tall/narrow CAPE:

Smaller upward acceleration

Water loading and entrainment more substantially weaken updraft

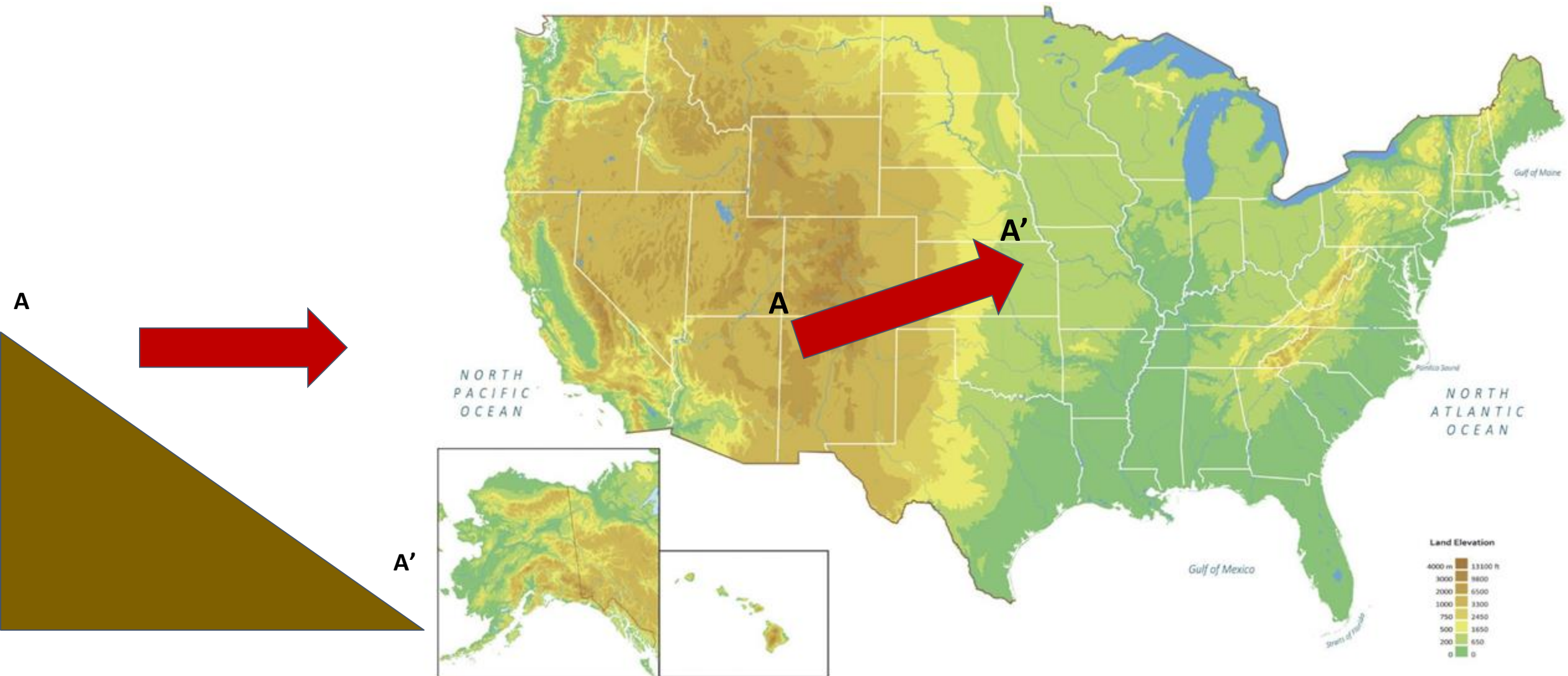
So, how do we get wider CAPE?



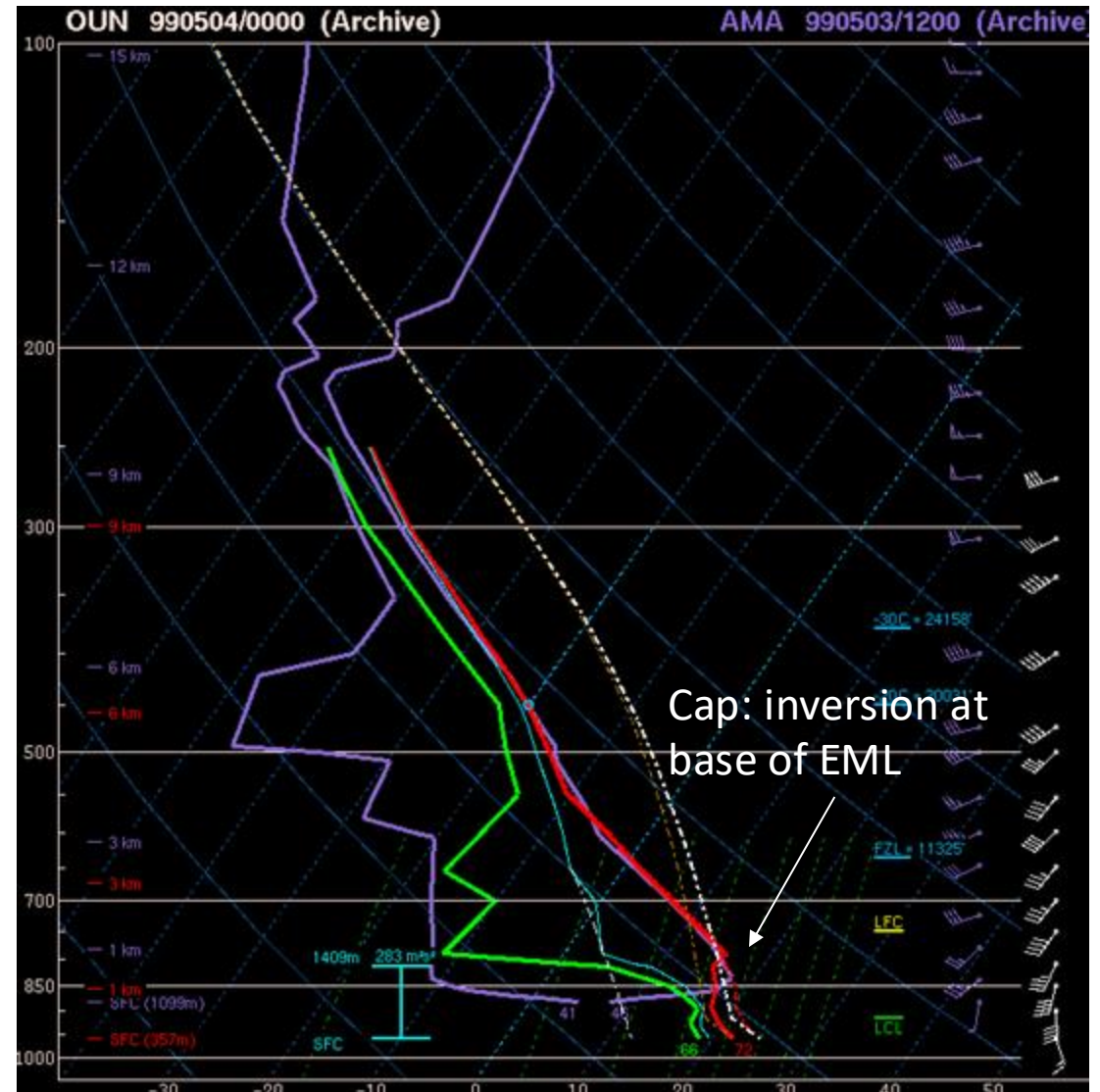
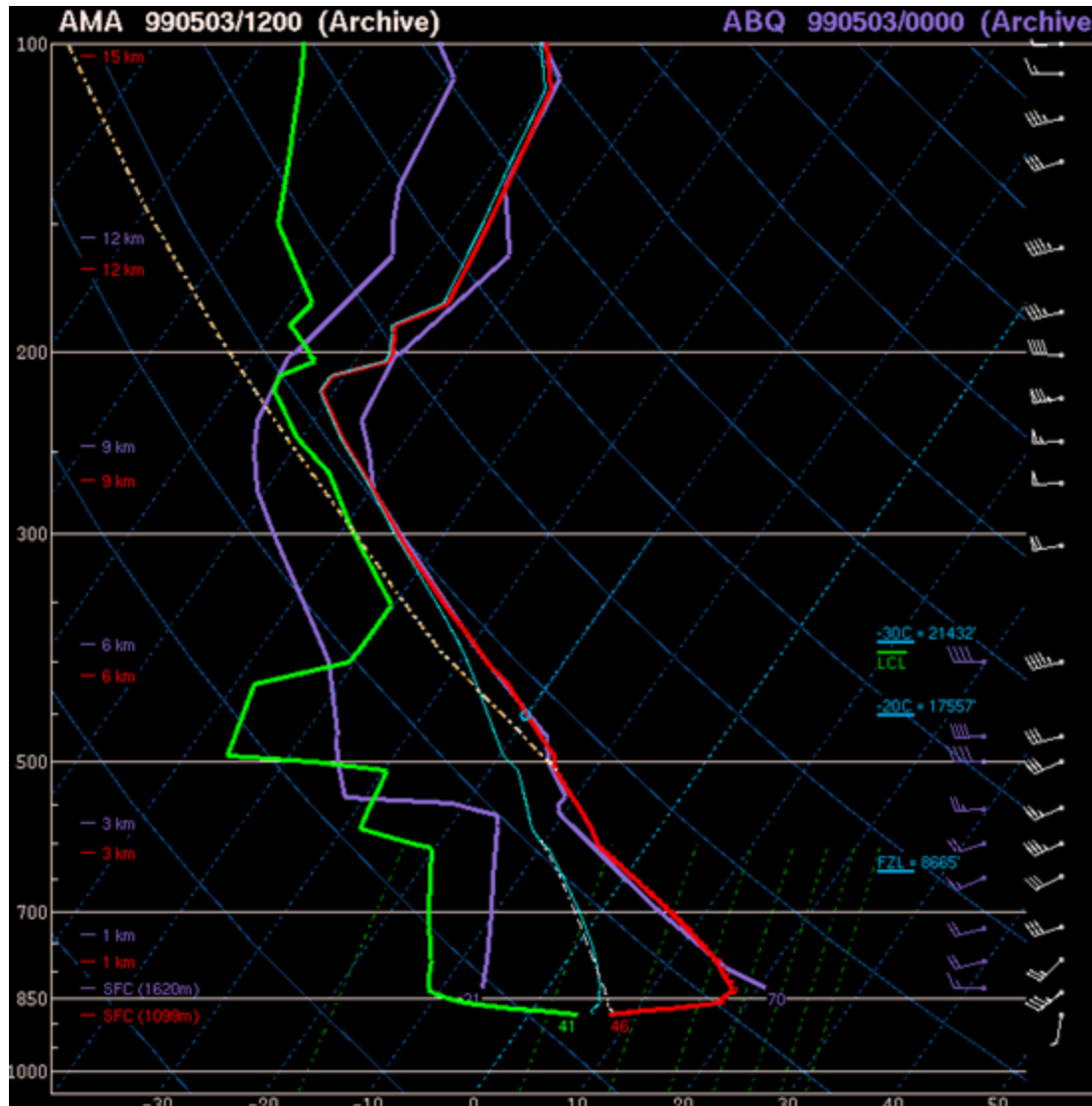
Instability: Elevated Mixed Layer (EML)

- **Deep mixed layer forms over Rockies and Sierra Madre in response to surface heating**
- **Stronger heating and/or cooler temperatures aloft results in deeper mixed layer**
- **Mixed layer is advected eastward and becomes elevated east of the Rockies and Sierra Madre (really, east of the dryline)**
- **Differential advection (eastward advection of EML above northward advection of moisture from GoM) created “loaded gun” profile**

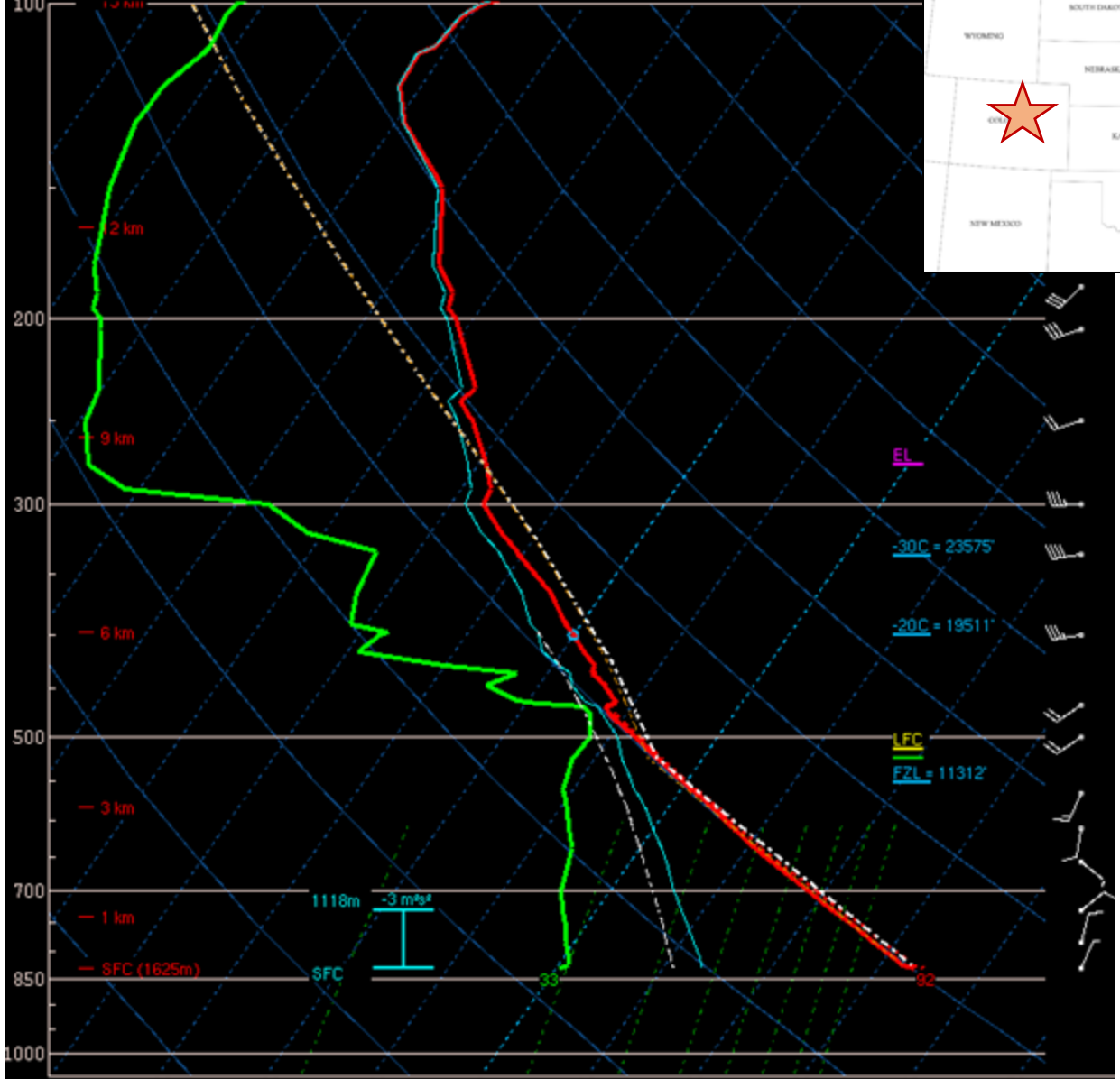
Elevated Mixed Layer (EML) Differential Advection



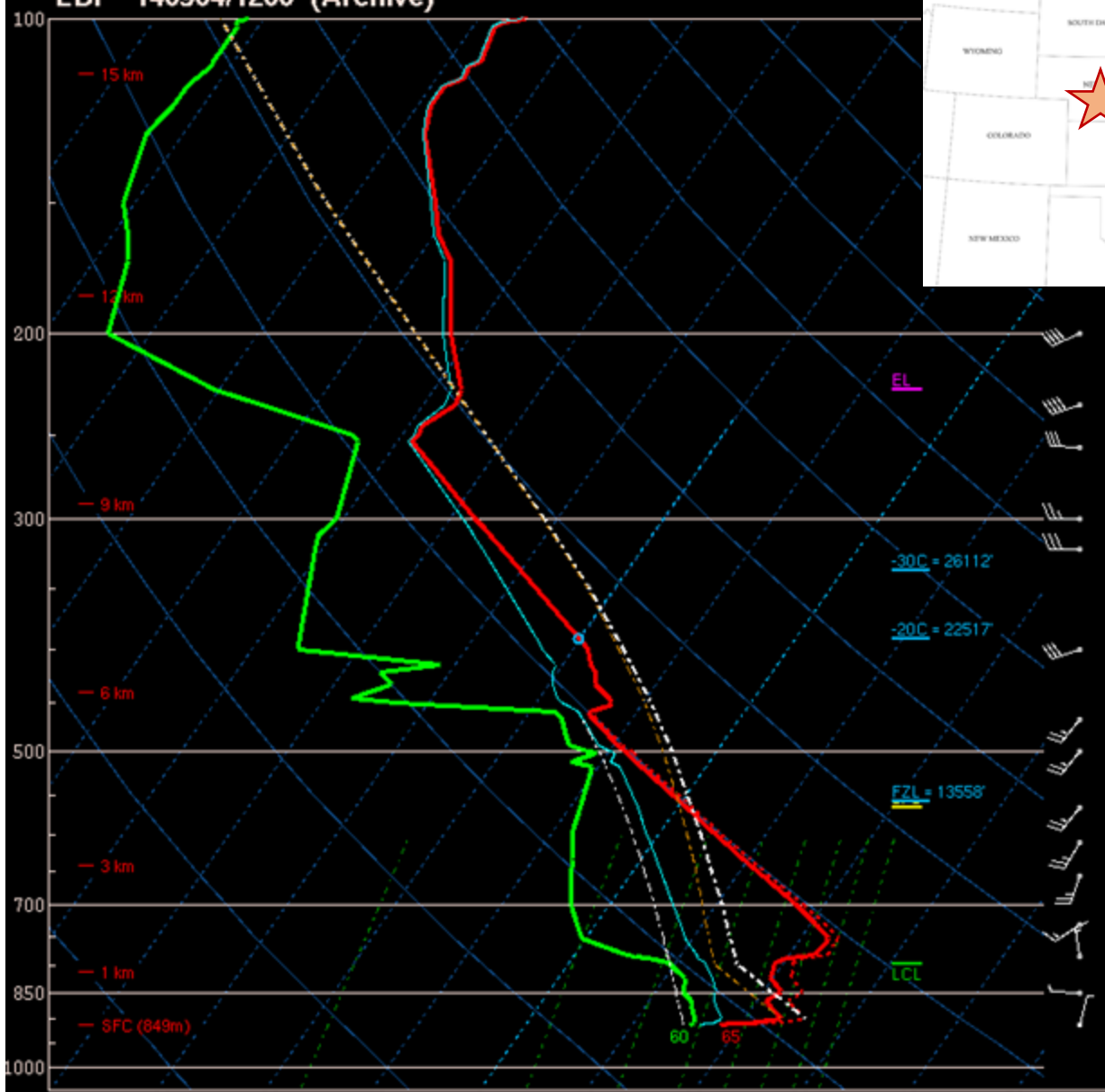
EML at OUN (0000 UTC 4 May 1999)



DNR 140904/0000 (Archive)

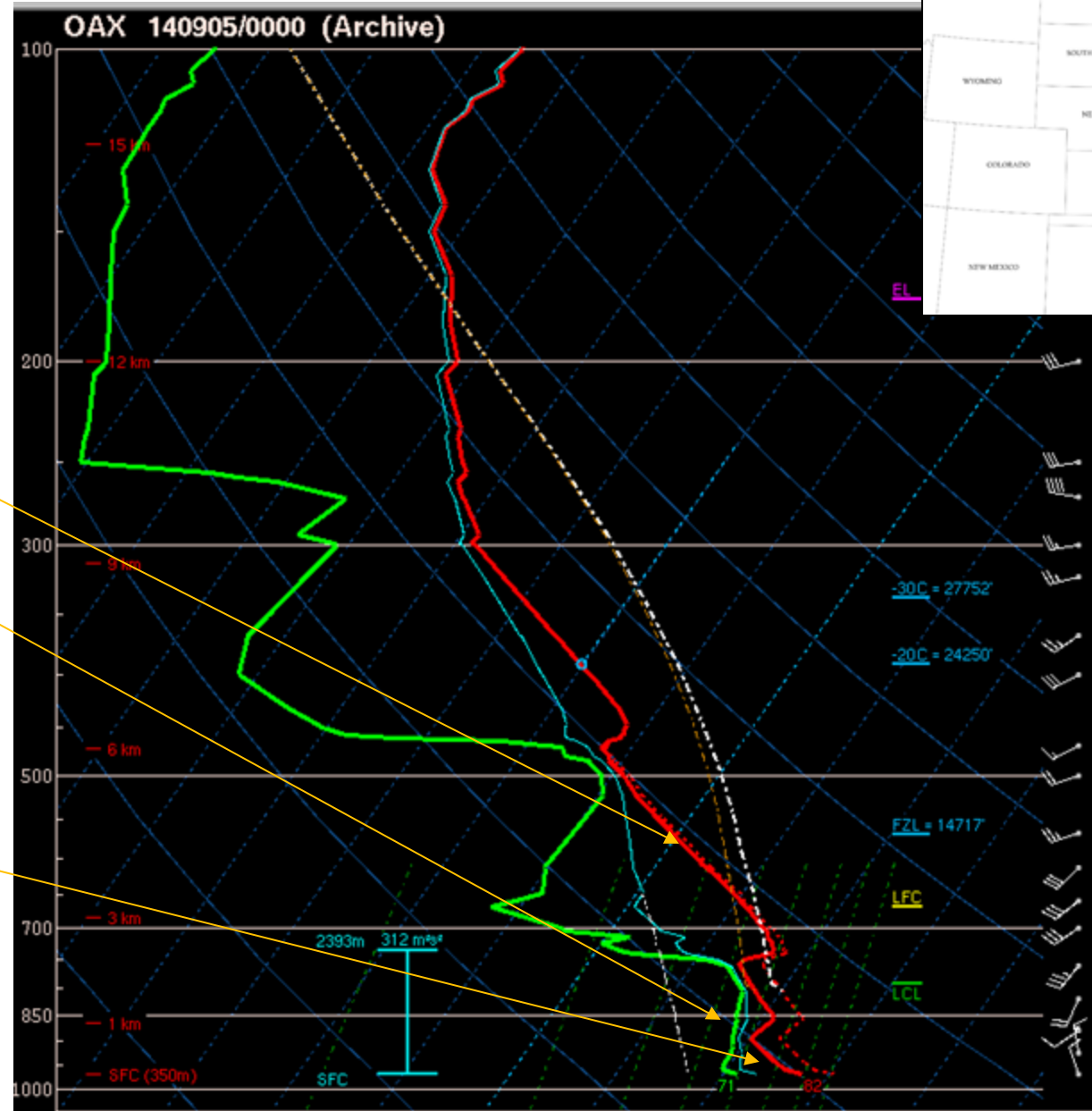


LBF 140904/1200 (Archive)



Cap strength is related to:

- EML θ
- Depth/quality of moisture
- Surface heating and mixing in moist sector



Lift

- **Synoptic-scale lift**

- 1-10 cm/s for many hours (6 hours ~2 km ascent (SFC-800 hPa))
- 6 hours to reach LFC
- Conditions environment for deep, moist convection over broad area
 - Lapse rate stretching
 - Removal of inhibition
- Layered clouds in regions of isentropic lift; some elevated storms can form

- **Mesoscale lift**

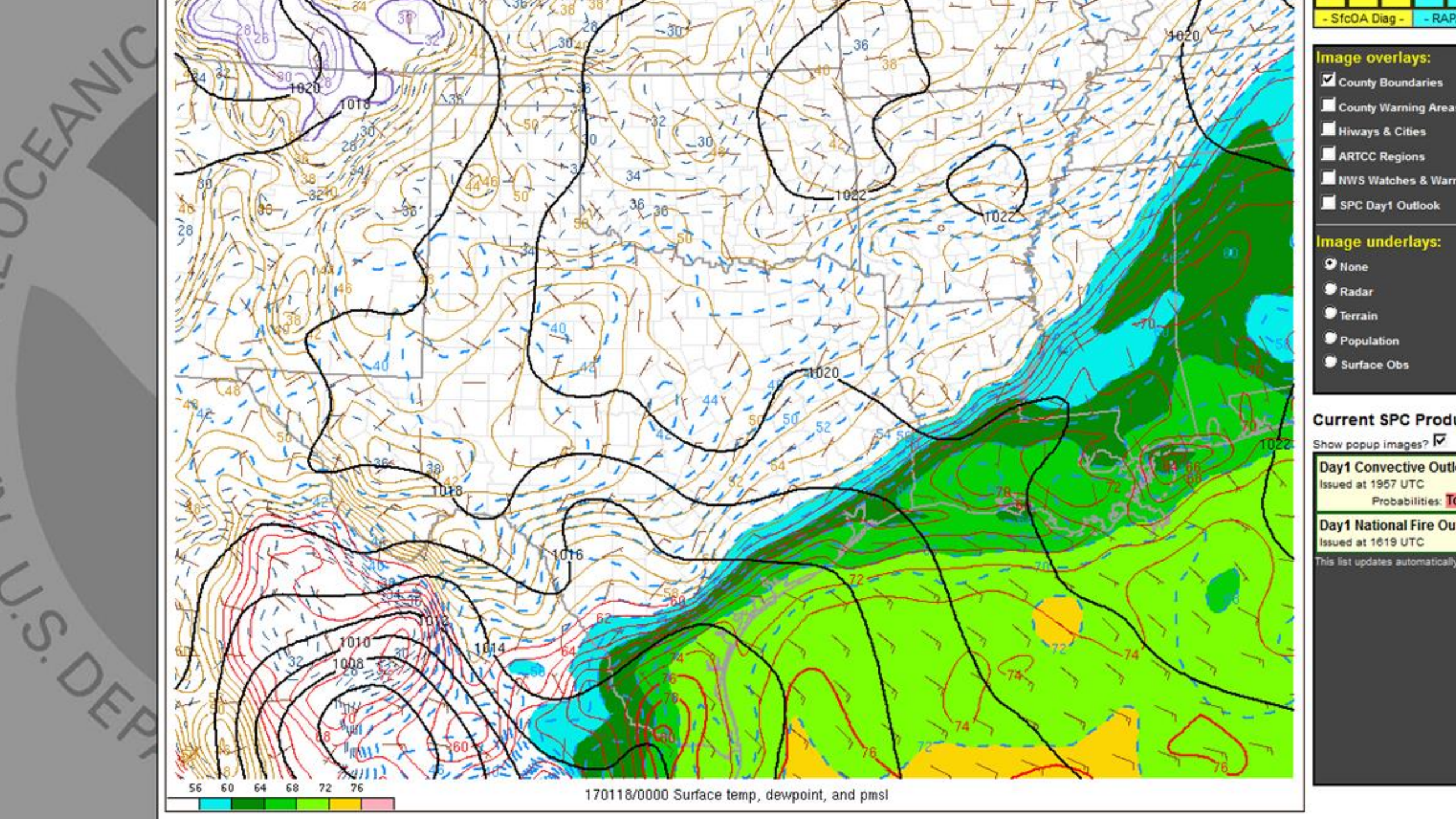
- 1 m/s for minutes to hours (1 hour ~3.6 km ascent (SFC-650 hPa))
- < 1 hour to reach LFC
- Narrow zones of ascent along boundaries and terrain features
- Direct initiation of thunderstorms

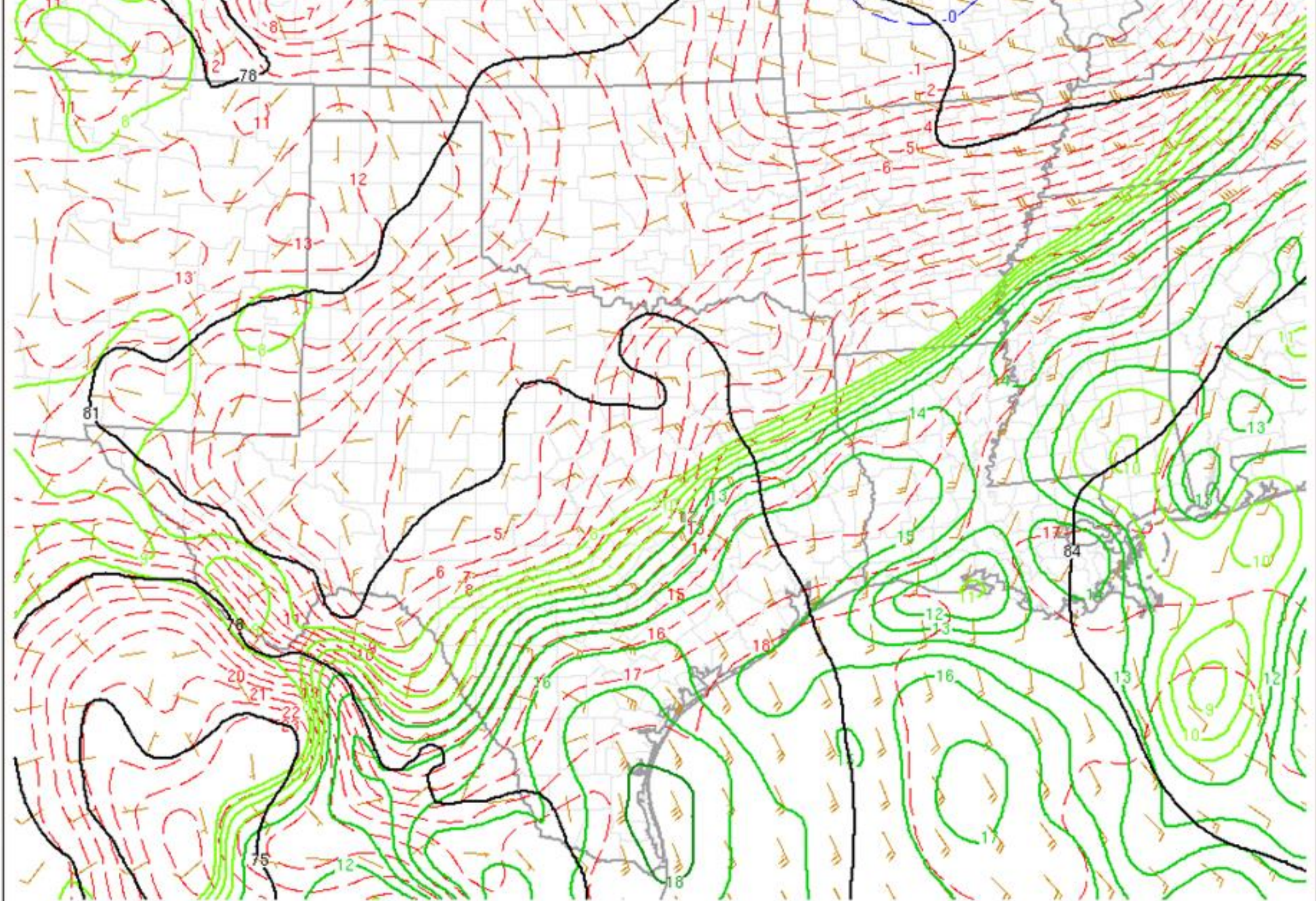
- **Storm-scale lift**

- 10 m/s for minutes (15 mins ~9 km of ascent (SFC-300 hPa))
- 5 minutes to reach LFC
- Storm maintenance and propagation (supercells, MCSs, squall lines, etc.)

Synoptic-Scale Lift

- **QG sources**
 - Warm advection/isentropic lift and differential vorticity advection
- **Jet streaks (also QG)**
 - Straight jet: ascent in right entrance and left exit regions
 - Curved jet: ascent downstream and along jet core





170118/0000V001 925mb hght/temp/dwpt/wind

- SfcOA Diag - - RAP

Image overlays:

- ☒ County Boundaries
- ☐ County Warning Area
- ☐ Hiways & Cities
- ☐ ARTCC Regions
- ☐ NWS Watches & Warnings
- ☐ SPC Day1 Outlook

Image underlays:

- ☒ None
- ☐ Radar
- ☐ Terrain
- ☐ Population
- ☐ Surface Obs

Current SPC Product

Show popup images? ☒

Day1 Convective Outlook

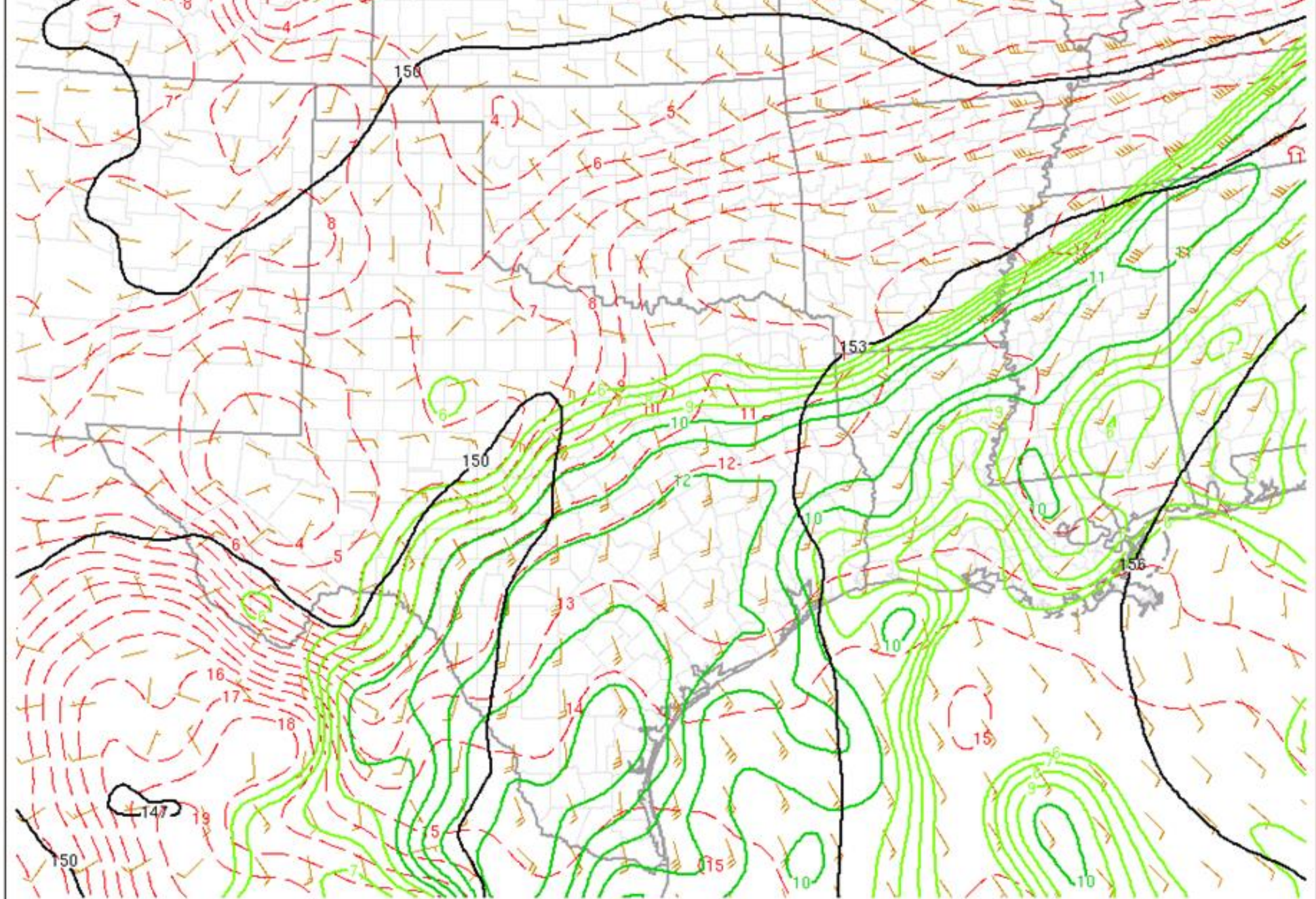
Issued at 1957 UTC

Probabilities: To

Day1 National Fire Outlook

Issued at 1819 UTC

This list updates automatically



170118/0000V001 850mb hght/temp/dwpt/wind

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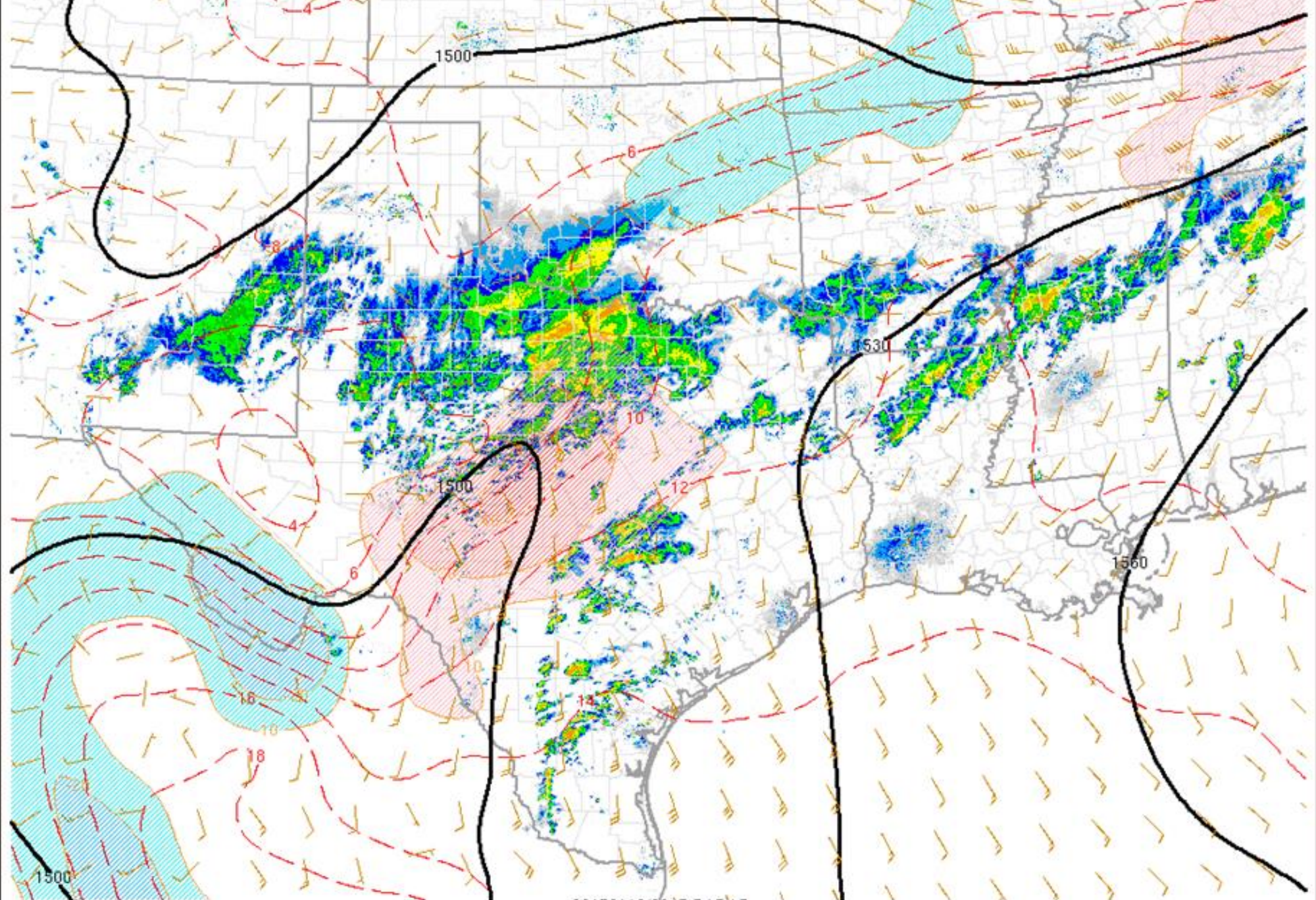
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Current SPC Product

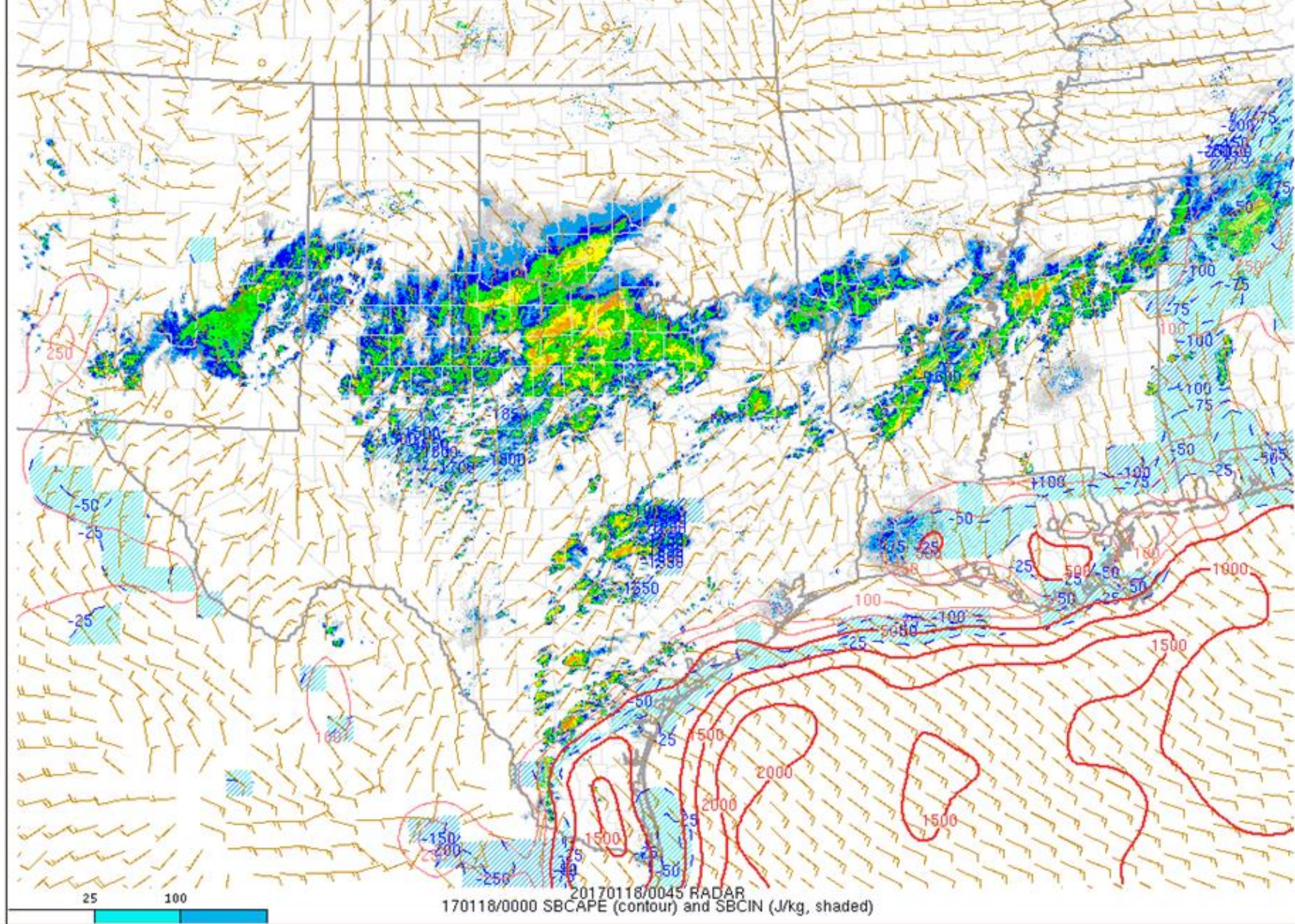
Show popup images? ☒

Day1 Convective Outlook

Issued at 1957 UTC
Probabilities: To

Day1 National Fire Outlook

Issued at 1819 UTC
This list updates automatically



- SfcOA Diag - - RAP

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Day1 Convective Outlook

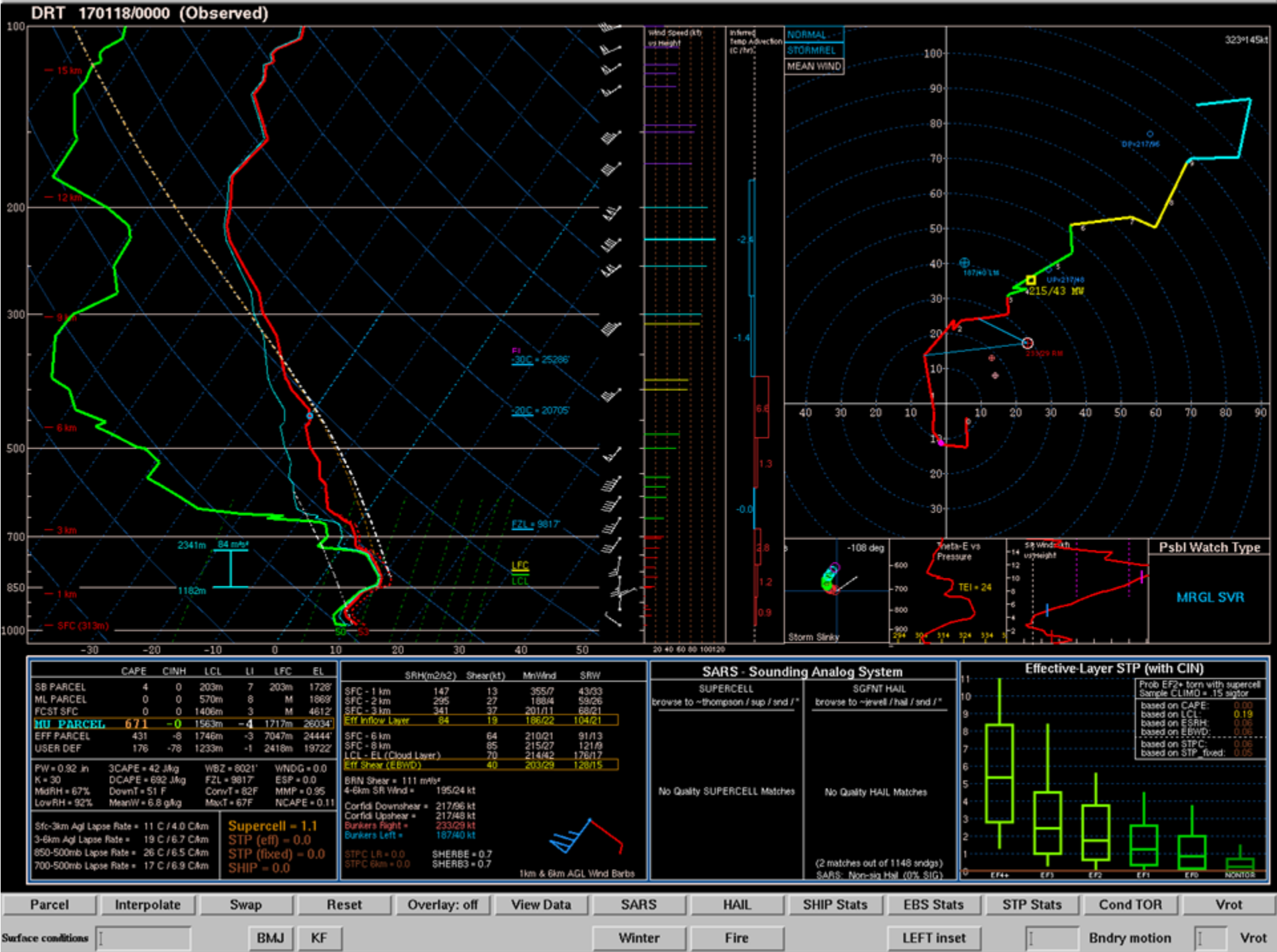
Issued at 1957 UTC

Probabilities: **TO**

Day1 National Fire Outlook

Issued at 1619 UTC

This list updates automatically



Synoptic-Scale Lift Summary

- **Gradual ascent over many hours**
 - Warm advection and differential vorticity advection are primary drivers
- **Primary role of QG ascent is to precondition environment for convection**
 - Can also initiate elevated storms in regions of warm advection
 - Most surface-based storms are triggered by mesoscale ascent

Mesoscale Lift

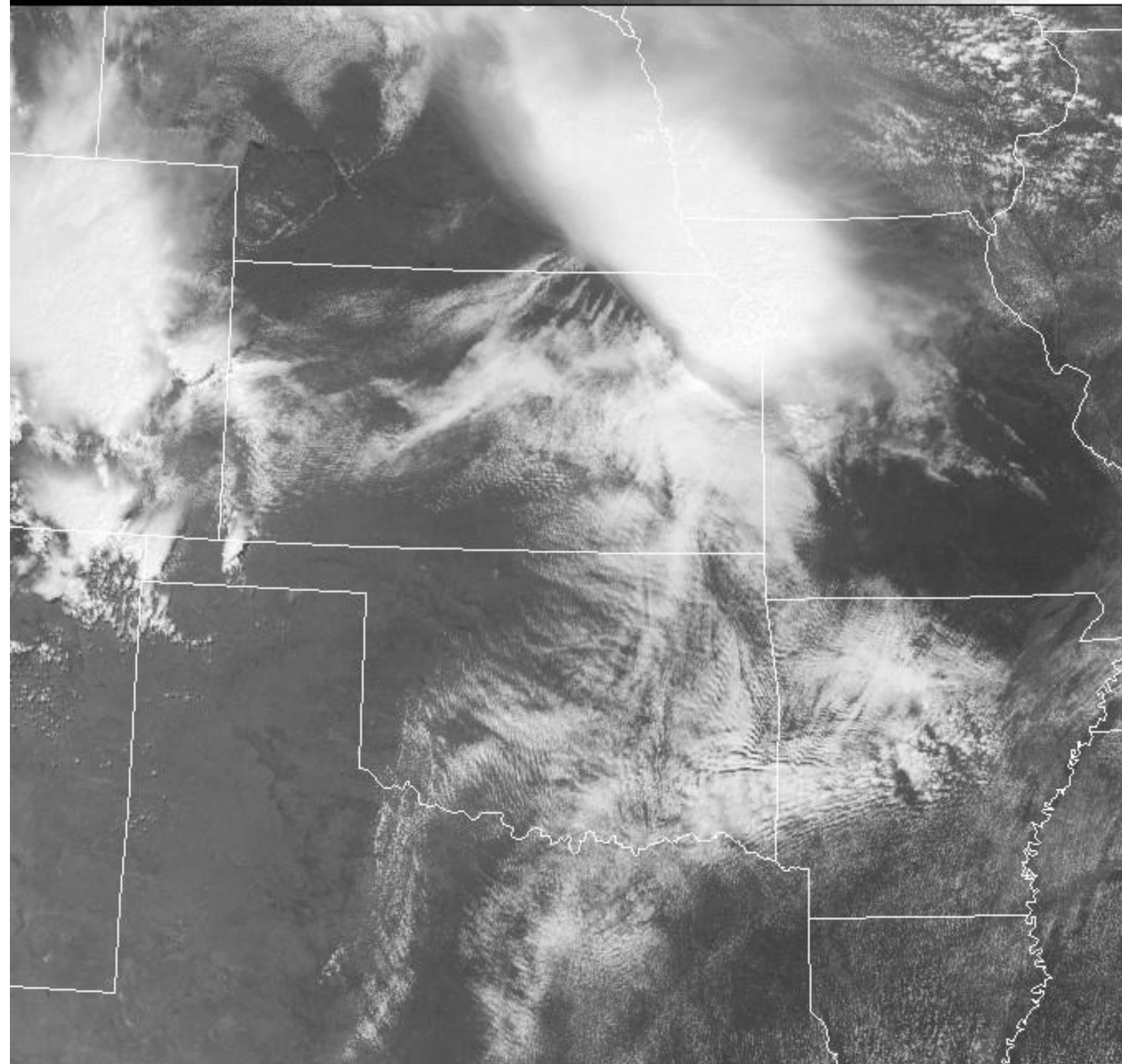
- Focused and narrow zones of lift for minutes to hours
- Fronts
 - Rising on warm side with frontogenesis
 - Isentropic ascent on sloped warm front
- Outflow boundary
 - Lift is governed by depth/strength of cold pool and low-level vertical wind shear [Rotunno-Klemp-Weisman (RKW) Theory]
- Dryline
 - Lift on dry/hot side (lower density air)
 - Lift governed by depth of mixing west of dryline and depth of moist layer east of dryline

1740 UTC Tue 24 May 2011

Visible Satellite

<http://adds.aviationweather.gov>

0 6 9 12 15 18 21 24 27 30 33 36 39 42 45 48 51 54 57 60 63 66 69 72 75 78 81 84 87 90 93 96 99

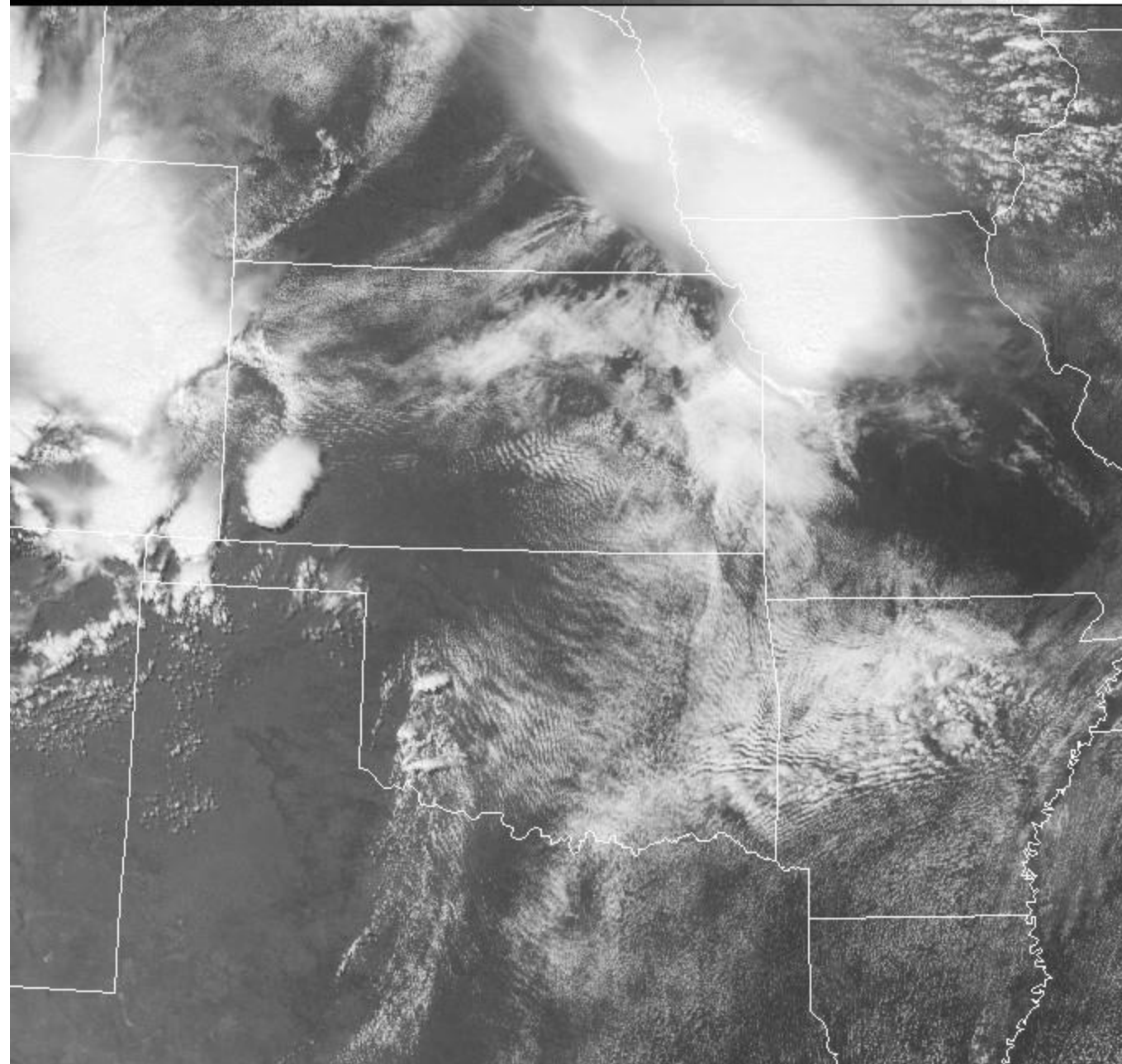


1845 UTC Tue 24 May 2011

Visible Satellite

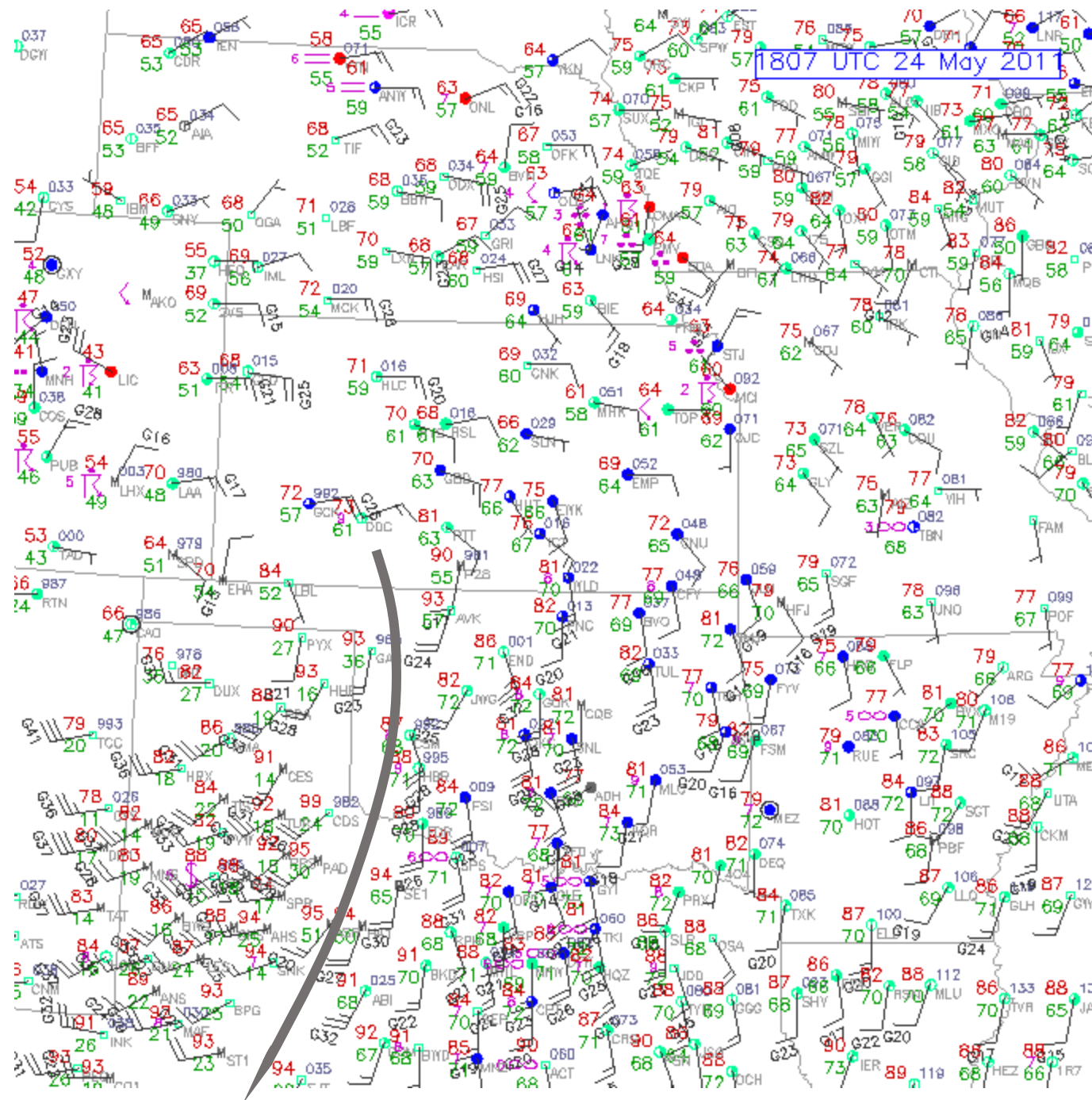
<http://adds.aviationweather.gov>

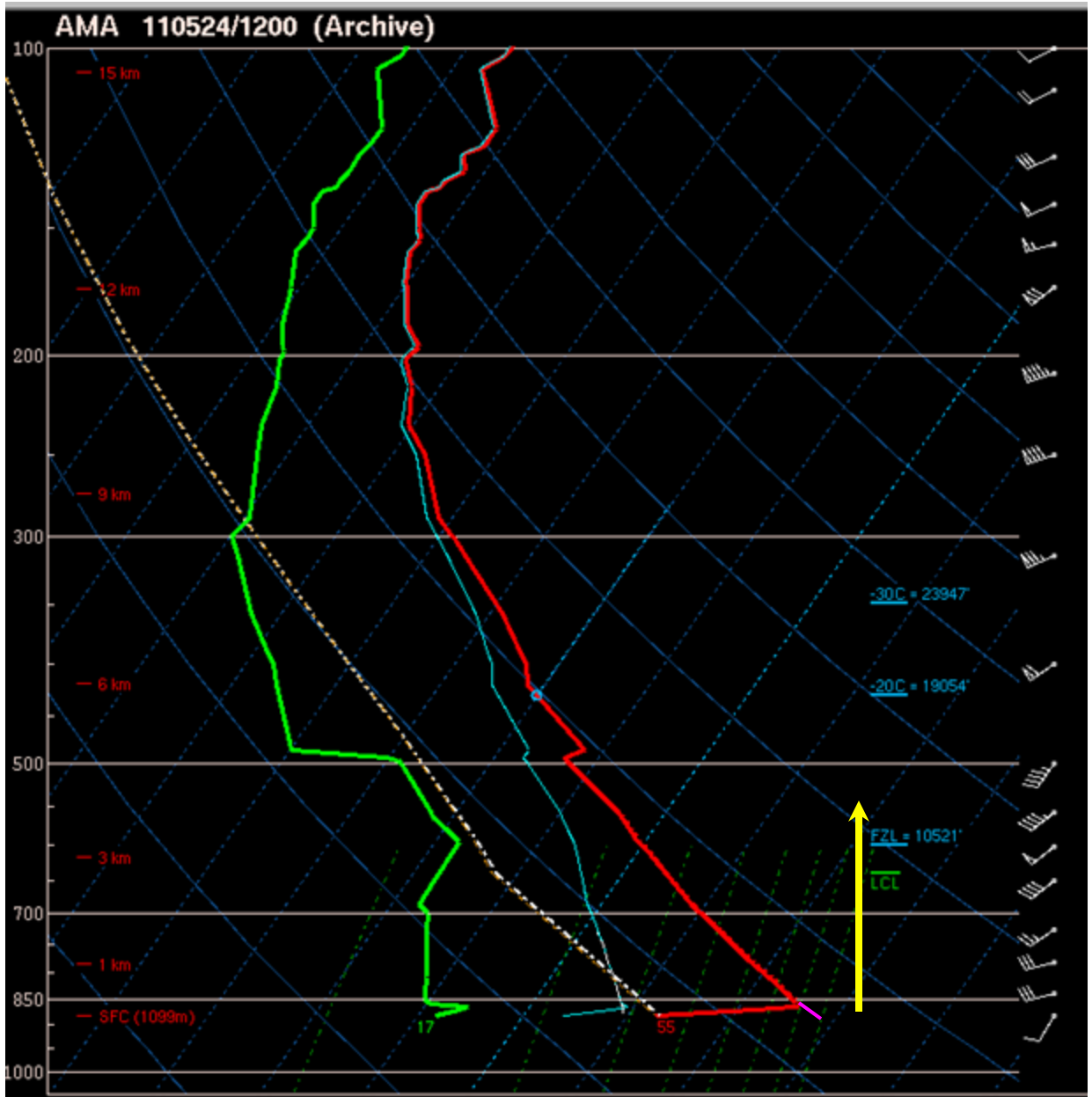
0 6 9 12 15 18 21 24 27 30 33 36 39 42 45 48 51 54 57 60 63 66 69 72 75 78 81 84 87 90 93 96 99

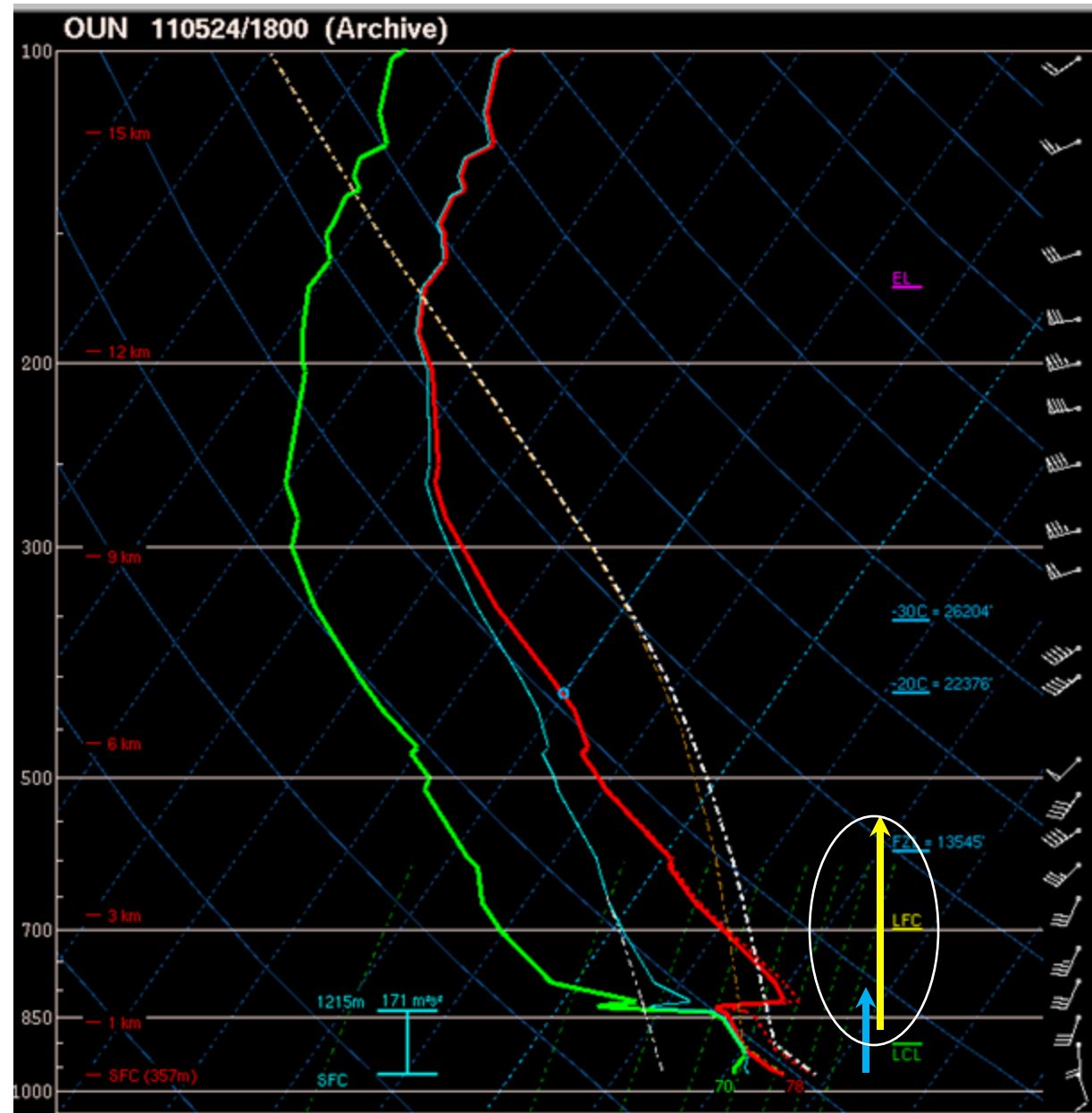


Dryline

- **Focused ascent along dryline can be ~ 1 m/s (~ 1 hour to reach LFC)**
 - Lift depends on depth of mixing west of dryline and depth of moist layer ahead of dryline
 - Maximum vertical motion along dryline scales with height of moist layer
- **Convection initiation linked to the residence time of air in the zone of ascent**
 - Winds above the surface parallel to dryline keeps air in zone of lift longer
- **Convection will initiate in points or bands**
 - Usually the mode of initiation for surface-based supercells in central/southern Plains

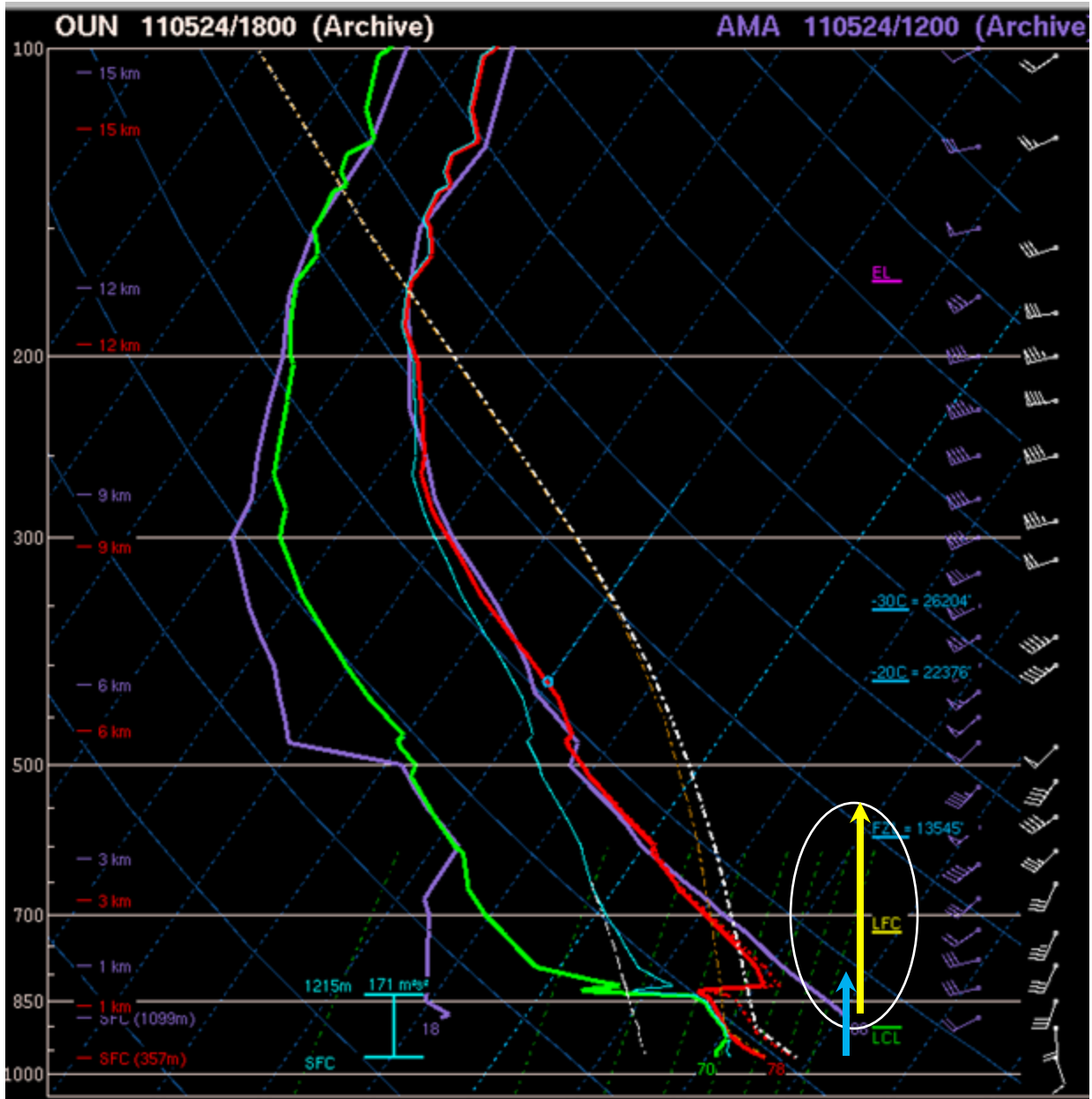


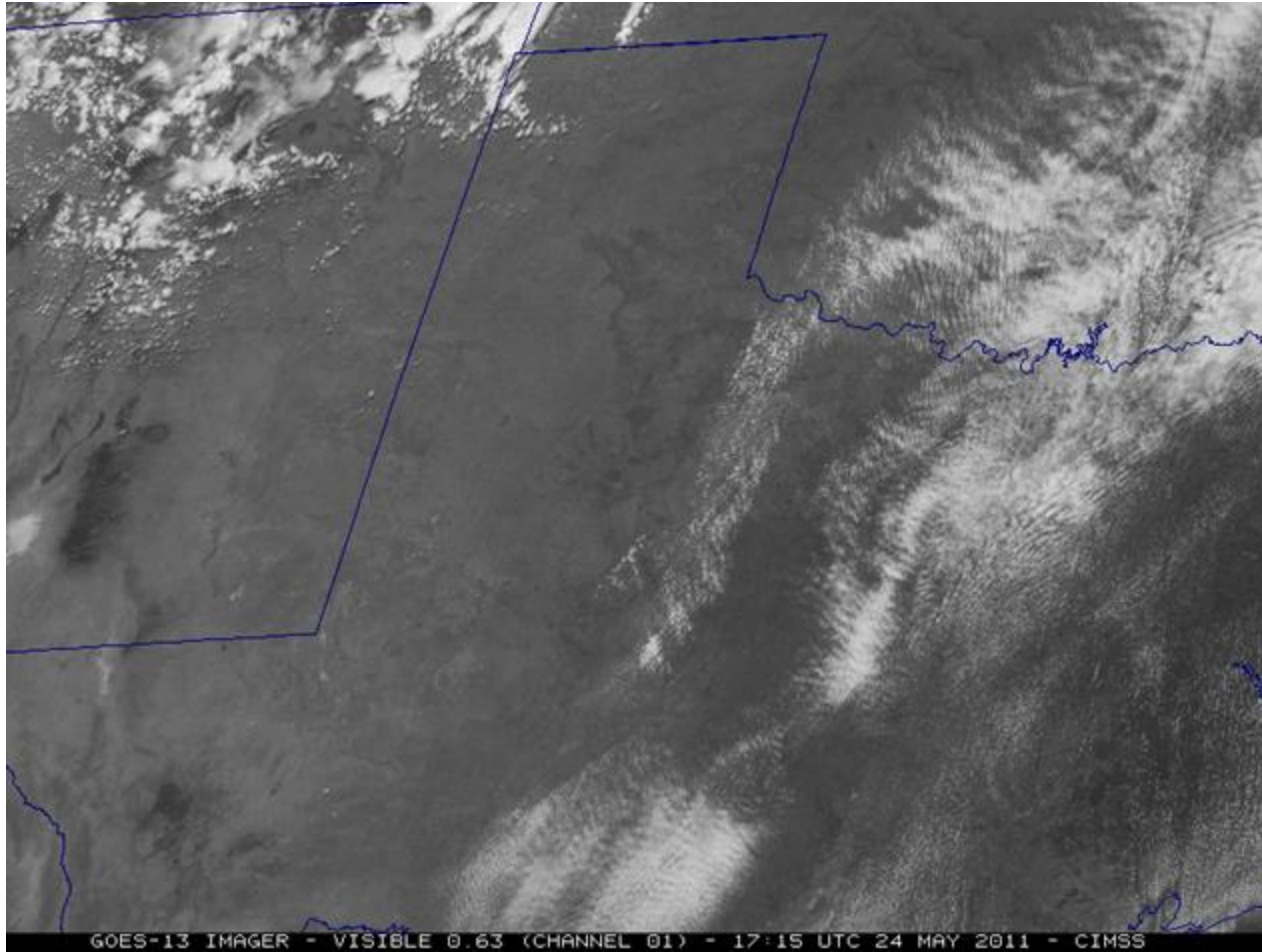




LFC is within the deep
dryline circulation

Keep parcels in the
zone of ascent – winds
within circulation SSW
(parallel to dryline)





GOES-13 IMAGER - VISIBLE 0.63 (CHANNEL 01) - 17:15 UTC 24 MAY 2011 - CIMSS

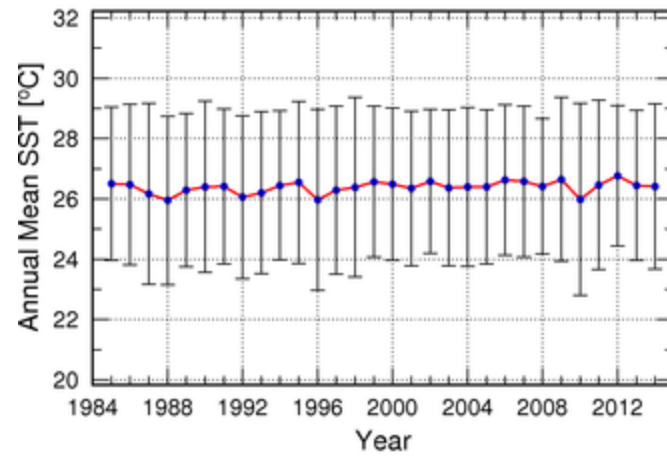
Moisture

- **Moisture in PBL needs to be large and deep enough for convection**
- **Return flow**
- **Inland vertical mixing/diurnal processes**
- **Evapotranspiration**

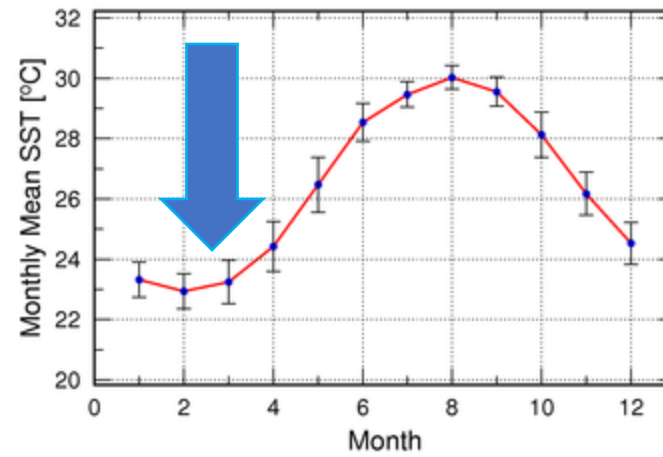
Low-level Moisture Return Flow

- Key elements include air mass modifications over warmer water and return flow trajectories
- Air mass modification is a multi-date process
- Ask yourself the following:
 - Where is the air coming from?
 - What are the underlying ocean characteristics?
 - What is the character of the returning moist layer?

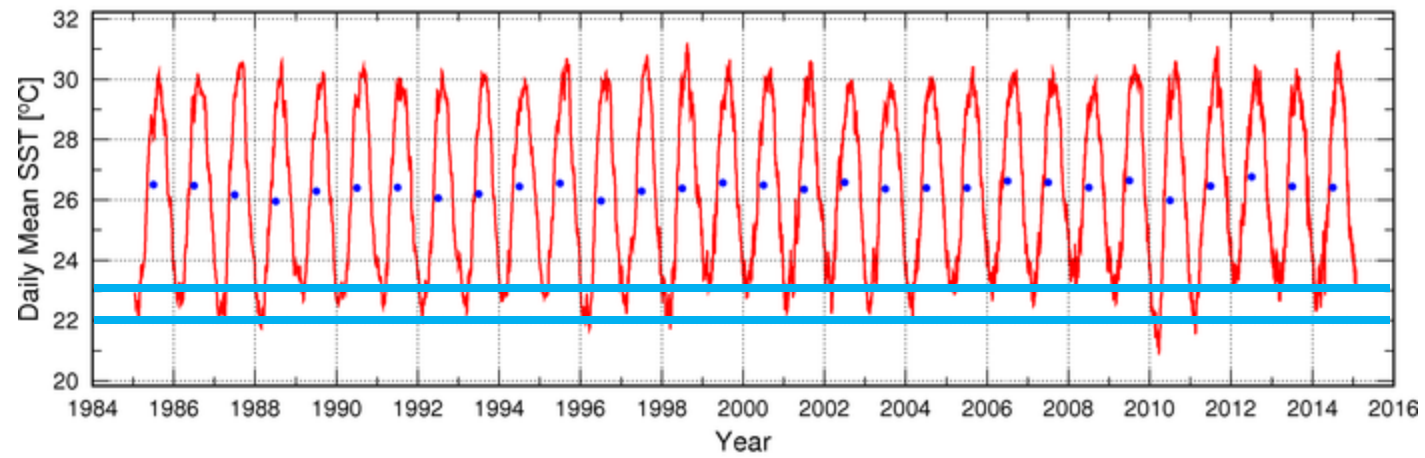
GOM Annual Mean SST



GOM Monthly Mean SST

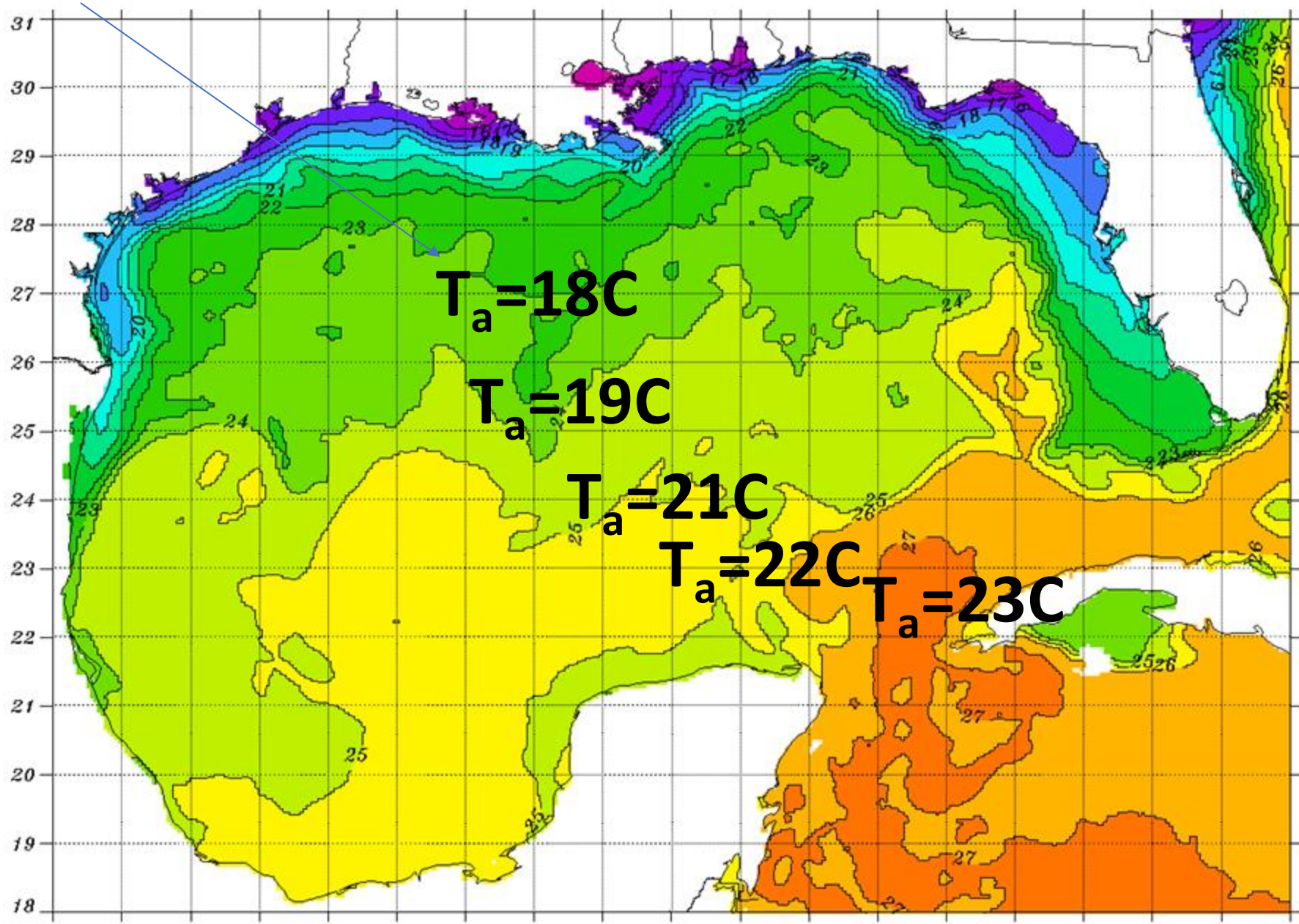


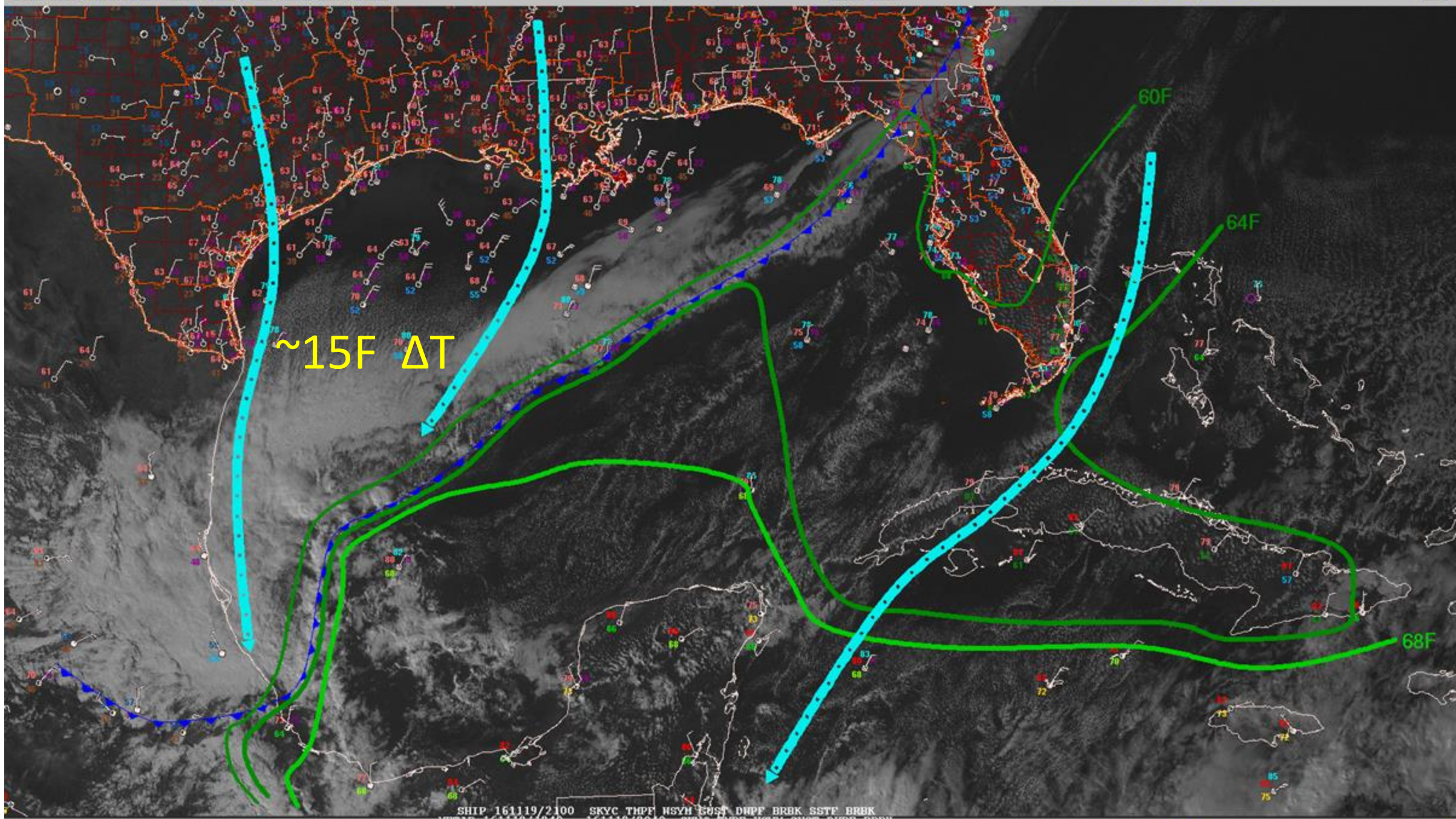
GOM Daily Mean SST

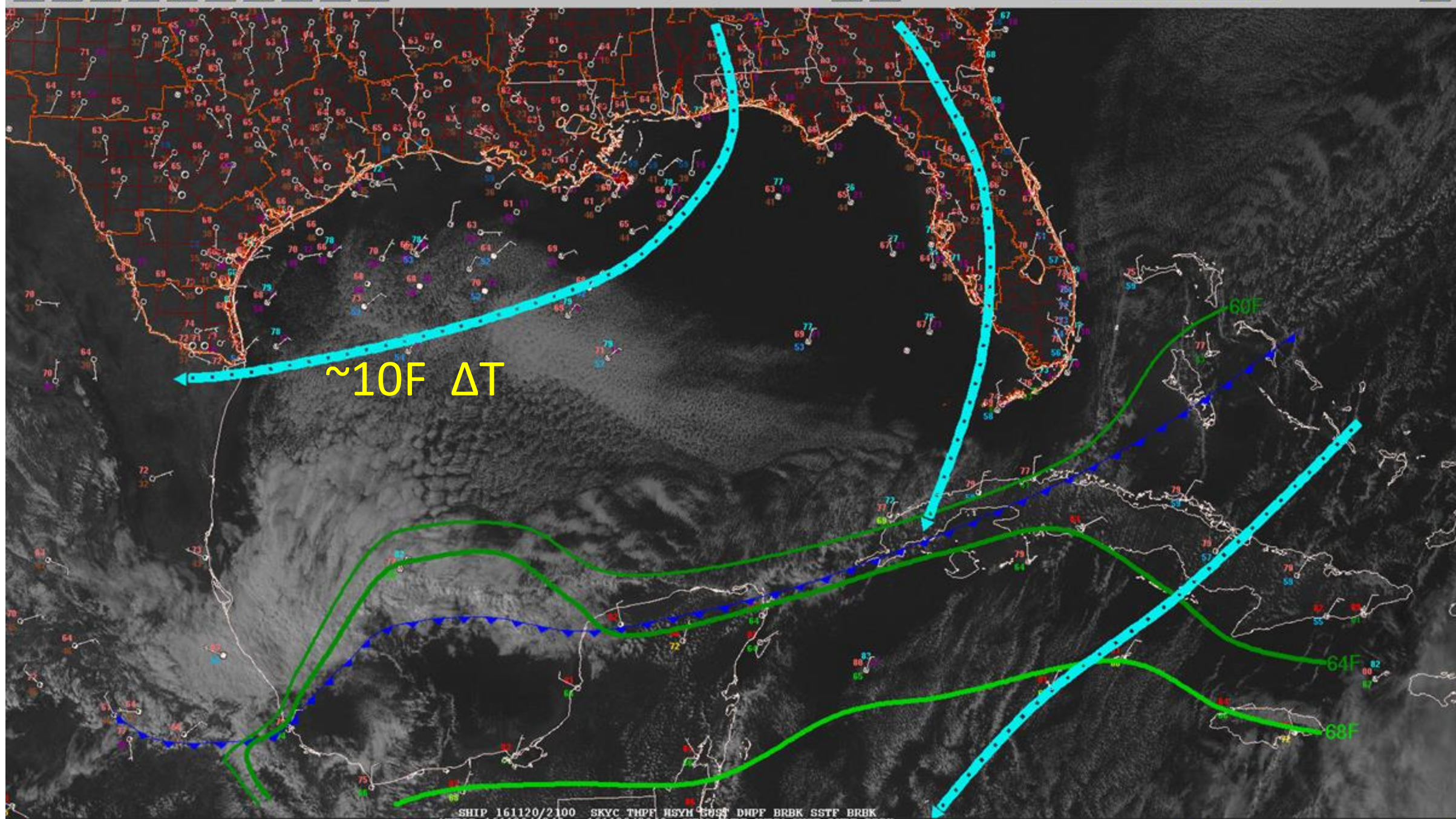


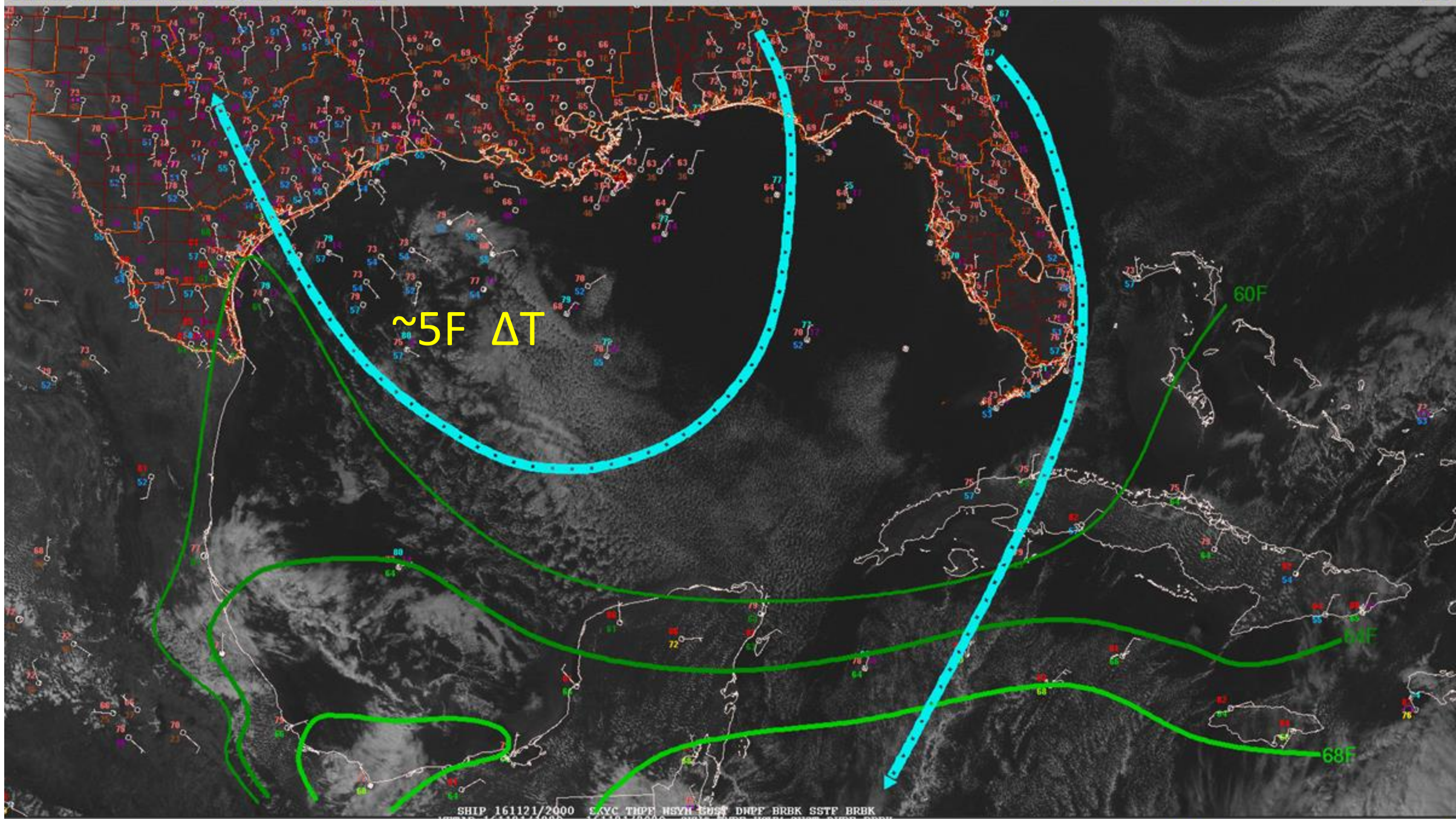
NOAA/NESDIS GEO-POLAR BLENDED 5 km SST ANALYSIS
FOR THE GULF OF MEXICO

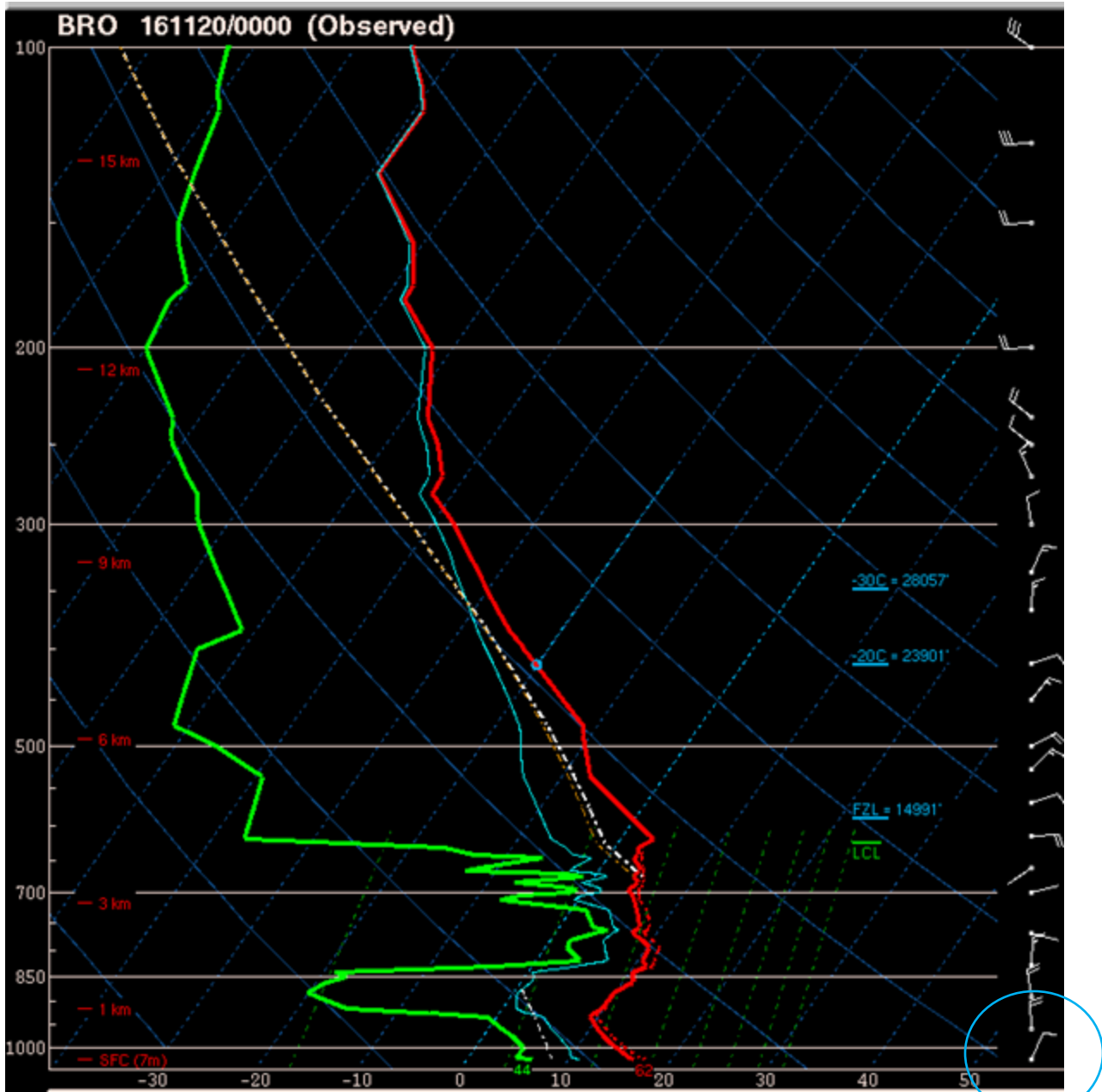
Air temperature after cold front passage

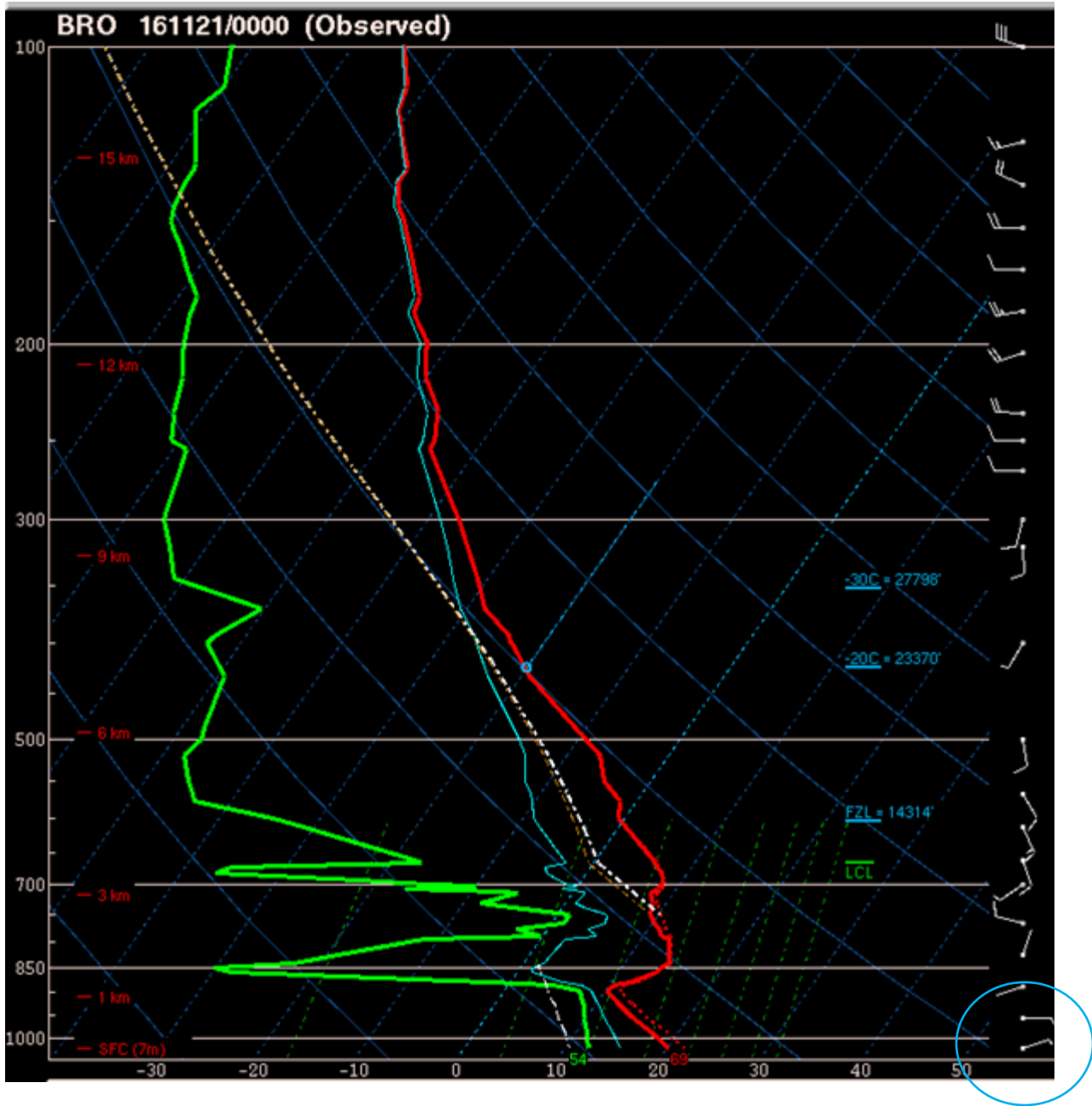


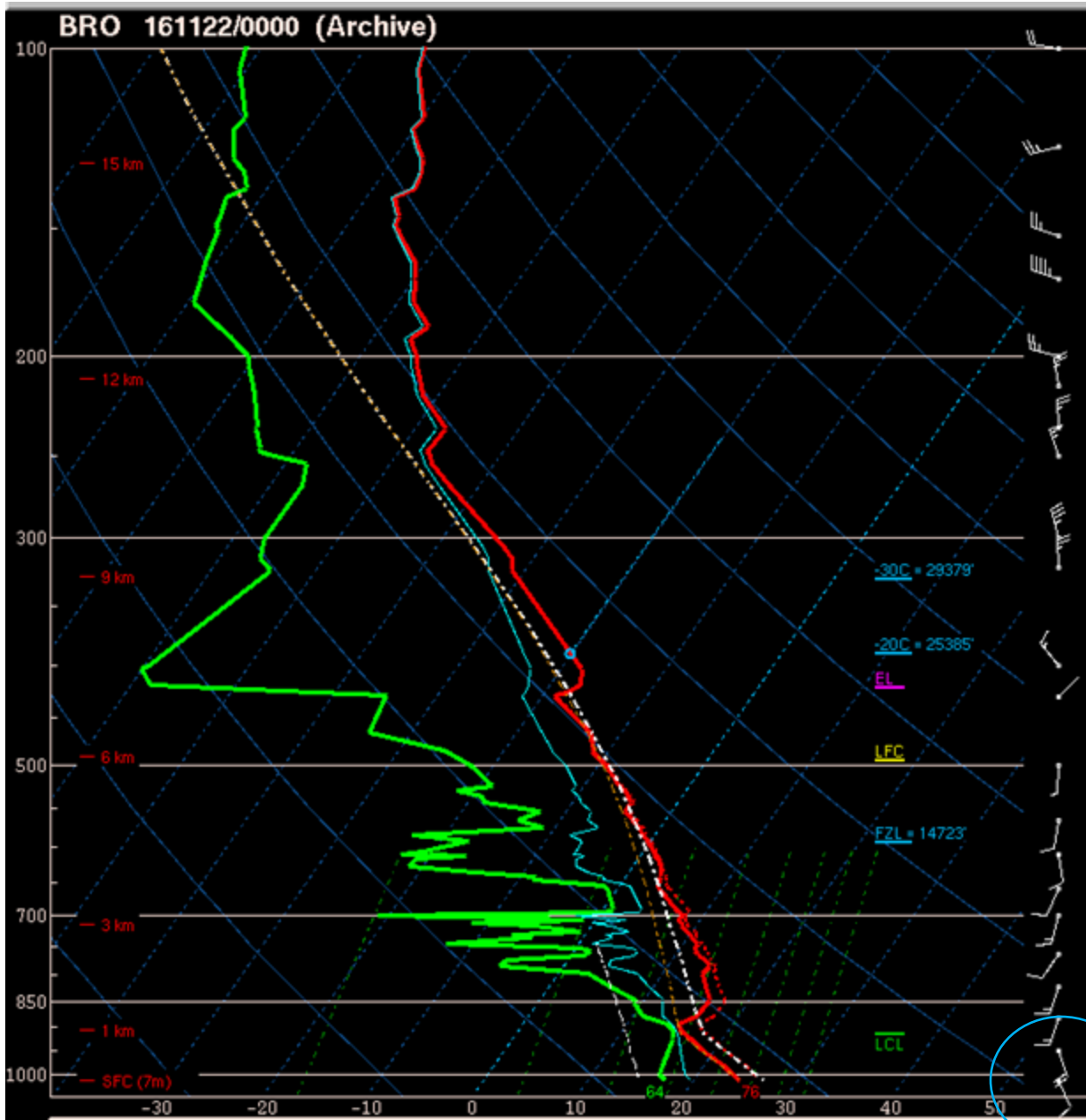


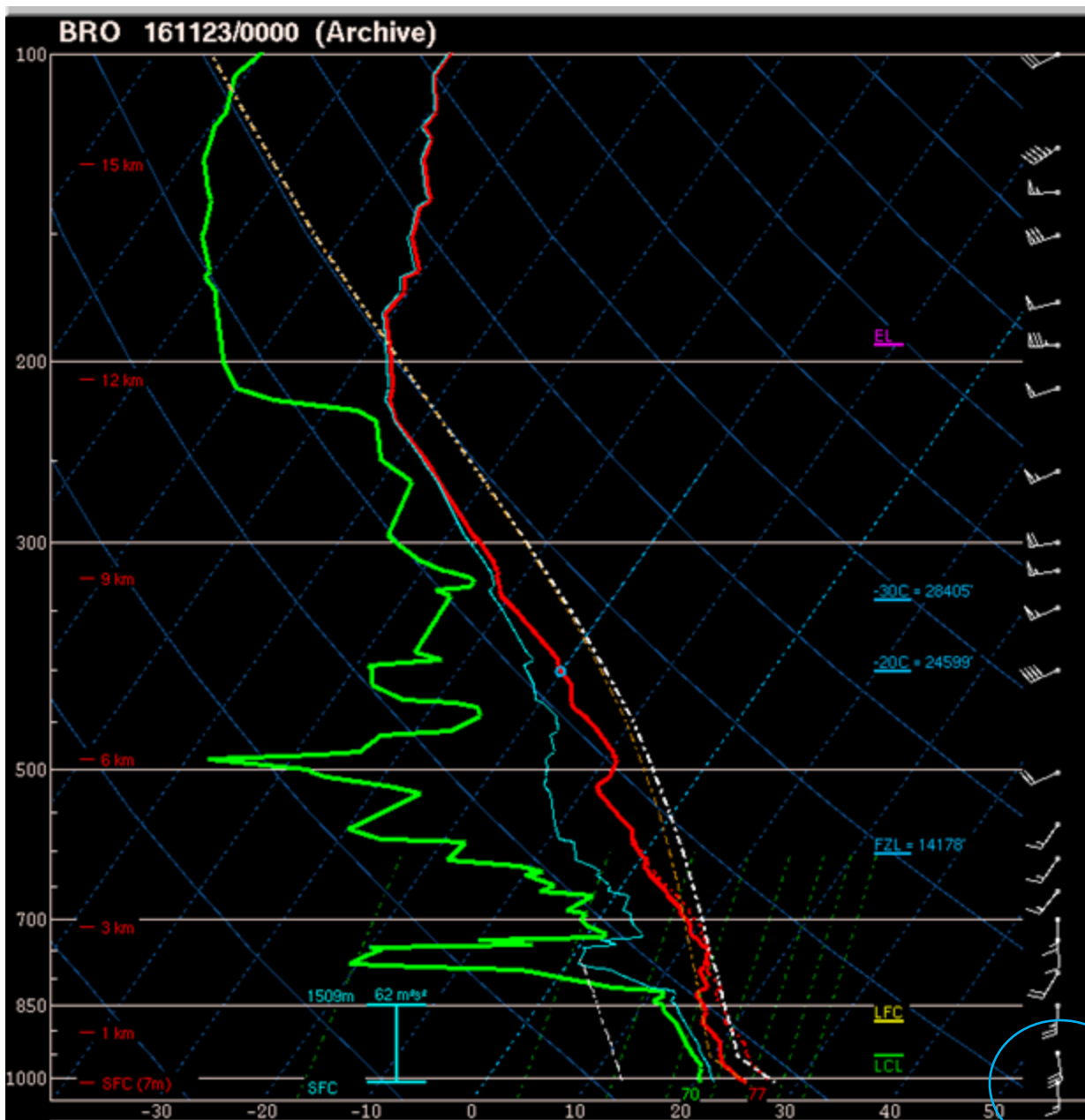






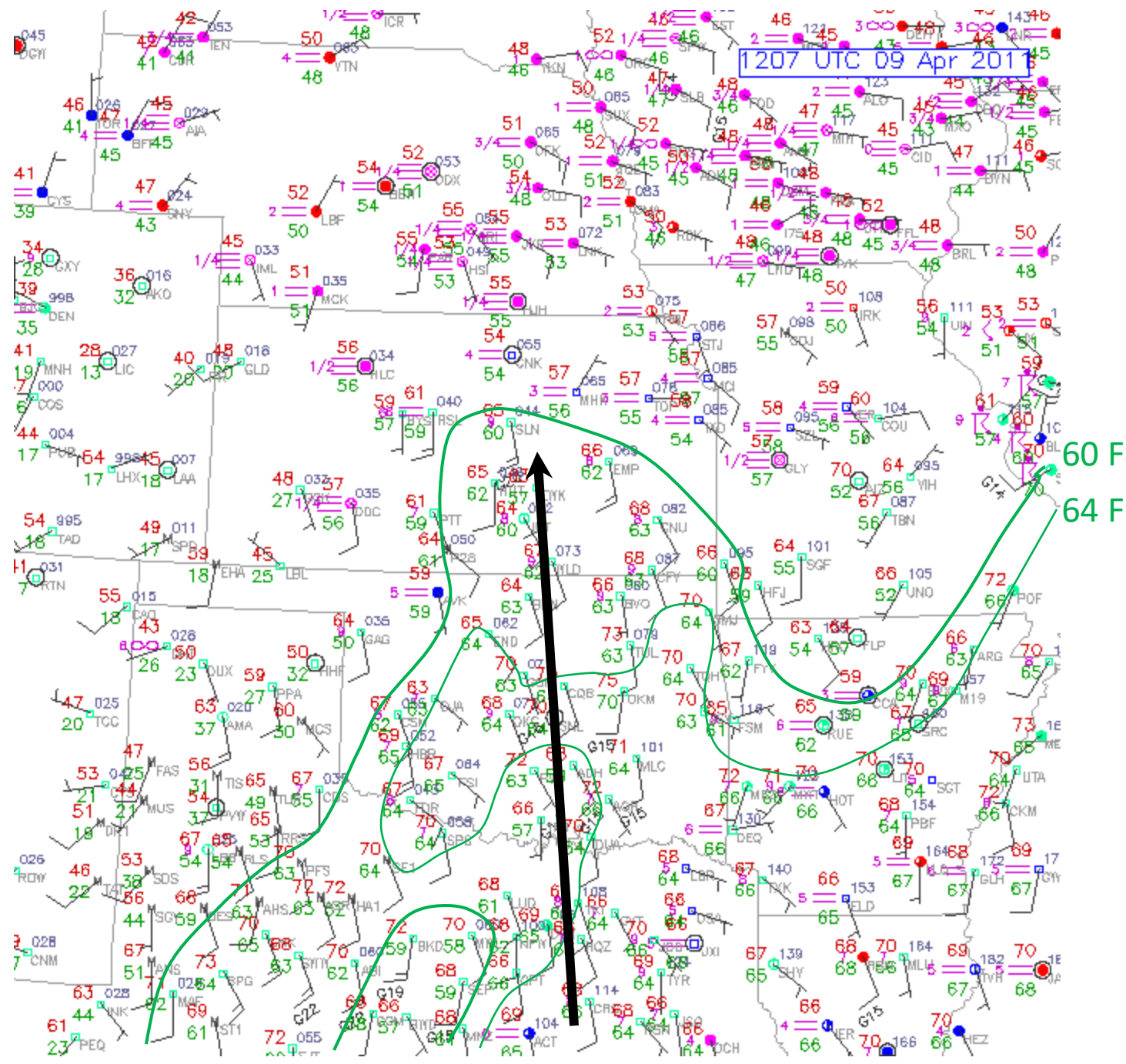


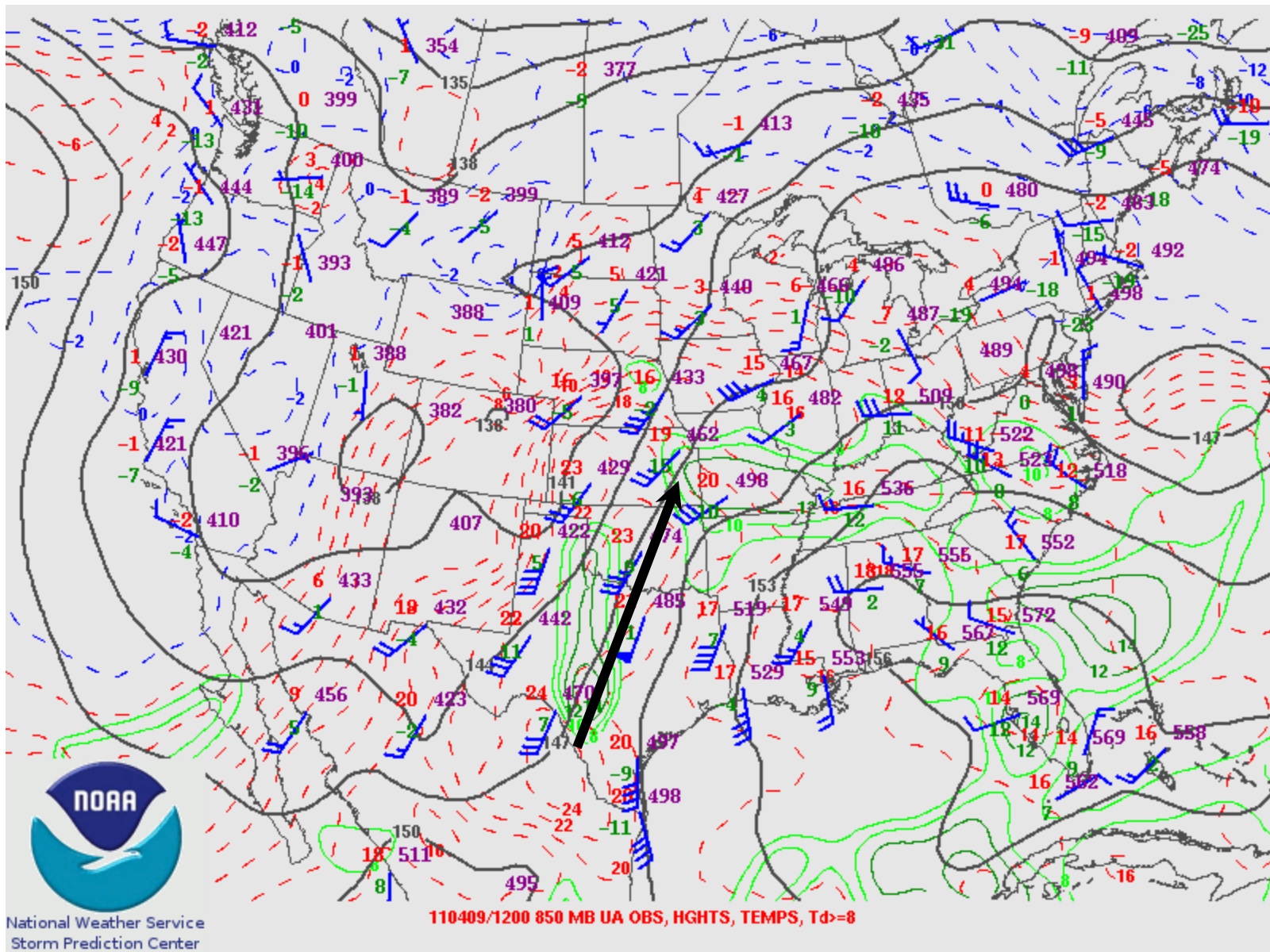


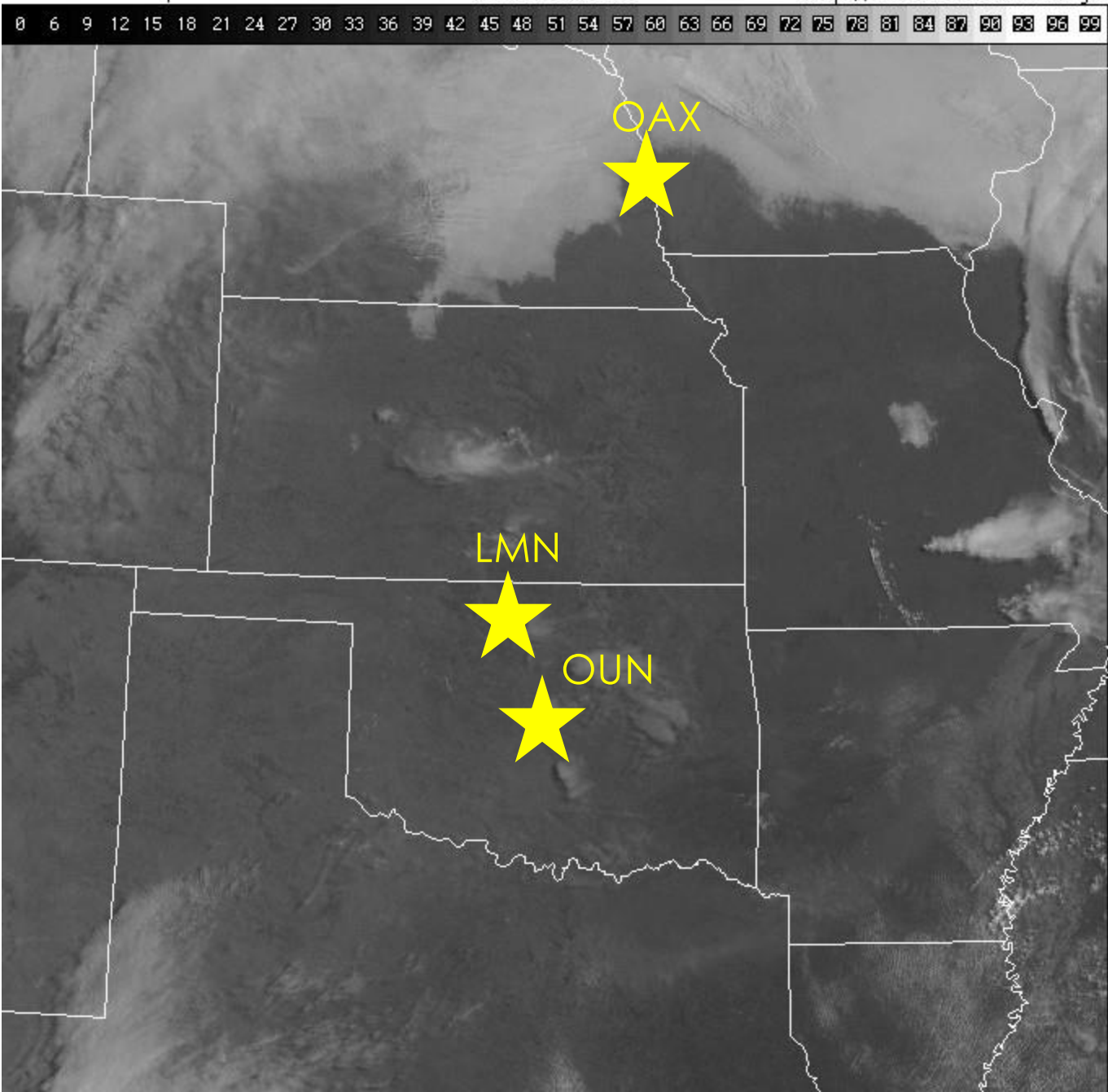


Inland Vertical Mixing

- **Surface moisture can decrease during the daytime when:**
 - Daytime heating/mixing extends deeper than original moist layer depth
 - Moisture advection is not enough to offset mixing
 - Compounded by dry soil, little green vegetation, and ongoing drought
- **Daytime mixing is governed by:**
 - Vertical moisture structure
 - Height and strength of lid/cap
 - Upstream moisture sources
 - Local moisture sources







U.S. Drought Monitor CONUS

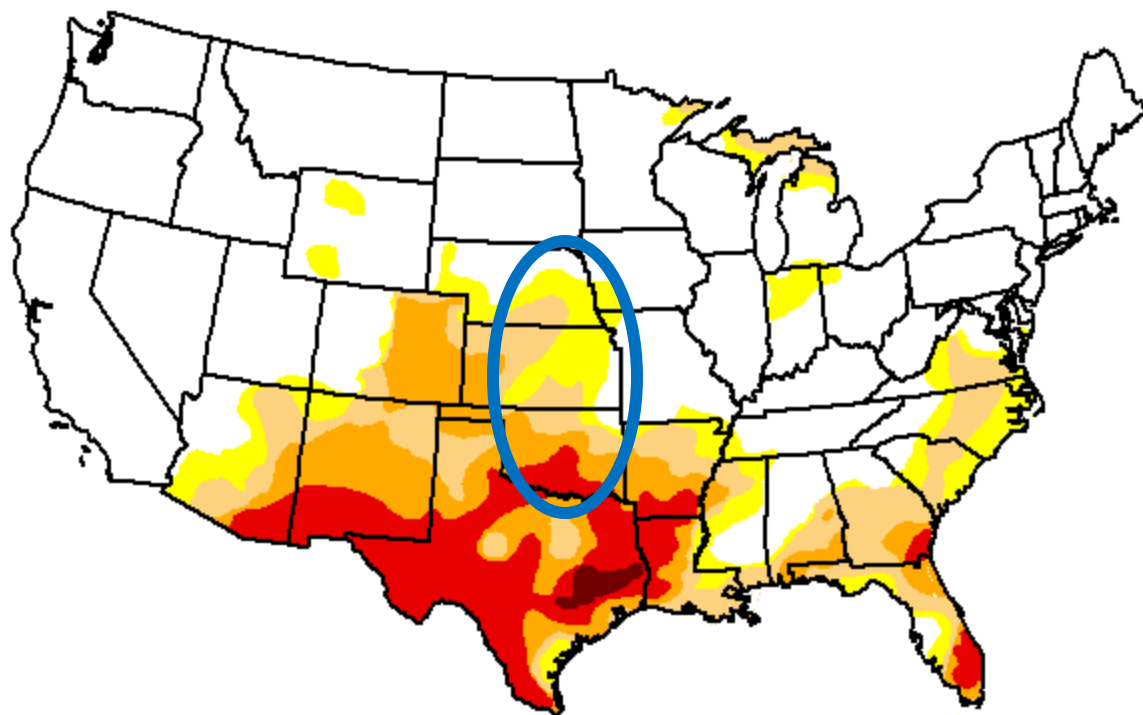
April 5, 2011

(Released Thursday, Apr. 7, 2011)

Valid 7 a.m. EST

Drought Conditions (Percent Area)

	None	D0-D4	D1-D4	D2-D4	D3-D4	D4
Current	62.46	37.54	28.75	18.17	8.75	0.42
Last Week <i>3/29/2011</i>	61.76	38.24	28.14	18.22	6.41	0.00
3 Months Ago <i>1/4/2011</i>	60.50	39.50	21.74	8.50	2.60	0.00
Start of Calendar Year <i>1/4/2011</i>	60.50	39.50	21.74	8.50	2.60	0.00
Start of Water Year <i>9/28/2010</i>	60.05	39.95	13.16	3.09	0.30	0.00
One Year Ago <i>4/6/2010</i>	71.57	28.43	9.10	2.00	0.00	0.00



Intensity:

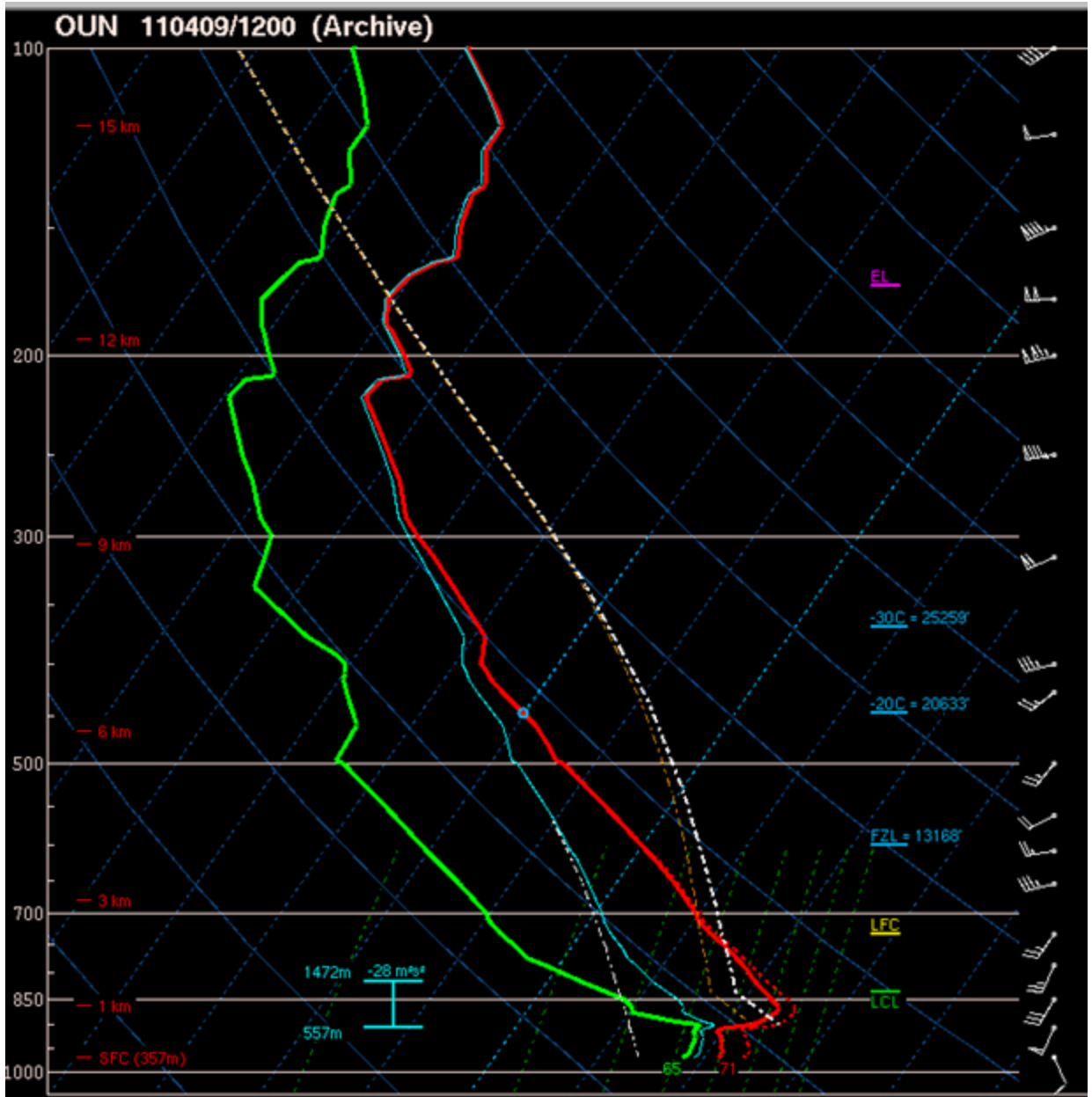
 D0 Abnormally Dry	 D3 Extreme Drought
 D1 Moderate Drought	 D4 Exceptional Drought
 D2 Severe Drought	

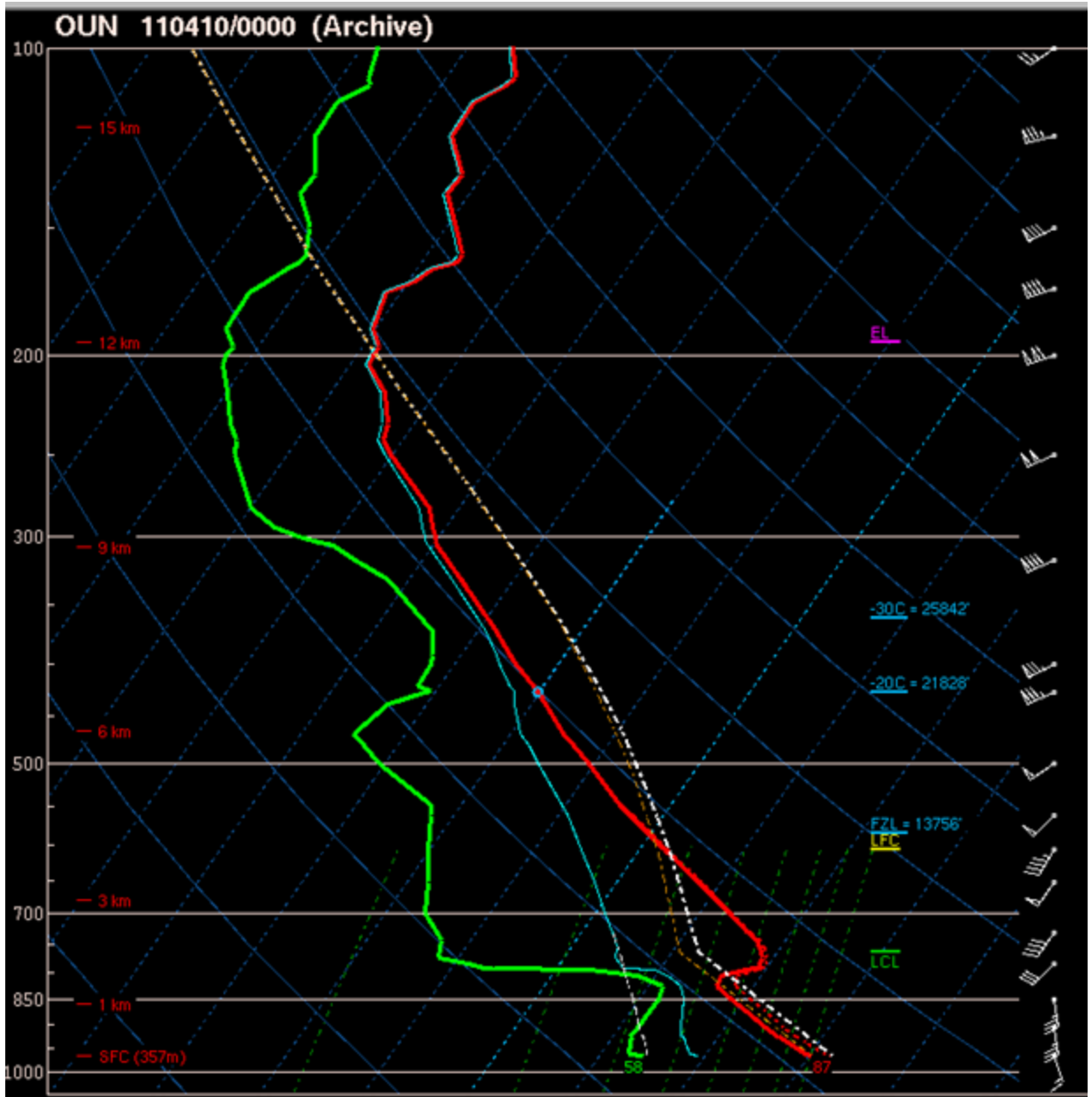
The Drought Monitor focuses on broad-scale conditions. Local conditions may vary. See accompanying text summary for forecast statements.

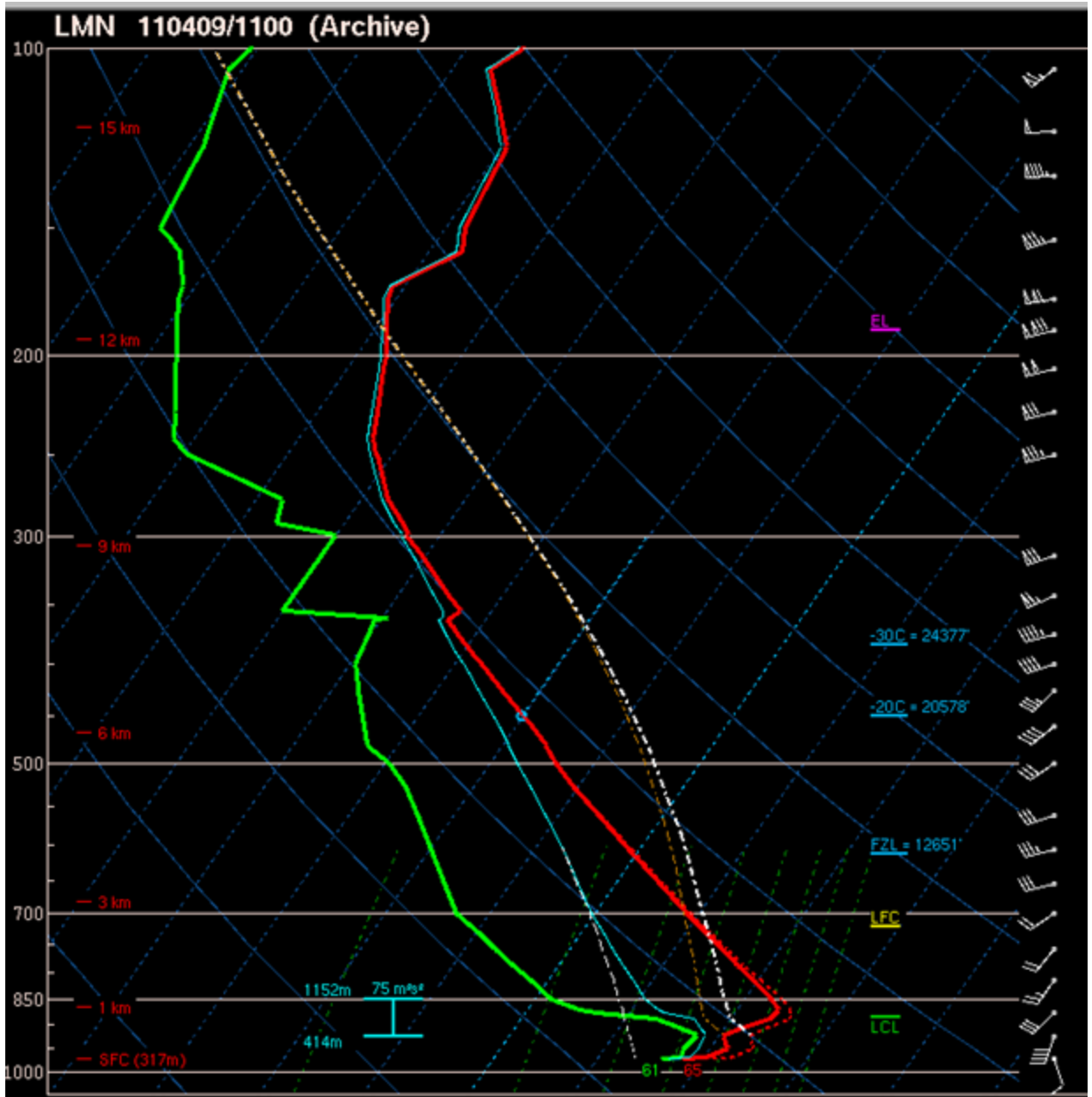
Author(s):

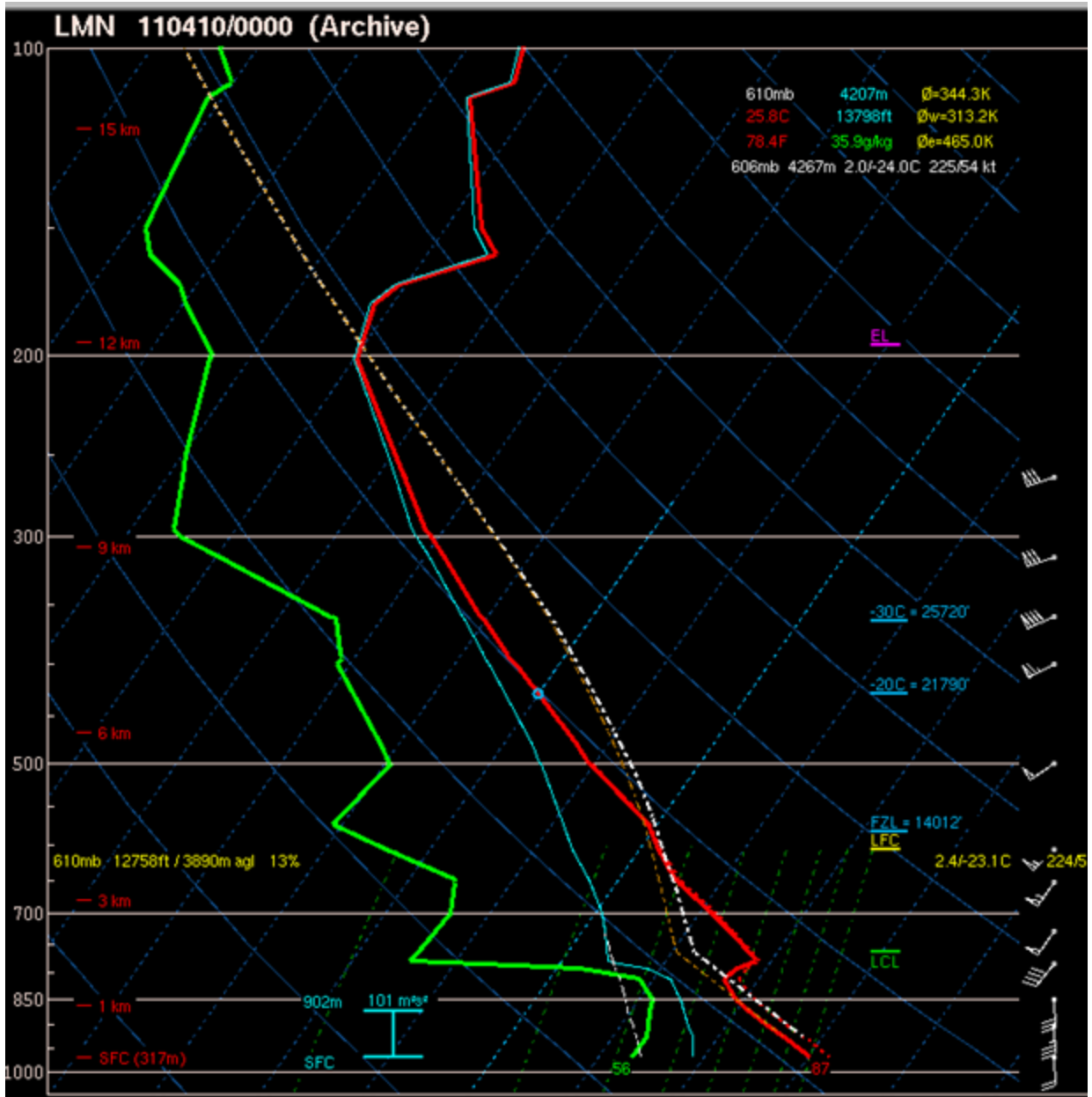
Mark Svoboda

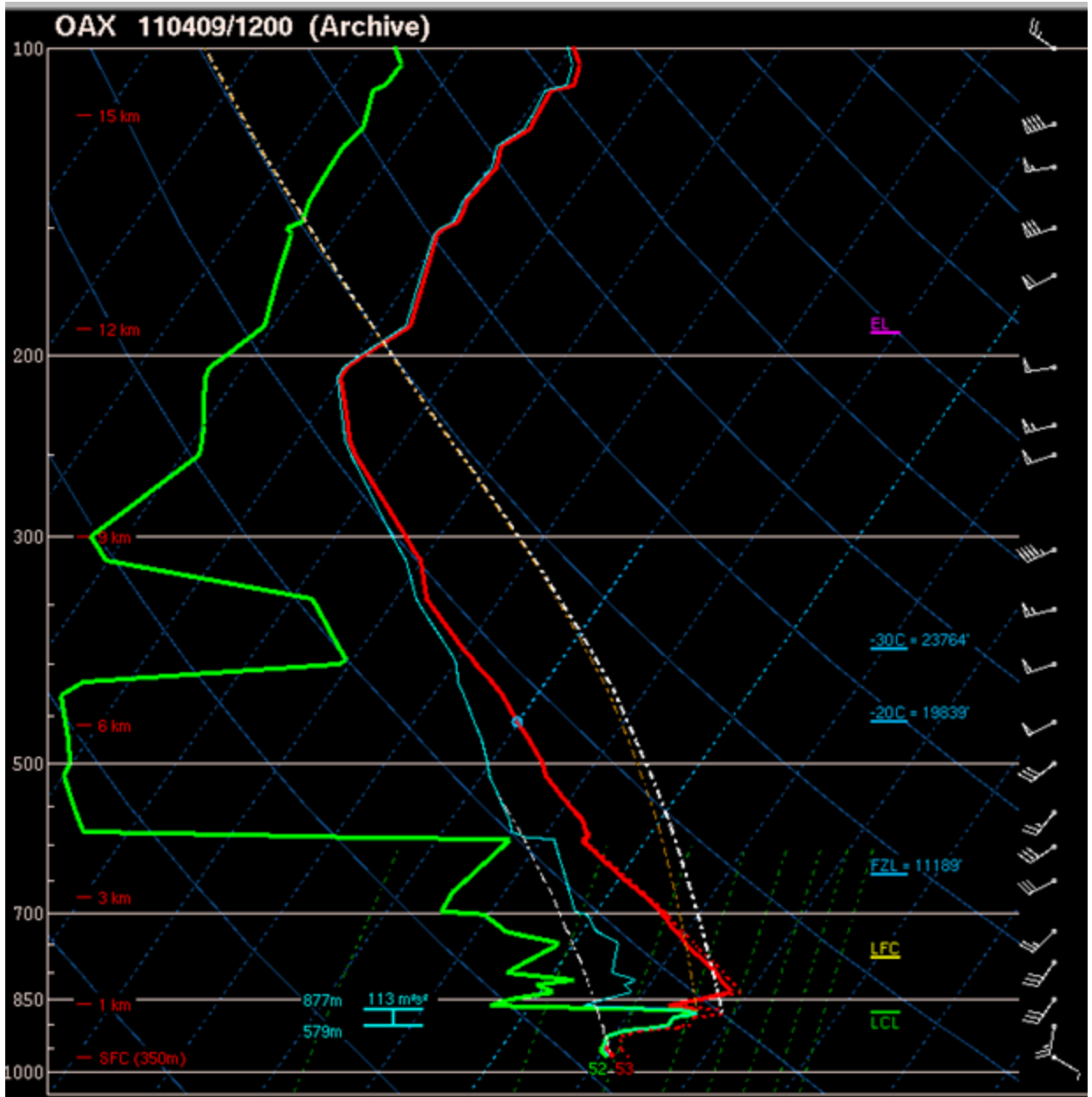
National Drought Mitigation Center

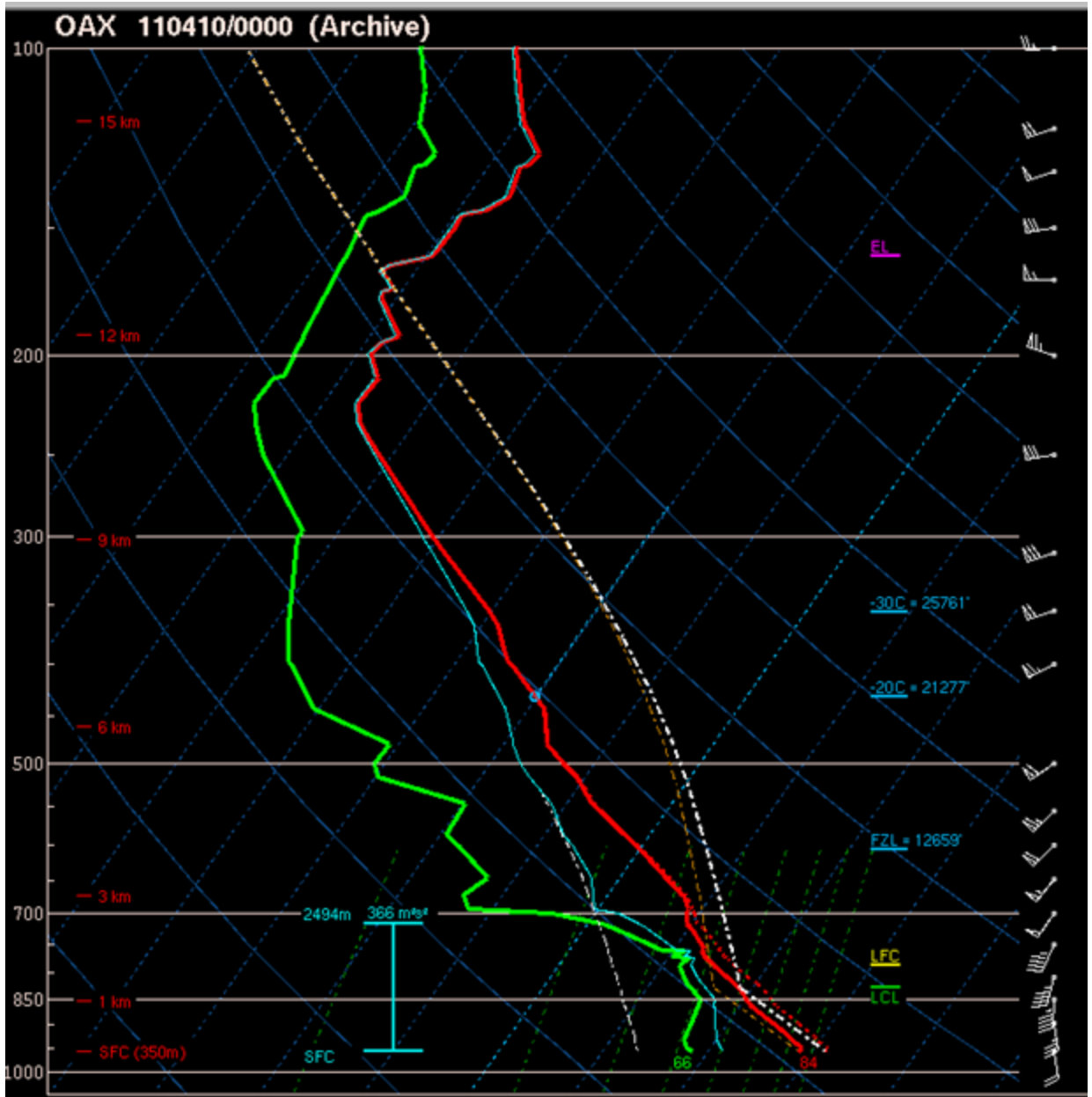












SPC Storm Reports for 04/09/11

Map updated at 1212Z on 04/19/11



- High Wind Report (65KT +)
- ▲ Large Hail Report (2" dia. +)

PRELIMINARY DATA ONLY

Evapotranspiration

- **Need moist soil and growing vegetation**
- **Plentiful rain previous 1-2 weeks**
- **Maturing crops (wheat, corn, or canola)**
- **Capped boundary layer to trap moisture; relatively weak winds**
- **Almost always a significant return flow contribution (in addition)**

U.S. Drought Monitor CONUS

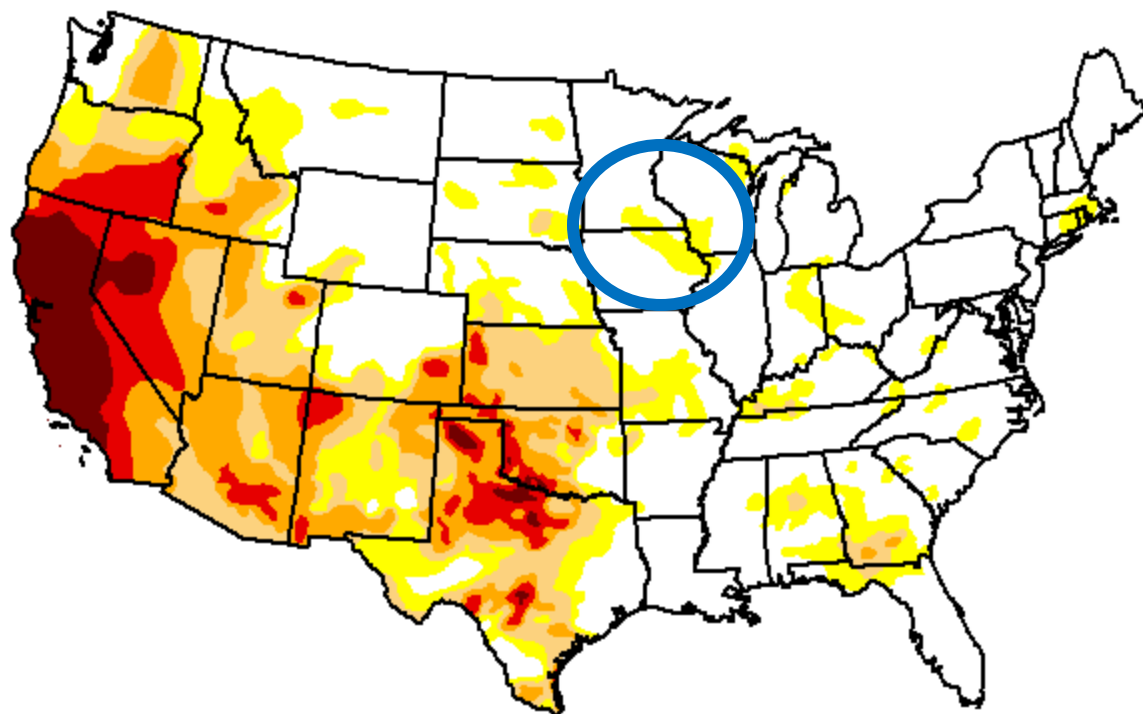
August 19, 2014

(Released Thursday, Aug. 21, 2014)

Valid 8 a.m. EDT

Drought Conditions (Percent Area)

	None	D0-D4	D1-D4	D2-D4	D3-D4	D4
Current	51.93	48.07	33.56	21.62	10.12	3.80
Last Week 8/12/2014	52.81	47.19	33.61	22.03	10.24	3.79
3 Months Ago 5/20/2014	52.36	47.64	38.12	28.30	14.47	4.99
Start of Calendar Year 12/01/2013	48.24	51.76	30.95	16.67	3.96	0.37
Start of Water Year 10/1/2013	39.57	60.43	41.21	20.70	3.06	0.29
One Year Ago 8/20/2013	40.02	59.98	45.61	32.23	10.54	1.32



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 D1 Moderate Drought	 D4 Exceptional Drought
 D2 Severe Drought	

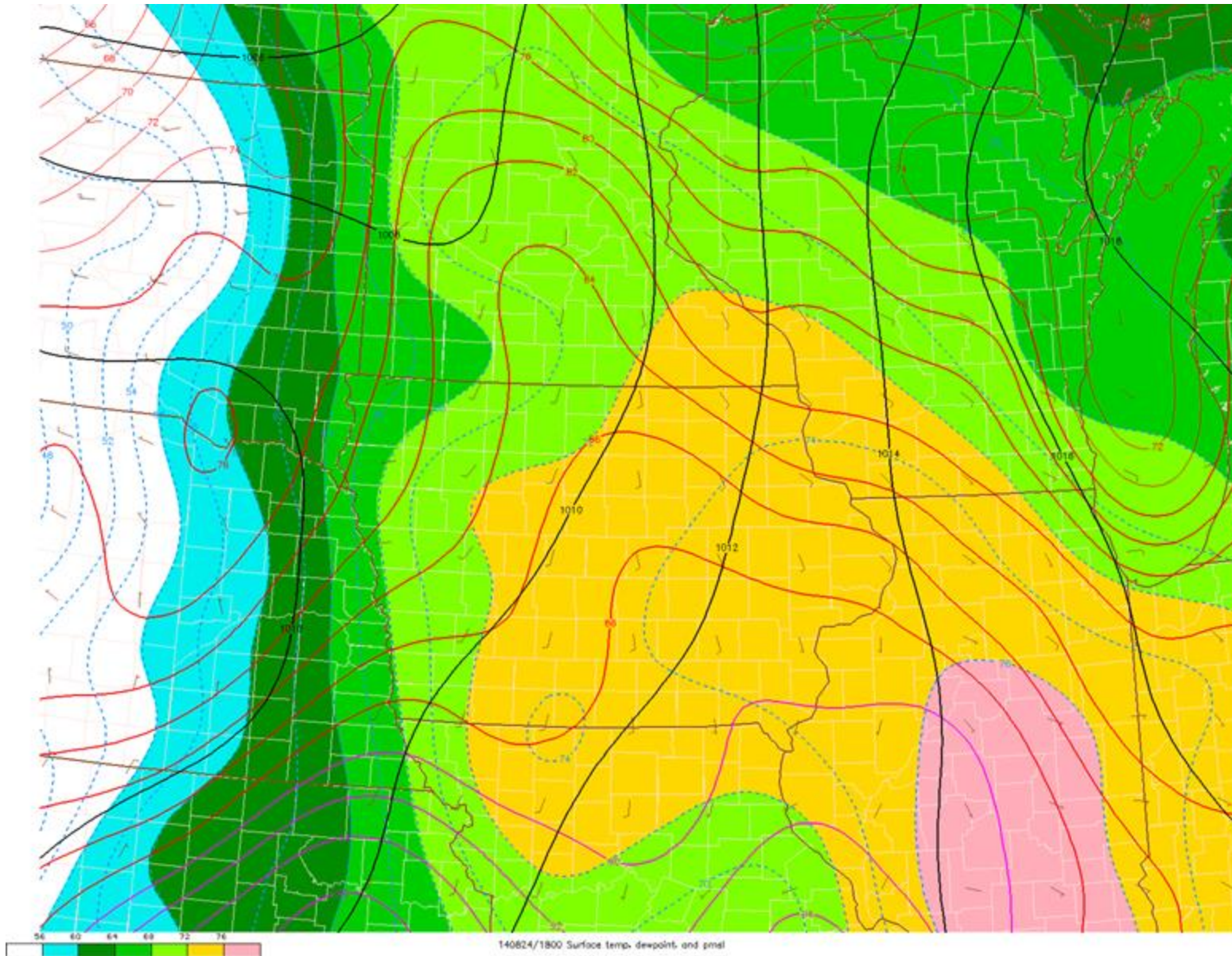
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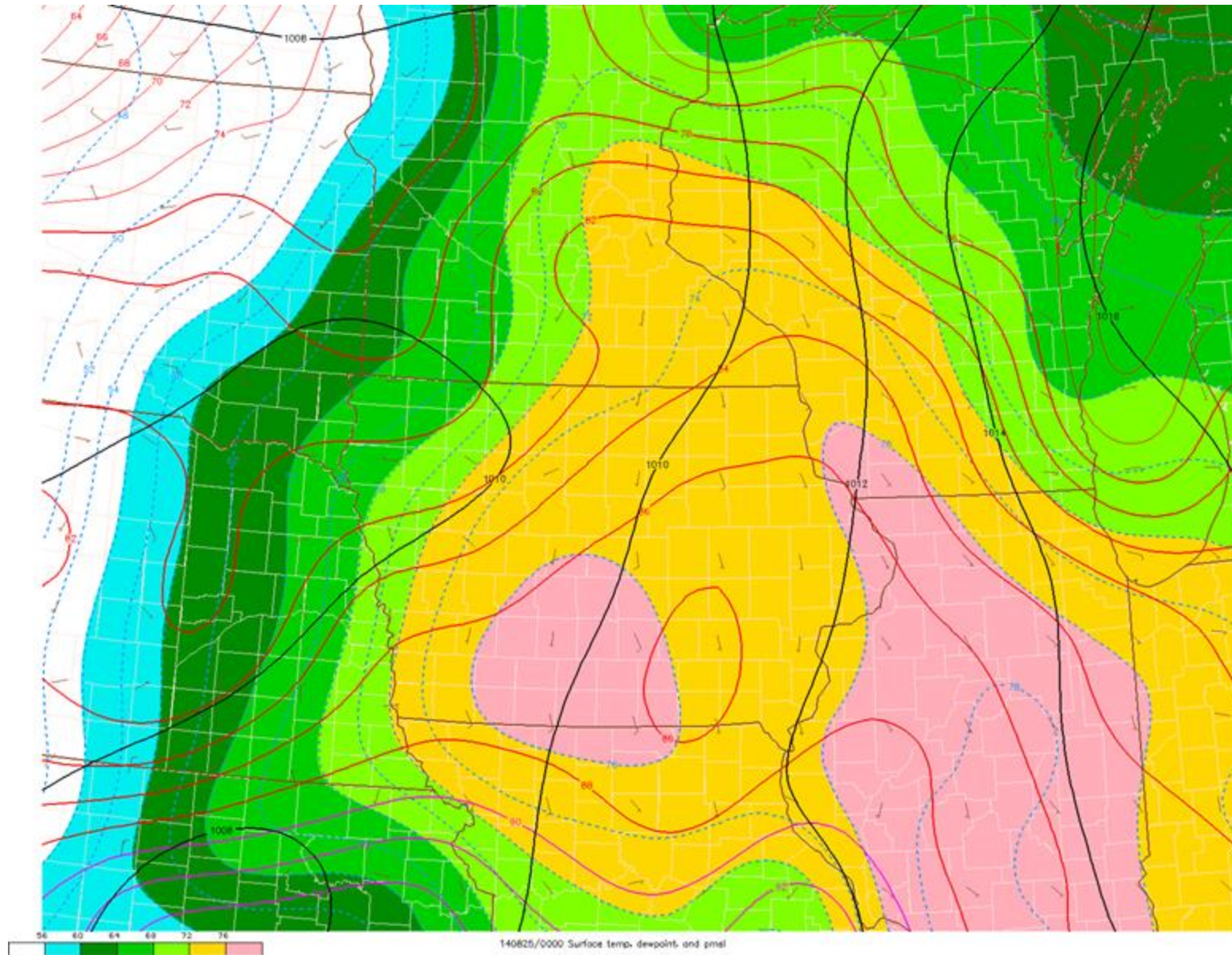
Author(s):

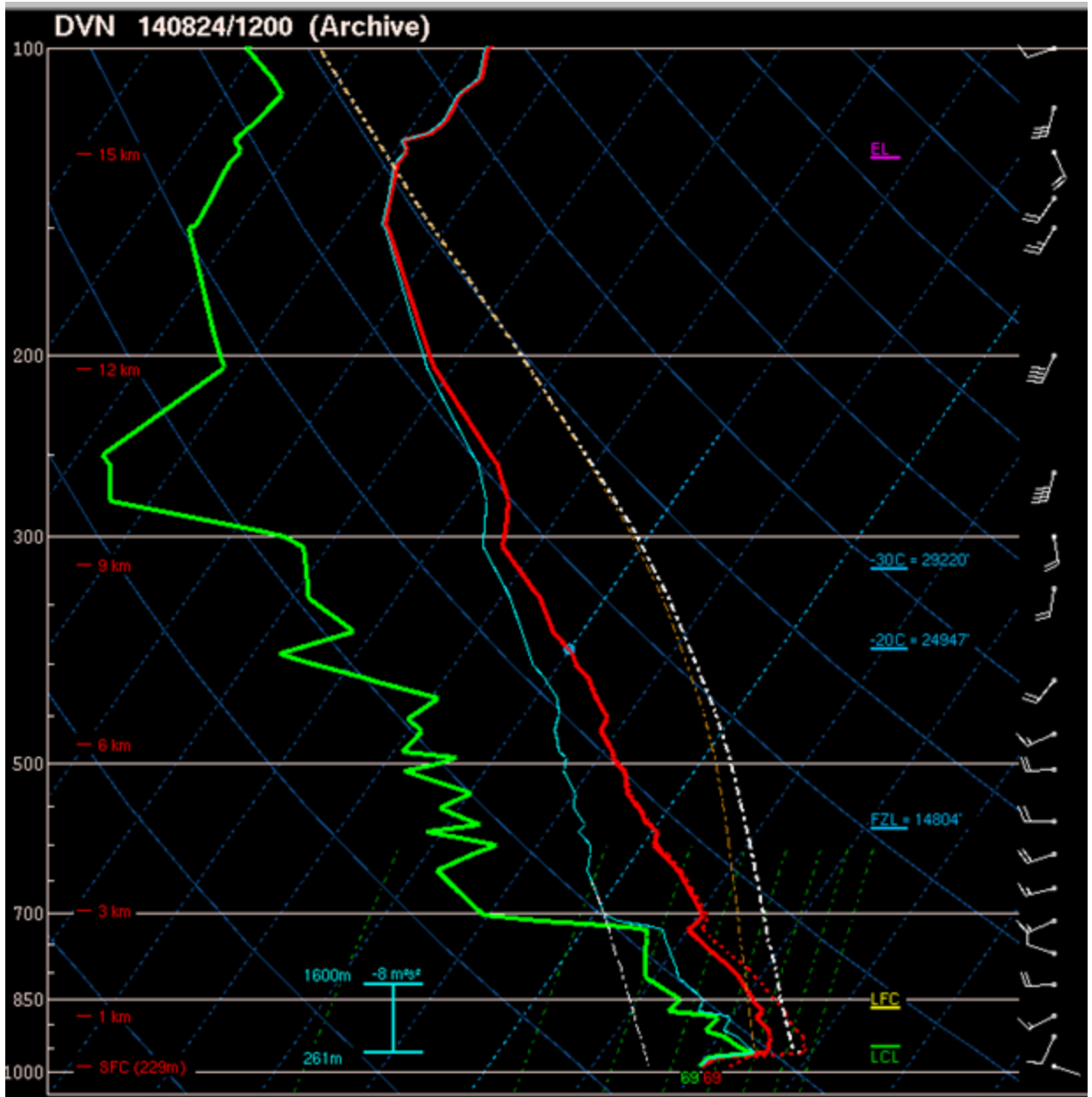
Richard Tinker

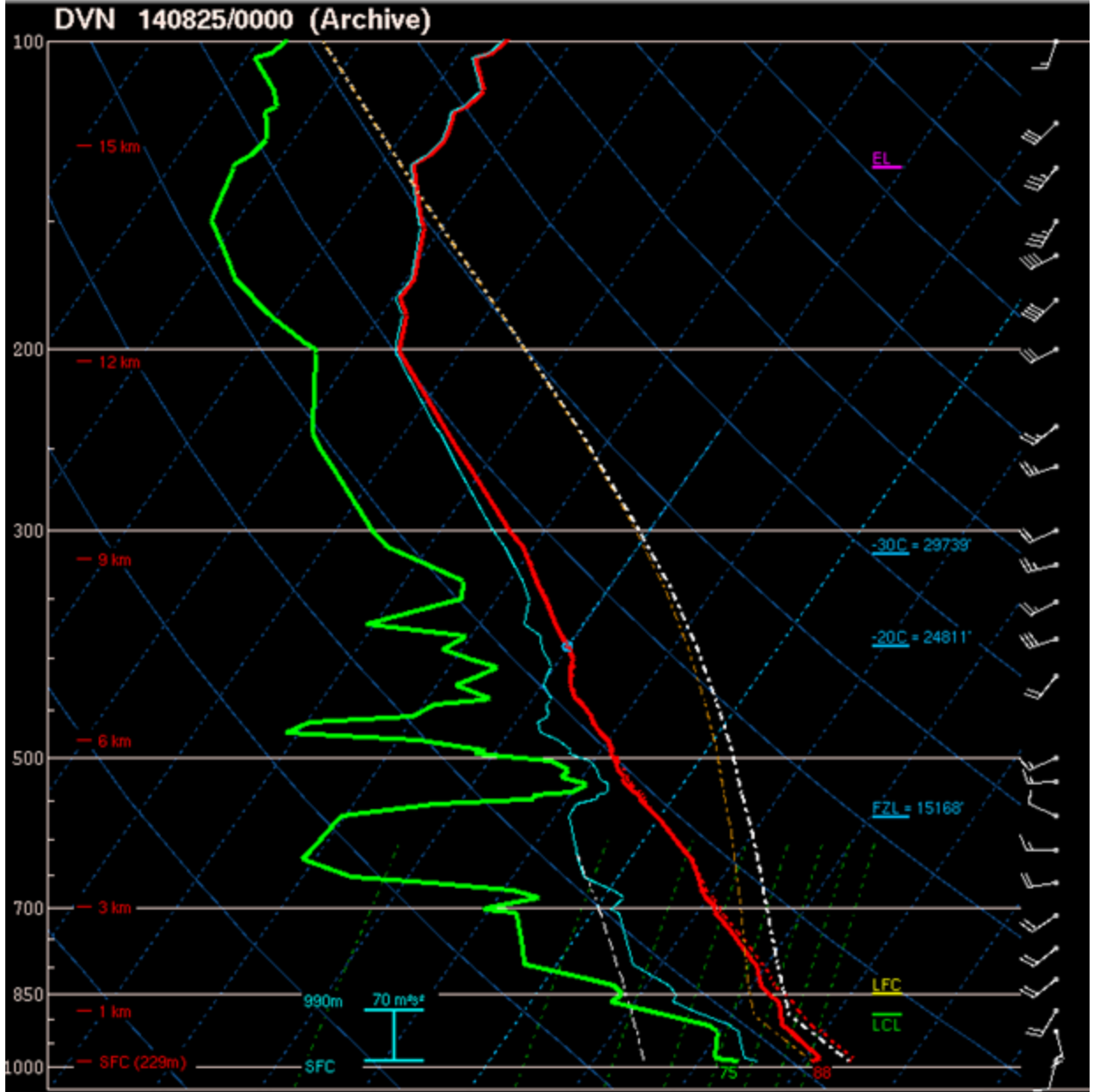
CPC/NOAA/NWS/NCEP

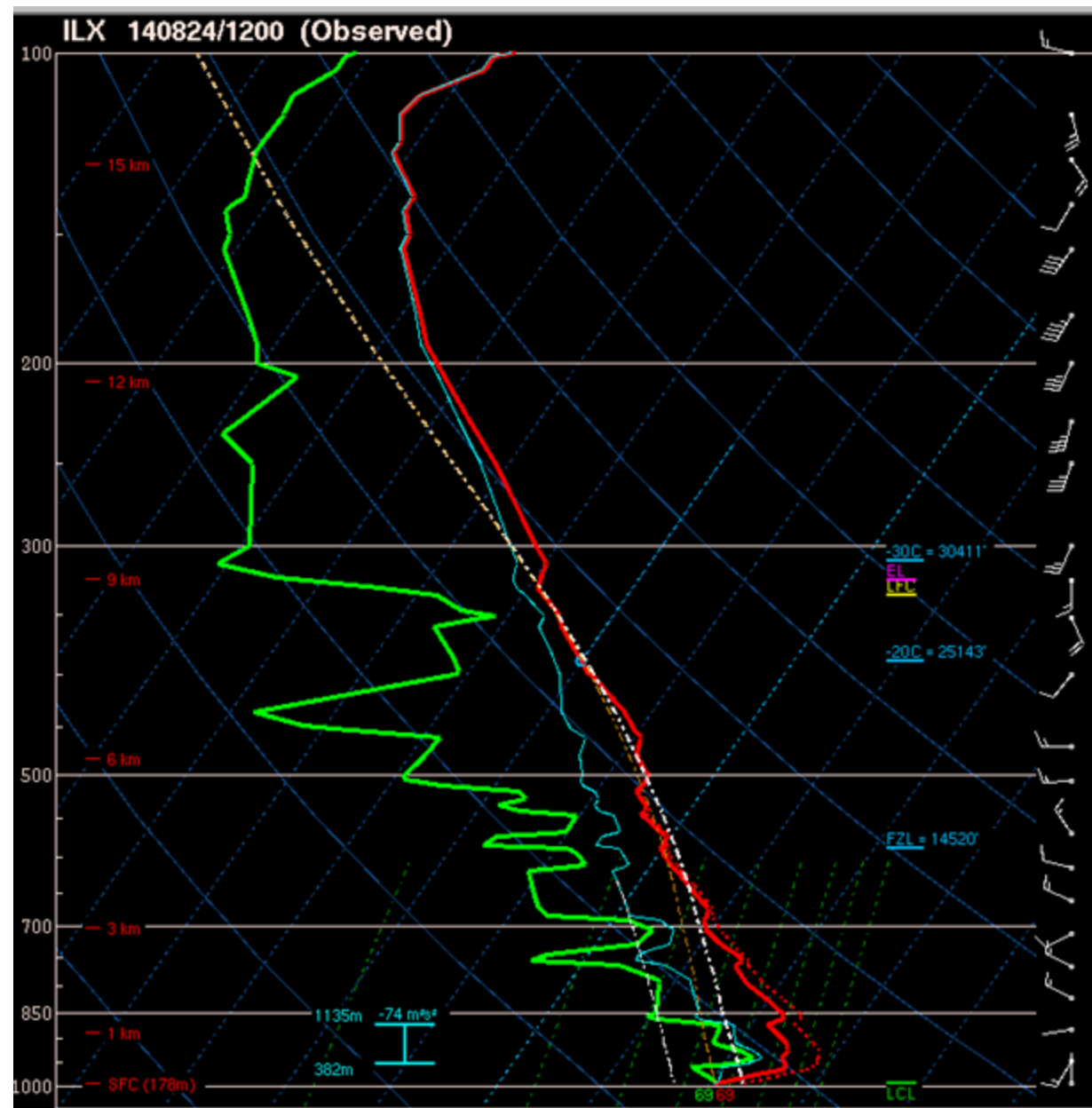


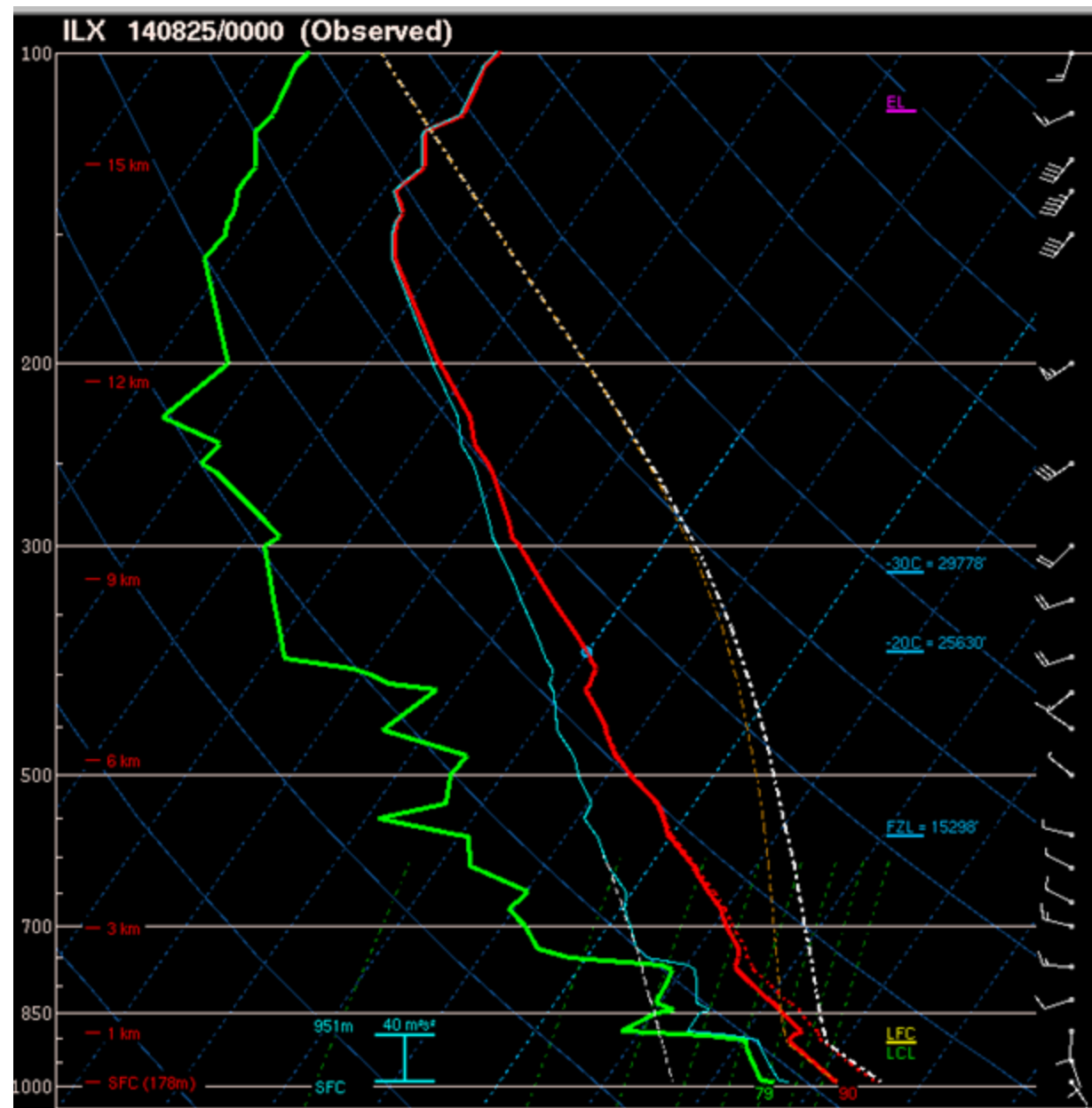






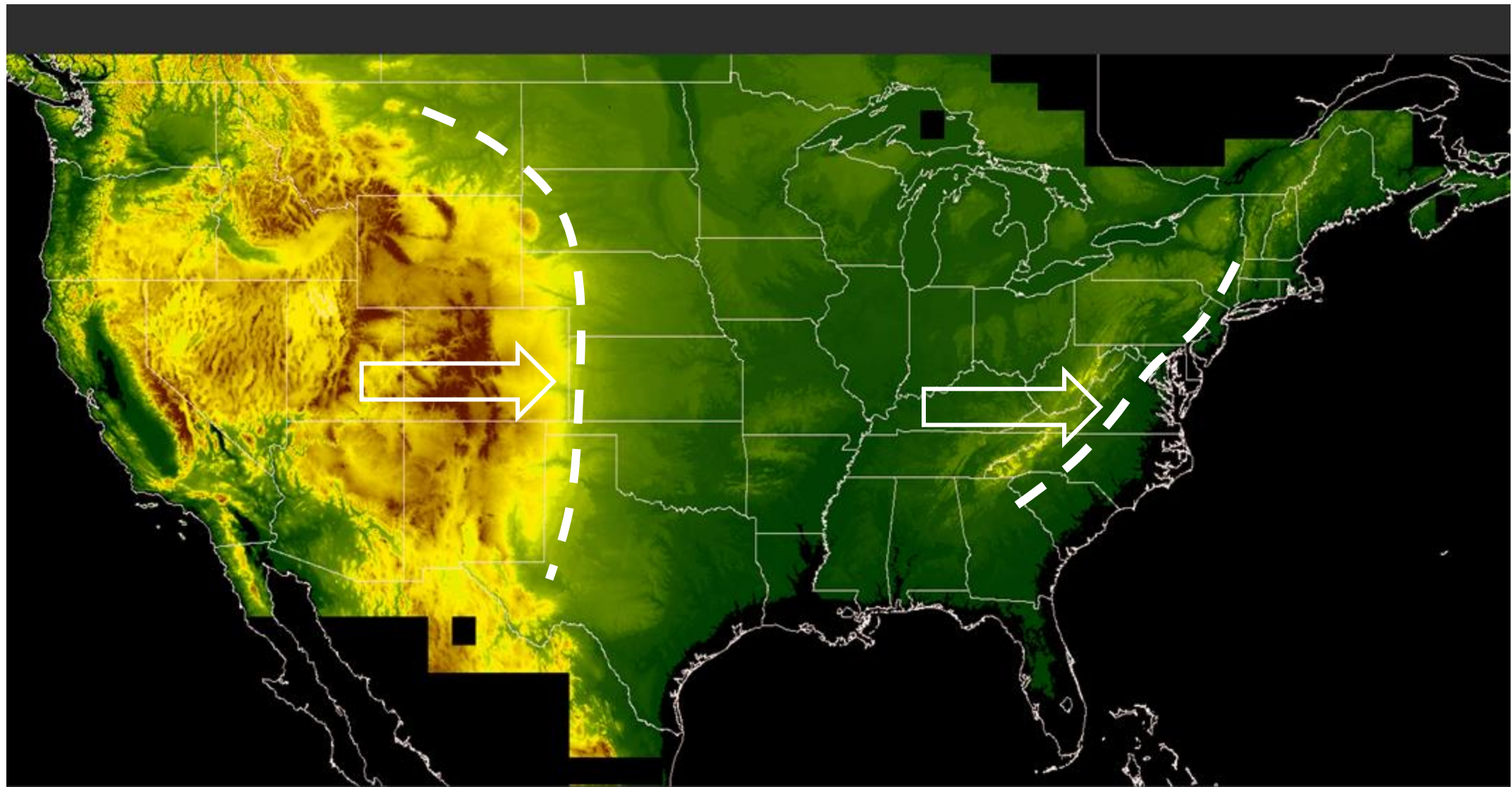




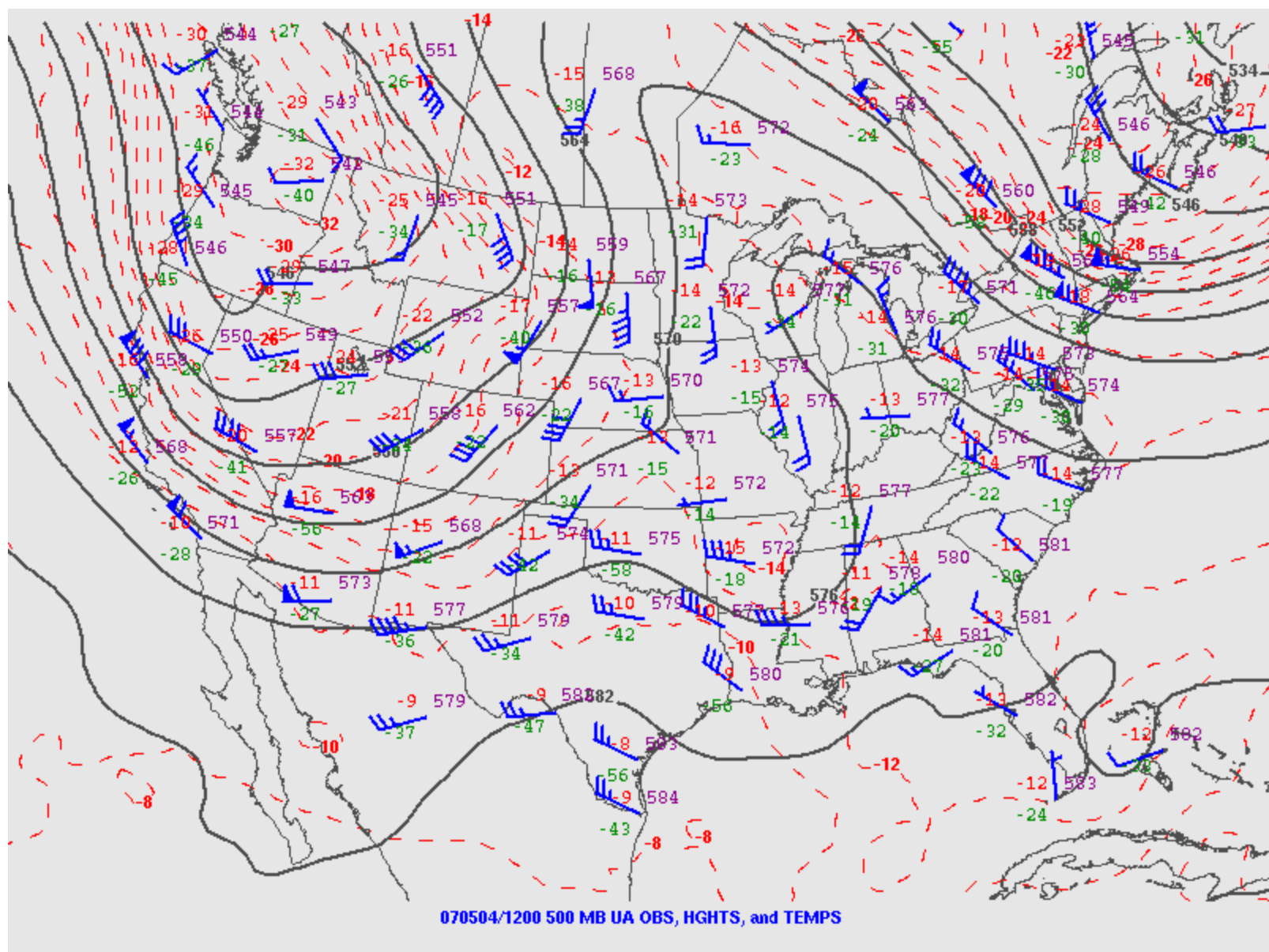


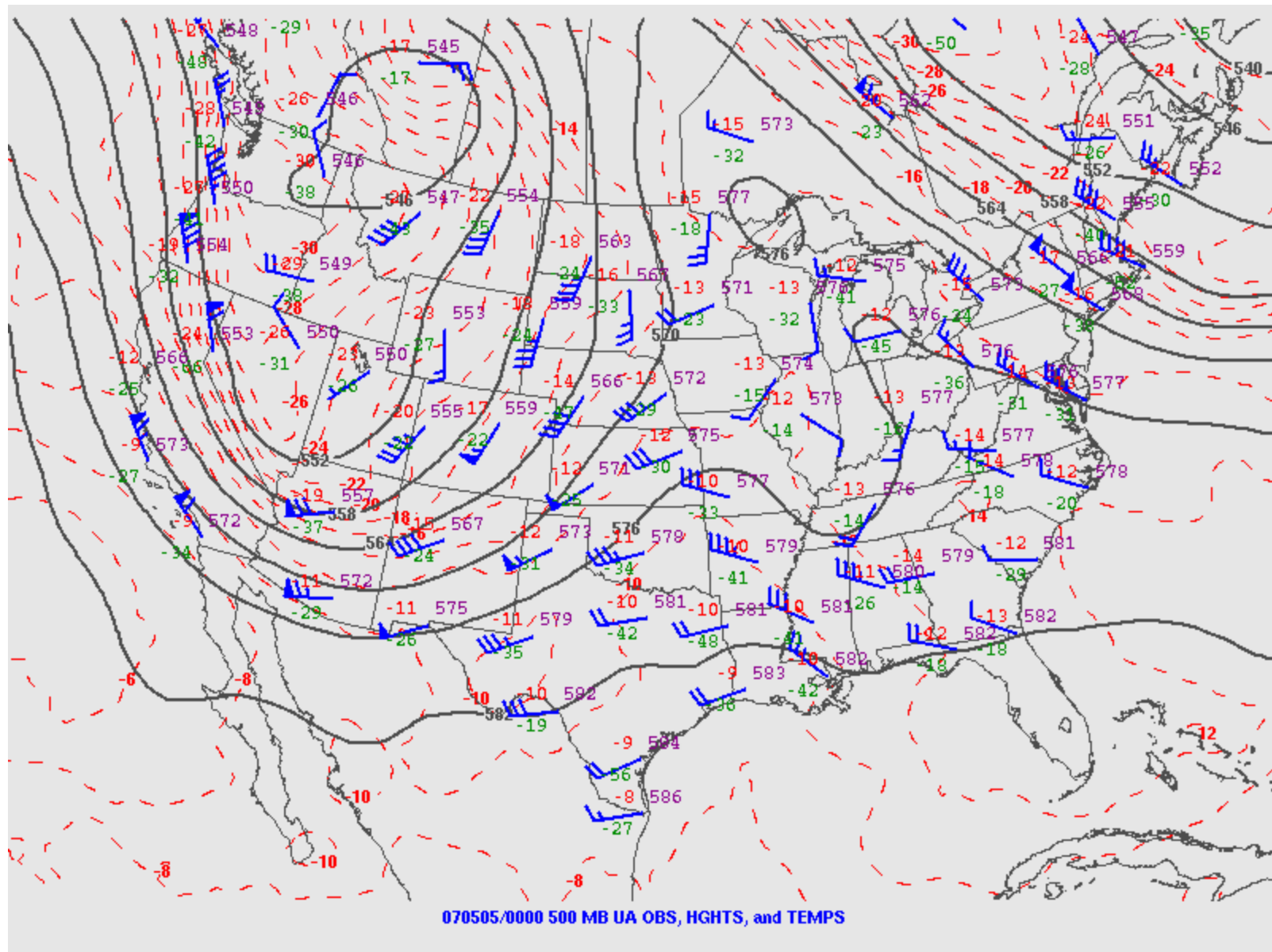
Lee Cyclogenesis

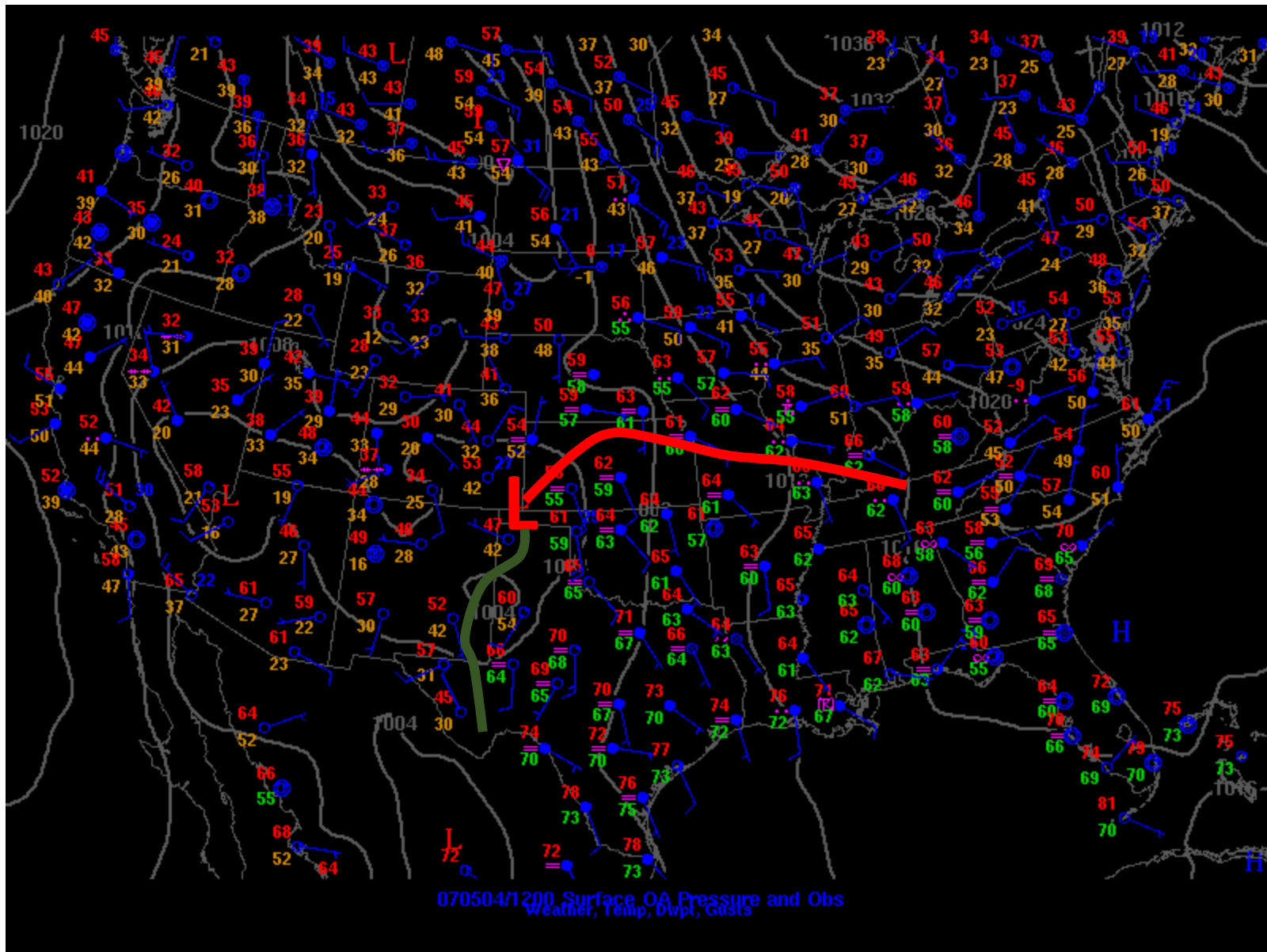
- **Midlevel flow crosses high terrain:**
 - Subsidence and warming E of mountains
 - Warming of column leads to “warm core” low formation
 - Strongest pressure falls with strongest flow crosses highest terrain
- **The lee trough/cyclone deepens before arrival of strongest Q-G forcing for ascent:**
 - Head start on differential advection, “loaded gun” sounding, and veering winds with height

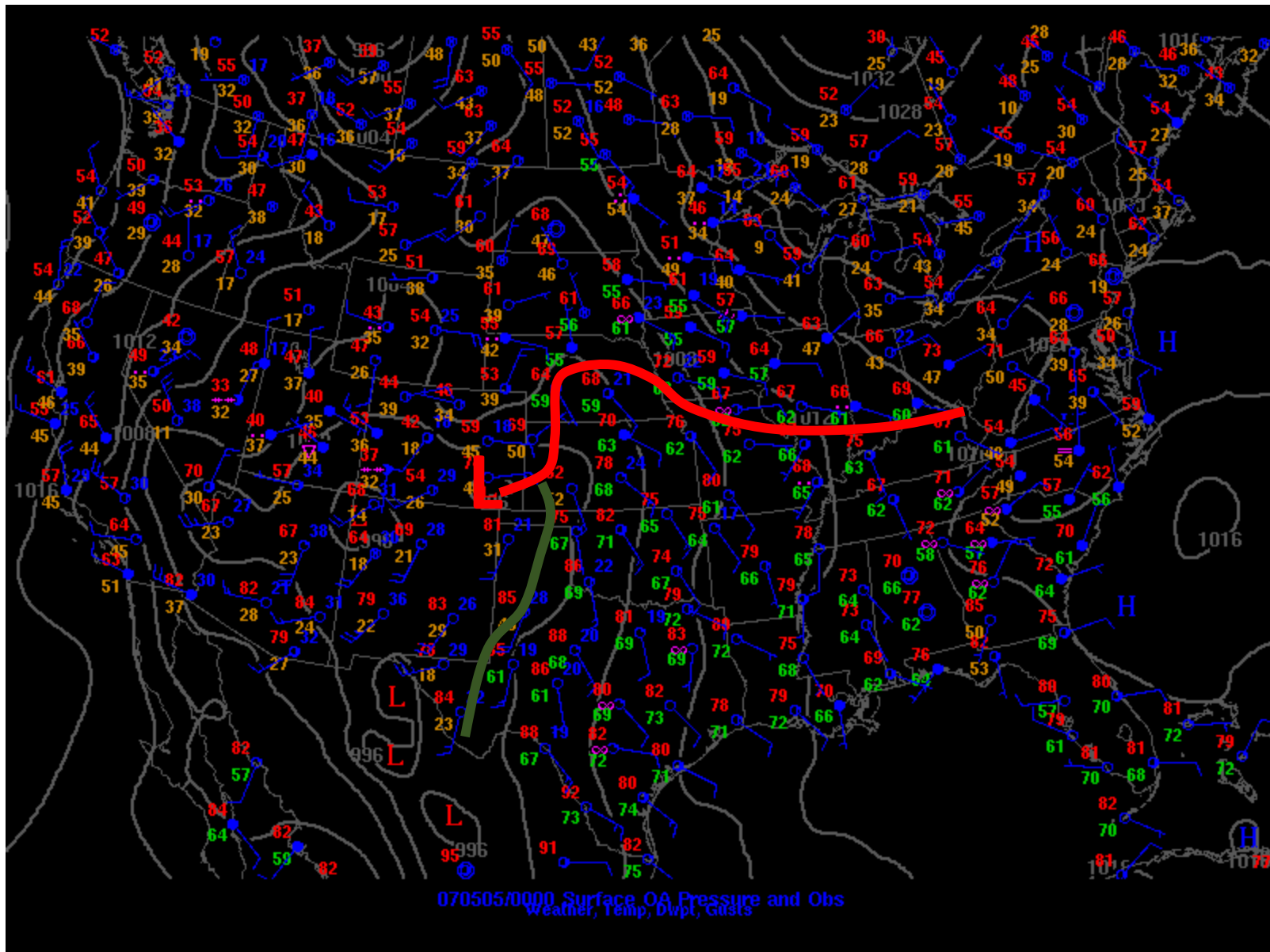


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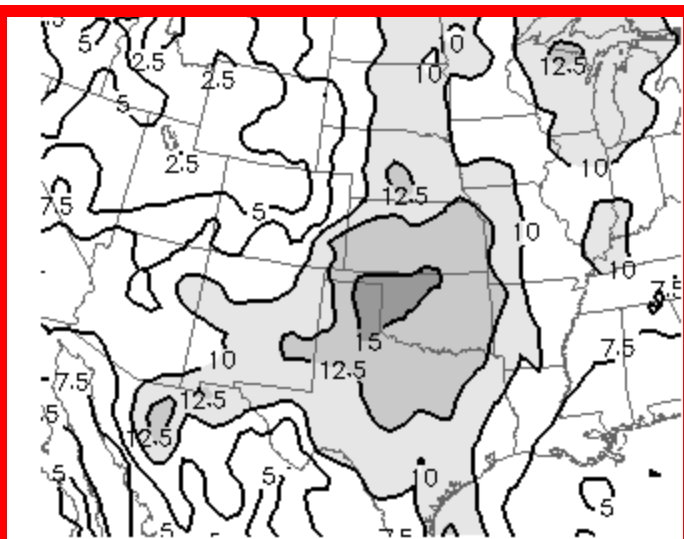




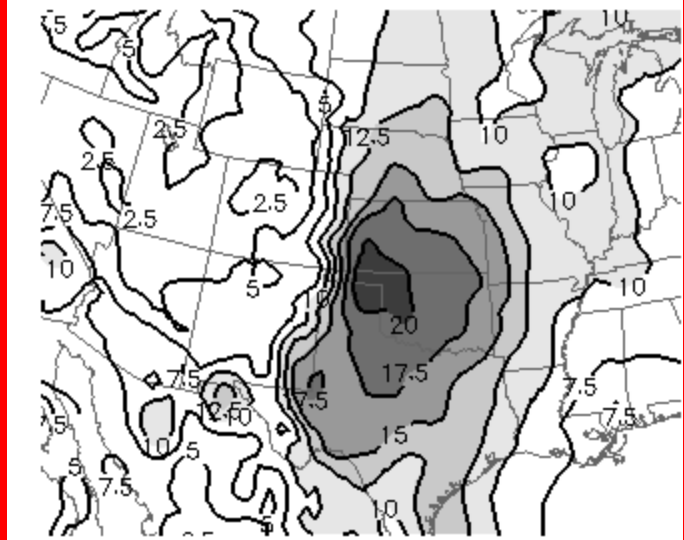
Nocturnal Low-Level Jet (LLJ)

- Plains often see early nocturnal low-level jet (LLJ) ramp up with lee cyclogenesis
- Related to two primary factors:
 - Boundary-layer decoupling and loss of surface friction (“inertial oscillation”)
 - Diurnal temperature variations over sloped terrain (thermal wind)
- Part of the process that can favor late evening/early overnight tornadoes (with favorable moisture/CAPE)

EF2+ supercell events

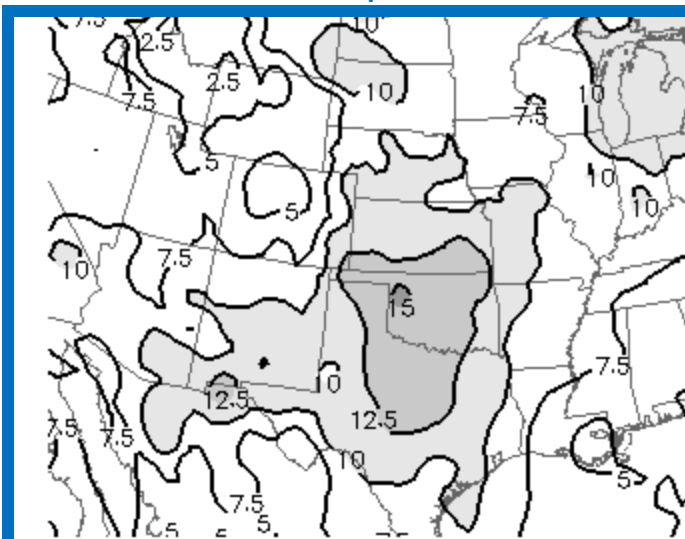


850 hPa isotachs at 0000 UTC (m s^{-1} SIGTOR)

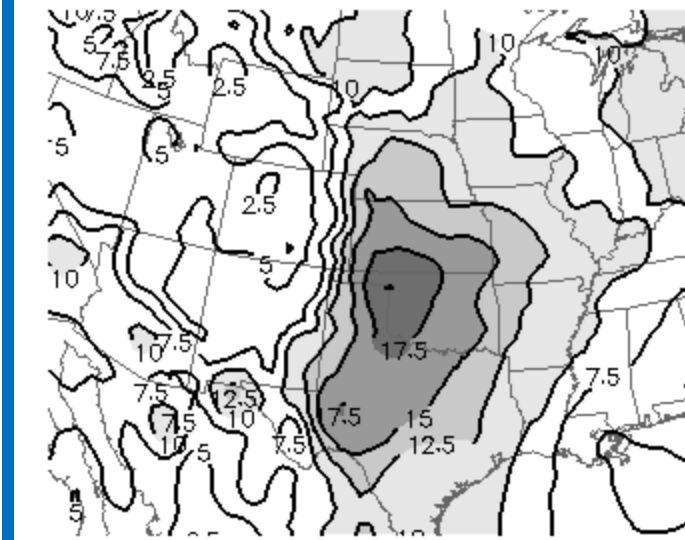


850 hPa isotachs at 0300 UTC (m s^{-1} SIGTOR)

Nontornadic supercell events



850 hPa isotachs at 0000 UTC (m s^{-1} NONTOR)



850 hPa isotachs at 0300 UTC (m s^{-1} NONTOR)

Lee Cyclogenesis Summary

- Lee cyclogenesis occurs where upper-level flow is perpendicular to terrain features (westerly flow for Rockies)
- Lee cyclogenesis more robust with lower static stability
- Lee cyclone helps drive low-level moisture return
- Nocturnal low-level jet associated with lee cyclone is driven by inertial oscillation and thermal wind → increased low-level shear and tornado threat in evening and after dark if adequate CAPE