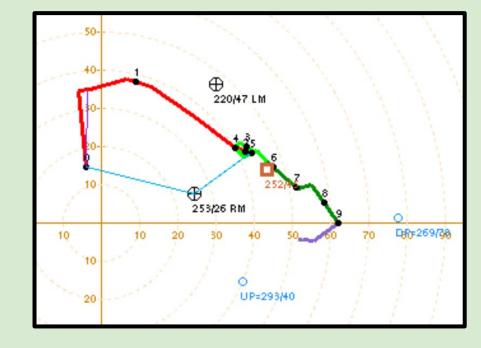
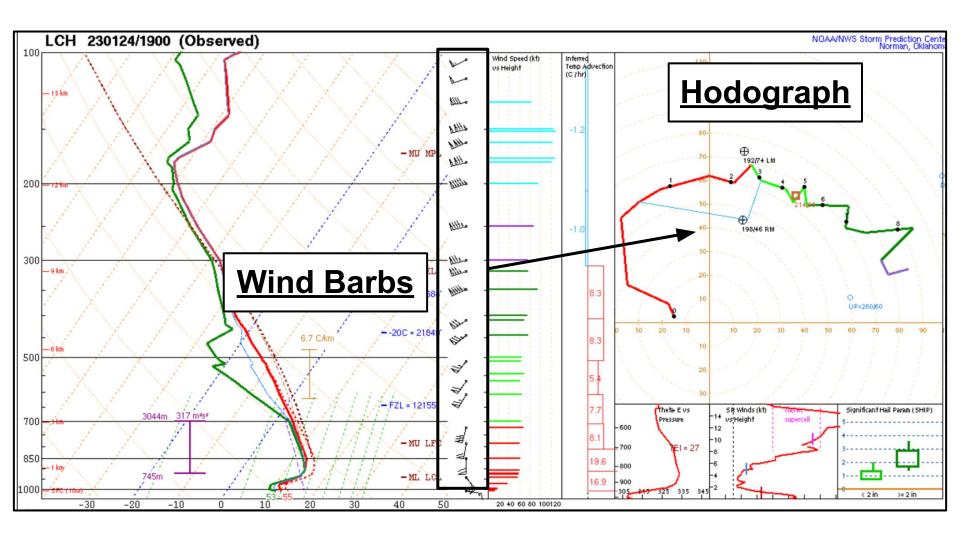
## Lesson 1: Hodograph Fundamentals



**Harry Weinman** – Meteorologist, Storm Prediction Center (harry.weinman@noaa.gov)

Cameron Nixon – Research Scientist, SPC / CIWRO (cameron.nixon@noaa.gov)



## What is a Hodograph?

<u>Hodograph</u>: Plot of vertical wind shear (change in wind speed/direction) with height

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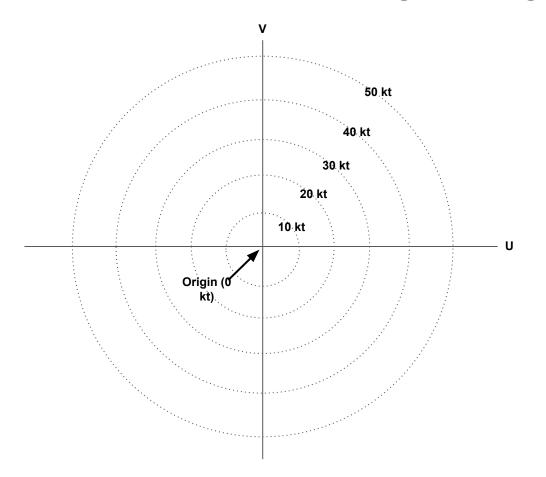
- Highly useful tool for assessing wind shear, storm-relative winds, and horizontal vorticity (more to come on this)

## What is a Hodograph?

<u>Hodograph</u>: Plot of vertical wind shear (change in wind speed/direction) with height

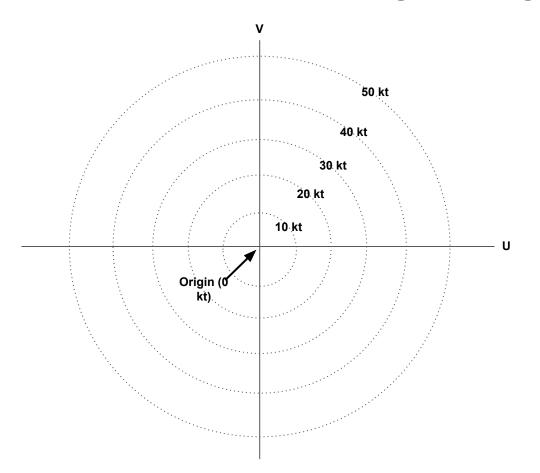
- Highly useful tool for assessing wind shear, storm-relative winds, and horizontal vorticity (more to come on this)

 Hodograph length and shape have direct implications on storm mode, evolution, and overall behavior (more to come on this)

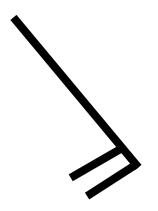


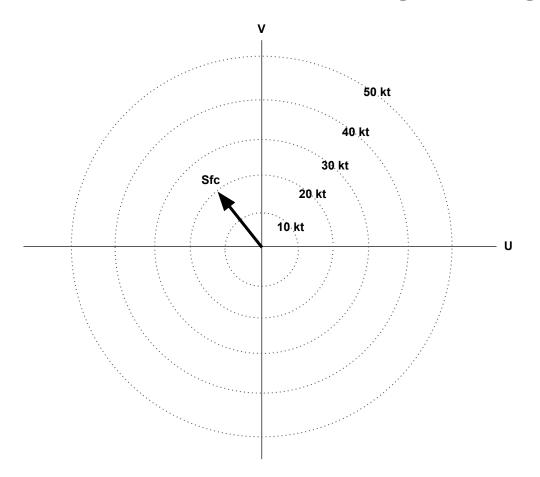
Wind speed increases radially outward from the origin

Axes represent zonal and meridional directions

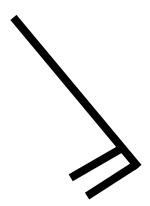


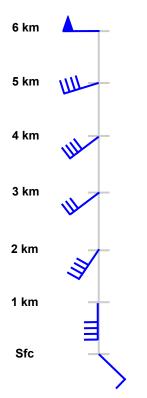
Given a southeasterly surface wind,



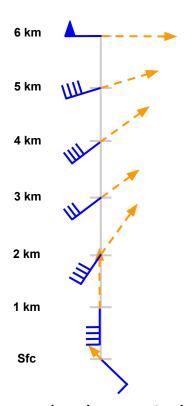


Start at the origin, then draw a wind vector toward the direction the wind is blowing!

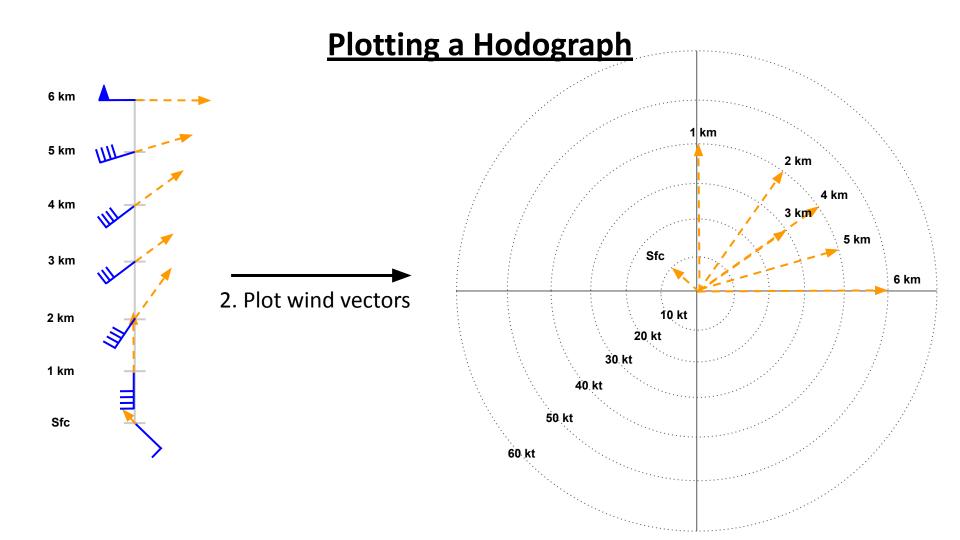




1. Take these wind barbs

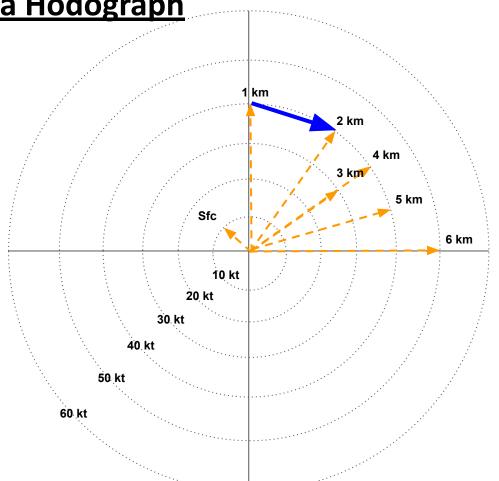


1. Take these wind barbs and convert them to vectors

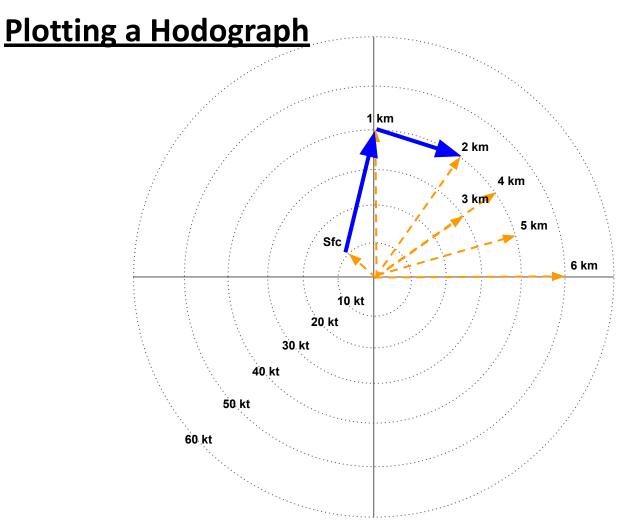


3. Connect the tips of the wind vectors

(This is the vertical shear vector!)

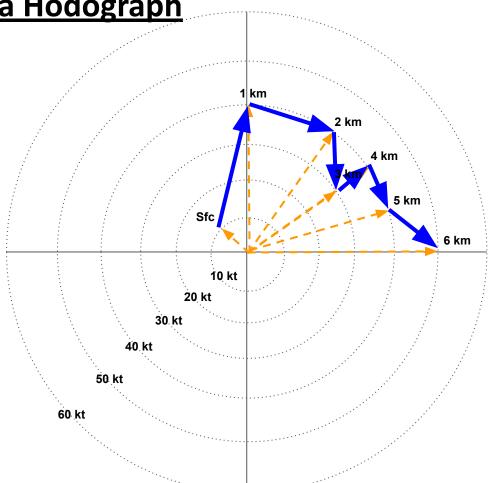


The longer the shear vector, the stronger the shear!

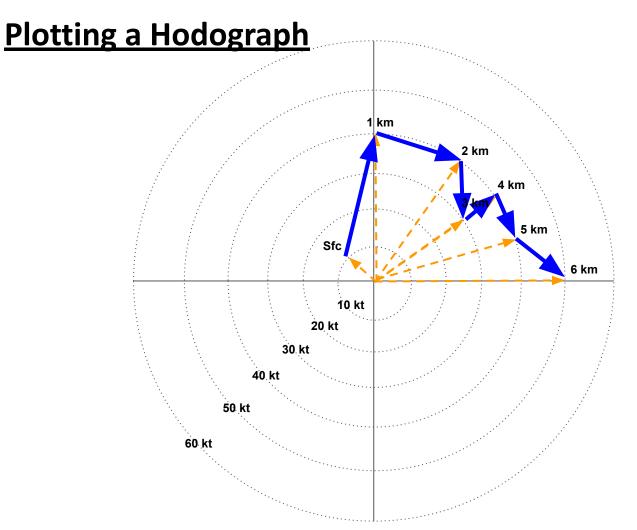


4. Draw all shear vectors

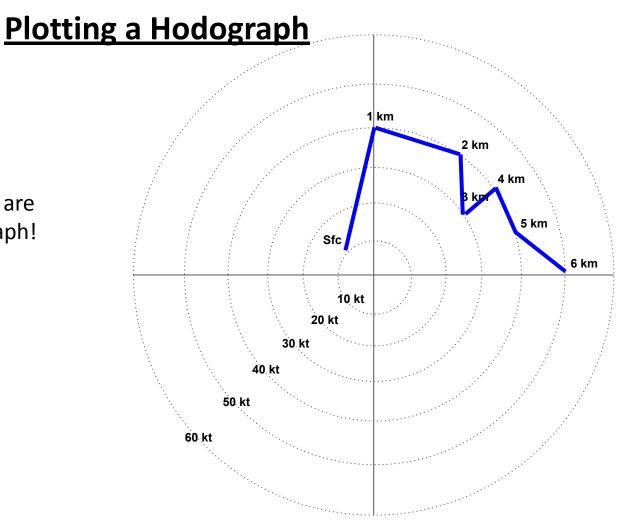
(This traces out the hodograph)



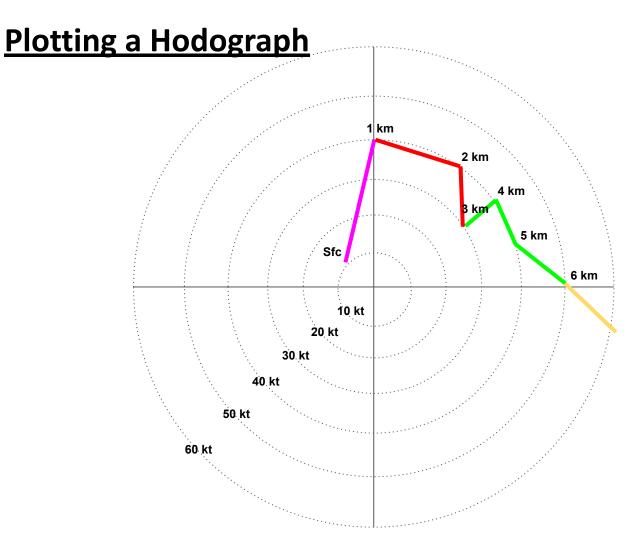
The hodograph is a plot of vertical shear with height!

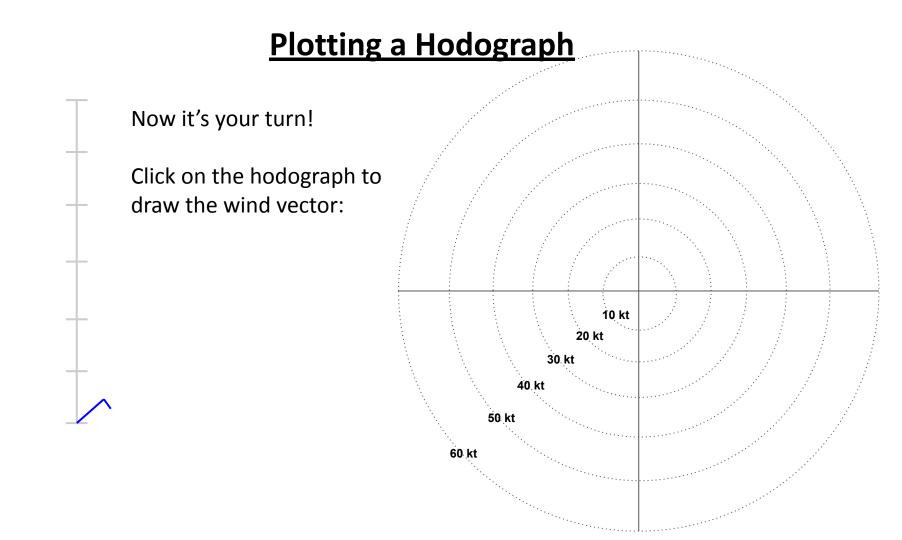


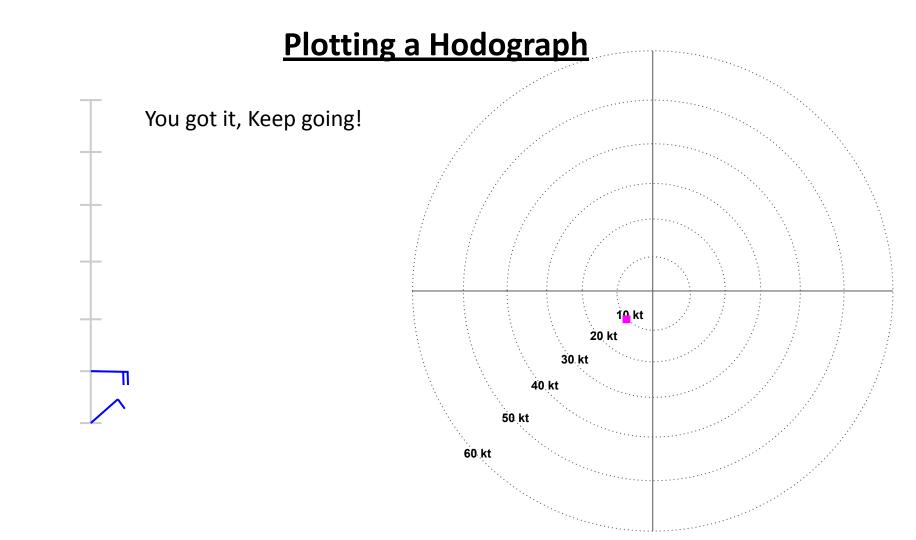
In practice, the wind vectors are not plotted, just the hodograph!

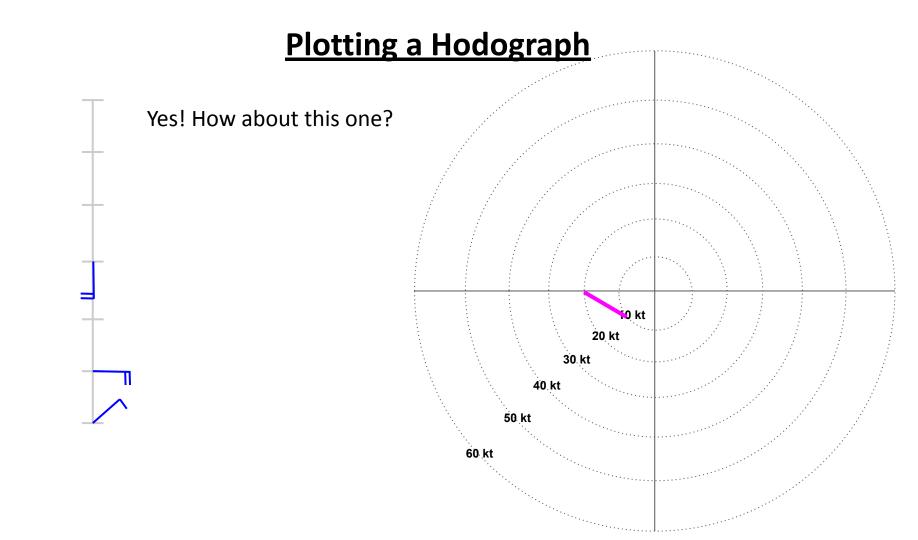


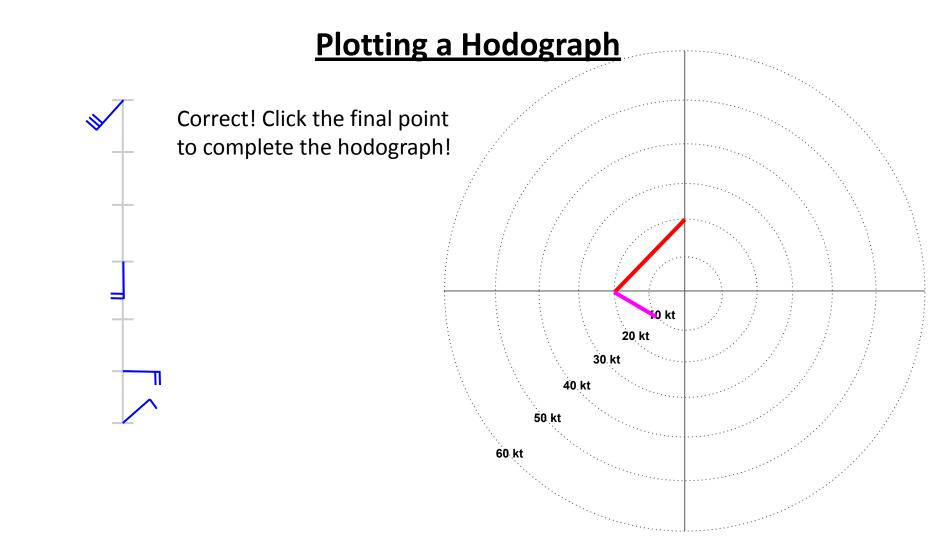
Hodographs are often color-coded by level:

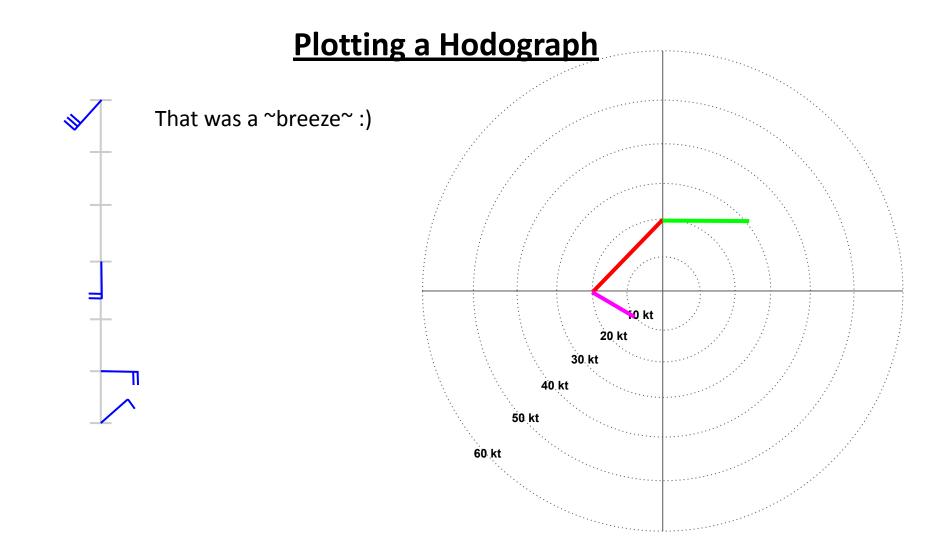






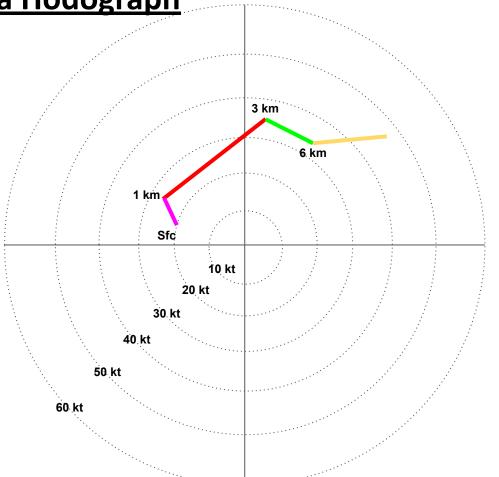






Where is the strongest vertical shear located in this hodograph?

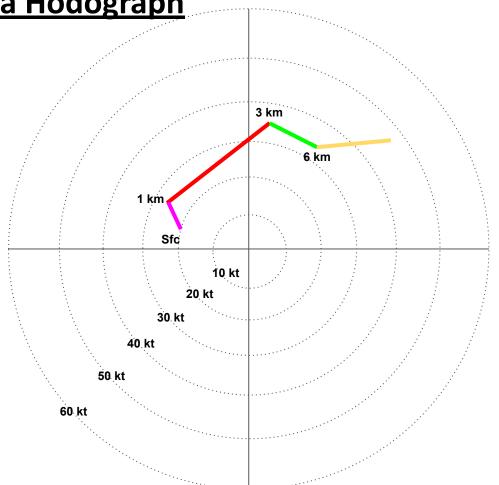
- a. 3-6 km layer
- b. 1-3 km layer
- c. Sfc-1 km layer



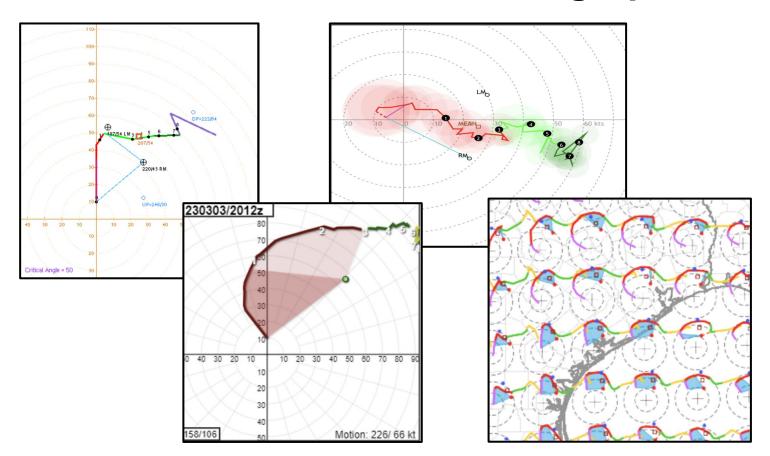
Where is the strongest vertical shear located in this hodograph?

- a. 3-6 km layer
- b. 1-3 km layer
- c. Sfc-1 km layer

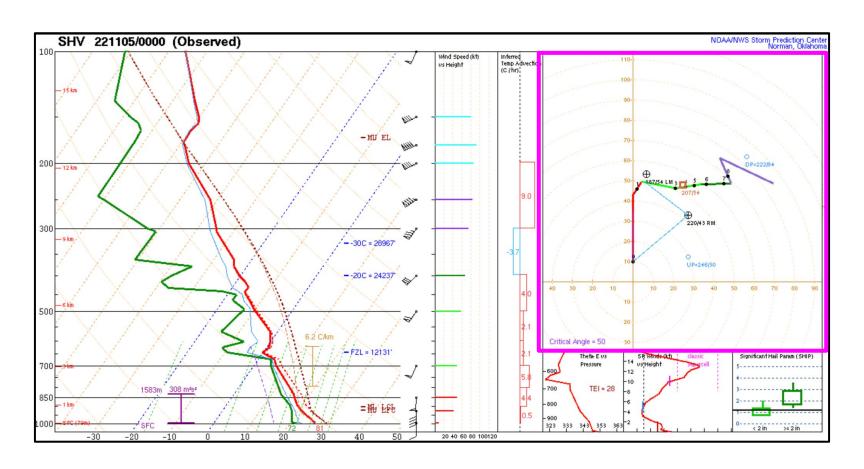
Correct!



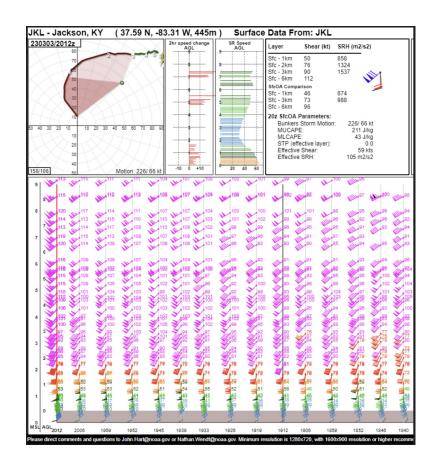
#### Where Can You Find Hodographs?

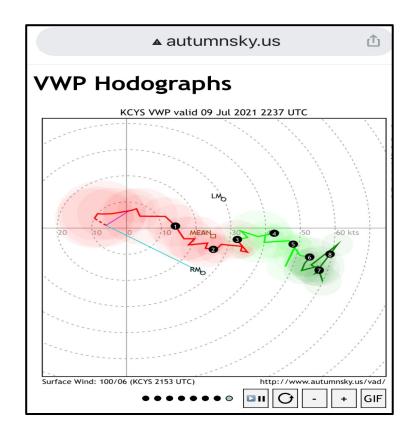


## **Observed Soundings**

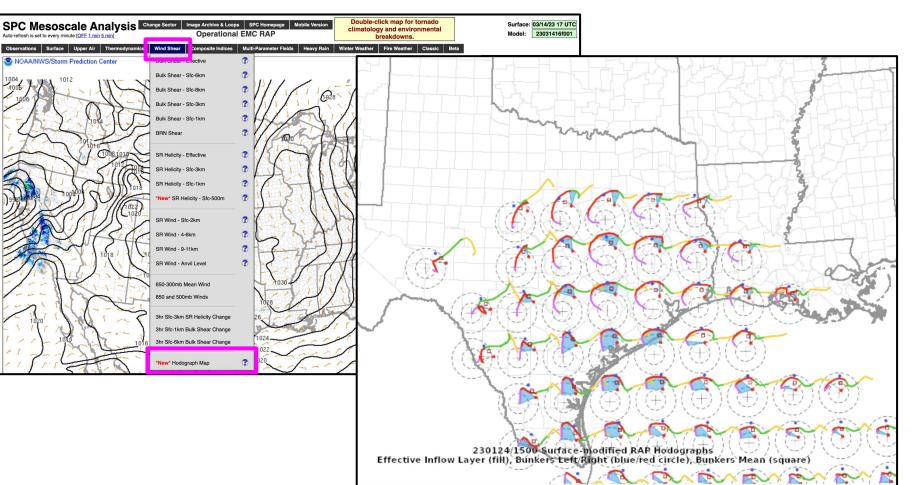


#### **VWP Data**

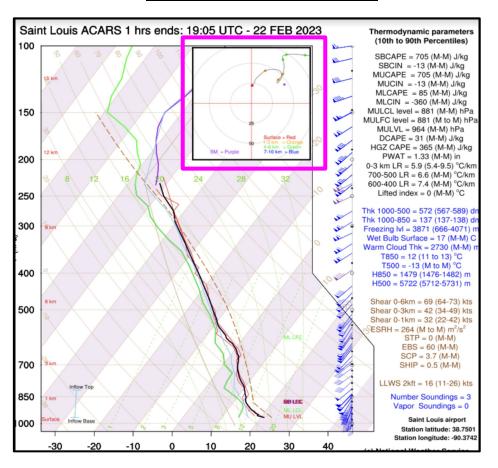




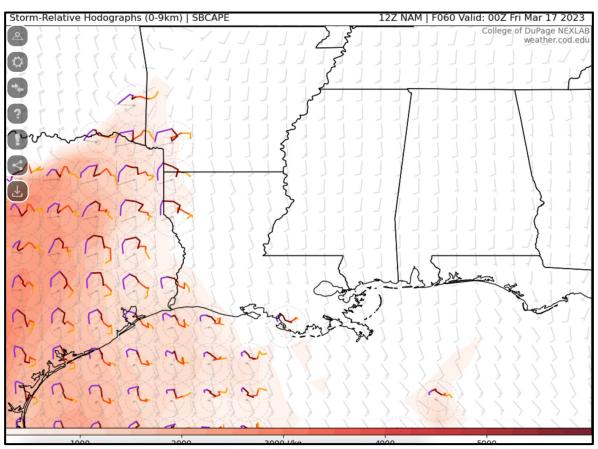
## **SPC Mesoanalysis**



#### **ACARS Data**



## **Model Data**



## **Hodograph Fundamentals Summary**

- 1. Hodograph is a plot of vertical wind shear with height.
- 2. The length and shape of the hodograph have direct implications on storm mode, evolution, and overall behavior (more to come on this).
- 3. Understand how a hodograph is plotted.

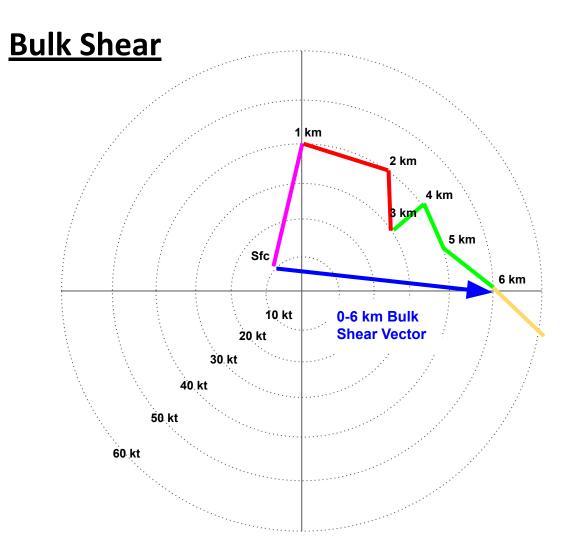
# Lesson 2: Bulk Shear

226/40 RM

Cameron Nixon – Research Scientist, SPC / CIWRO (cameron.nixon@noaa.gov)

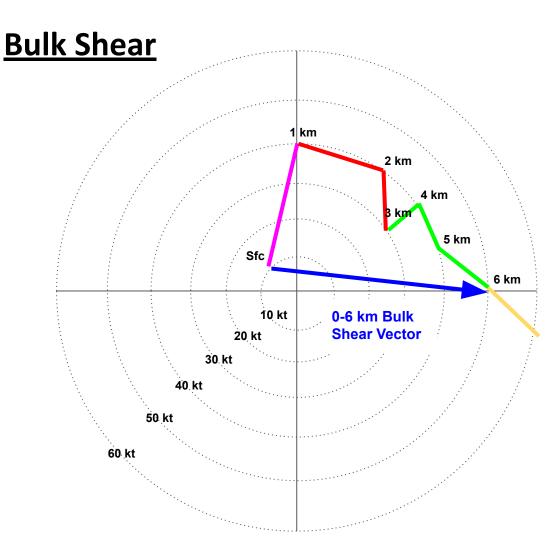
Harry Weinman – Meteorologist, Storm Prediction Center (harry.weinman@noaa.gov)

**Bulk Shear** is the difference in wind between *any* two levels on the hodograph

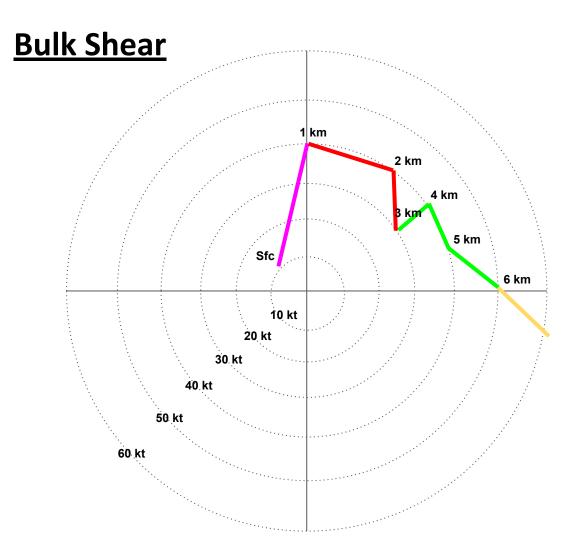


**Bulk Shear** is the difference in wind between *any* two levels on the hodograph

In this case, we have ~60 kts of 0-6 km bulk shear

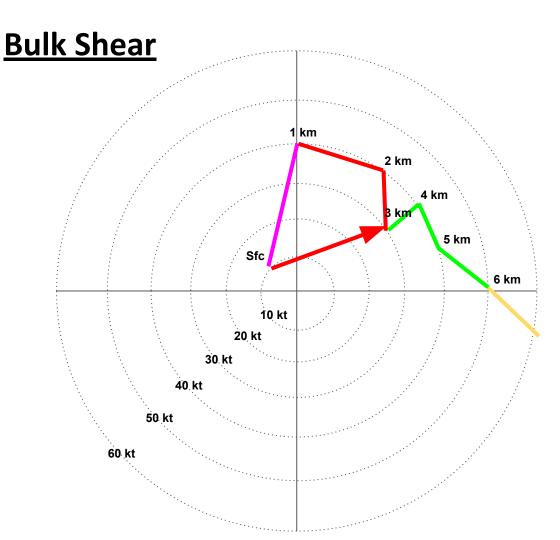


Draw a 0-3 km shear vector!

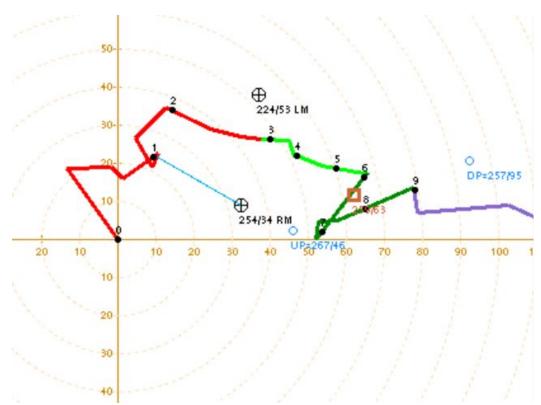


Draw a 0-3 km shear vector!

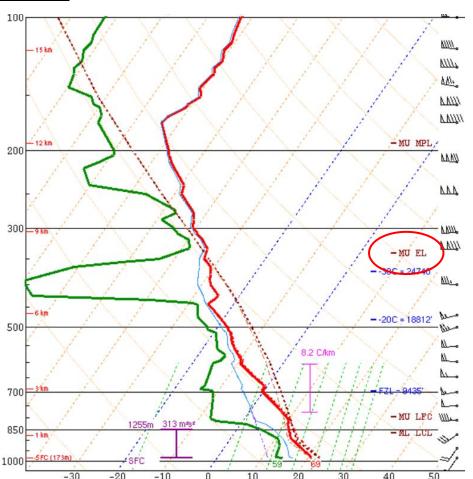
Excellent!!



How much of the hodograph matters for convective storms?

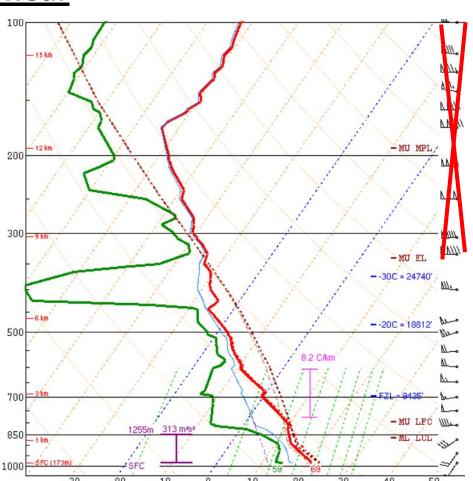


Generally, shear only matters up to the equilibrium level.



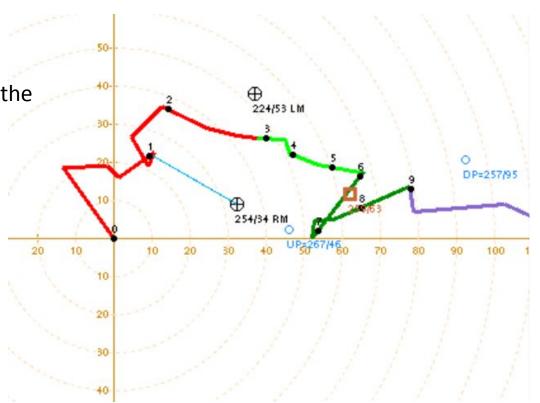
Generally, shear only matters up to the equilibrium level.

In this case, we can neglect shear above 8 km.



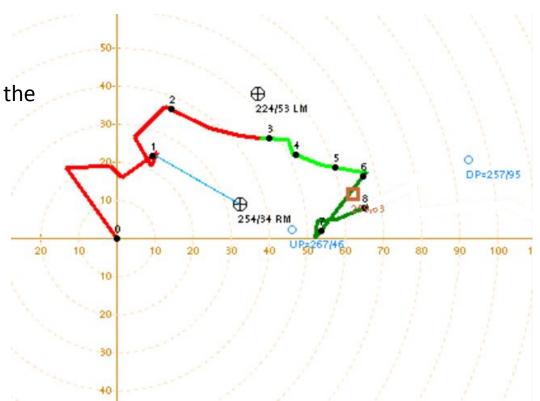
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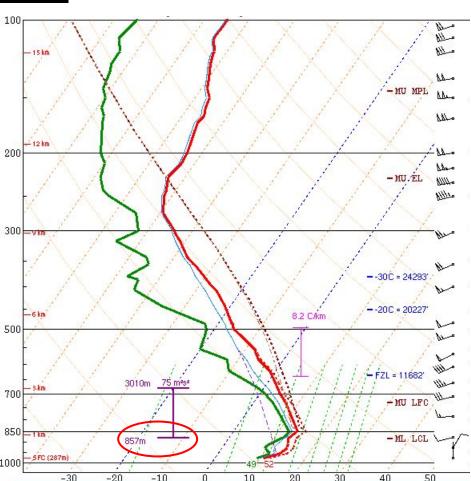


Generally, shear only matters up to the equilibrium level.

In this case, we can neglect shear above 8 km.

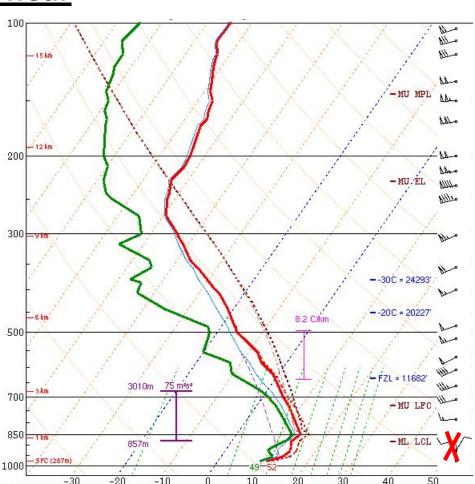


Generally, shear only matters above the **effective inflow base**.



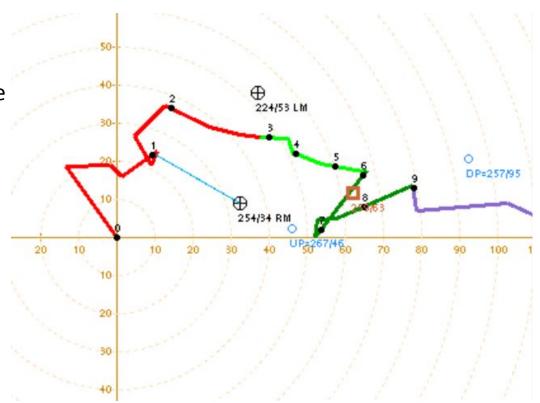
Generally, shear only matters above the **effective inflow base**.

In this case, we can neglect shear below 800 m



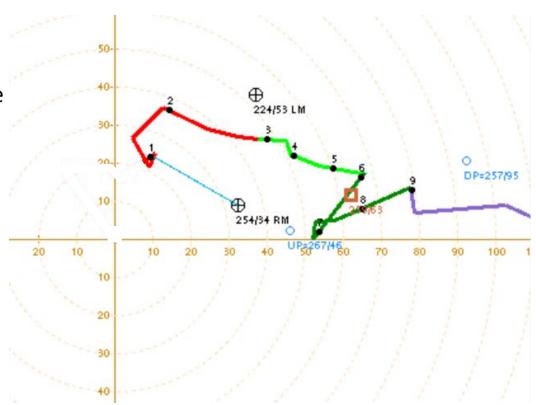
Generally, shear only matters above the **effective inflow base**.

In this case, we can neglect shear below 1 km.



Generally, shear only matters above the **effective inflow base**.

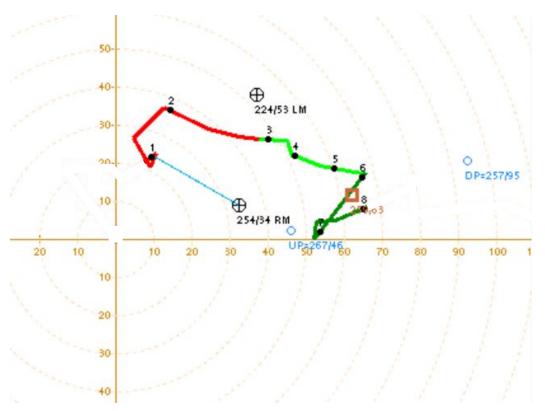
In this case, we can neglect shear below 1 km.



**Effective Bulk Shear** takes into account *both* the equilibrium

level and effective inflow base.

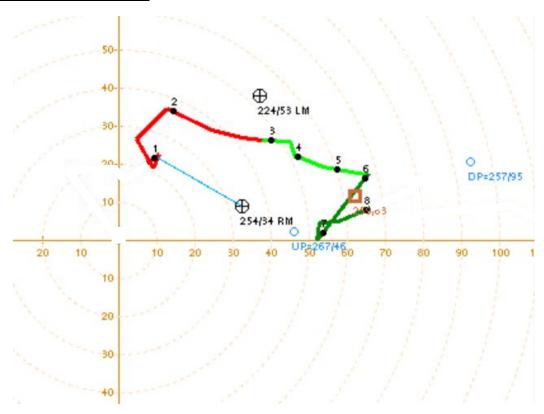
# **Bulk Shear**



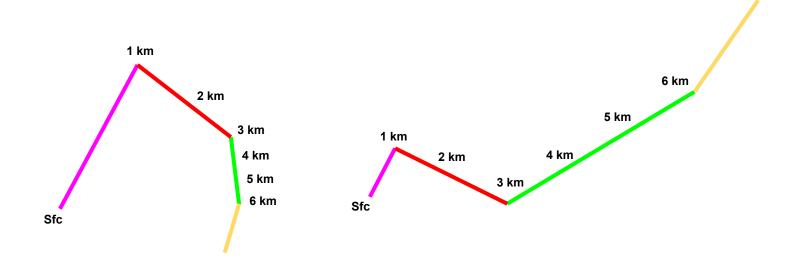
**Effective Bulk Shear** takes into account *both* the equilibrium level and effective inflow base.

It's useful in a wide variety of scenarios, including shallow or elevated storms.

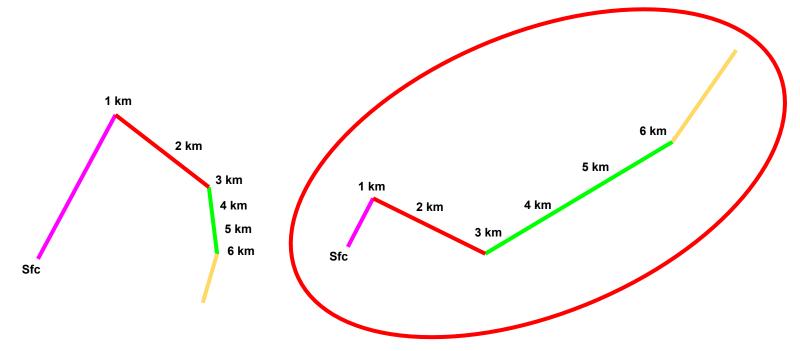
#### **Bulk Shear**



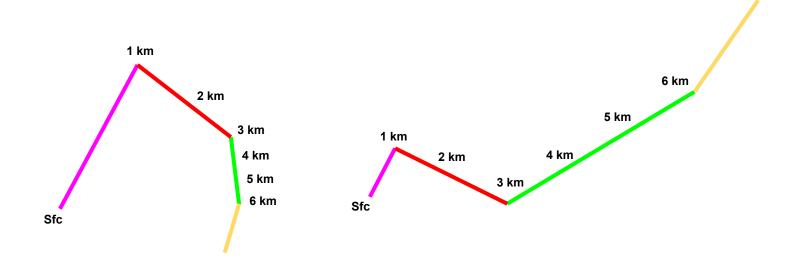
Which hodograph has stronger 0-6 km shear?



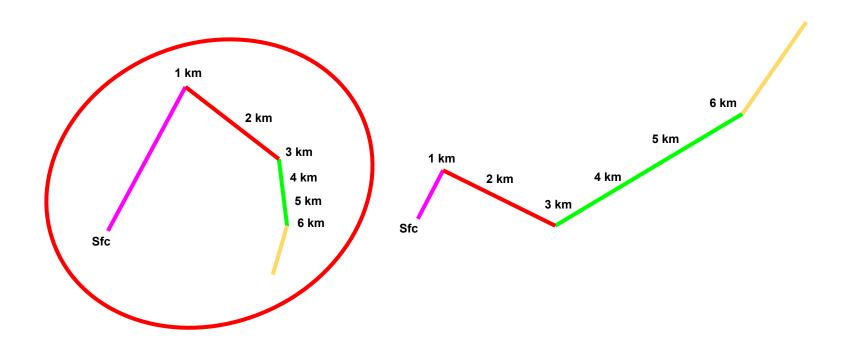
Which hodograph has stronger 0-6 km shear?



Which hodograph has stronger 0-1 km shear?

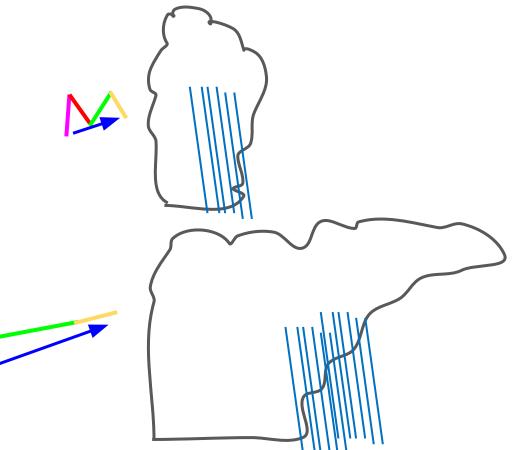


#### Nice job!



Stronger bulk shear can create convective storms that are:

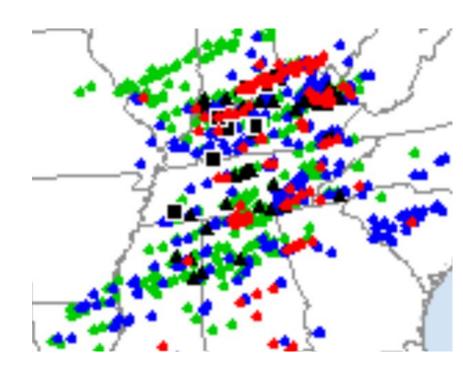
- Stronger
- Larger
- More organized
- Longer-lived



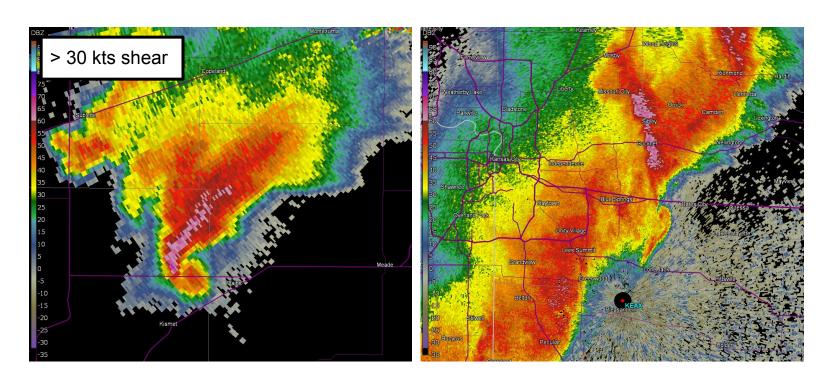
Stronger bulk shear can create convective storms that are:

- Stronger
- Larger
- More organized
- Longer-lived

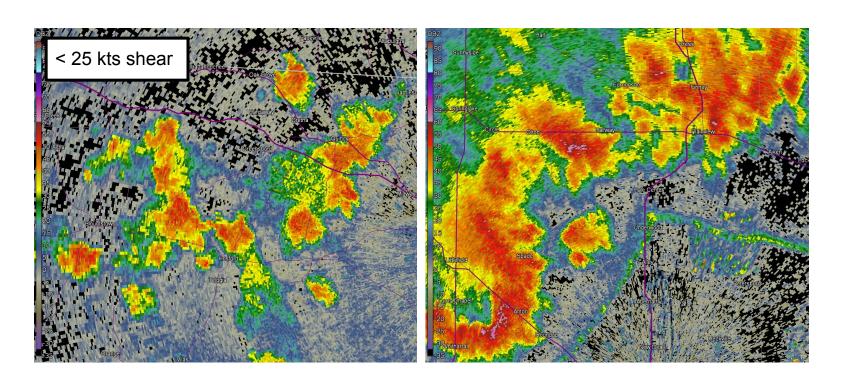
...and more hazardous!



Stronger shear can support supercells and organized convective systems



Weak shear may only support brief, weakly-organized storms



# **Bulk Shear Summary**

- 1. Bulk shear is the difference in wind between two levels
- 2. Shear matters to storms between the Eff. Inflow Base and EL
- 3. Effective bulk shear is a good "go-to" bulk shear parameter
- 4. Stronger shear can support more hazardous storms

# Lesson 3: Hodograph Shape

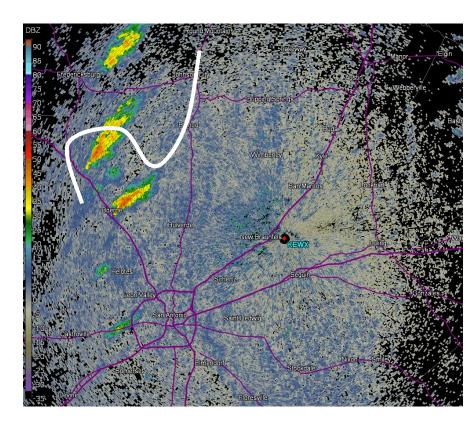
Left-Mover Right-Mover

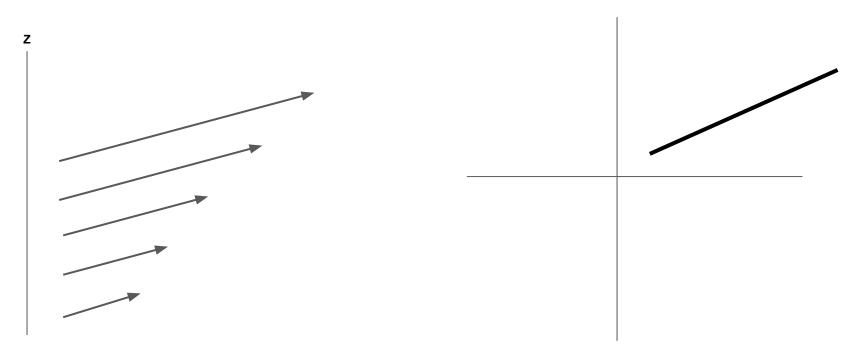
**Harry Weinman** – Meteorologist, Storm Prediction Center (harry.weinman@noaa.gov)

Cameron Nixon – Research Scientist, SPC / CIWRO (cameron.nixon@noaa.gov)

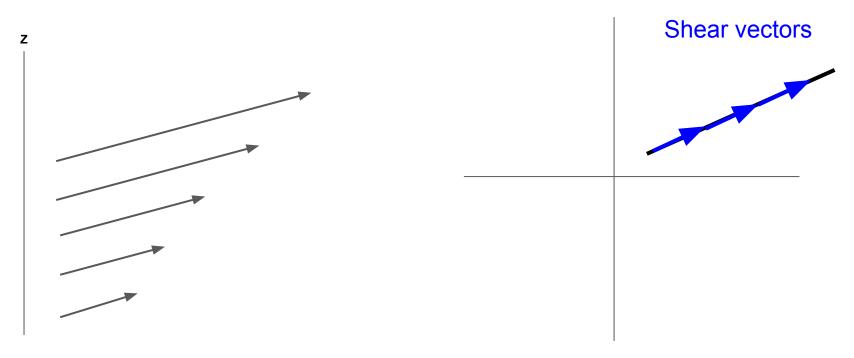
#### **Hodograph Shapes**

The shape of the hodograph can help us determine what types of storms are favored and how they will evolve

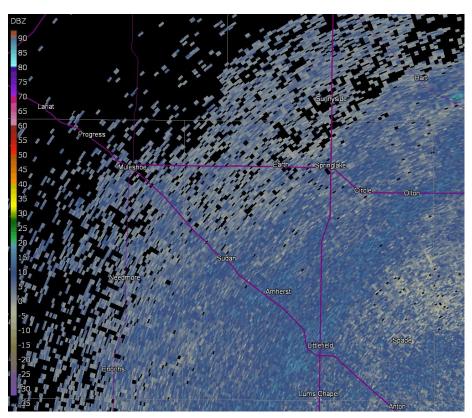




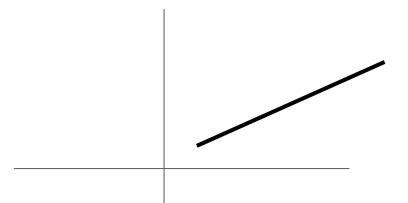
Unidirectional wind profile, straight hodograph

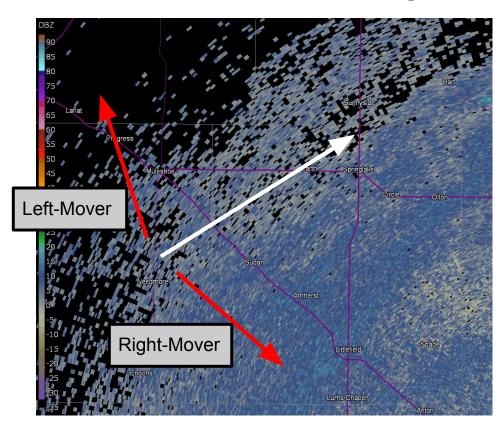


Unidirectional wind profile, straight hodograph

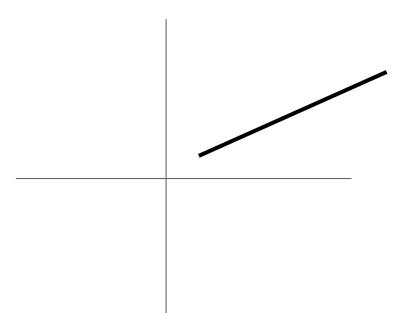






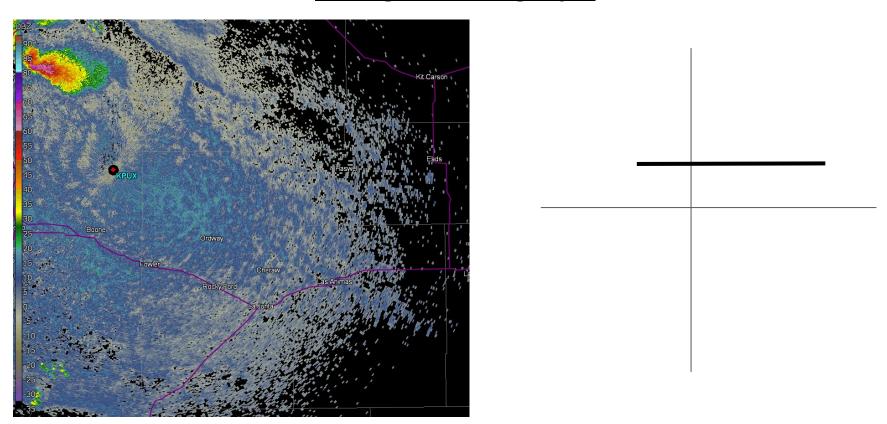






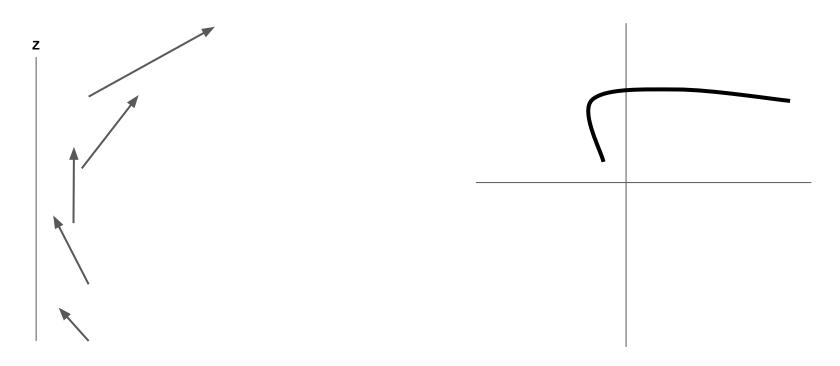


Veering wind profile, still a straight hodograph!



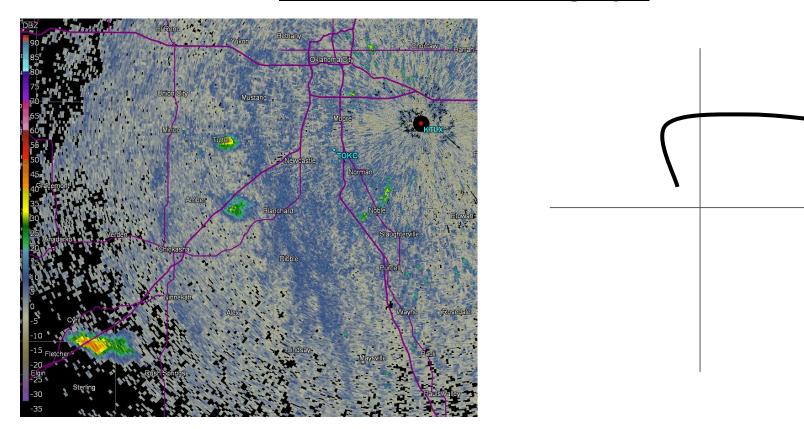
Supports both left-moving and right-moving storms (and many cell interactions!)

#### **Quarter-Circle Hodograph**



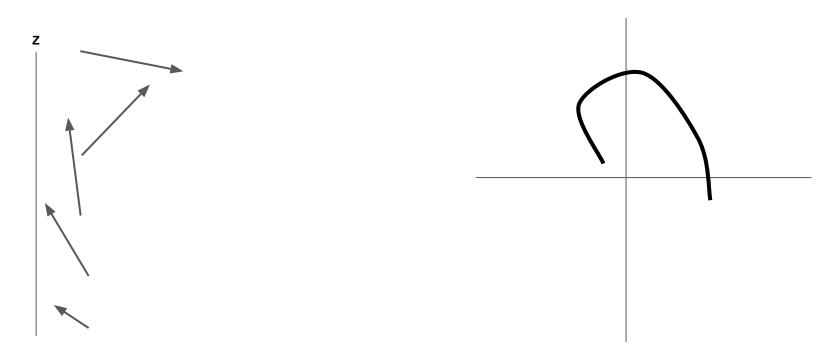
Weakly veering wind profile, clockwise-curved hodograph. Found in warm-air advection regimes.

# **Quarter-Circle Hodograph**



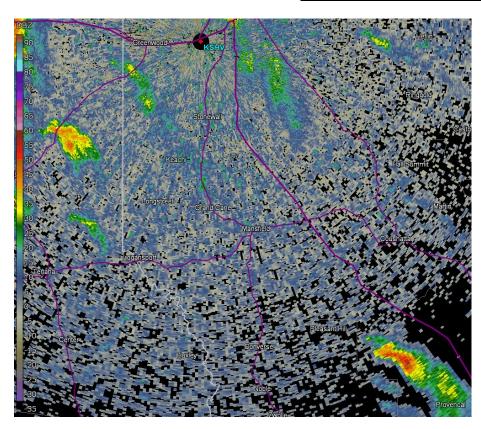
Supports splitting storms, with dominant right-movers

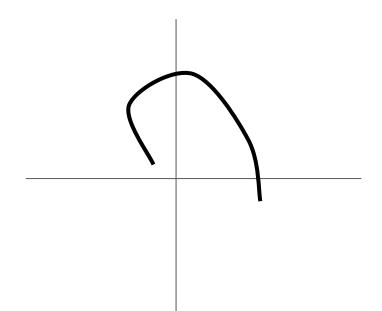
#### **Half-Circle Hodograph**



Strongly veering wind profile, clockwise-curved hodograph. Found in strong warm-air advection regimes and warm-core systems.

# **Half-Circle Hodograph**





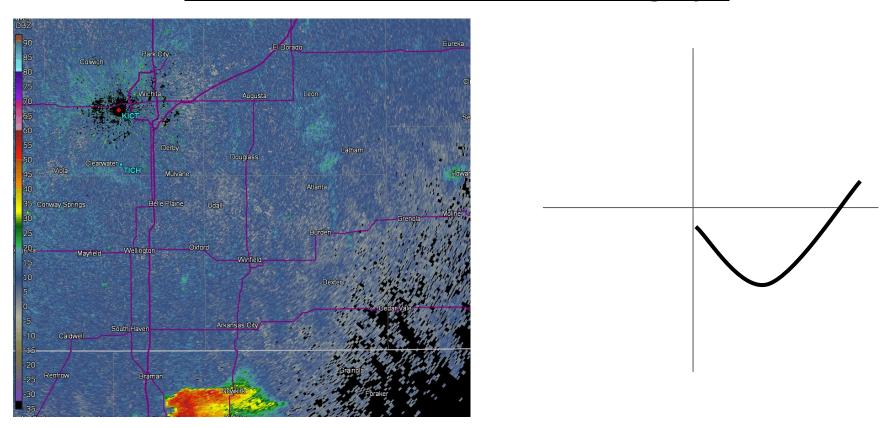
Typically supports right-moving storms only

#### **Counter-Clockwise Curved Hodograph**

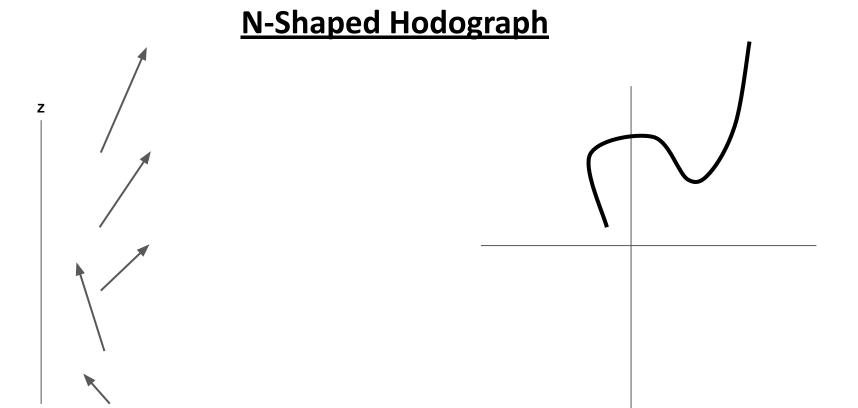


Backing wind profile, counter-clockwise curved hodograph. Found in cold-air advection regimes.

#### **Counter-Clockwise Curved Hodograph**



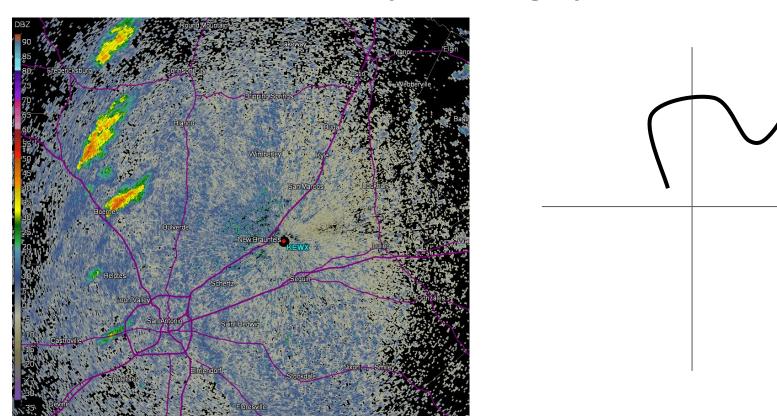
Supports splitting storms, with dominant left-movers



Veer-back wind profile.

Found in the wakes of midlevel troughs, CFAs, and in cold-core systems.

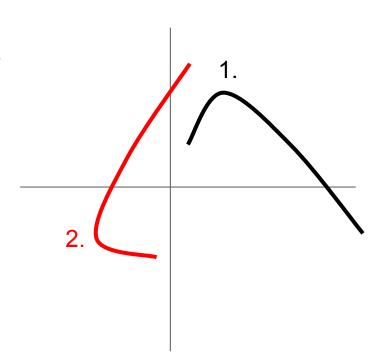
## **N-Shaped Hodograph**



Supports splitting storms, with dominant right-movers

Compared to hodograph #1, how might storms behave differently with hodograph #2 (all else being equal)?

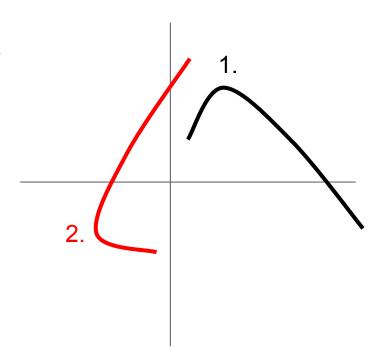
- a. Storms would split faster
- b. Left-movers would be more dominant
- c. Storms wouldn't behave much differently



Compared to hodograph #1, how might storms behave differently with hodograph #2 (all else being equal)?

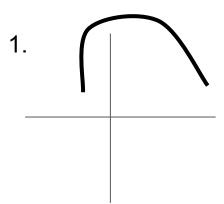
- a. Storms would split faster
- b. Left-movers would be more dominant
- c. Storms wouldn't behave much differently

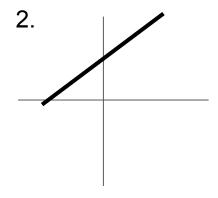
You got it!



Which hodograph would favor a right-moving storm?

- a. Hodograph 2
- b. Hodograph 1
- Neither of these hodographs favor RM

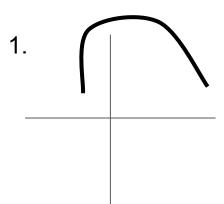


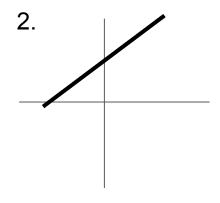


Which hodograph would favor a right-moving storm?

- a. Hodograph 2
- b. Hodograph 1
- c. Neither of these hodographs favor RM

**Great work!** 





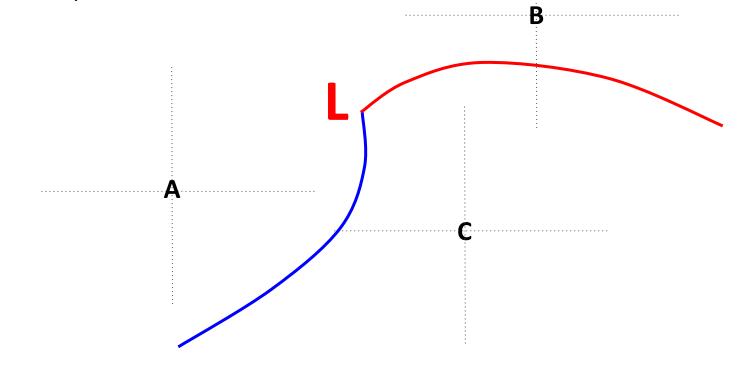
Synoptic pattern is important!

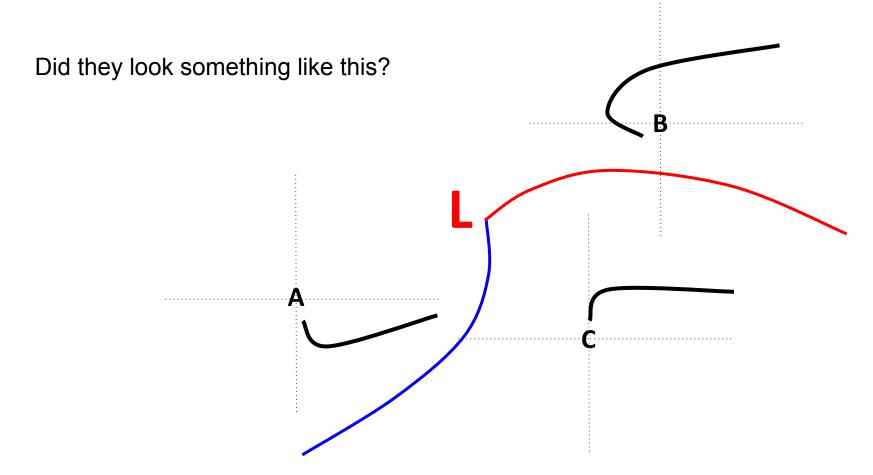
Warm-air advection: veering wind profile

**Cold-air advection:** backing wind profile

Think thermal wind!

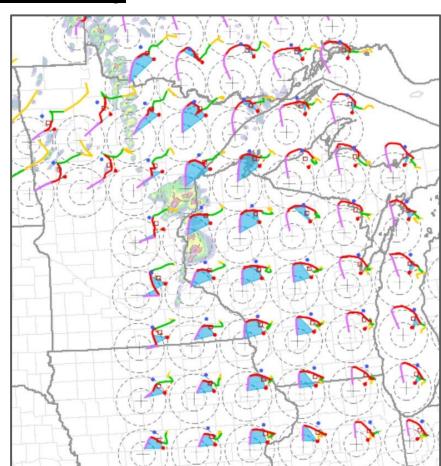
How would you expect a hodograph to look at each point?



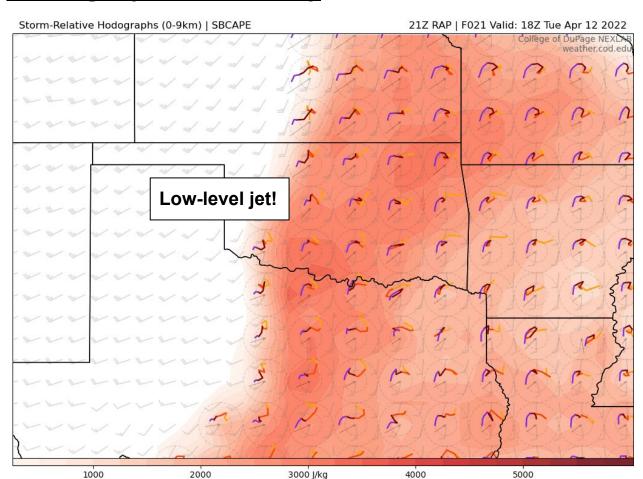


Hodograph shape can vary considerably over a given area

Use a hodograph map for this!



Hodograph shape can change quickly over time!



# **Hodograph Shape Summary**

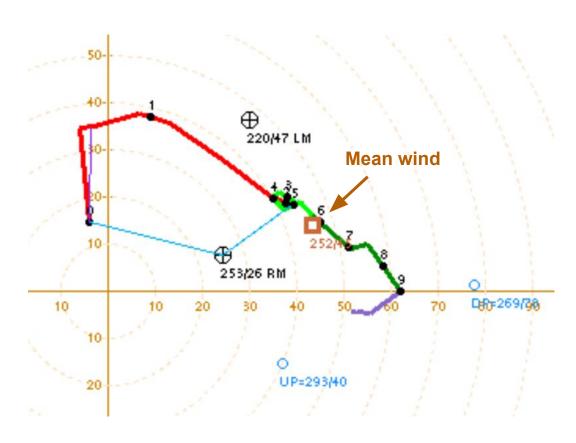
- 1. Hodograph shape is directly related to the vertical wind profile.
- 2. It is the shape of the hodograph that matters, not the orientation of the hodograph itself (all else being equal).
- Straight hodographs favor splitting storms, clockwise curvature favors RM, counterclockwise curvature favors LM.
- 4. Synoptic and mesoscale pattern/features influence hodograph shape.

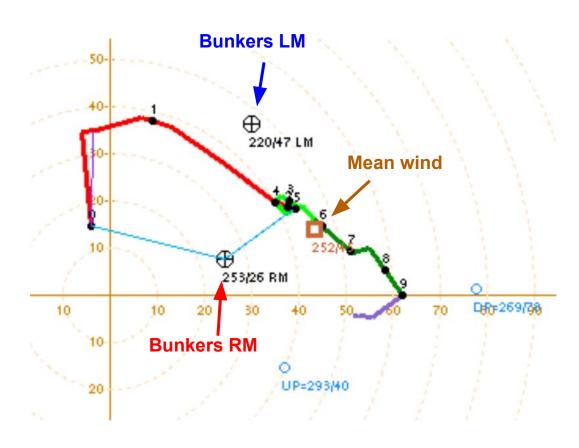
# Lesson 4: Storm Motion

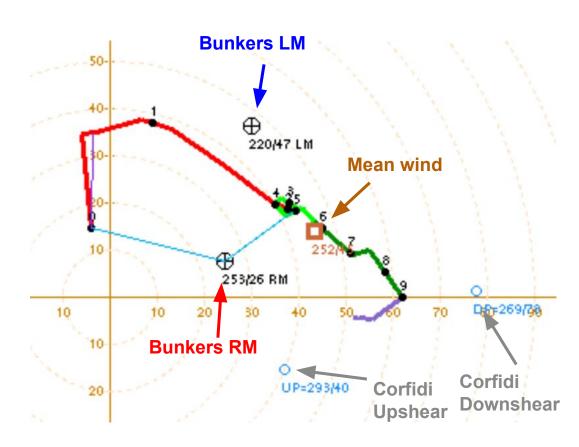
**Bunkers RM** 253/26 RM

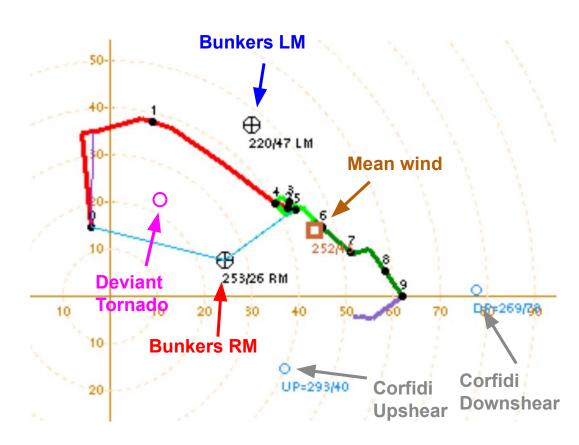
**Cameron Nixon** – Research Scientist, SPC / CIWRO (cameron.nixon@noaa.gov)

Harry Weinman – Meteorologist, Storm Prediction Center (harry.weinman@noaa.gov)



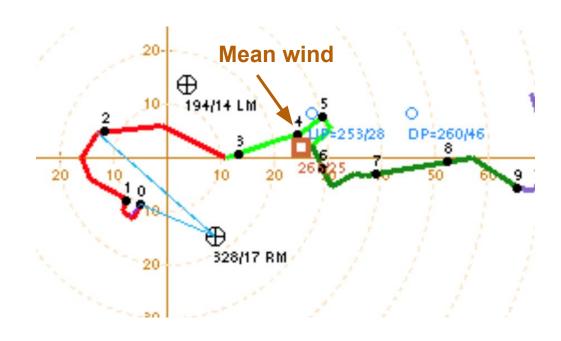






#### **Mean Wind**

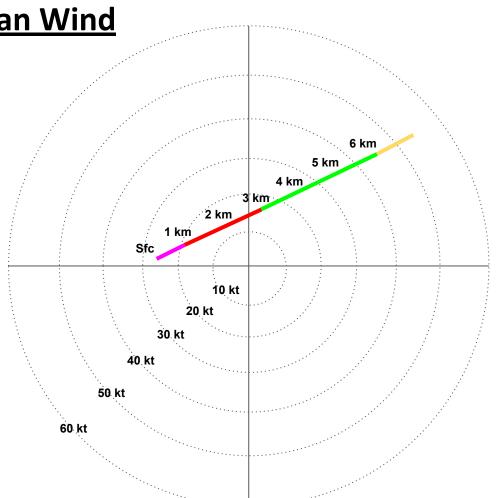
Initially, a good approximation for storm motion is the mean wind



**Mean Wind** 

How to find the mean wind:

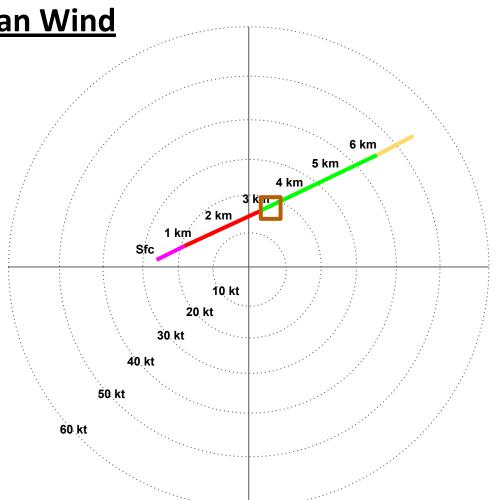
Take the average of points along the hodograph (in this case, we'll do 0-6 km)



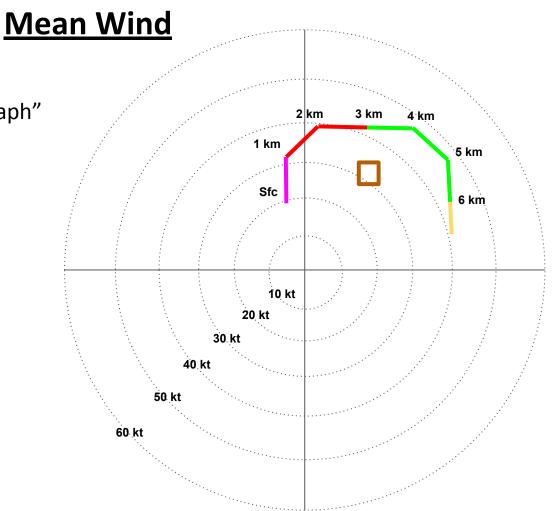
**Mean Wind** 

How to find the mean wind:

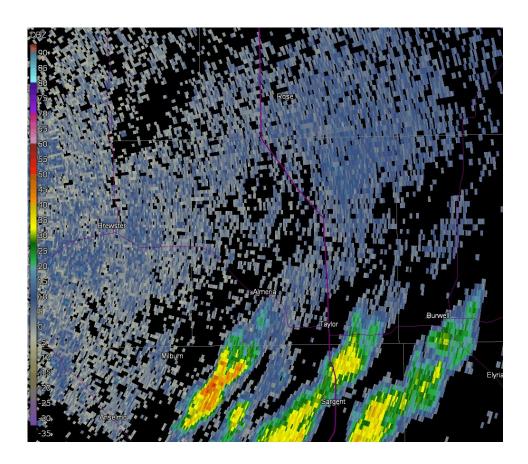
Take the average of points along the hodograph (in this case, we'll do 0-6 km)

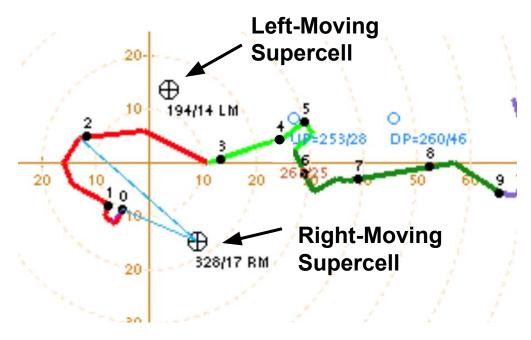


For a curved hodograph, the mean wind may lie "off the hodograph"



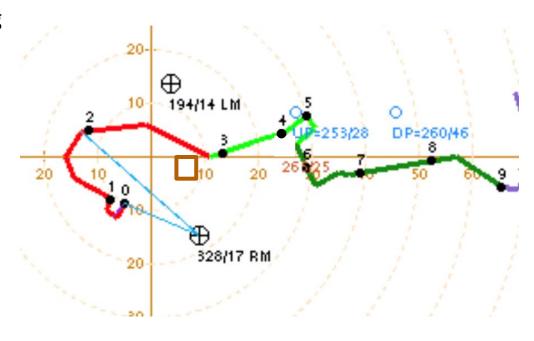
Supercells can propagate deviantly from the mean wind, due to their internal dynamics



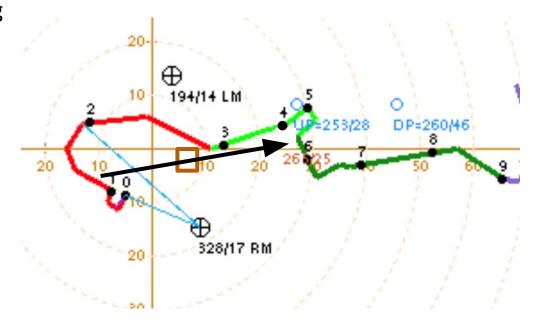


We can predict supercell motion using the Bunkers method:

1. Find the mean wind

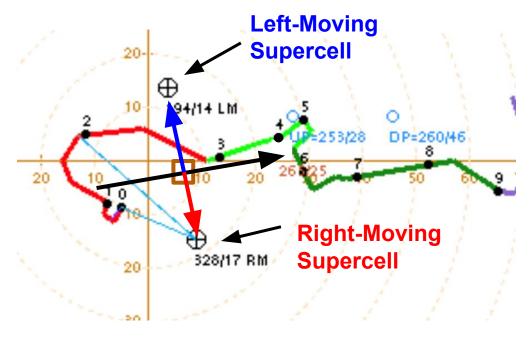


- 1. Find the mean wind
- 2. Draw the deep-layer shear vector\*



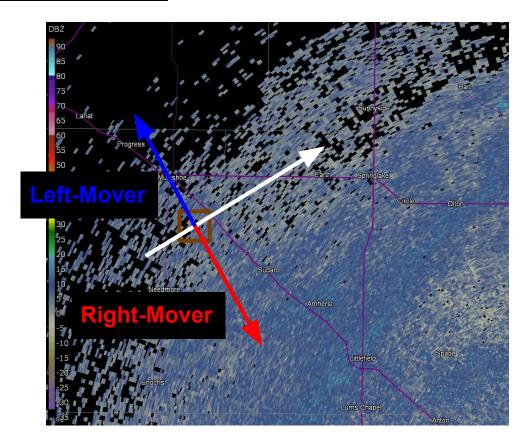
<sup>\*</sup>these are approximations

- 1. Find the mean wind
- 2. Draw the deep-layer shear vector\*
- 3. Deviate left / right from this vector (by about 15 kts\*)



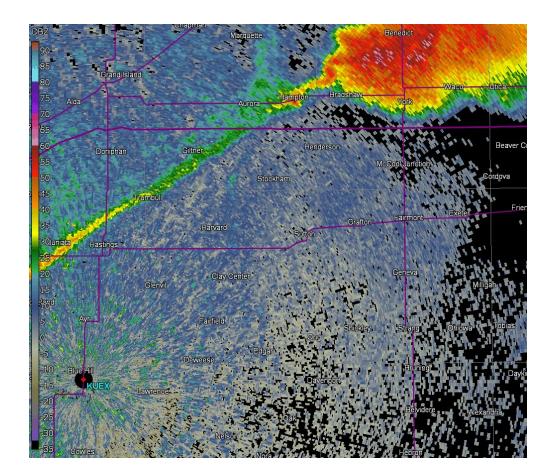
<sup>\*</sup>these are approximations

- 1. Find the mean wind
- 2. Draw the deep-layer shear vector\*
- 3. Deviate left / right from this vector (by about 15 kts\*)



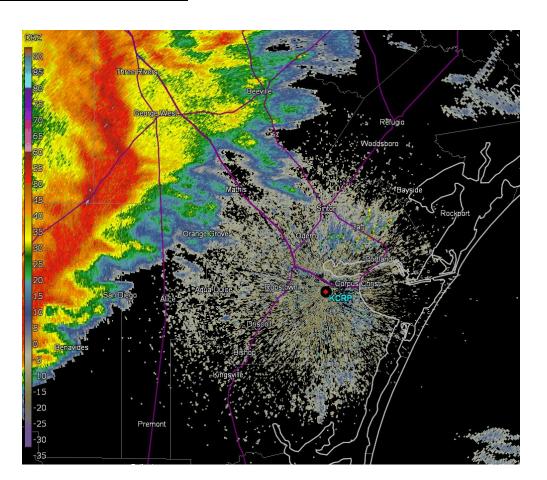
<sup>\*</sup>these are approximations

Deviant storm motion also associated with:
-boundaries



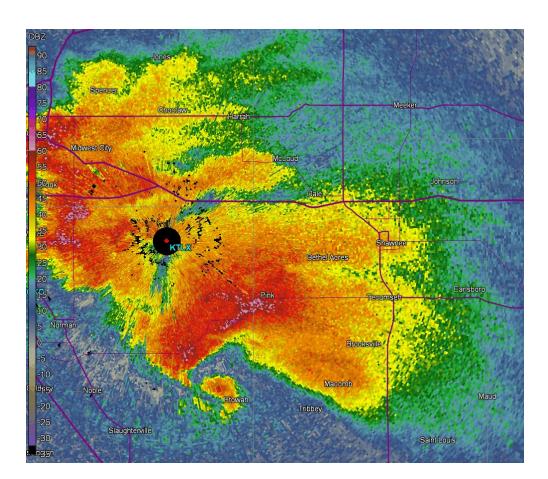
Deviant storm motion also associated with:

- -boundaries
- -cold pools



Deviant storm motion also associated with:

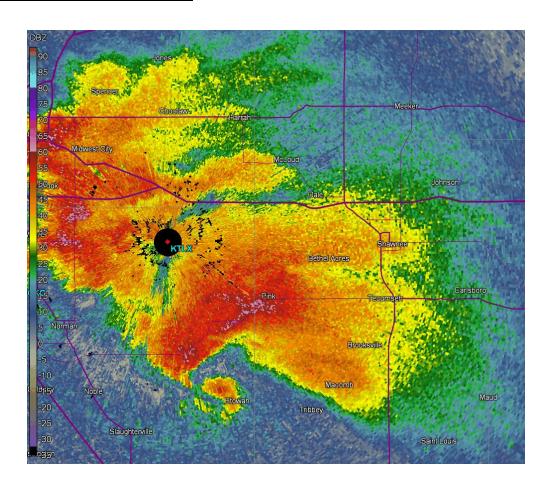
- -boundaries
- -cold pools
- -cell interactions



Deviant storm motion also associated with:

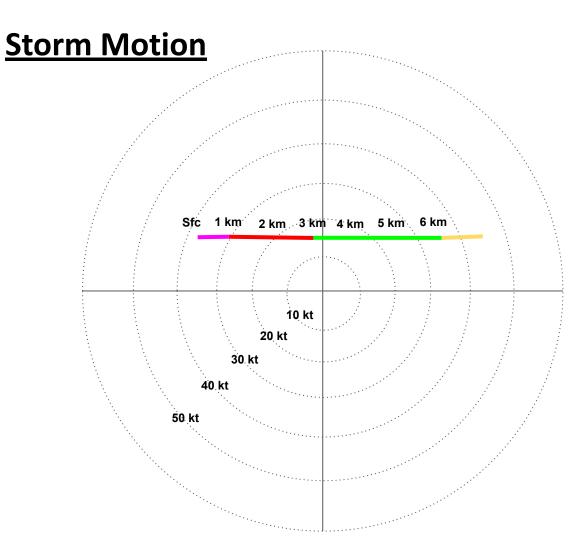
- -boundaries
- -cold pools
- -cell interactions

These can **not** be predicted by the hodograph!



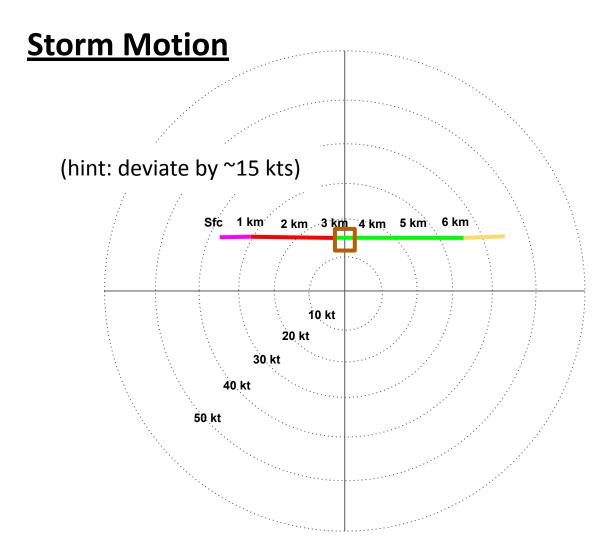
Now you try:

Estimate the mean wind for this hodograph:



Now you try:

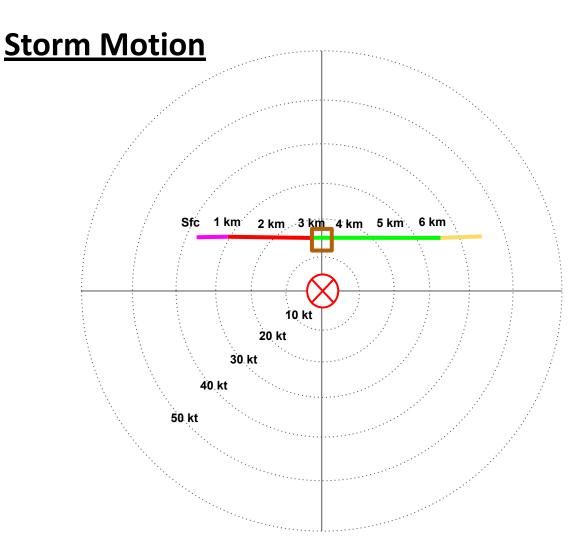
Now, estimate the motion of a *right-moving* supercell:



Now you try:

Now, estimate the motion of a *right-moving* supercell:

You were "right":)



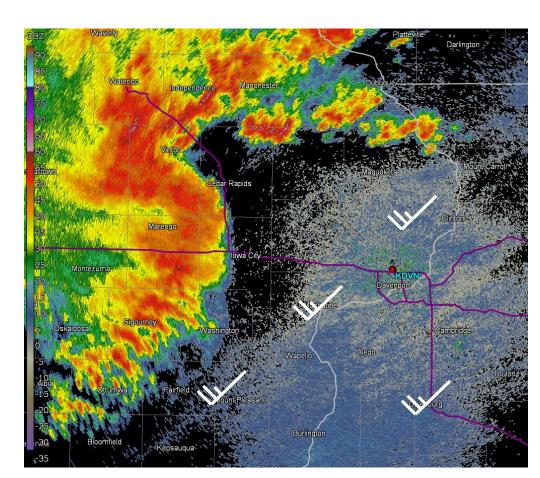
#### **Storm Motion**

What about this scenario?

You're using Bunkers' storm motion to estimate the potential path of an intensifying bow echo.

Is this a good idea?

- a. Yes
- b. No



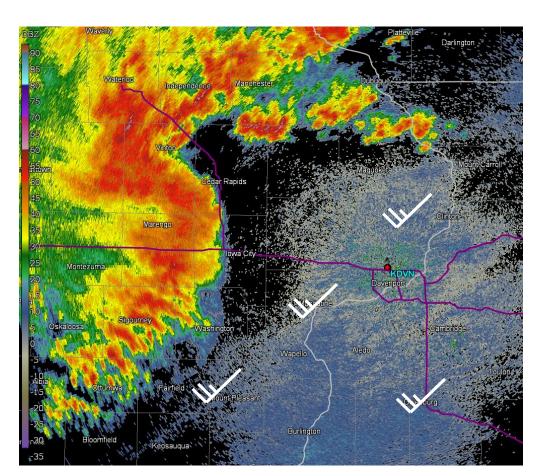
#### **Storm Motion**

What about this scenario?

You're using Bunkers' storm motion to estimate the potential path of an intensifying bow echo.

Is this a good idea?

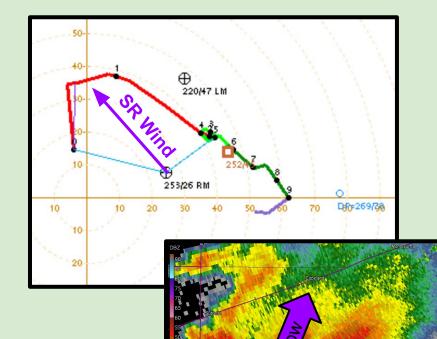
- a. Yes
- b. No



# **Storm Motion Summary**

- 1. Initially, the mean wind is a good estimate of storm motion
- 2. Bunkers' storm motion is an estimation of supercell motion
- Boundaries, convective systems, and cell interactions may also affect storm motion

# Lesson 5: Storm-Relative Winds

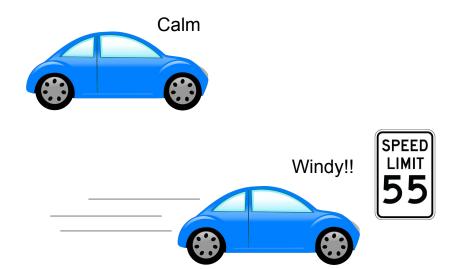


**Cameron Nixon** – Research Scientist, SPC / CIWRO (cameron.nixon@noaa.gov)

Harry Weinman – Meteorologist, Storm Prediction Center (harry.weinman@noaa.gov)

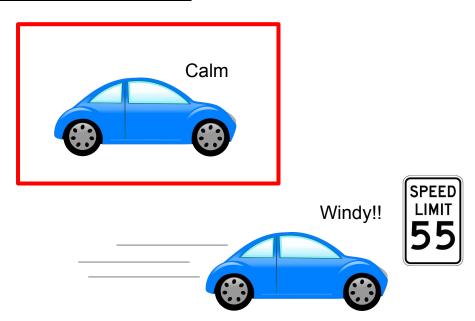
#### What are storm-relative winds?

These are the winds a storm actually feels when it's moving!



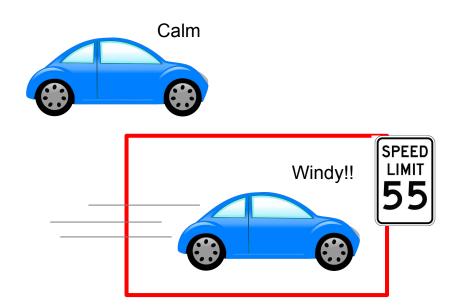
#### What are storm-relative winds?

These are the winds a storm actually feels when it's moving!



#### What are storm-relative winds?

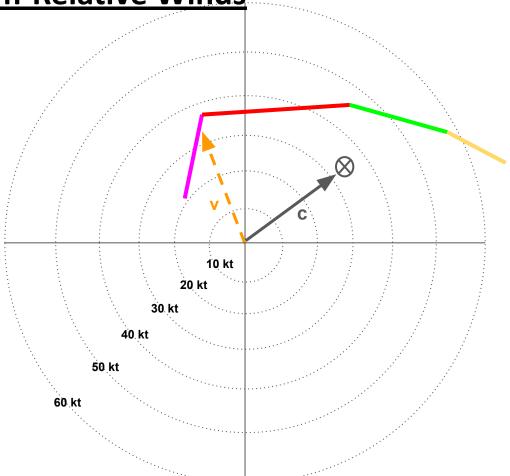
These are the winds a storm actually feels when it's moving!



How do we find storm-relative winds?

**V = Ground Relative Wind** 

**C** = Storm Motion

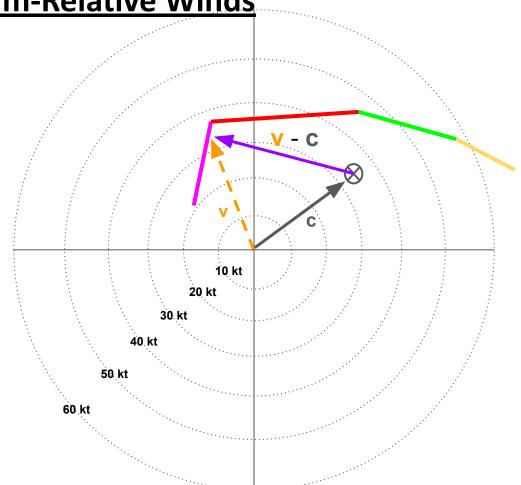


How do we find storm-relative winds?

**V = Ground Relative Wind** 

**C** = Storm Motion

V - C = Storm Relative Wind!



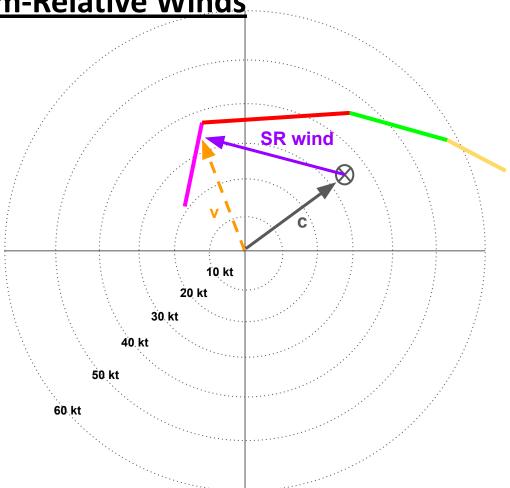
How do we find storm-relative winds?

V = Ground Relative Wind

**C** = Storm Motion

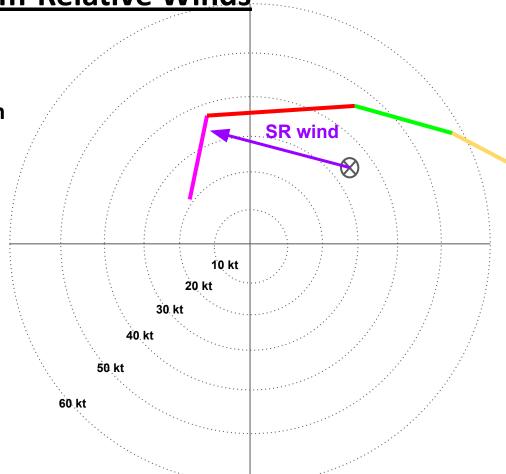
V - C = Storm Relative Wind!

(we are subtracting out storm motion!)



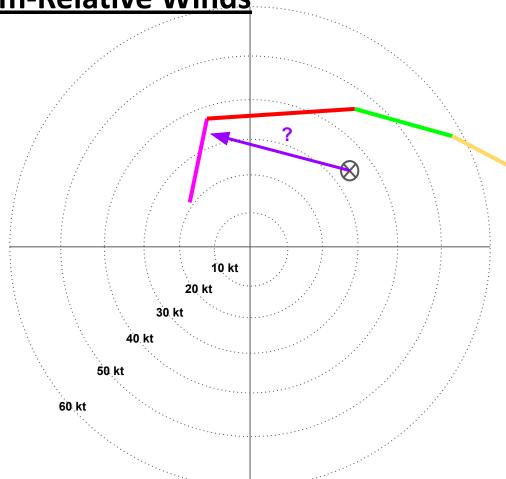
How do we find storm-relative winds?

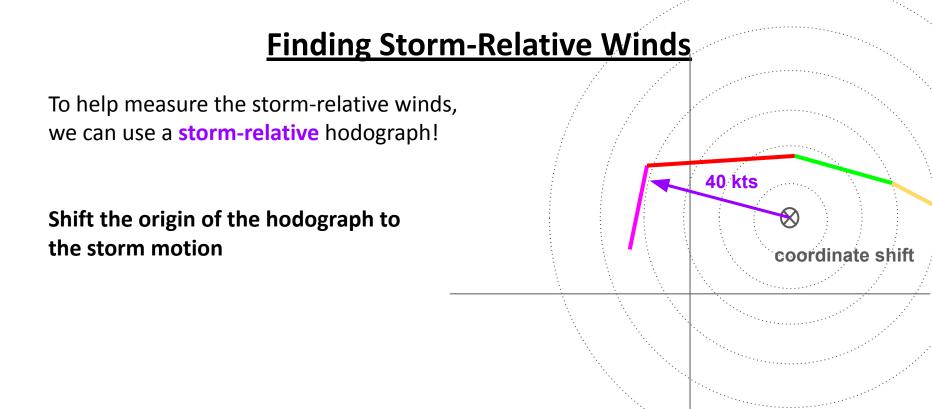
Simply draw a vector from **storm motion** to desired point on the hodograph



How strong is this storm-relative wind?

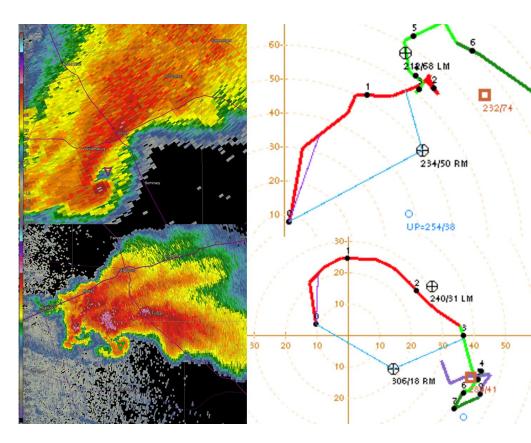
It's hard to tell





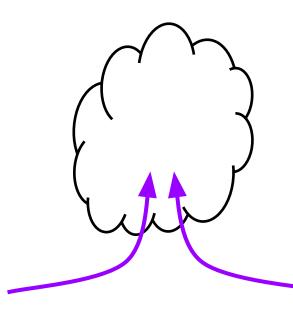
Why is **storm-relative wind** important?

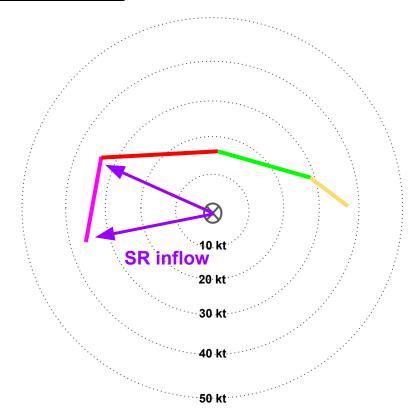
It directly governs storm structure!



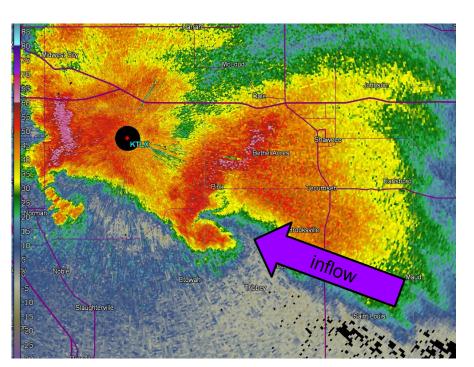
## and Storm Structure

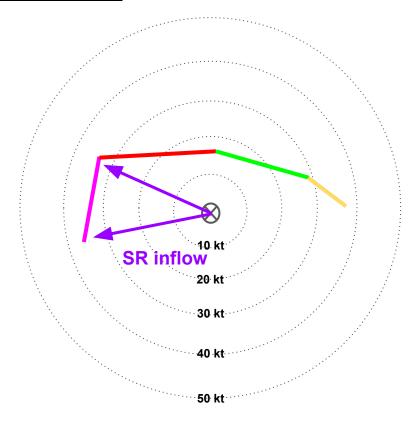
**SR Inflow** is the wind in the lower levels of a storm, where mass enters the storm



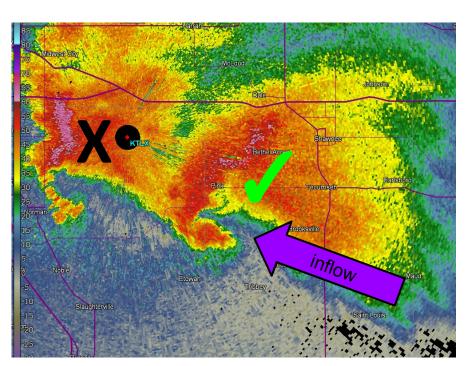


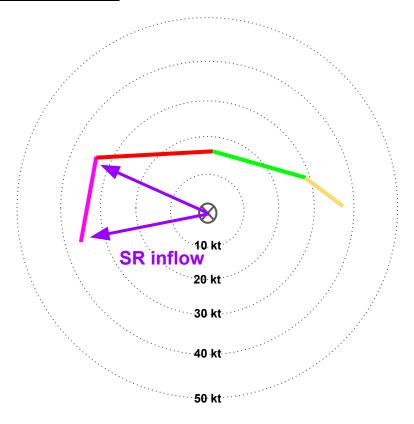
The direction of inflow can help determine whether or not a storm has access to "clean", unstable air





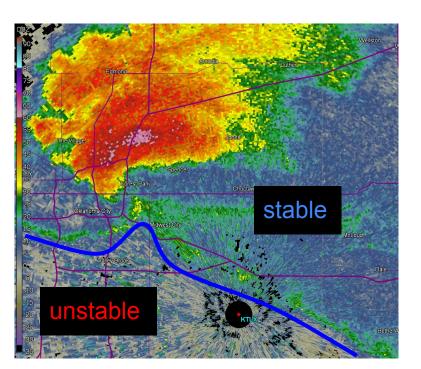
The direction of inflow can help determine whether or not a storm has access to "clean", unstable air

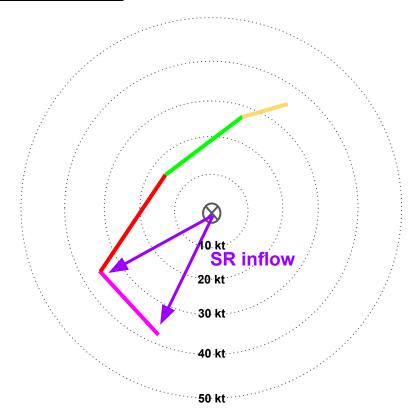




This storm is "anchored" to a boundary. Is its inflow air from the unstable or stable side?

a. Stable b. Unstable

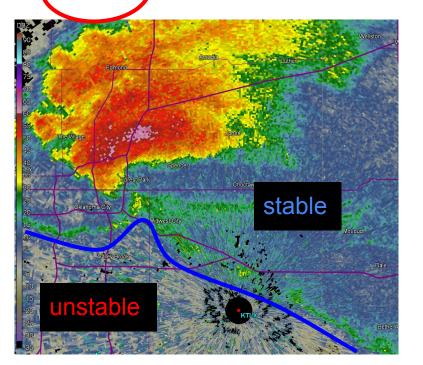


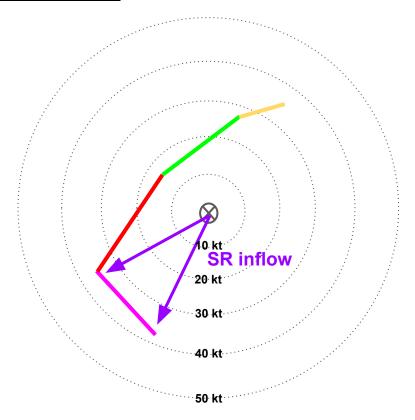


This storm is "anchored" to a boundary. Is its inflow air from the unstable or stable side?

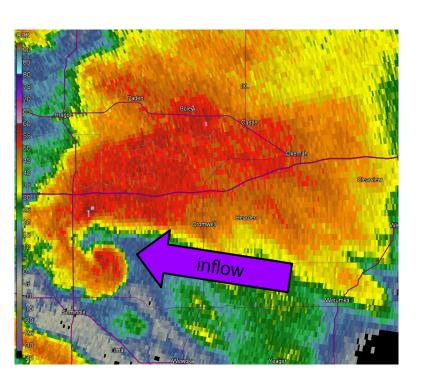
a. Stable

b. Unstable





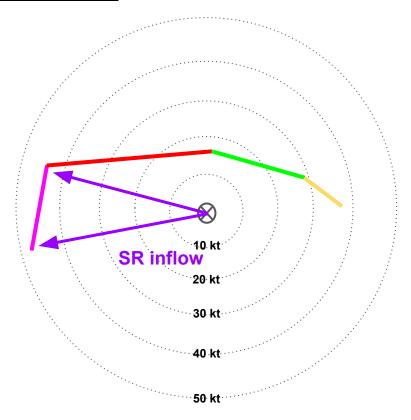
Stronger inflow can sustain larger storms with more precipitation



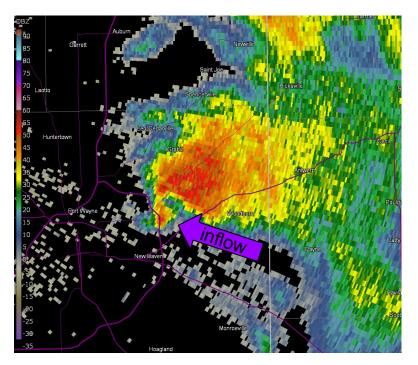


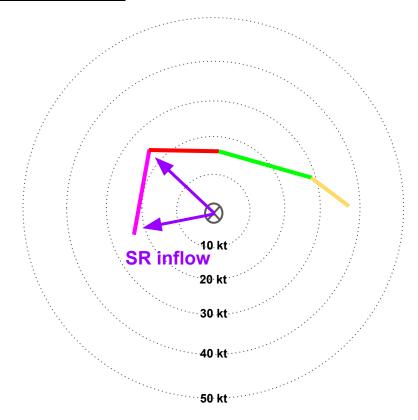
Inflow > 40 kts can support particularly large supercells





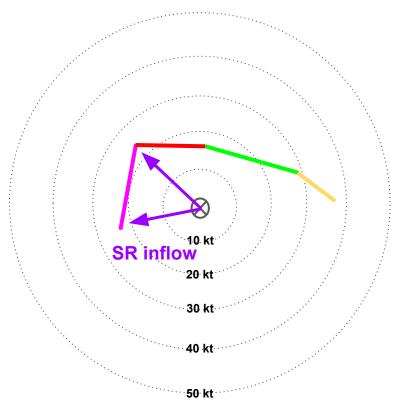
Weaker inflow often only supports smaller updrafts or "mini-supercells" with less precipitation



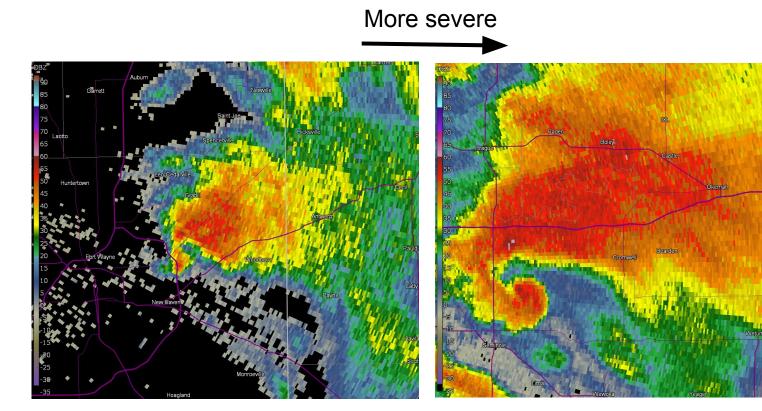


Inflow < 25 kts is often associated with mini-supercells





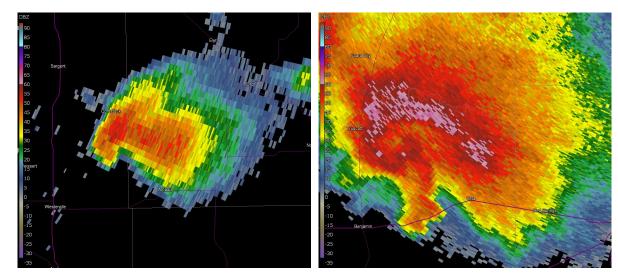
Generally, larger storms can produce more significant hazards

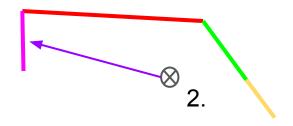


Think you've got it figured out?

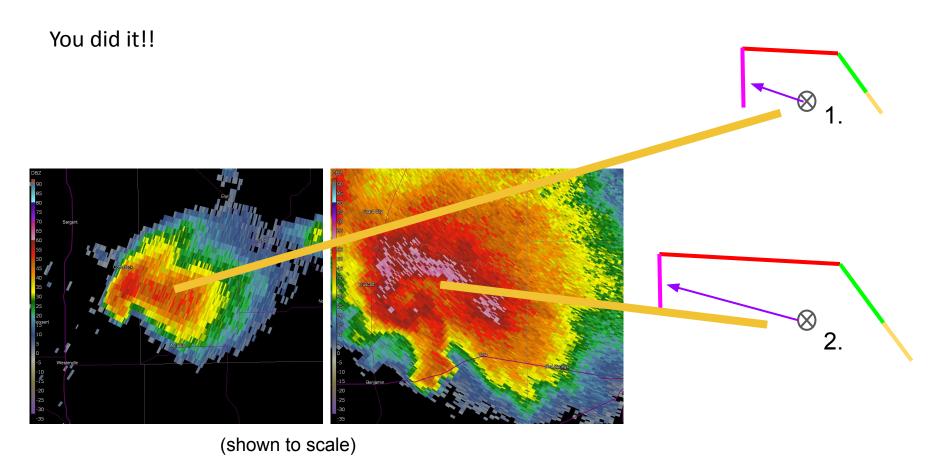
Match the supercell with its most likely hodograph:



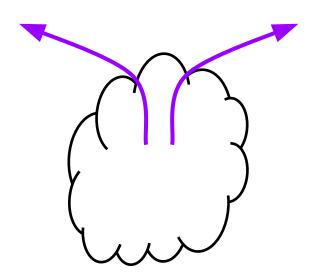


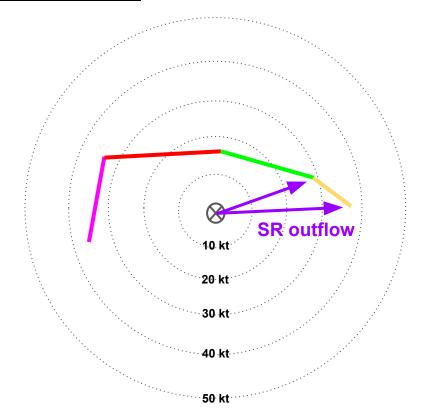


(shown to scale)

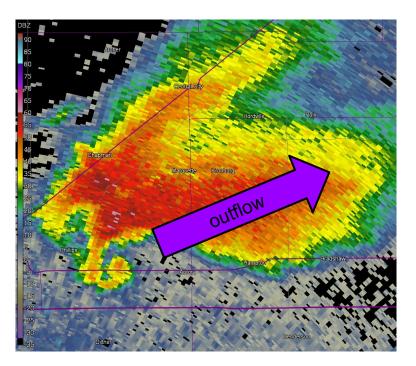


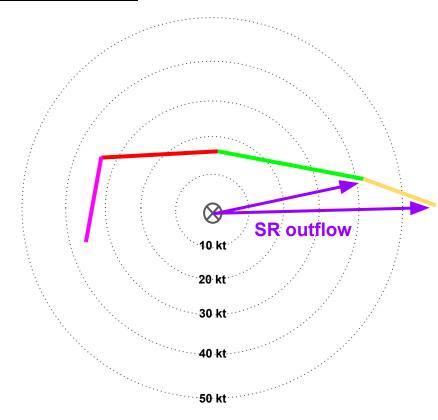
**SR Outflow** is the wind in the upper levels of a storm, where mass exits the storm





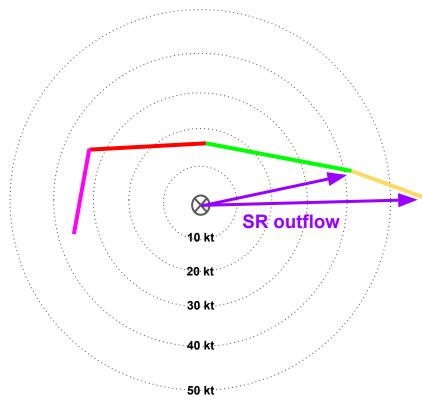
Stronger outflow can improve precipitation ventilation, supporting better updraft/downdraft separation



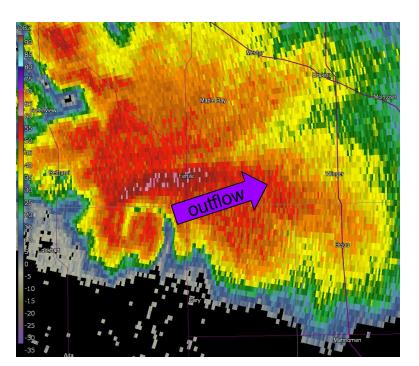


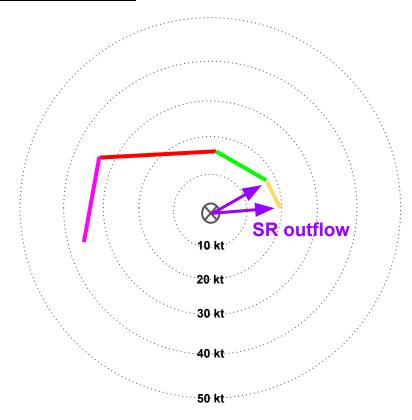
Outflow > 40 kts is particularly strong





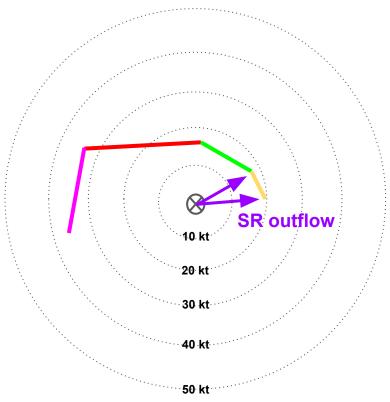
Weaker outflow can limit precipitation ventilation, inducing downdrafts close to the updraft



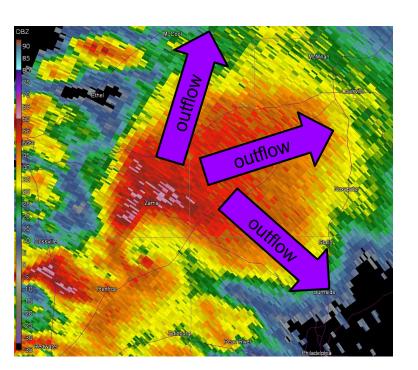


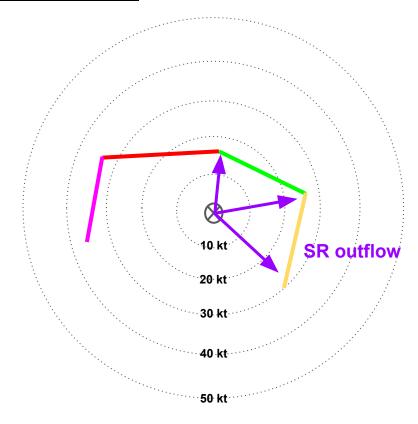
Outflow < 20 kts is particularly weak





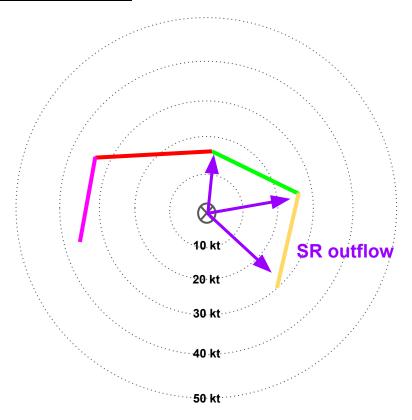
Multidirectional outflow can disperse precipitation fallout, inducing broader, more obstructive downdrafts



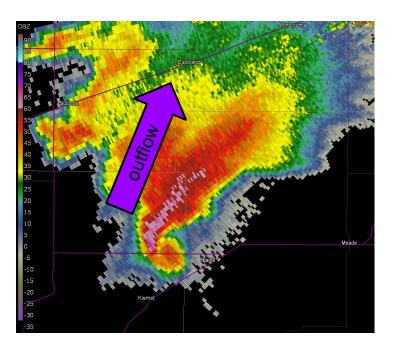


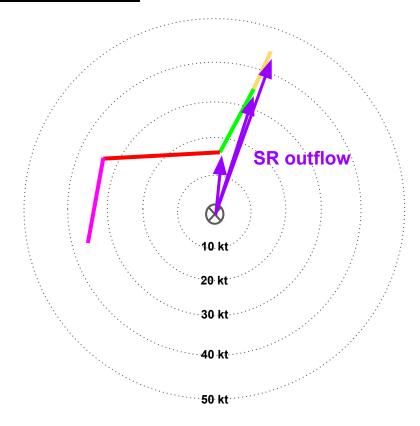
Excessive veering can make a storm appear more high-precipitation





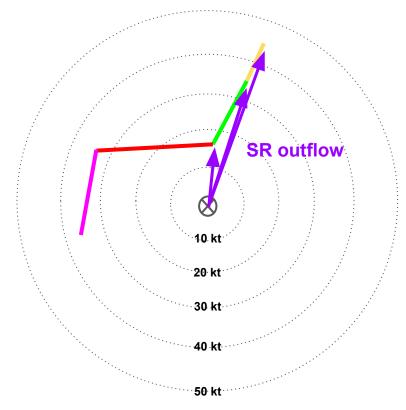
Unidirectional outflow can focus precipitation fallout, supporting less obstructive downdrafts





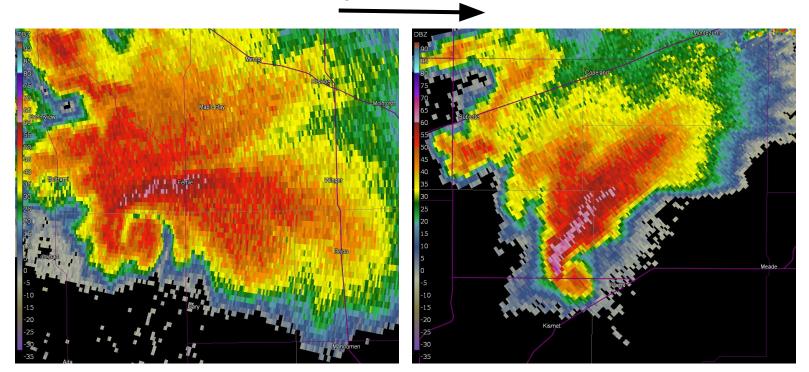
Backing shear aloft can make a storm appear more low-precipitation





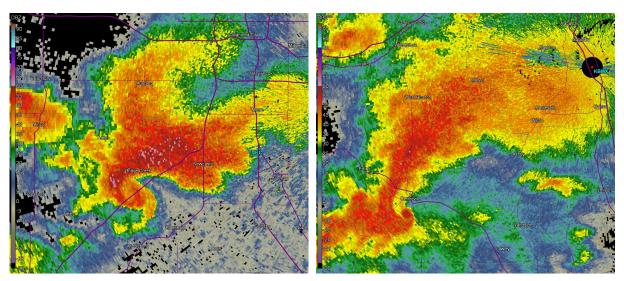
Generally, less obstructive downdrafts can support longer-lived tornadoes.

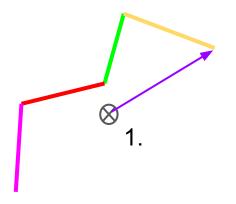
Longer-lived tornadoes



Here's an example for you!

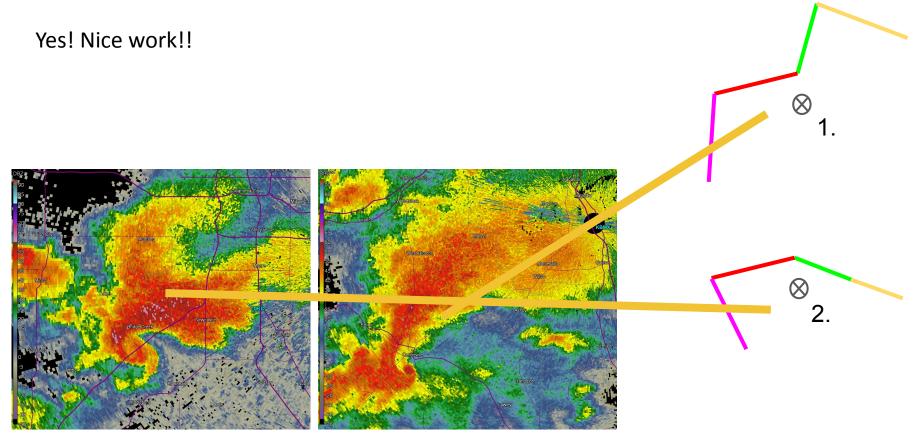
Match the supercell with its most-likely hodograph:







(shown to scale)

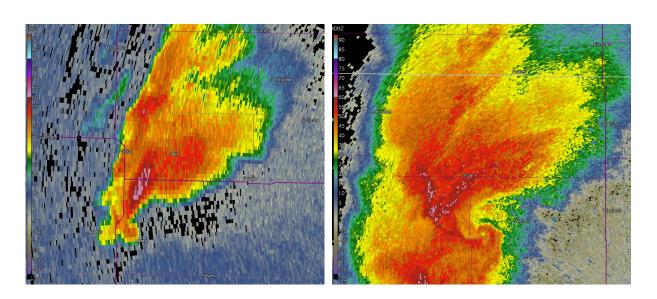


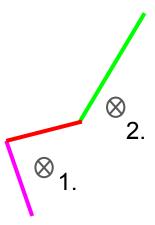
(shown to scale)

#### **Storm-Relative Inflow and Outflow**

This was the same storm, just two hours apart.

Match the radar presentation with its most likely storm motion:





#### **Storm-Relative Inflow and Outflow**

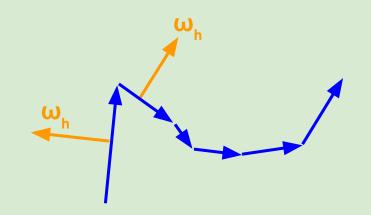
That's right!!

You're relatively good at this:)

### **Storm-Relative Winds Summary**

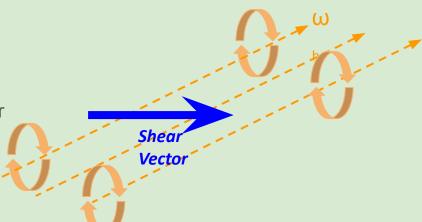
- 1. Storm-relative winds govern storm structure
- Storm-relative inflow can affect a storm's size, and how much precipitation it can produce
- Storm-relative outflow can affect precipitation ventilation, downdraft placement, and tornado maintenance

# Lesson 6: Horizontal Vorticity

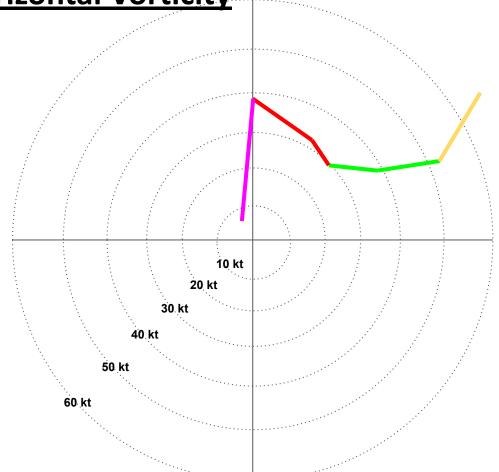


**Harry Weinman** – Meteorologist, Storm Prediction Center (harry.weinman@noaa.gov)

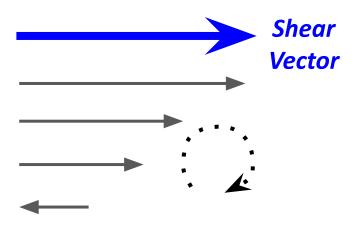
Cameron Nixon – Research Scientist, SPC / CIWRO (cameron.nixon@noaa.gov)

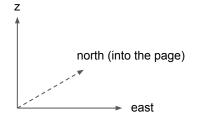


How do we find horizontal vorticity on the hodograph?

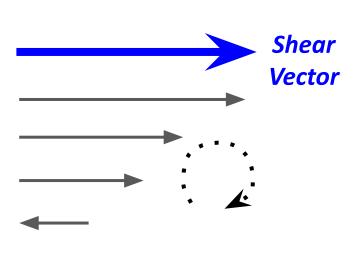


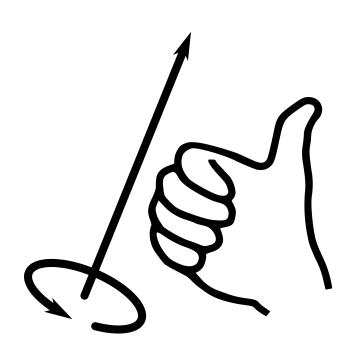
#### **Vertical Wind Shear**

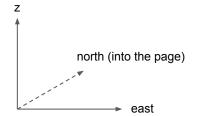


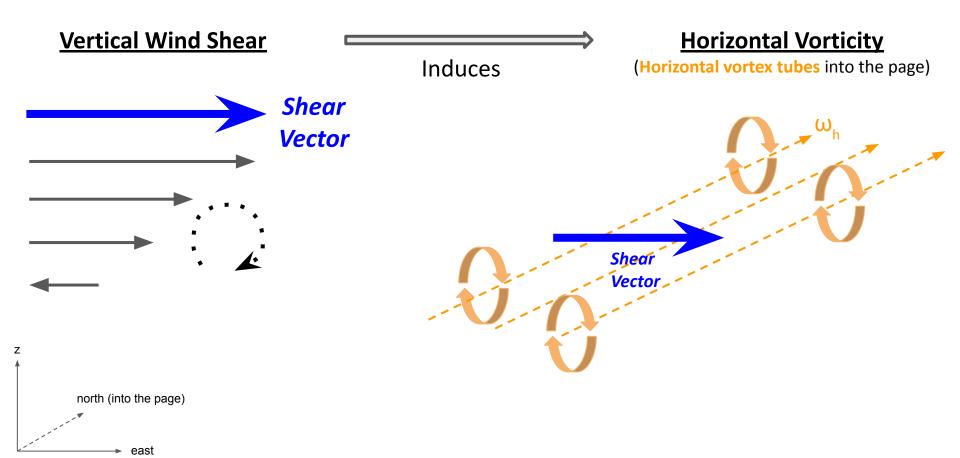


#### **Vertical Wind Shear**



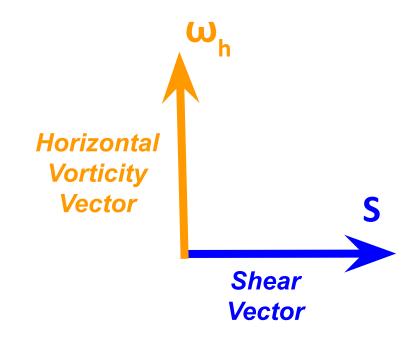




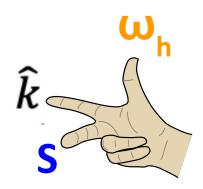


#### **Horizontal vorticity vectors**

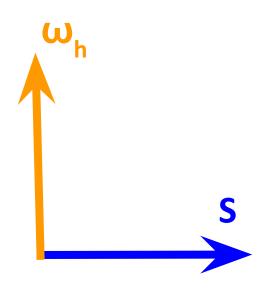
are always perpendicular and to the left of the shear vectors



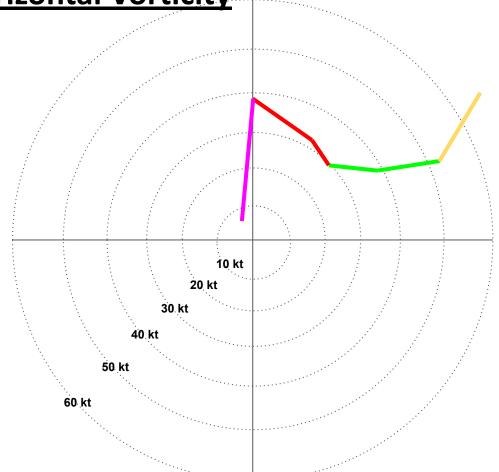
Explained using the right-hand rule!



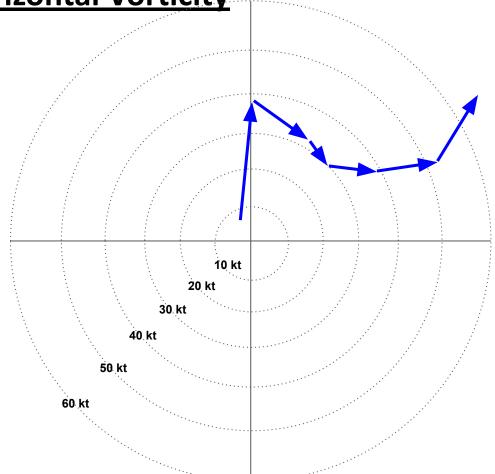
$$\omega = \hat{k} \times S$$



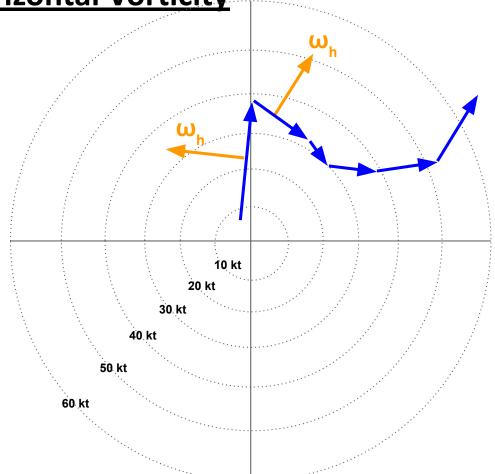
How do we find horizontal vorticity on the hodograph?



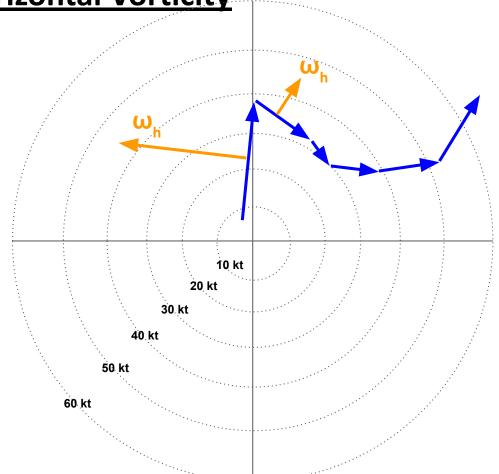
Remember, the hodograph is a plot of shear vectors



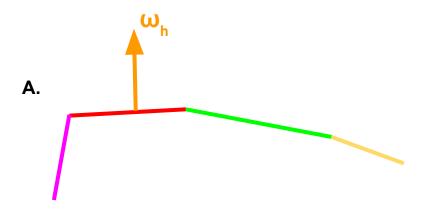
Therefore, horizontal vorticity vectors are perpendicular and to the left of the hodograph!

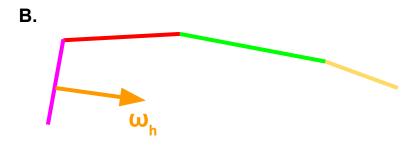


The stronger the **shear** in any given layer, the stronger the **horizontal vorticity!** 



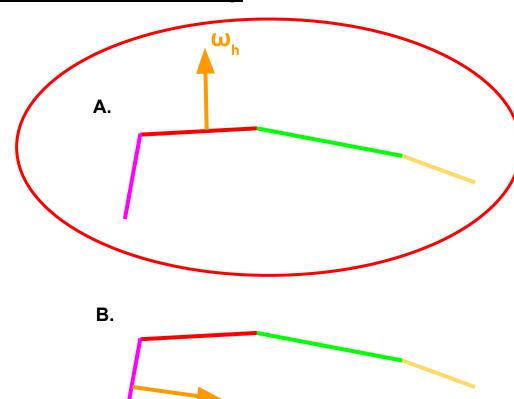
Which hodograph has the correctly drawn horizontal vorticity vector?



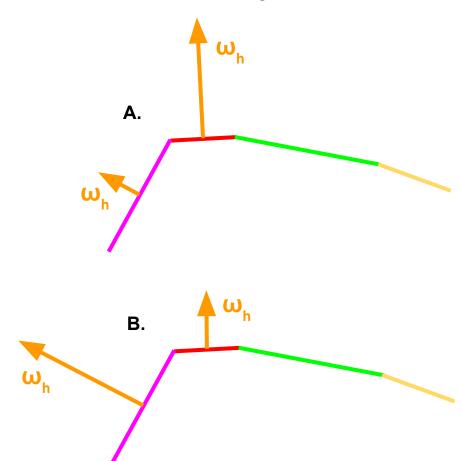


Which hodograph has the correctly drawn horizontal vorticity vector?

Great work!

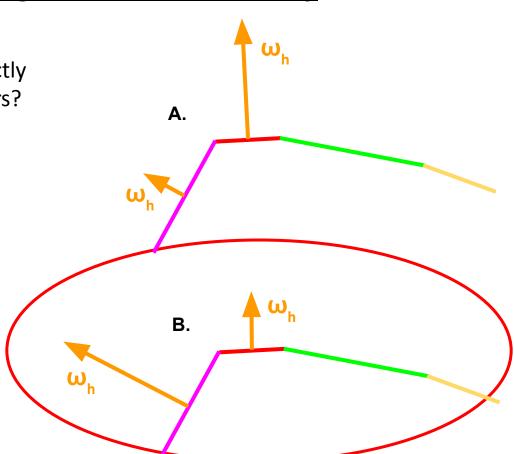


Which hodograph has the correctly drawn horizontal vorticity vectors?

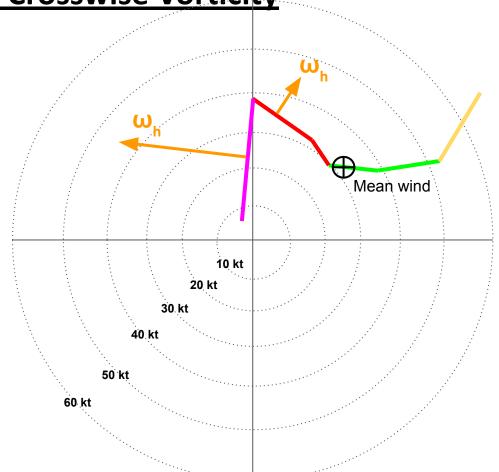


Which hodograph has the correctly drawn horizontal vorticity vectors?

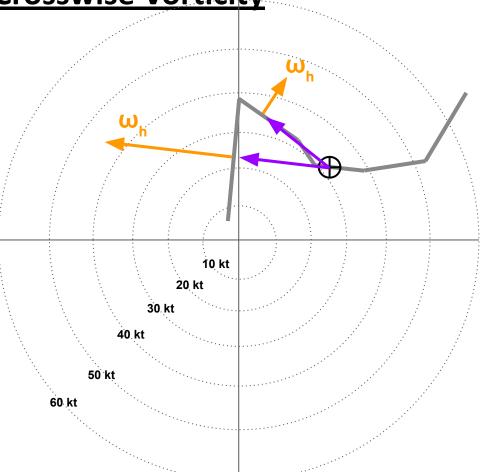
You got it!



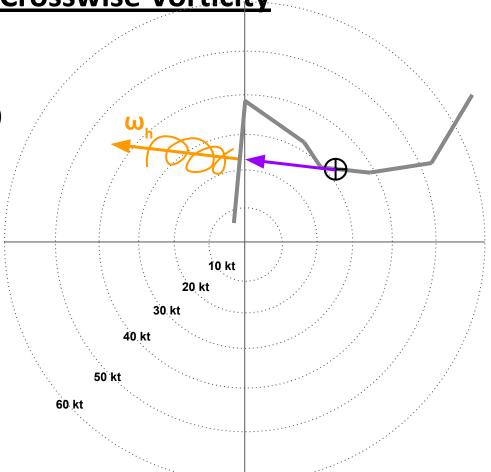
We want to know how a **storm** interacts with this horizontal vorticity



Let's compare the horizontal vorticity vectors with the storm-relative wind vectors

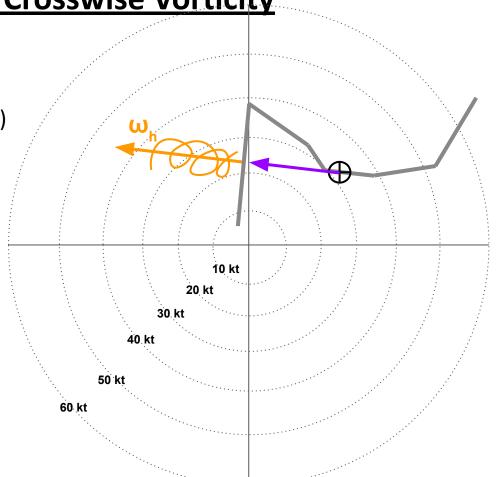


Here, storm-relative wind is parallel to horizontal vorticity (the two are aligned)

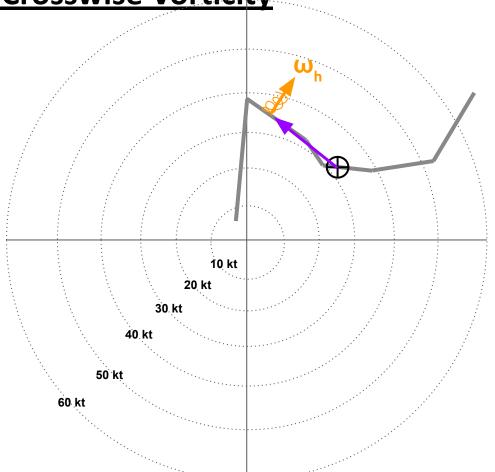


Here, storm-relative wind is parallel to horizontal vorticity (the two are aligned)

This is called **streamwise** vorticity!

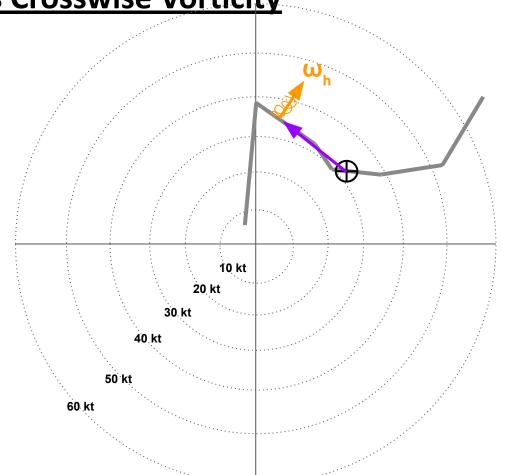


Here, storm-relative wind is perpendicular to horizontal vorticity (the two are NOT aligned)



Here, storm-relative wind is perpendicular to horizontal vorticity (the two are NOT aligned)

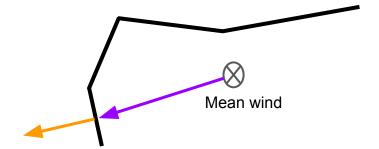
This is called **crosswise** vorticity.

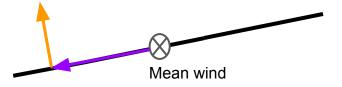


Hodograph shape (and storm motion) affects streamwise vorticity:

**Streamwise** vorticity!

**Crosswise vorticity** 





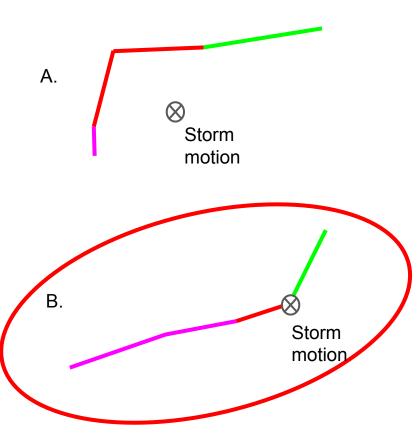
Which hodograph contains more crosswise vorticity?





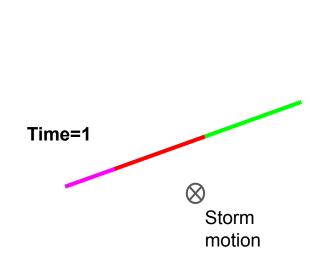
Which hodograph contains more crosswise vorticity?

You got it!



What happens from Time=0 to Time=1?

- a. Storm loses streamwise vorticity
- Storm acquires a greater streamwise vorticity component and can rotate
- c. No change, hodograph remains straight



Storm

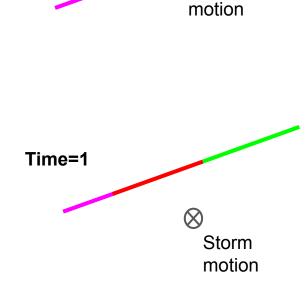
motion

Time=0

What happens from Time=0 to Time=1?

- a. Storm loses streamwise vorticity
- Storm acquires a greater streamwise vorticity component and can rotate
- c. No change, hodograph remains straight

Yes!



Storm

Time=0

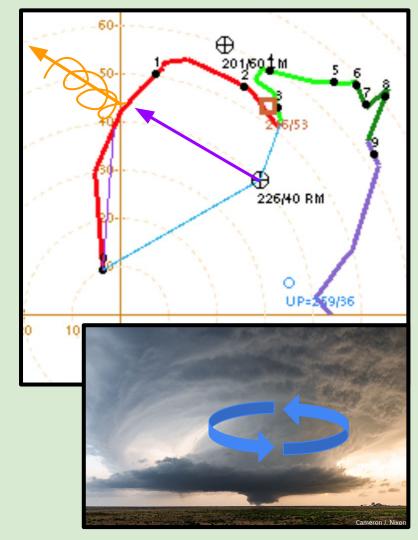
### **Horizontal Vorticity Summary**

- 1. Horizontal vorticity is always perpendicular and to the left of the shear and therefore the hodograph.
- 2. The stronger the shear, the stronger the vorticity in that layer.
- 3. When SR winds are parallel to horizontal vorticity, the vorticity is streamwise. When SR winds are perpendicular to horizontal vorticity, the vorticity is crosswise.
- 4. Both hodograph shape and storm motion influence development of streamwise and crosswise vorticity.

# Lesson 7: Streamwise Vorticity and SRH

**Cameron Nixon** – Research Scientist, SPC / CIWRO (cameron.nixon@noaa.gov)

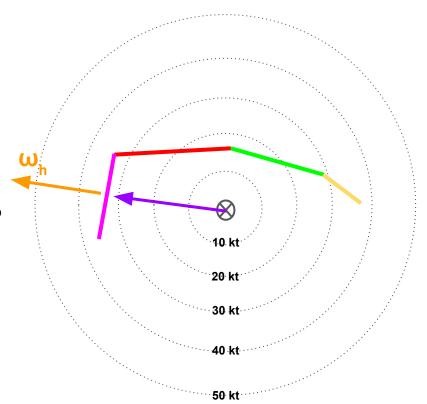
Harry Weinman – Meteorologist, Storm Prediction Center (harry.weinman@noaa.gov)



#### **Streamwise Vorticity**

There are two ways to describe streamwise vorticity:

- 1. How streamwise is the vorticity?
- 2. How much streamwise vorticity is there?



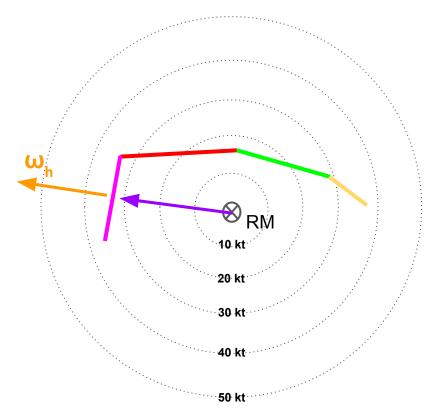
**Streamwise Vorticity** 

## and Mesocyclogenesis

#### **Streamwiseness**

**Streamwiseness** is how parallel the **storm-relative wind** is to the **horizontal vorticity** 

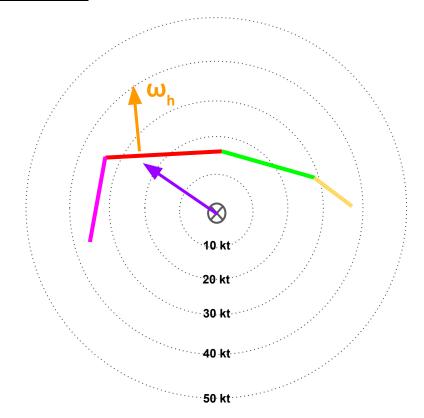
Here, vorticity is purely streamwise.



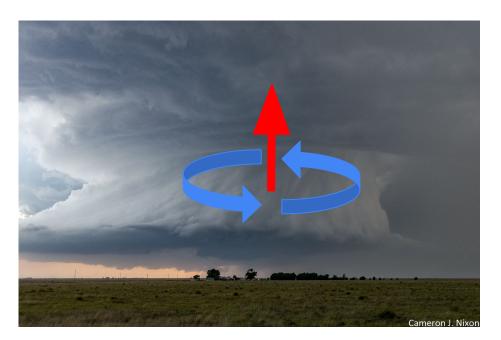
#### **Streamwiseness**

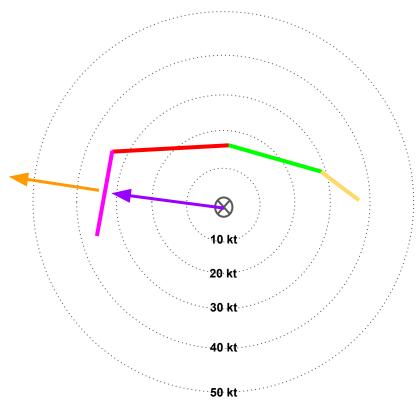
**Streamwiseness** is how parallel the **storm-relative wind** is to the **horizontal vorticity** 

Here, vorticity is less streamwise. (it's at an angle to the inflow)

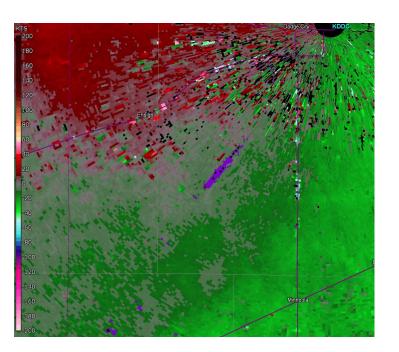


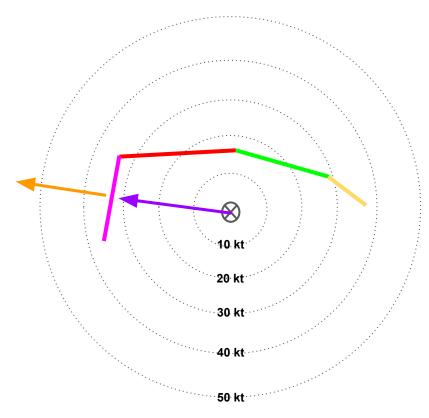
With more streamwiseness, vorticity tends to be aligned with the updraft, so mesocyclogenesis can be faster.





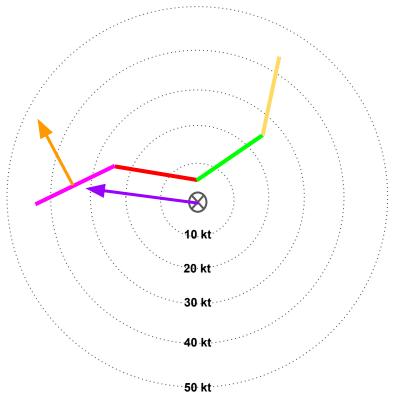
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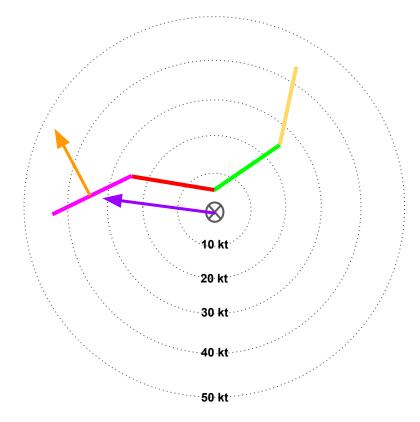
With less streamwiseness, vorticity tends to be dislocated from the updraft, so mesocyclogenesis may be slower.





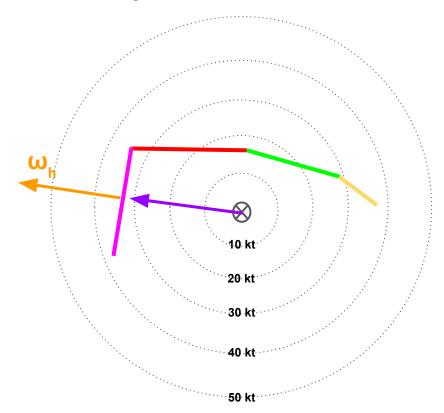
With less streamwiseness, vorticity tends to be dislocated from the updraft, so mesocyclogenesis may be slower.





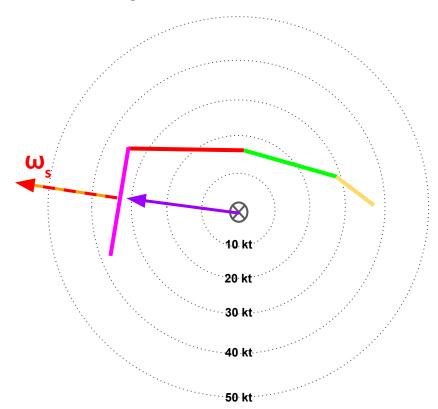
But streamwiseness doesn't tell the whole story...

The **magnitude** of streamwise vorticity is important!

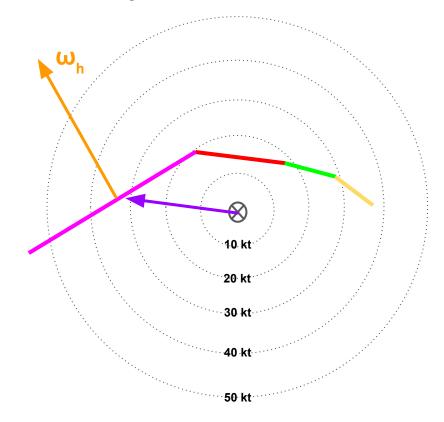


The **magnitude** of streamwise vorticity is important!

Here, there is .03 s<sup>-1</sup> of streamwise vorticity

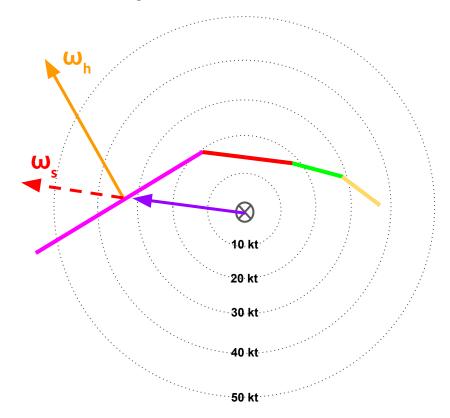


Even if horizontal vorticity is not purely streamwise, there may still be a streamwise component.



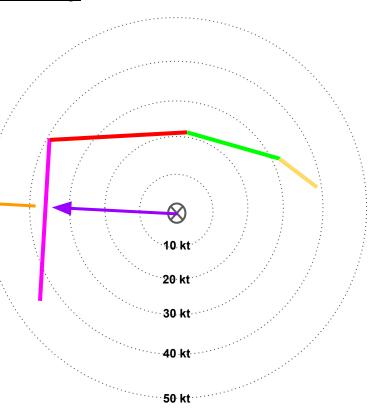
Even if horizontal vorticity is not purely streamwise, there may still be a streamwise component.

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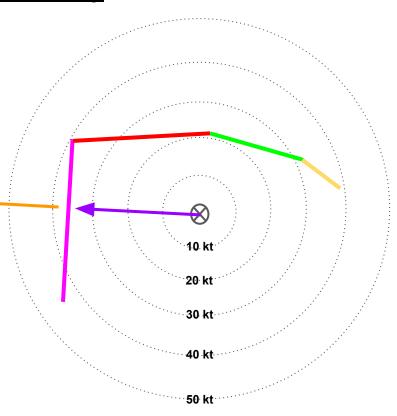
With a greater magnitude of streamwise vorticity, mesocyclones can be **stronger** and more persistent.





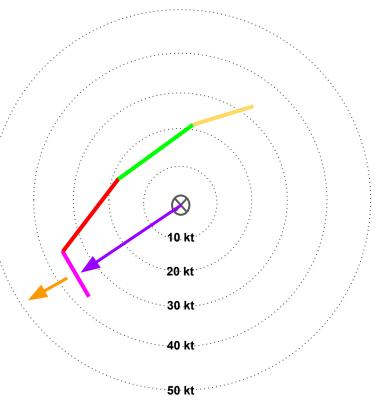
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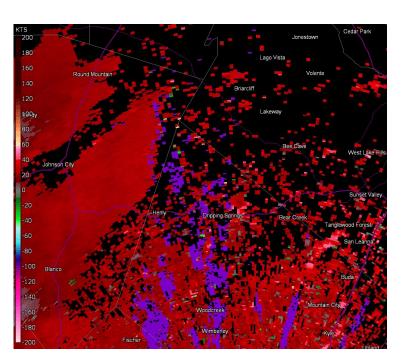


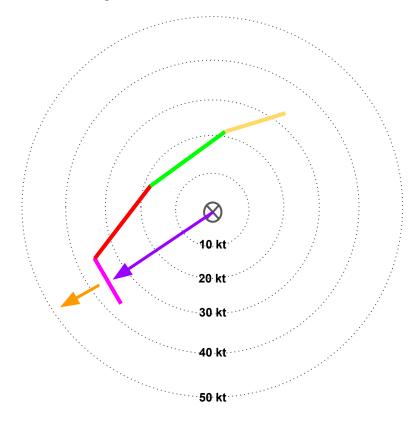
With a lesser magnitude of streamwise vorticity, mesocyclones are generally weaker and more transient.



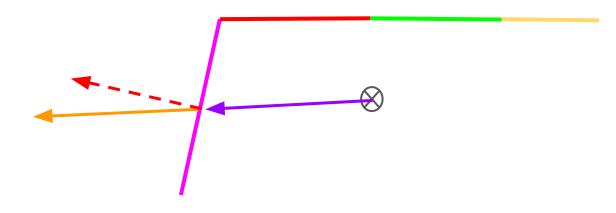


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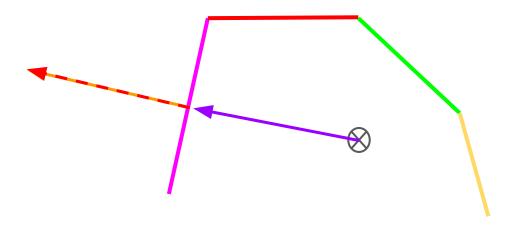




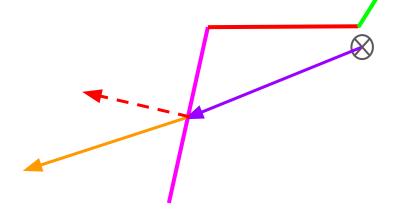
Keep in mind: storm motion affects streamwise vorticity



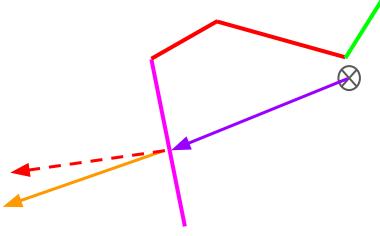
Larger hodograph curvature can maximize streamwise vorticity



More backing shear aloft can minimize streamwise vorticity



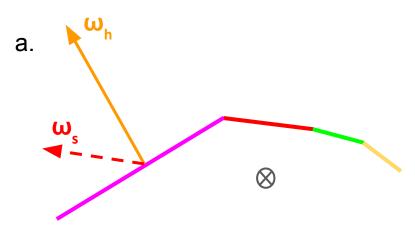
Larger low-level curvature can increase streamwise vorticity



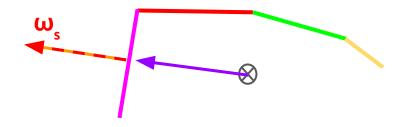
Quiz time!

Which of these two hodographs has a greater magnitude of **streamwise vorticity**?

- a. a
- b. b
- c. they both have the same



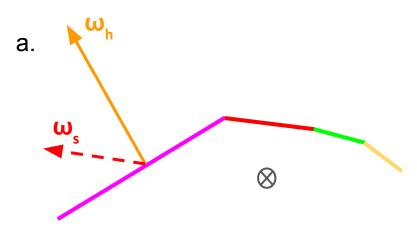
b.



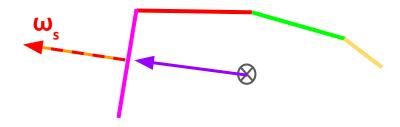
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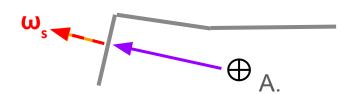
b.

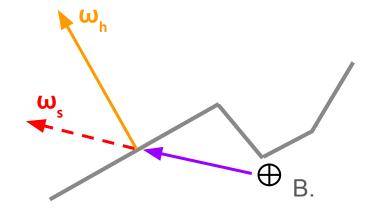


Quiz time!

Compared to hodograph A, hodograph B shows:

- a. A lesser magnitude of streamwise vorticity, but this vorticity is more streamwise
- b. A greater magnitude of streamwise vorticity, but this vorticity is less streamwise



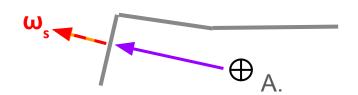


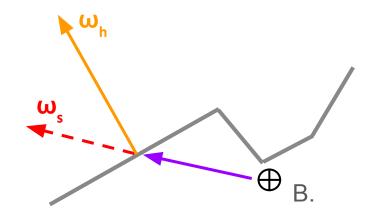
Quiz time!

Compared to hodograph A, hodograph B shows:

- a. A lesser magnitude of streamwise vorticity, but this vorticity is more streamwise
- A greater magnitude of streamwise vorticity,
   but this vorticity is less streamwise

You're very (stream)wise:)



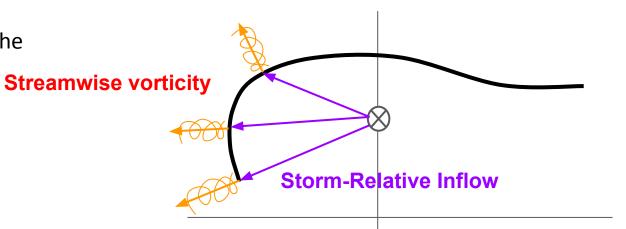


The magnitude of streamwise vorticity can be approximated using **SRH**.

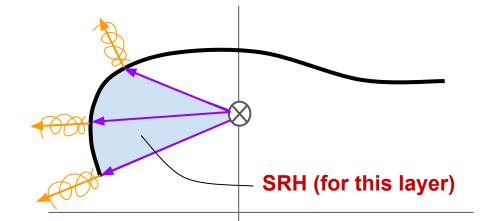
Storm Relative Helicity (SRH) = 
$$\int_{Zb}^{Vt} (inflow) \cdot (horizontal \ vorticity) \ dz$$

$$\Rightarrow |inflow| \cdot |streamwise \ vorticity|$$

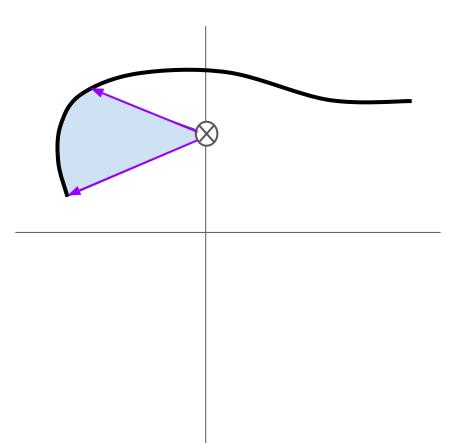
What does this look like on the hodograph?



Just take the area under the hodograph swept out by the storm-relative winds over some depth!

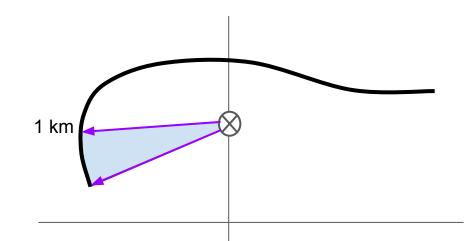


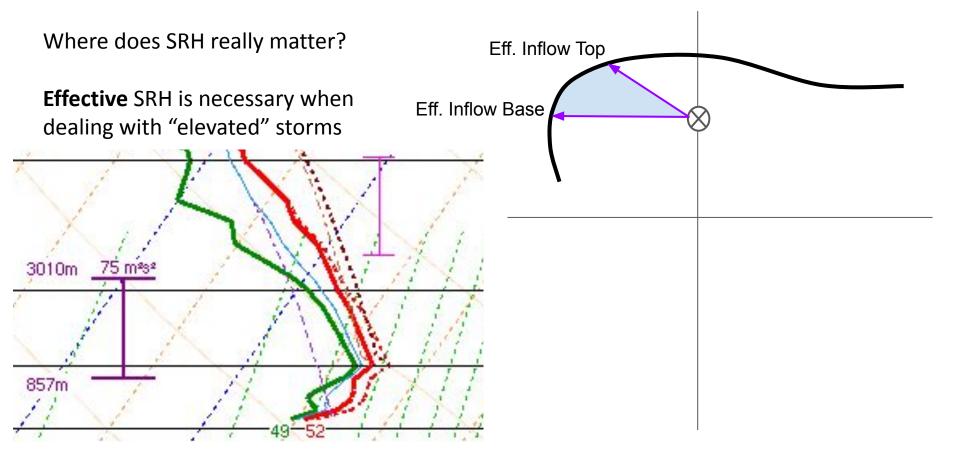
Where does SRH really matter?



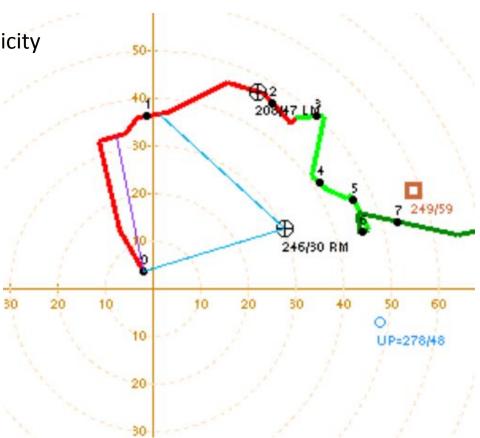
Where does SRH really matter?

SRH in the lowest kilometer is most relevant for mesocyclogenesis



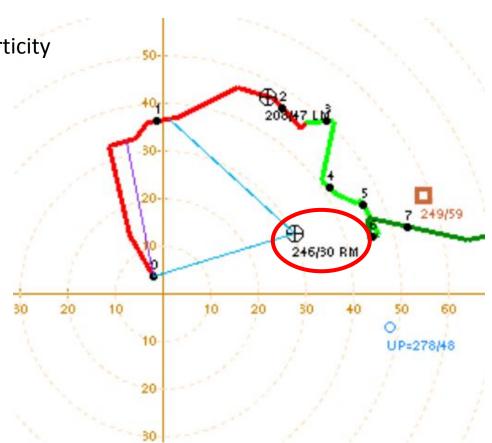


Remember, all measures of streamwise vorticity (including SRH) depend on storm motion.



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In fact, most calculations of SRH use Bunkers Right Storm Motion

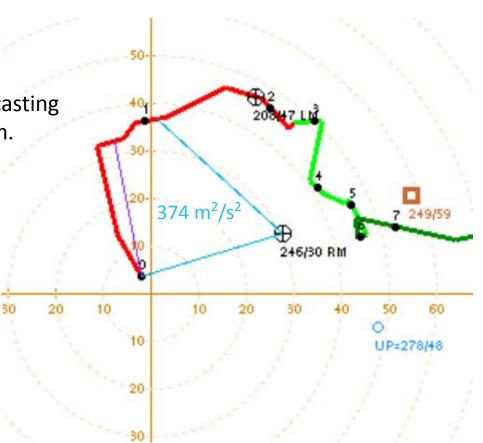


What about this scenario?

There is lots of 0-1 km SRH, and you're forecasting a slow, southward-moving convective system.

Are embedded tornadoes possible?

- a. Yes,  $374 \text{ m}^2/\text{s}^2$  is a lot of SRH
- b. Be careful, use the right storm motion



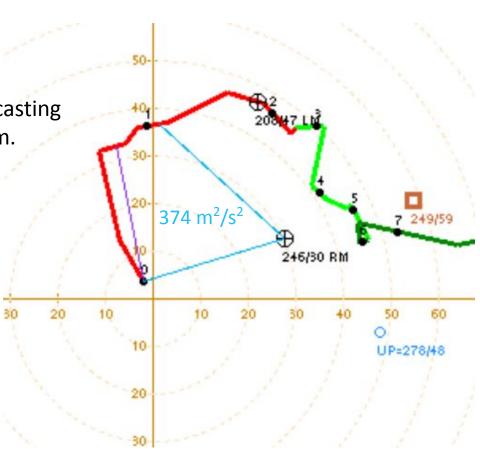
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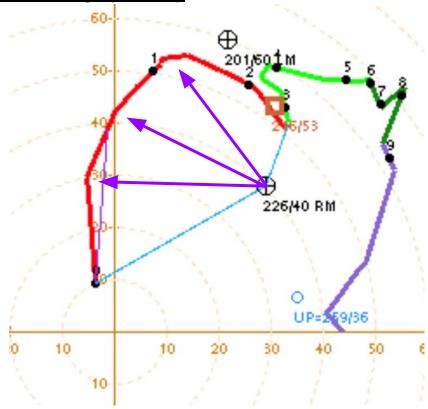
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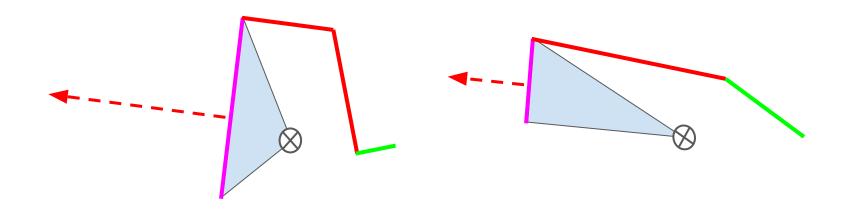
Good choice.



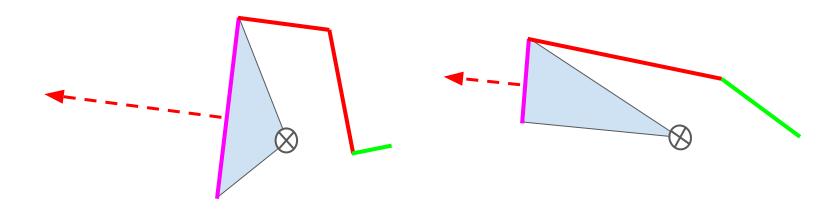
Also keep in mind, SRH is *not just* the streamwise vorticity, but also the strength of the *storm-relative winds*!



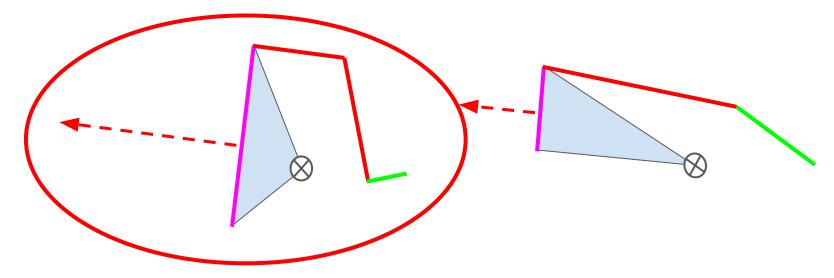
So, despite having equal SRH, two hodographs can have different magnitudes of streamwise vorticity.



Which hodograph would you expect to be most conducive to strong low-level mesocyclones?

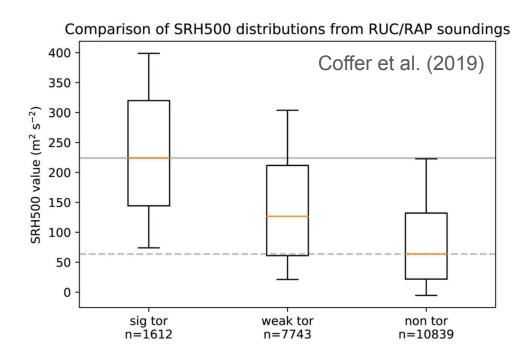


Which hodograph would you expect to be most conducive to strong low-level mesocyclones?



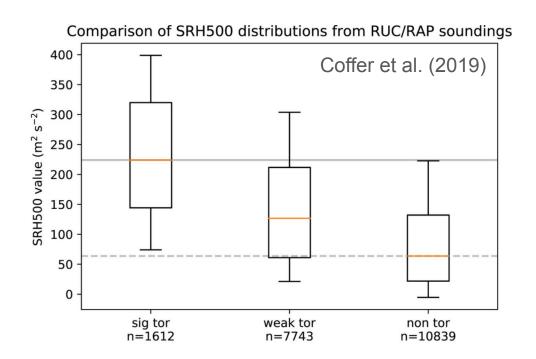
That's right! Streamwise vorticity and inflow strength both contribute to SRH

SRH is a useful and convenient parameter to assess the potential for mesocyclones and tornadoes



SRH is a useful and convenient parameter to assess the potential for mesocyclones and tornadoes

Just keep in mind that streamwise vorticity and inflow strength can affect storms in different ways!

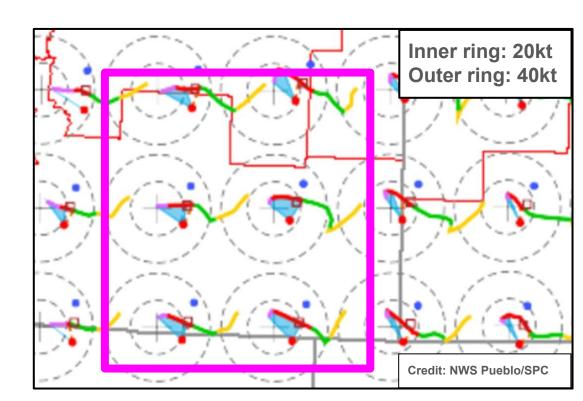


# **Streamwise Vorticity Summary**

- 1. The streamwiseness of vorticity can affect how quickly an updraft rotates
- 2. The magnitude of streamwise vorticity can affect how strong and persistent a mesocyclone can become
- 3. Storm-Relative Helicity (SRH) is an estimate of streamwise vorticity

2100z: Ample buoyancy exists over your CWA, and storms are beginning to initiate. Looking at the hodograph map, what is *most likely* from initial activity?

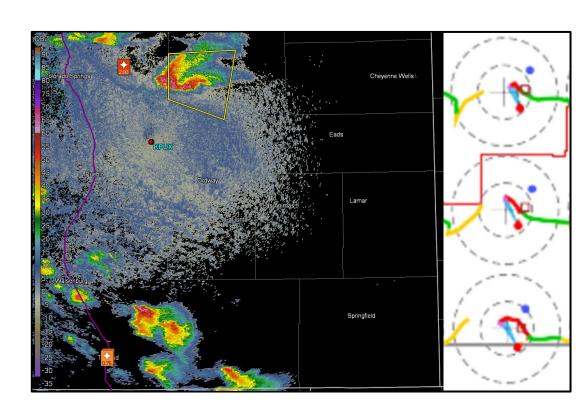
- a. Weak multicells (wind)
- b. Splitting supercells (hail+wind)
- c. Strong, right-moving supercells (tor+hail+wind)



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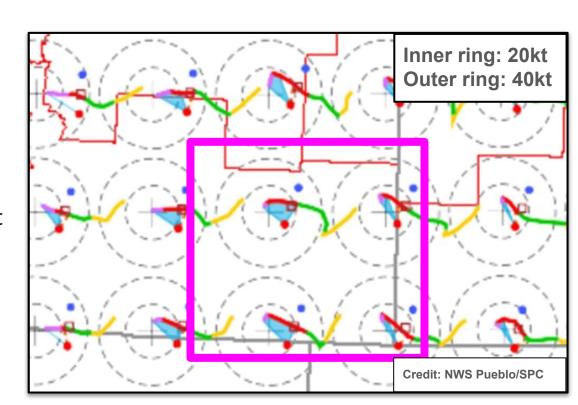
- a. Weak multicells (wind)
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Correct!



2300z: Now, you examine forecast hodographs. What should you expect over the next couple hours?

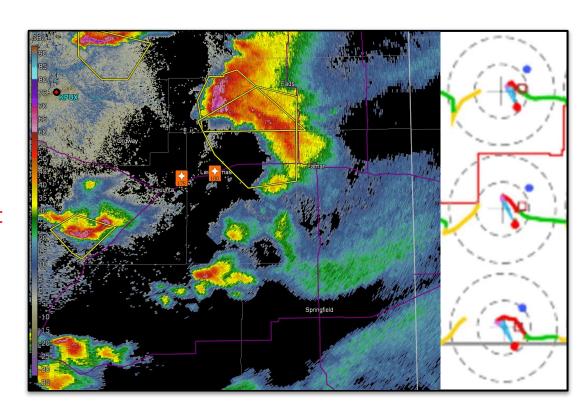
- a. Storms continue to split, and grow upscale (wind)
- b. Left-movers become dominant (hail+wind)
- c. Right-movers become dominant (tor+hail+wind)



2300z: Now, you examine forecast hodographs. What should you expect over the next couple hours?

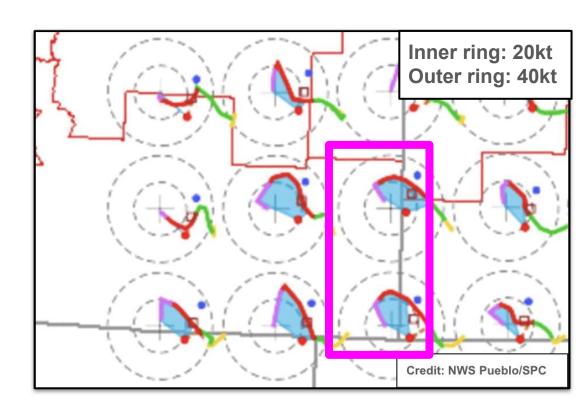
- a. Storms continue to split,
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- c. Right-movers become dominant (tor+hail+wind)

Correct!



0000z: The cap is setting in, and CIN is greater than -200 J/kg in some locations! What do you think will happen next?

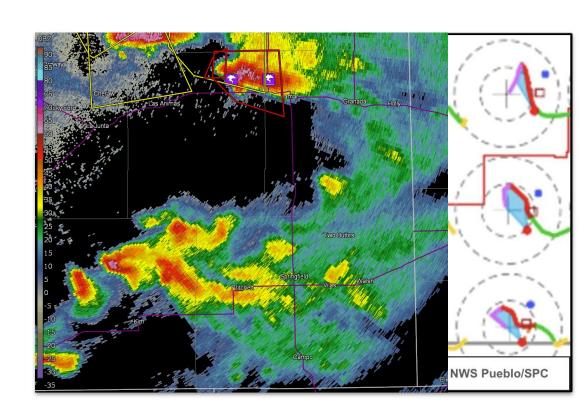
- a. All storms should dissipate. That's a lot of CIN!
- b. Some storms will dissipate, but supercells can persist
- c. All storms should persist



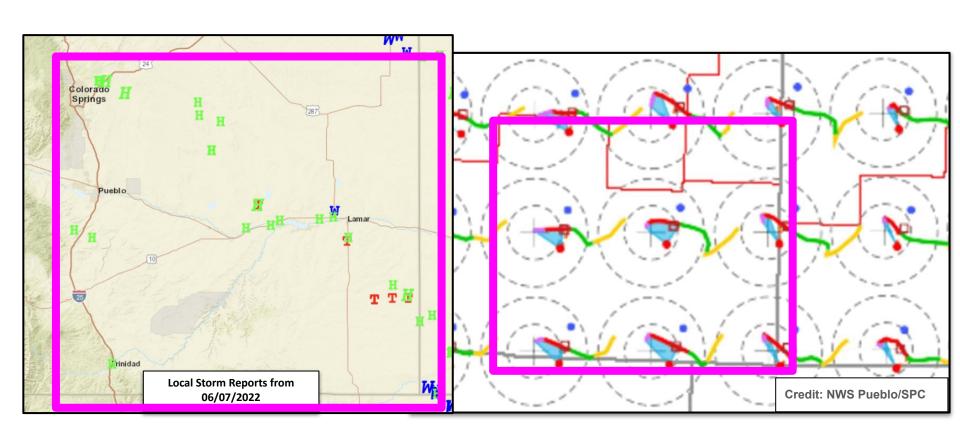
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- a. All storms should dissipate. That's a lot of CIN!
- b. Some storms will dissipate, but supercells can persist
- c. All storms should persist

Good call!



#### Summary



#### **Straight Hodograph**

Horizontal Vorticity
Vectors
Initial Updraft
(before RM and LM form)
Inflow Vectors

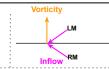
#### When an initial updraft forms...

normal to vorticity vectors).

Purely *crosswise vorticity* (no SRH) for initial updraft.

 Initial updraft not rotating because none of the inflow is aligned with the vorticity (inflow vectors.)

3. The initial updraft draws the horizontal vorticity upward, resulting in two new updrafts: one to the right of the shear vector, the other to the left -- see helow



#### After initial updraft tilts vorticity into the vertical, creating two new updrafts...

 New updrafts to the left and right of deep-shear vector gain rotation as they are now aligned with the vorticity tilted into the vertical. There is also an added contribution to updraft rotation from the now non-zero SRH (inflow vectors and vorticity vectors have a parallel/antiparallel component).

- 2. Both updrafts are neither enhanced nor suppressed with this hodograph shape, with mirror-image cell splitting.
- 3. The survival of both left and right members can introduce higher potential for cell interactions, fostering the potential for upscale growth.

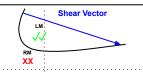
#### **Single Concavity Hodograph**



Within a layer containing hodograph concavity, dynamic lifting/suppression occurs on the concave/convex side of the initial updraft, relative to the deep shear vector (the vector can be

transposed anywhere on the coordinate system).

- 2. The concave side of this hodograph favors dynamic lifting to the right of the initial updraft relative to the shear vector -- i.e., dominant right-mover (following Bunkers motion). The convex side of this hodograph yields left-of-shear dynamic suppression -- i.e., vanishing left-mover.
  - **3.** Also, area swept out between inflow vectors and the hodograph (i.e., SRH) increases with the hodograph bow, resulting in much more efficient mesocyclogenesis.

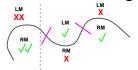


1. The concave side of this hodograph favors left-of-shear dynamic lifting -- i.e., dominant left-mover (following Bunkers motion). The convex side of this hodograph yields right-of-shear dynamic

suppression -- i.e., vanishing right-mover.

2. Negative SRH given counter-clockwise hodograph bow favors efficient meso-anticyclone-enhancement for left-mover.

#### Multi-Inflection Hodograph



- 1. Both right and left members are enhanced **and** suppressed at alternating levels in the vertical -- separated by / inflection points.
- 2. Lift √ for **right member** in lower and upper levels of hodograph, suppression **X** for **right member** at middle levels.
- Lift √ for left member in middle levels of hodograph, suppression X for left member at lower and upper levels.
- 4. SRH beneath the second inflection is reduced by the hodograph concavity alternation, and weak storm-relative flow between the first and second inflections results in updraft seeding and precipitation drag suppressing the updraft.

5. Both members of the split survive in a degraded

manner, sometimes yielding convective

clustering/upscale growth.

**6.** The larger the looping action, the more induced vertical motion.

Note: RM and LM as marked on hodographs do

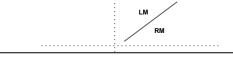
specifically owing to the updraft-in-shear effect.

√√ means more lift than √

xx means more suppression than x

not indicate a storm motion vector (use Bunkers supercell motion). However, RM and LM make reference to the updraft members on either side of the deep-shear vector (right and left, respectively) that potentially experience additional dynamic lifting/subsidence.

#### Additional Examples



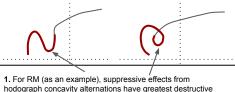
hodograph.

2. This hodograph favors splitting mirror-image cells, with no enhancement or suppression for either

member.

1. Unidirectional shear profile, though still a veering

**3.** This has less suppressive effects compared to a looping hodograph with multiple inflections.



effects on updrafts when they're in the effective-inflow layer (as above).

2. SRH is much reduced with this hodograph shape, in addition to the dynamic suppressive effects fostering upscale growth.



- For RM (as an example), these concavity alternations will be much less detrimental on the updraft, as they are located well above the effective-inflow layer.

  The source assign (where under a set to indepth is set.)
  - The source region (where updrafts get their start) is not disrupted by dynamic subsidence -- owing to the favorable single/dominant concavity seen in the lower and middle levels.