

Severe Weather Ingredients

**Material by Tom Galarneau Andrew Lyons
and Rich Thompson**

Ingredients for Organized Severe Thunderstorms

- **Instability (Lapse Rates)**
- **Lift (QG, mesoscale, convective scale)**
- **Moisture (return flow, soil moisture, evapotranspiration)**
- **Vertical wind shear**

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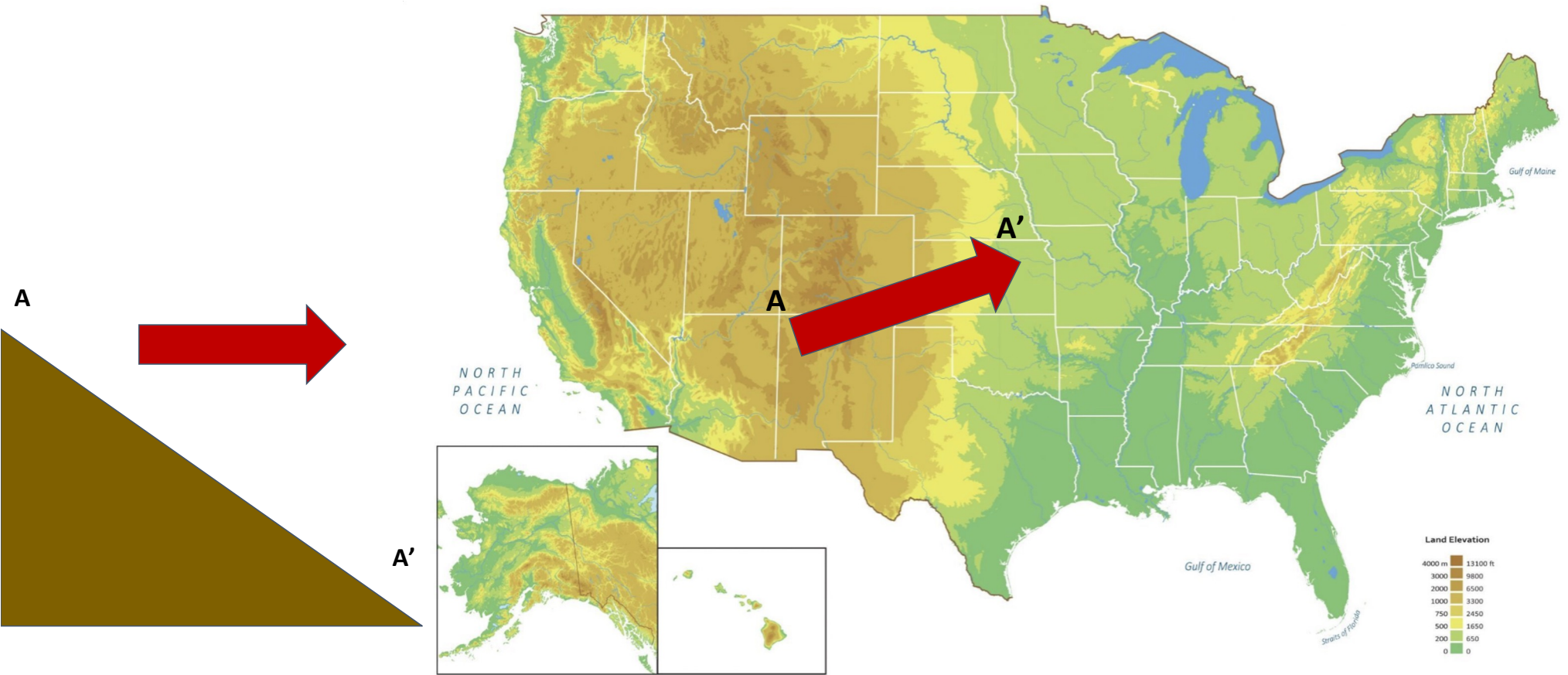
This lecture

- a separate lecture

Instability: Elevated Mixed Layer (EML)

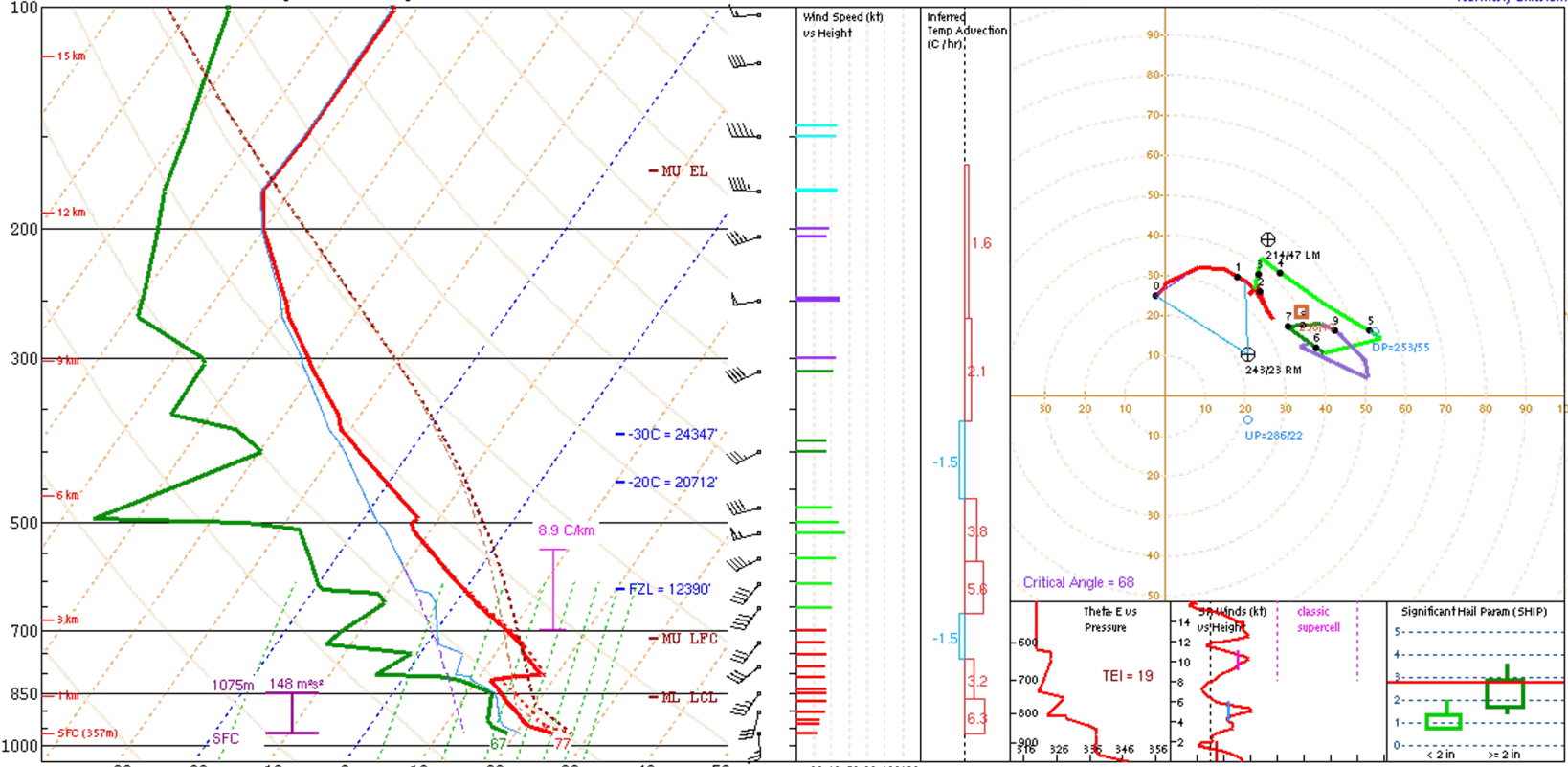
- **Deep mixed layer forms over Rockies and Sierra Madre in response to surface heating**
- **Stronger heating and/or cooler temperatures aloft results in deeper mixed layer**
- **Mixed layer is advected eastward and becomes elevated east of the Rockies and Sierra Madre (really, east of the dryline)**
- **Differential advection (eastward advection of EML above northward advection of moisture from GoM) created “loaded gun” profile**

Elevated Mixed Layer (EML) Differential Advection



OUN 220429/1800 (Observed)

NOAA/NWS Storm Prediction Center
Norman, Oklahoma



PARCEL	CAPE	CINH	LCL	LI	LFC	EL
SURFACE	3444	-36	752m	-10	2501m	41986'
MIXED LAYER	1736	-178	952m	-6	3301m	38863'
FCST SURFACE	2376	-102	1306m	-7	2950m	40712'
MU (964 mb)	3444	-36	752m	-10	2501m	41986'

PW = 1.01 in	3CAPE = 5 J/kg	WBZ = 9225'	WINDG = 0.0
K = 19	DCAPE = 1284 J/kg	FZL = 12390'	ESP = 0.0
MidRH = 30%	DownT = 56 F	ConvT = 93F	MMP = 0.97
LowRH = 79%	MeanW = 12.3 g/kg	MaxT = 80F	NCAPE = 0.33
SigSevere = 37509 m3/s3			

3fc-3km Agl Lapse Rate = 6.1 C/km	Supercell = 9.9 Left Supercell = -1.0 STP (eff layer) = 0.2 STP (fix layer) = 2.2 Sig Hail = 2.7
3-6km Agl Lapse Rate = 8.1 C/km	
850-500mb Lapse Rate = 6.4 C/km	
700-500mb Lapse Rate = 8.6 C/km	

	SRH(m2/s2)	Shear(kt)	MnWind	SRW
SFC - 1 km	139	21	192/30	141/24
SFC - 3 km	161	26	211/31	163/17
Eff Inflow Layer	148	23	193/30	142/24
SFC - 6 km	42	221/34	184/16	
SFC - 8 km	37	223/34	190/15	
LCL - EL (Cloud Layer)	48	233/37	216/15	
Eff Shear (EBWD)	39	222/34	185/15	
BRN Shear = 63 m/s ²				
4-6km SR Wind = 247/22 kt				
Storm Motion Vectors				
Bunkers Right = 243/23 kt				
Bunkers Left = 214/47 kt				
Corfidi Downshear = 253/55 kt				
Corfidi Upshear = 266/22 kt				

*** BEST GUESS PRECIP TYPE ***

Rain.
Based on sfc temperature of 77.4 F.

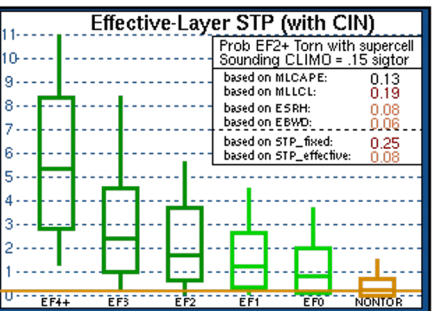
SARS - Sounding Analogs

SUPERCELL	SGFNT HAIL
06042500.OUN	3.00
94060500.HON	2.75
90082600.BIS	1.75

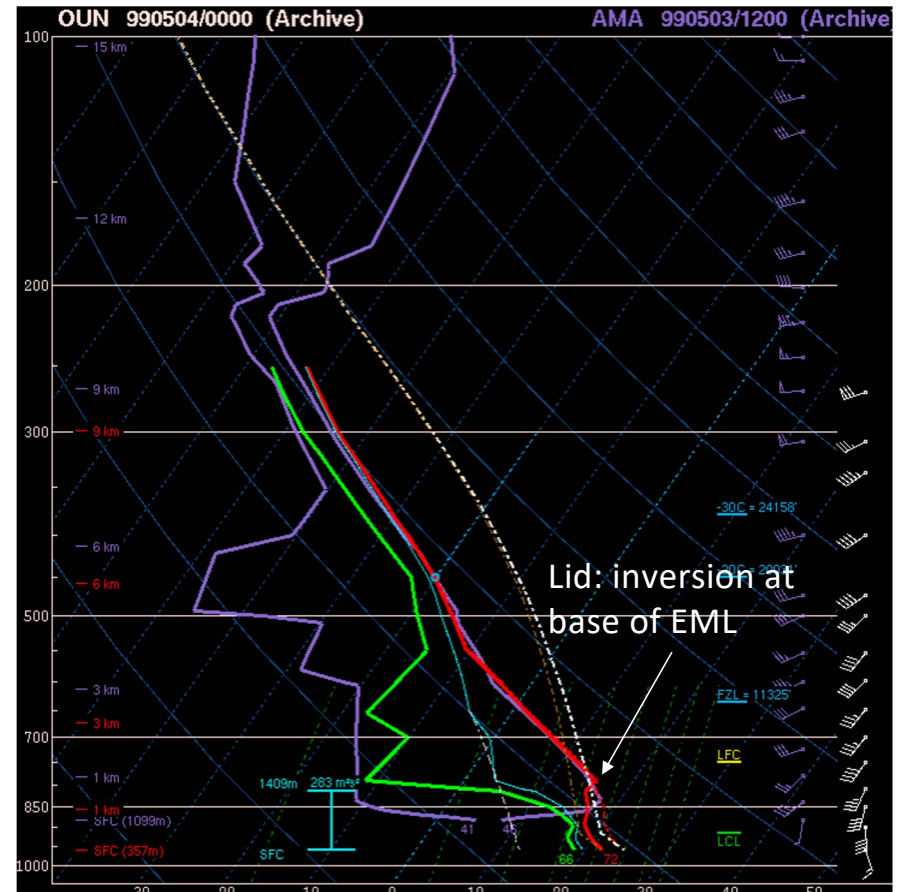
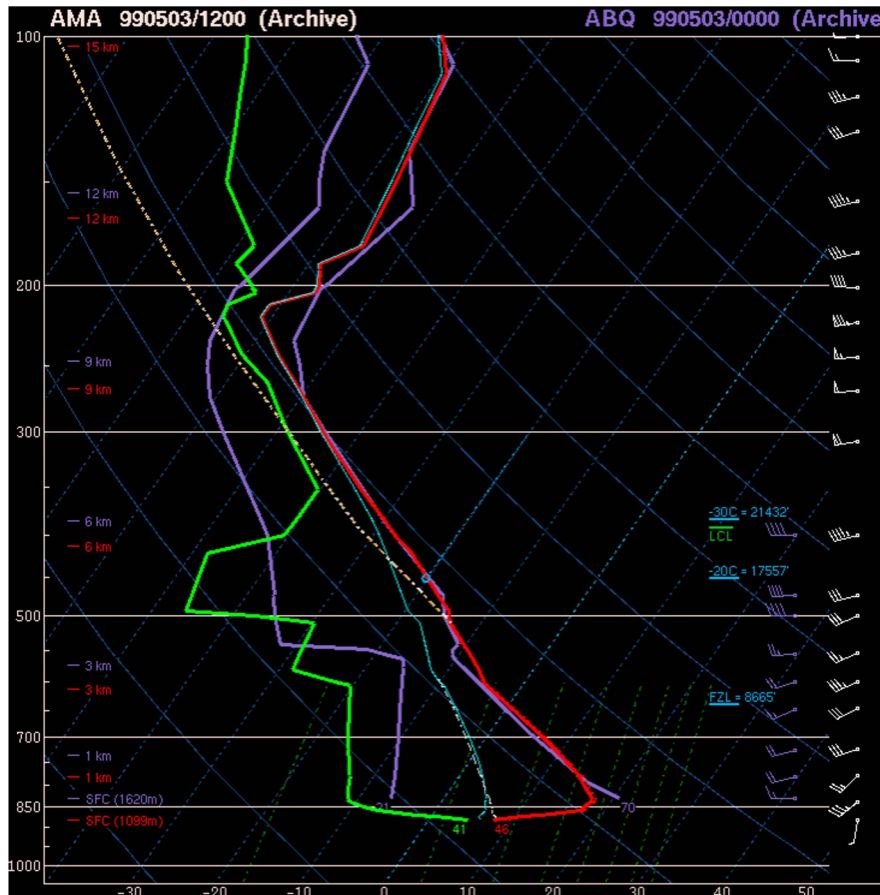
No Quality Matches

(12 loose matches) SARS: 75% TOR

(231 loose matches) SARS: 72% SIG



EML at OUN (0000 UTC 4 May 1999)



Climatology of Lid Occurrence

- Series of papers by Lanicci and Warner (1991)
 - *Weather and Forecasting* (June issue)
- Distribution of maximum surface potential temperature
 - What is the “normal” dry adiabat value by month over the high terrain of the west?
 - See their Fig. 6 for April (36-40 C), May (40-44 C), and June (48-52 C)

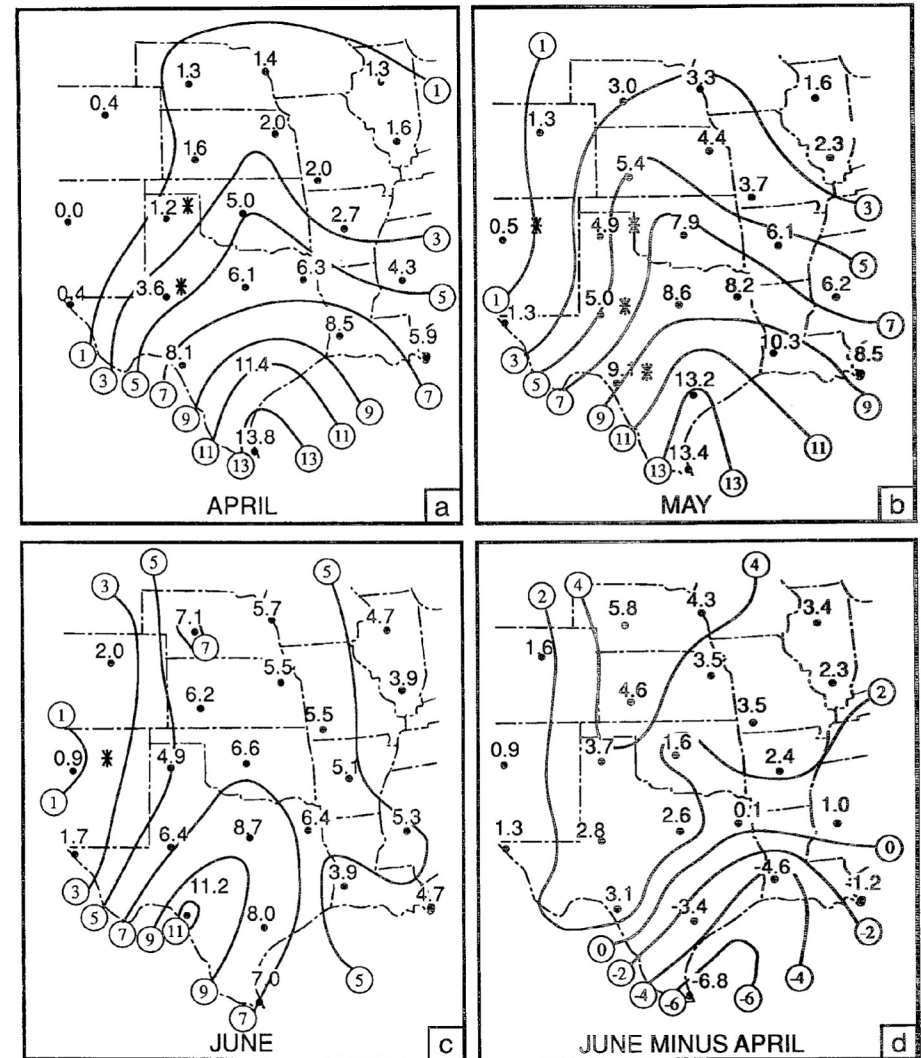


FIG. 5. Mean number of lid days over the region (see text for the definition of a “lid day”). An asterisk (*) next to the station means that the 1200 UTC lid frequency is greater than or equal to twice the 0000 UTC frequency. Charts are shown for a) April, b) May, and c) June. The change in lid days from April to June is shown by the difference field in (d). Lanicci and Warner (1991)

Climatology of Surface θ

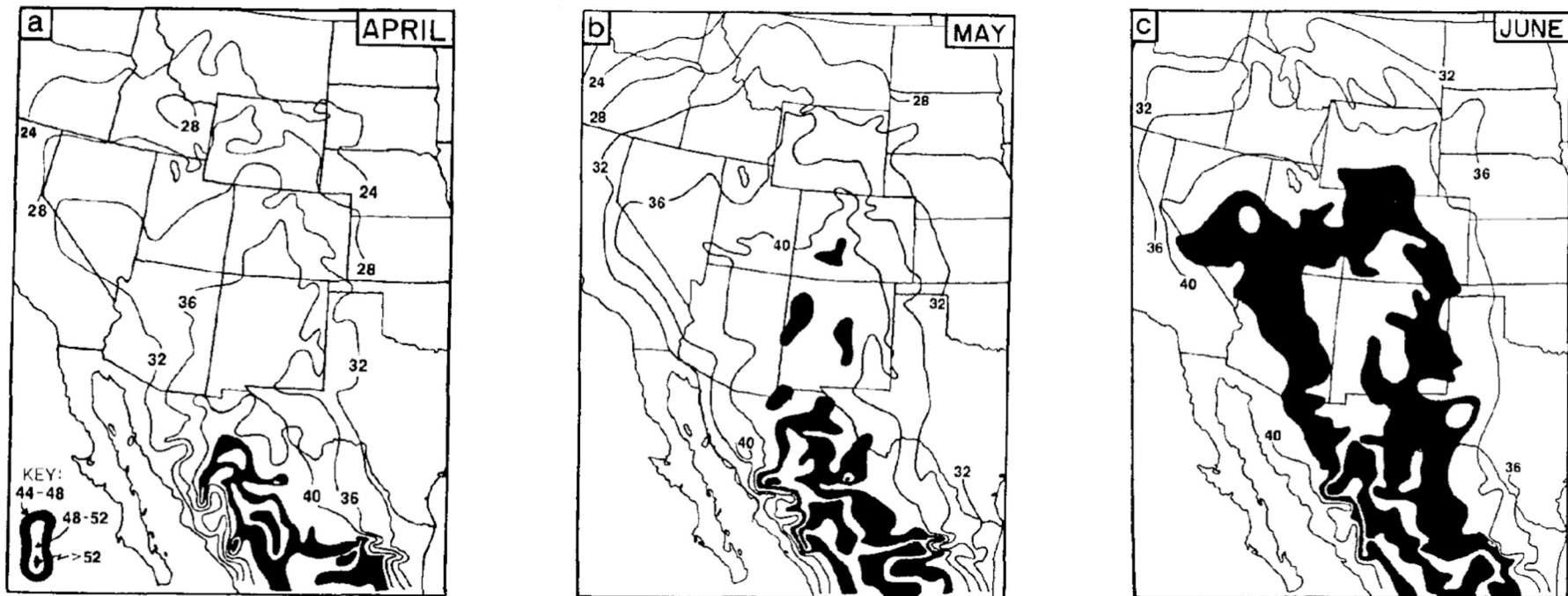
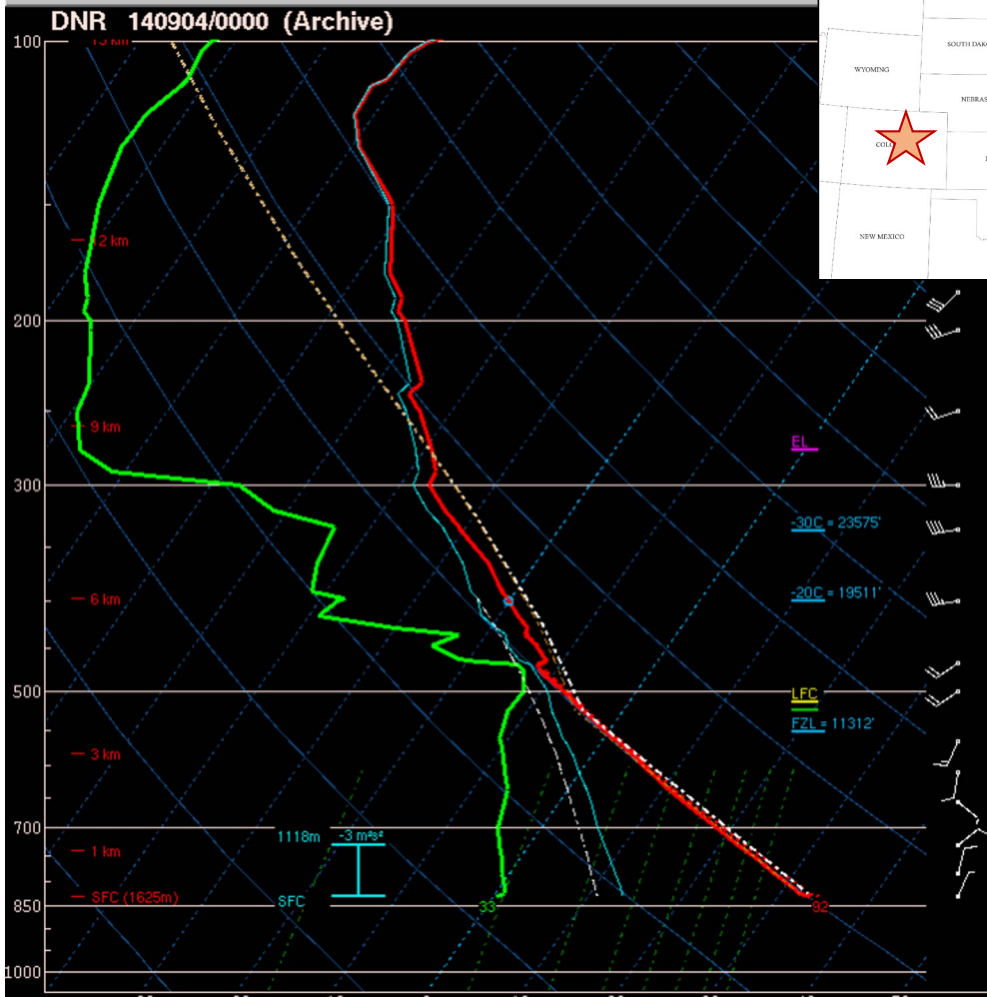
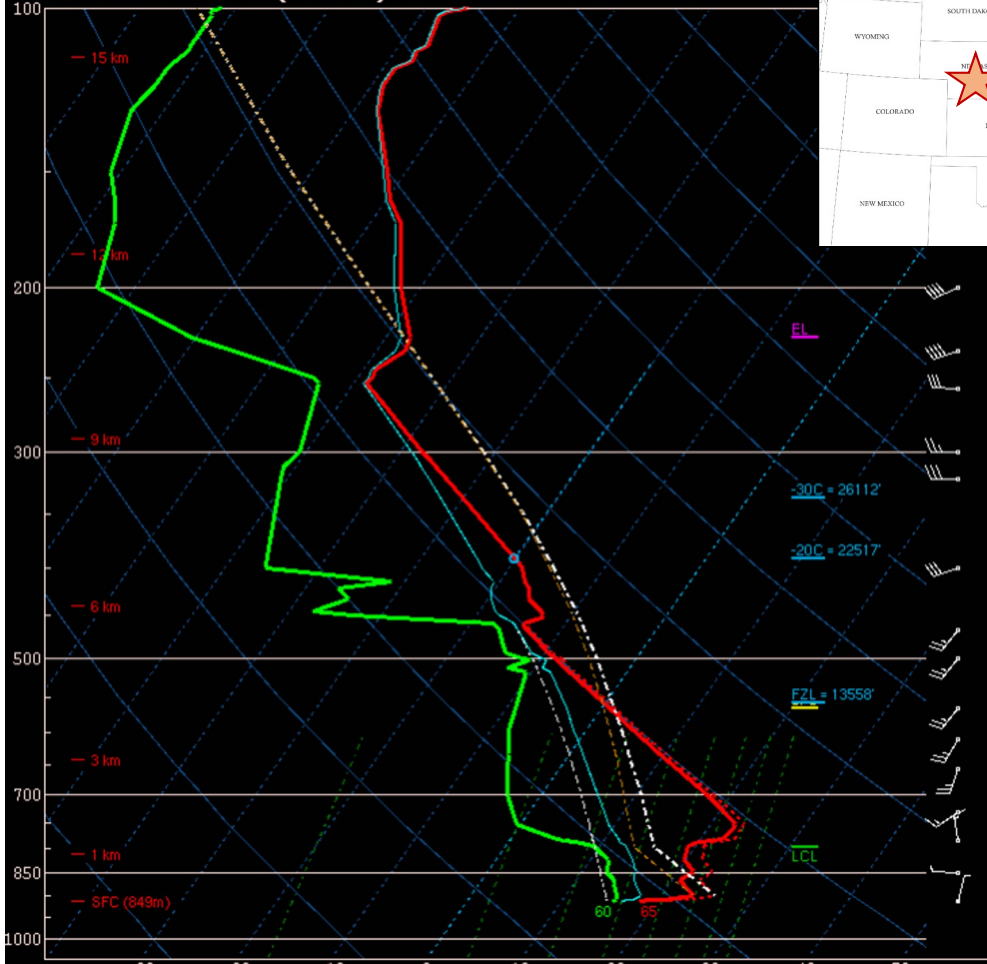


FIG. 6. Climatological mean maximum surface potential temperature analysis (in °C) over northern Mexico and the western U.S. for a) April, b) May, and c) June. Values greater than or equal to 44°C are alternately highlighted in 4°C intervals as shown in the lower left corner of chart (a).

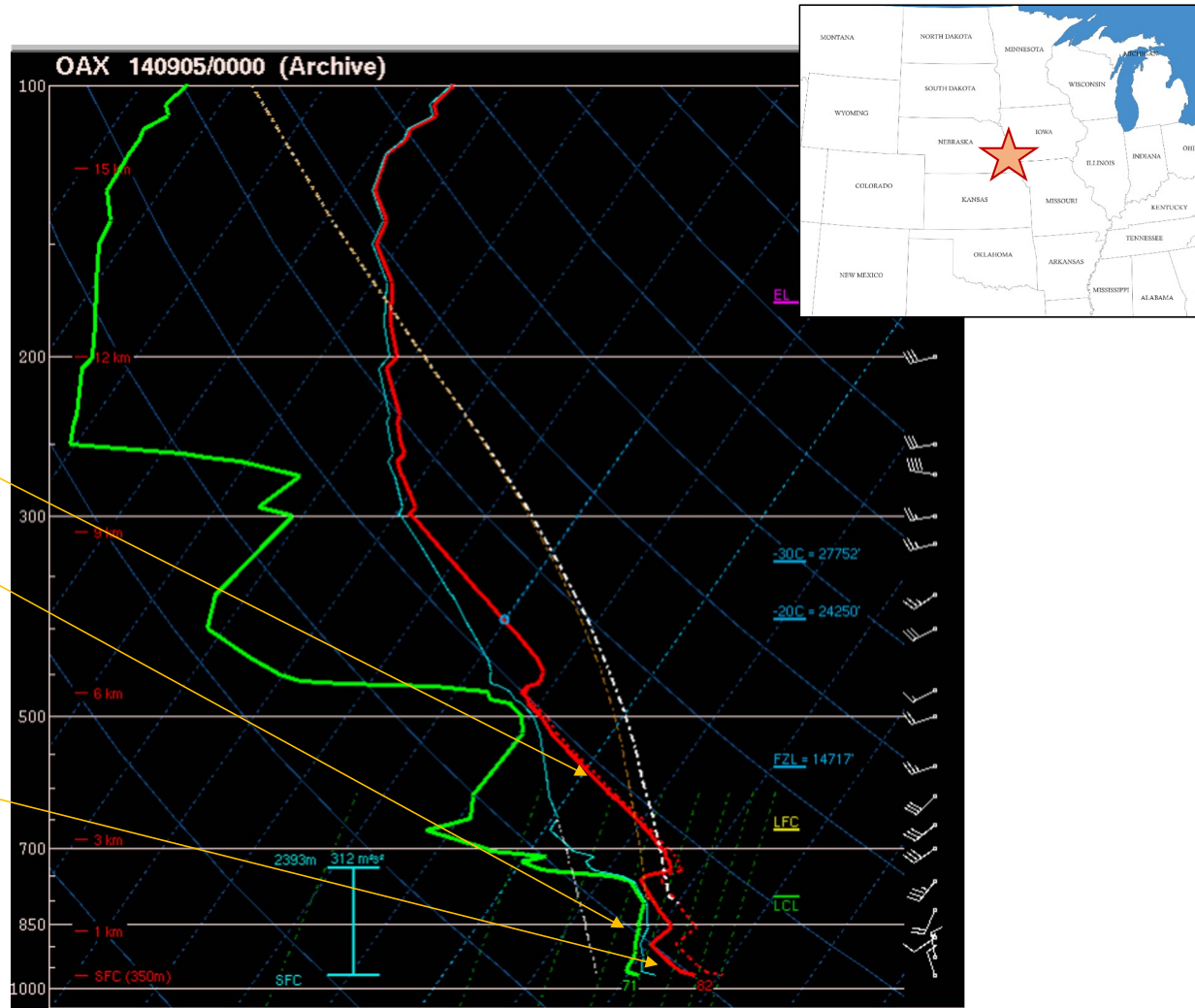


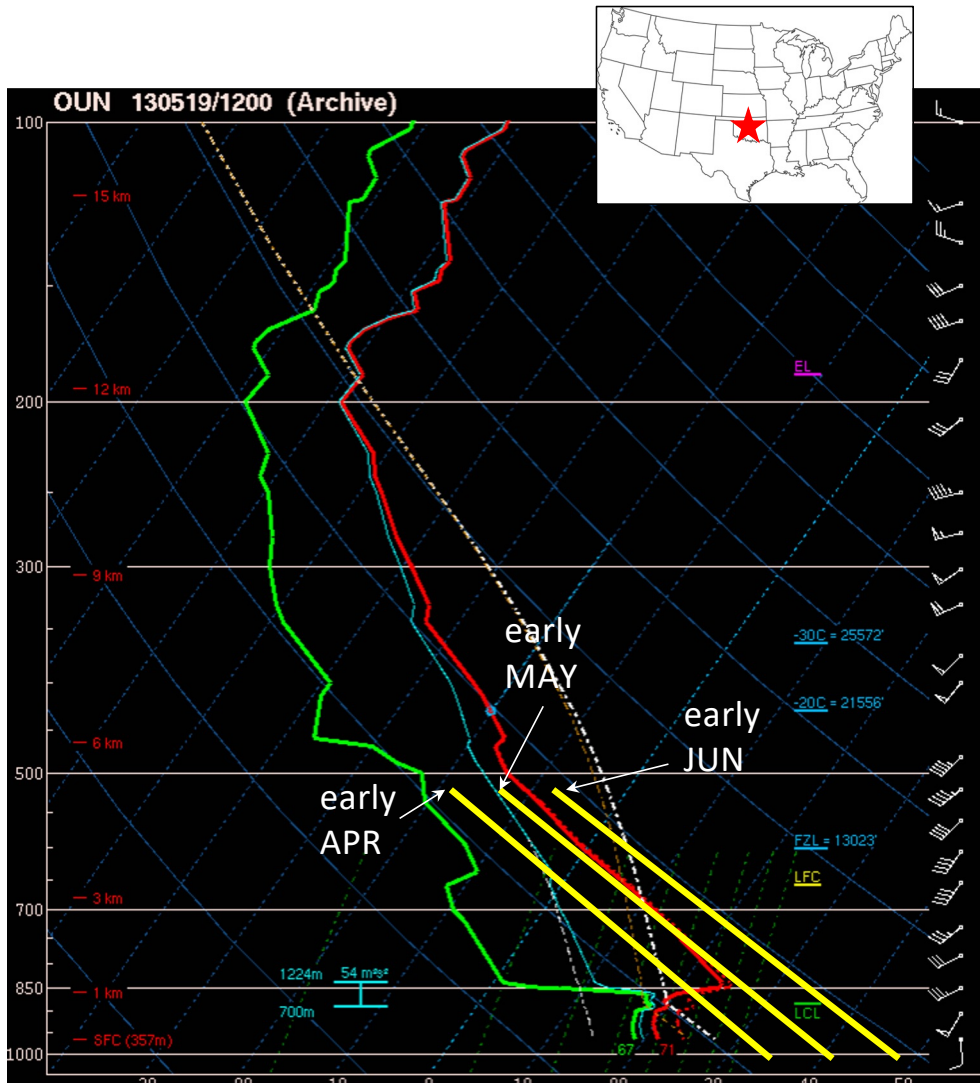
LBF 140904/1200 (Archive)



- Lid strength is related to:

- EML θ
- Depth/quality of moisture
- Surface heating and mixing in moist sector





EML Potential Temperature

- EML θ values 36-40C in April, 40-44C in May, 48-52C in June
- Warmer than normal surface temperatures over Rockies can lead to very warm EML and strong lid (or, cap) in Plains
- Cooler than normal surface temperatures over Rockies might suggest not much EML formation
- EML develops over a few days (like moisture return)

Lift

- **Synoptic-scale lift**

- 1-10 cm/s for many hours (6 hours \Rightarrow ~2 km ascent (SFC-800 hPa))
- 6 hours to reach LFC
- Conditions environment for deep, moist convection over broad area
 - Lapse rate stretching
 - Removal of inhibition
- Layered clouds in regions of isentropic lift; some elevated storms can form

- **Mesoscale lift**

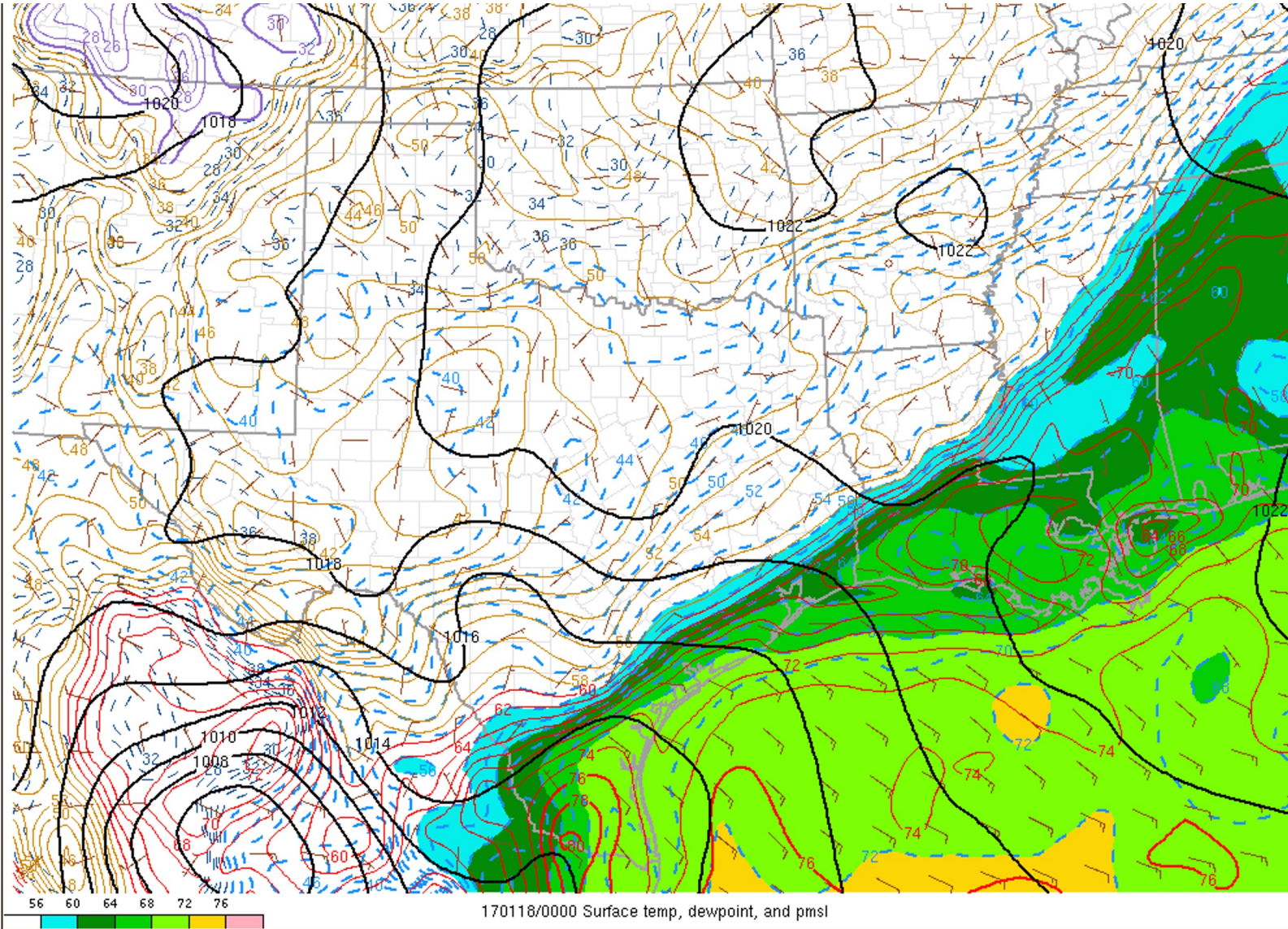
- 1 m/s for minutes to hours (1 hour \Rightarrow ~3.6 km ascent (SFC-650 hPa))
- < 1 hour to reach LFC
- Narrow zones of ascent along boundaries and terrain features
- Direct initiation of thunderstorms

- **Storm-scale lift**

- 10 m/s for minutes (15 mins \Rightarrow ~9 km of ascent (SFC-300 hPa))
- 5 minutes to reach LFC
- Storm maintenance and propagation (supercells, MCSs, squall lines, etc.)

Synoptic-Scale Lift

- **QG sources**
 - Warm advection/isentropic lift and differential vorticity advection
- **Jet streaks (also QG)**
 - Straight jet: ascent in right entrance and left exit regions
 - Curved jet: ascent downstream and along jet core



170118/0000 Surface temp, dewpoint, and pmsl

- SfcOA Diag - - RAP

Image overlays:

- County Boundaries
- County Warning Area
- Hiways & Cities
- ARTCC Regions
- NWS Watches & Warr
- SPC Day1 Outlook

Image underlays:

- None
- Radar
- Terrain
- Population
- Surface Obs

Current SPC Prod

Show popup images?

Day1 Convective Outl

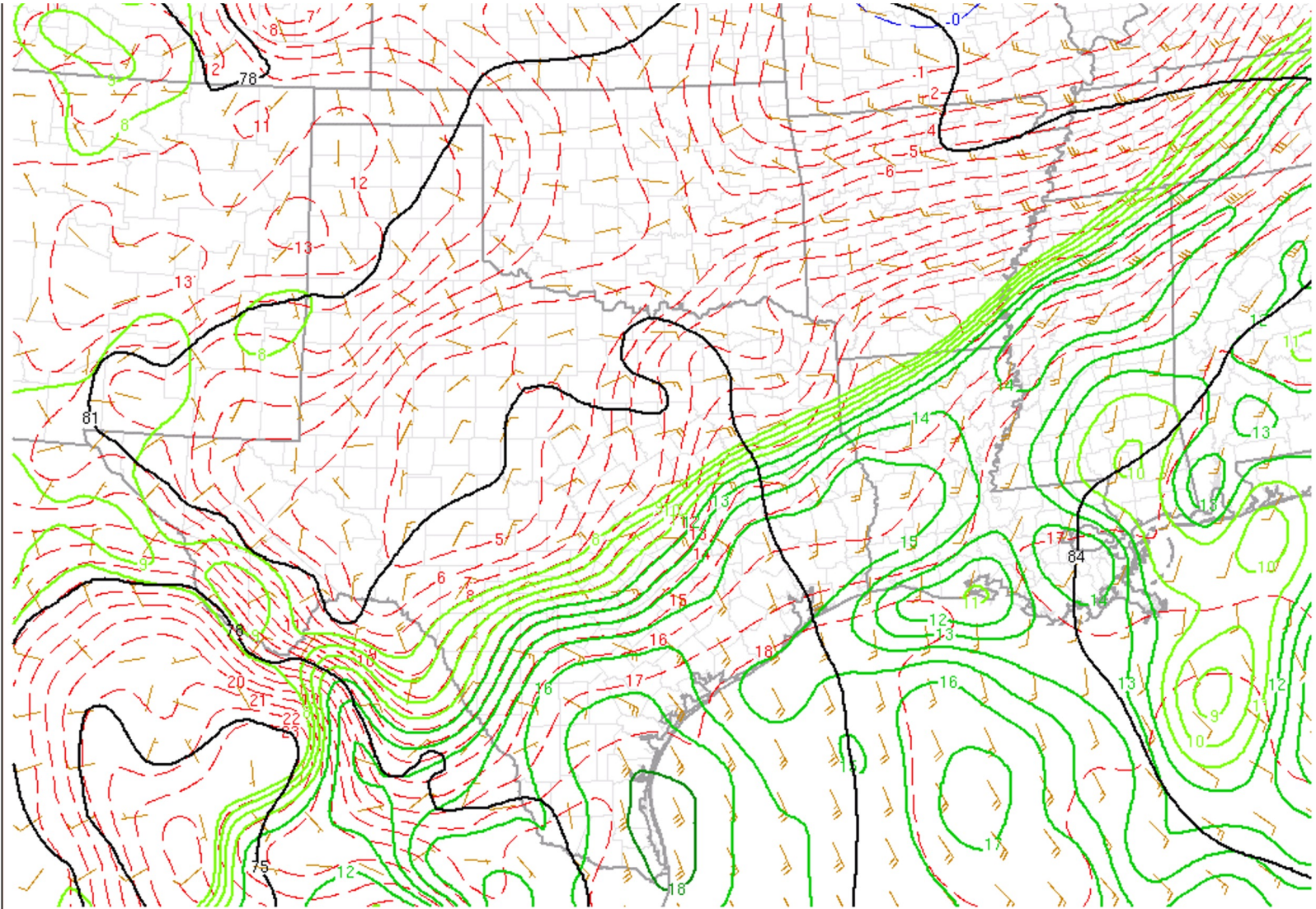
Issued at 1957 UTC

Probabilities: **T**

Day1 National Fire Ou

Issued at 1819 UTC

This list updates automaticl



170118/0000V001 925mb hght/temp/dwpt/wind

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Image underlays:

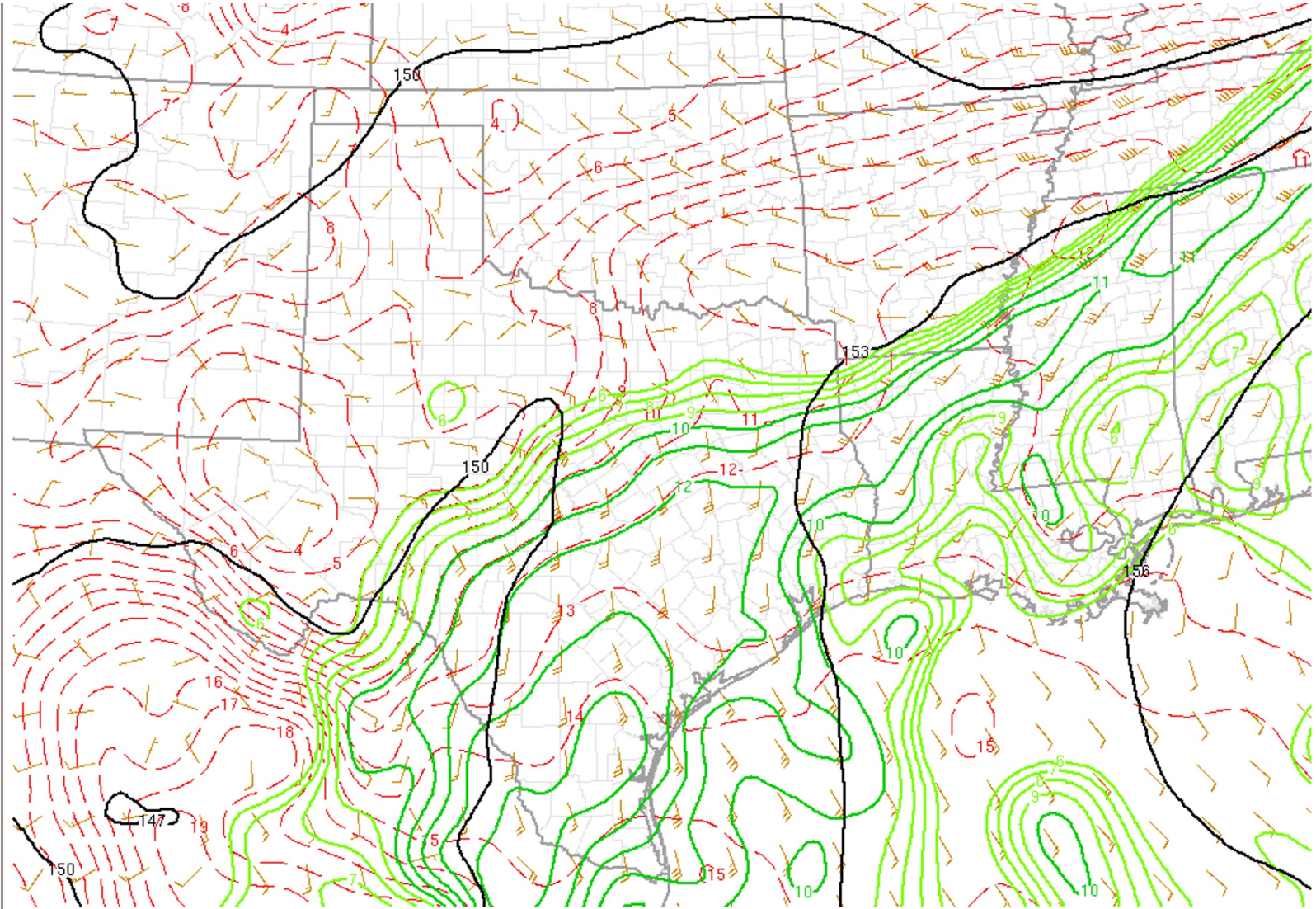
- None
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- Population
- Surface Obs

Current SPC Prod

Show popup images?

Day1 Convective Outl
Issued at 1957 UTC
Probabilities: **T**

Day1 National Fire Ou
Issued at 1819 UTC
This list updates automatically



170118/0000V001 850mb hght/temp/dwpt/wind

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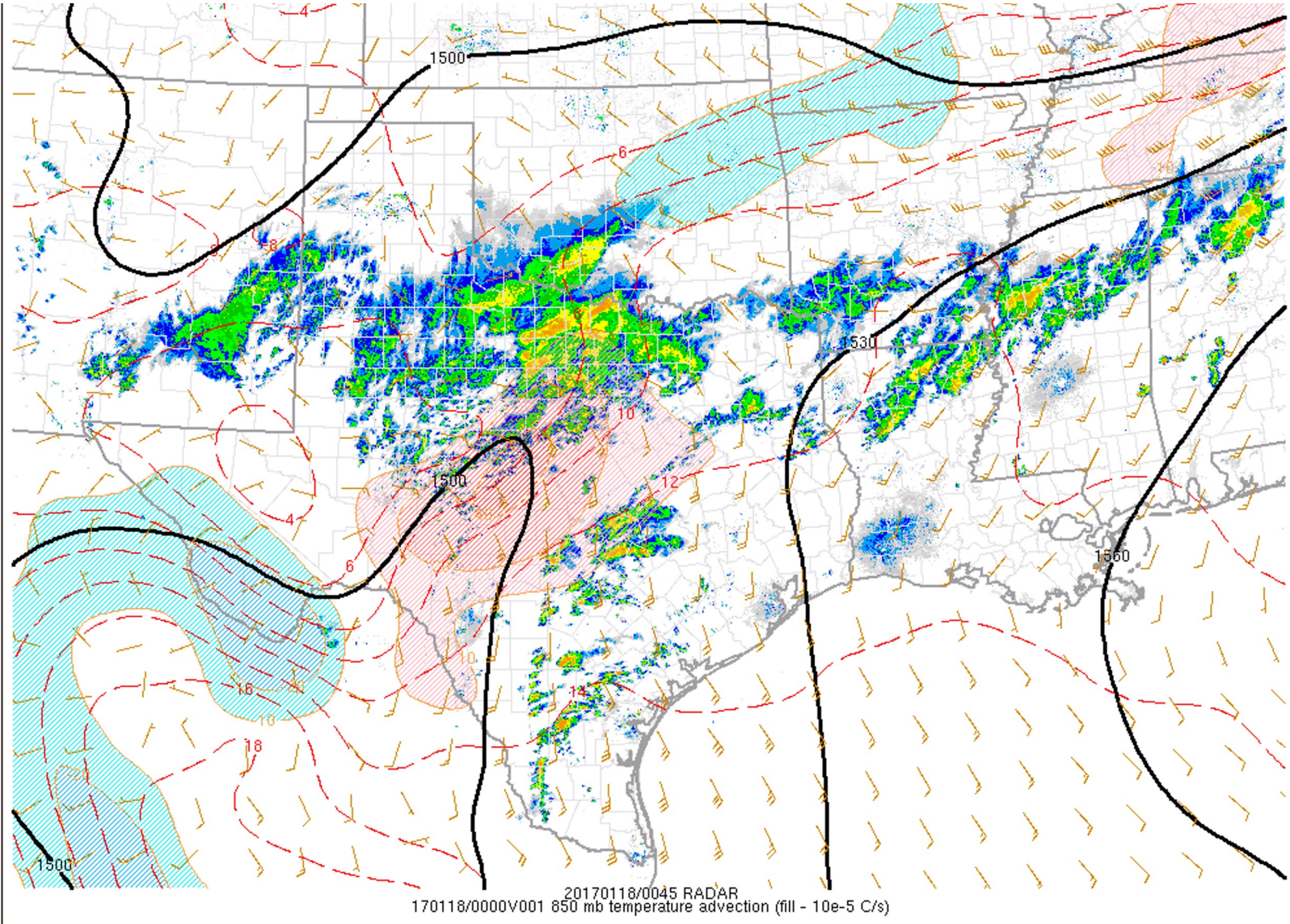
Current SPC Prod

Show popup images?

Day1 Convective Outl
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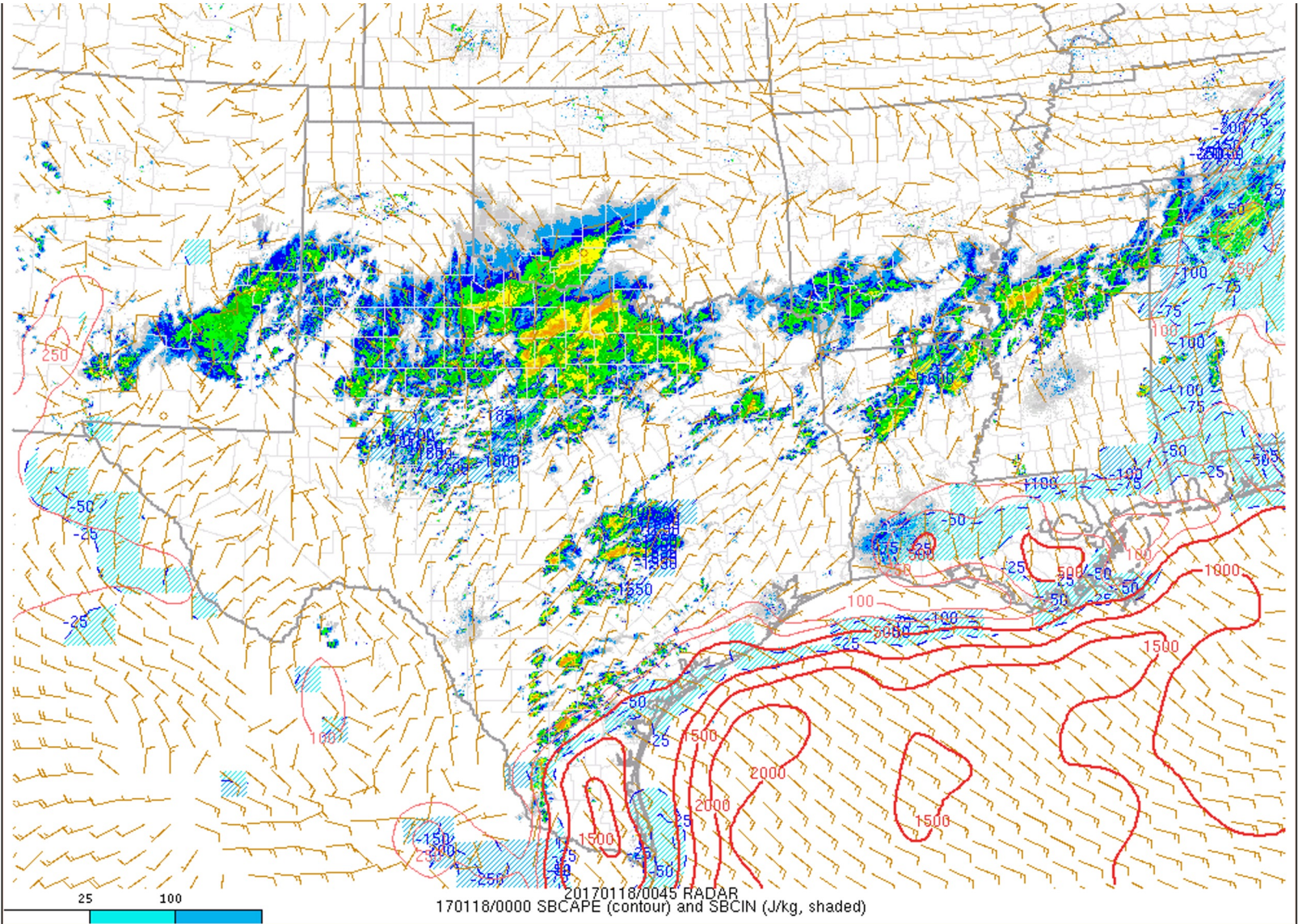
Current SPC Product

Show popup images?

Day1 Convective Outlook
Issued at 1957 UTC
Probabilities: **T**

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Issued at 1819 UTC

This list updates automatically



- SfcOA Diag - - RAP -

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- NWS Watches & Warr
- SPC Day1 Outlook

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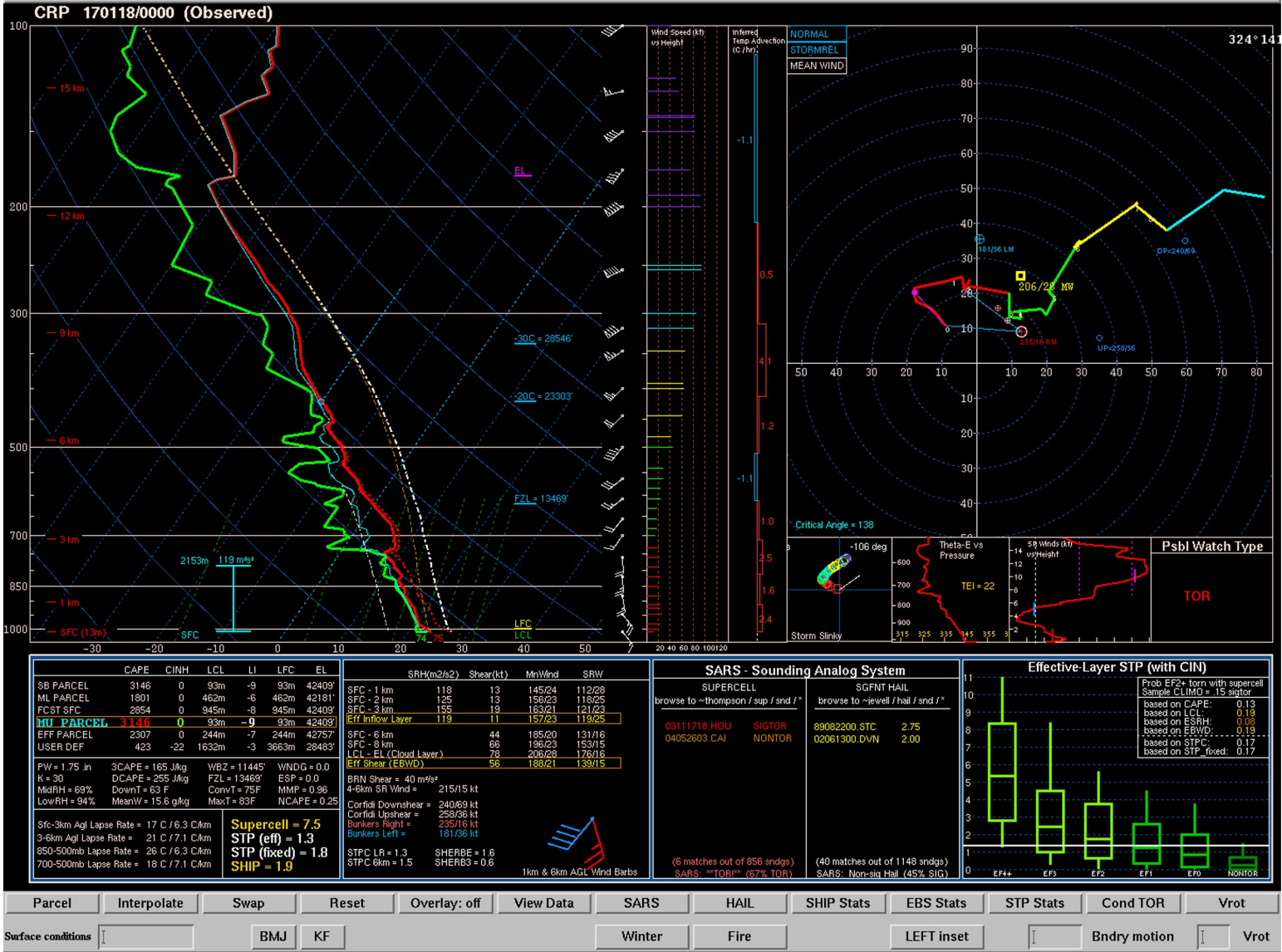
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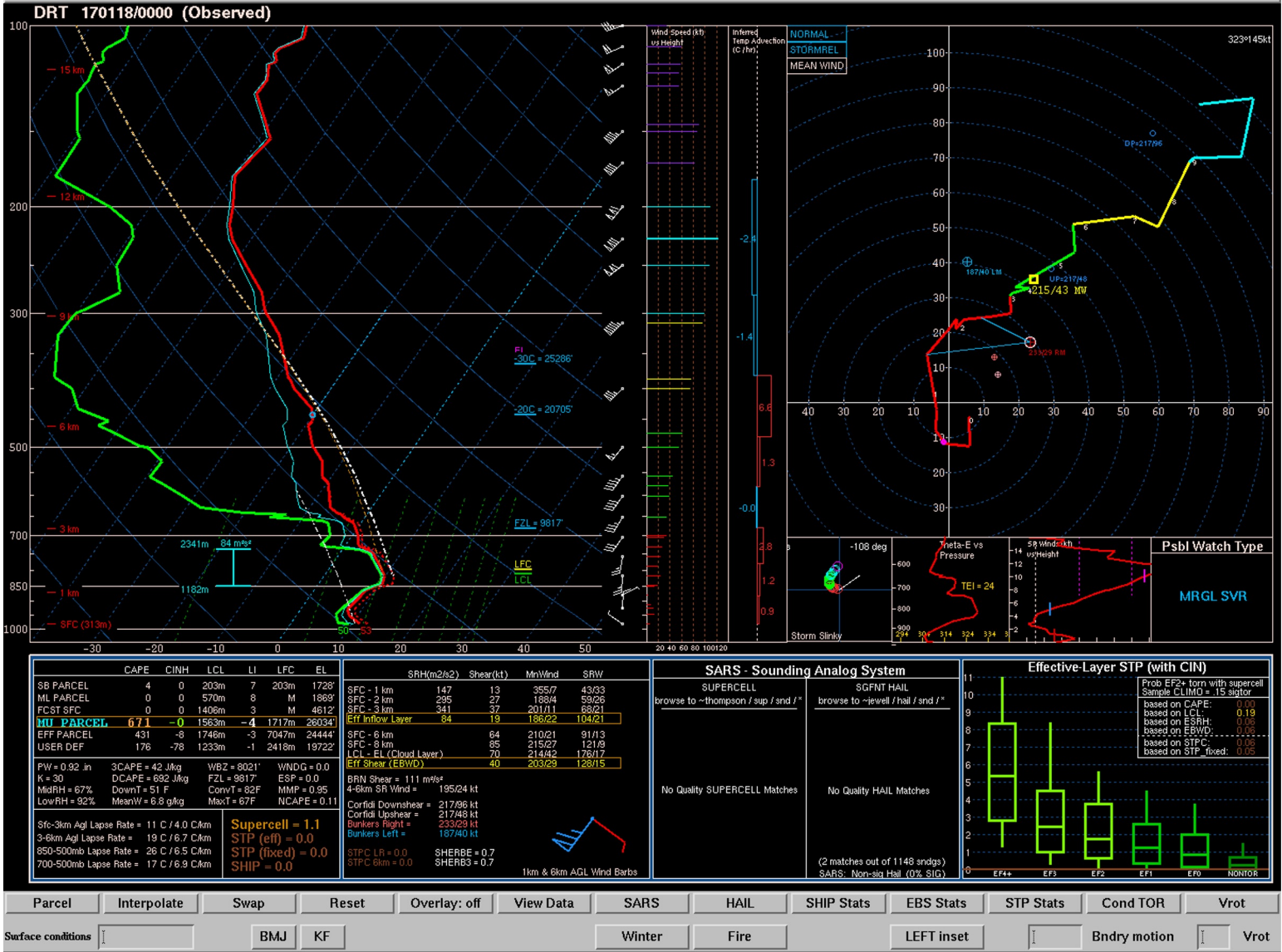
Current SPC Prod

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Day1 Convective Outl
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This list updates automatically





Synoptic-Scale Lift Summary

- **Gradual ascent over many hours**
 - Warm advection and differential vorticity advection are primary drivers
- **Primary role of QG ascent is to precondition environment for convection**
 - Can also initiate elevated storms in regions of warm advection
 - Most surface-based storms are triggered by mesoscale ascent

Mesoscale Lift

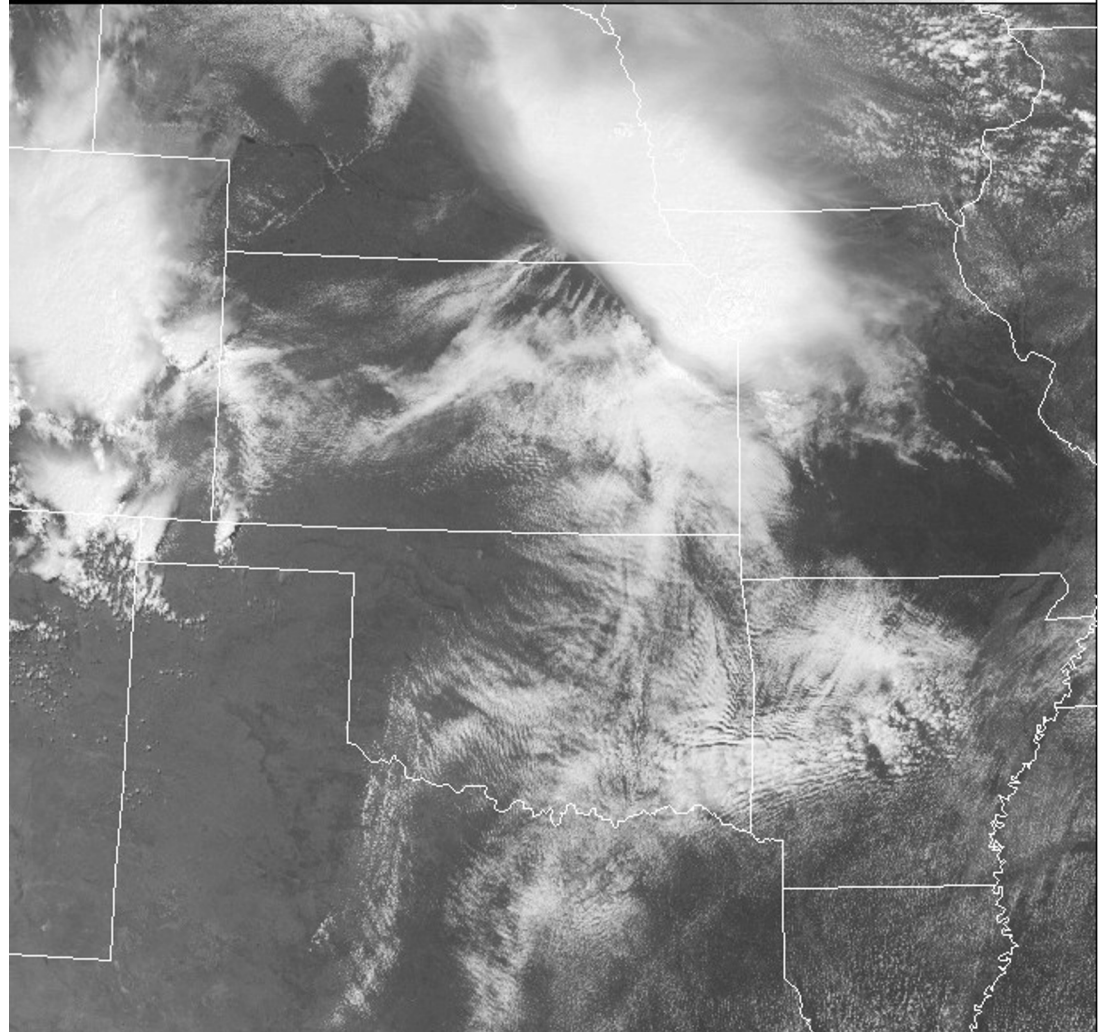
- Focused and narrow zones of lift for minutes to hours
- Fronts
 - Rising on warm side with frontogenesis
 - Isentropic ascent on sloped warm front (“overrunning”); similar to synoptic lift except in focused bands
- Outflow boundary
 - Lift is governed by depth/strength of cold pool and low-level vertical wind shear [Rotunno-Klemp-Weisman (RKW) Theory]
- Dryline
 - Lift on dry/hot side due to frontogenesis
 - Lift also governed by depth of mixing west of dryline and depth of moist layer east of dryline

1740 UTC Tue 24 May 2011

Visible Satellite

<http://adds.aviationweather.gov>

0 6 9 12 15 18 21 24 27 30 33 36 39 42 45 48 51 54 57 60 63 66 69 72 75 78 81 84 87 90 93 96 99

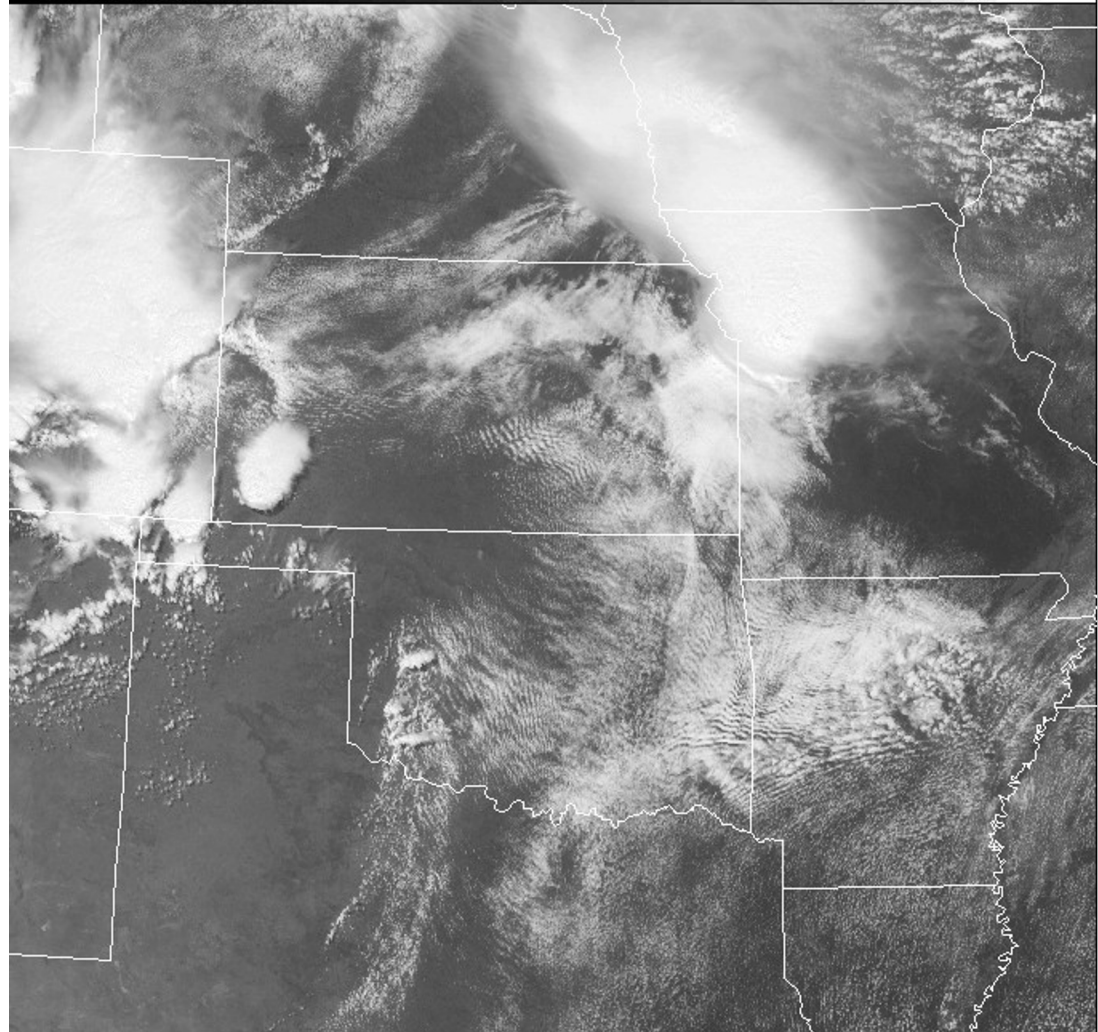


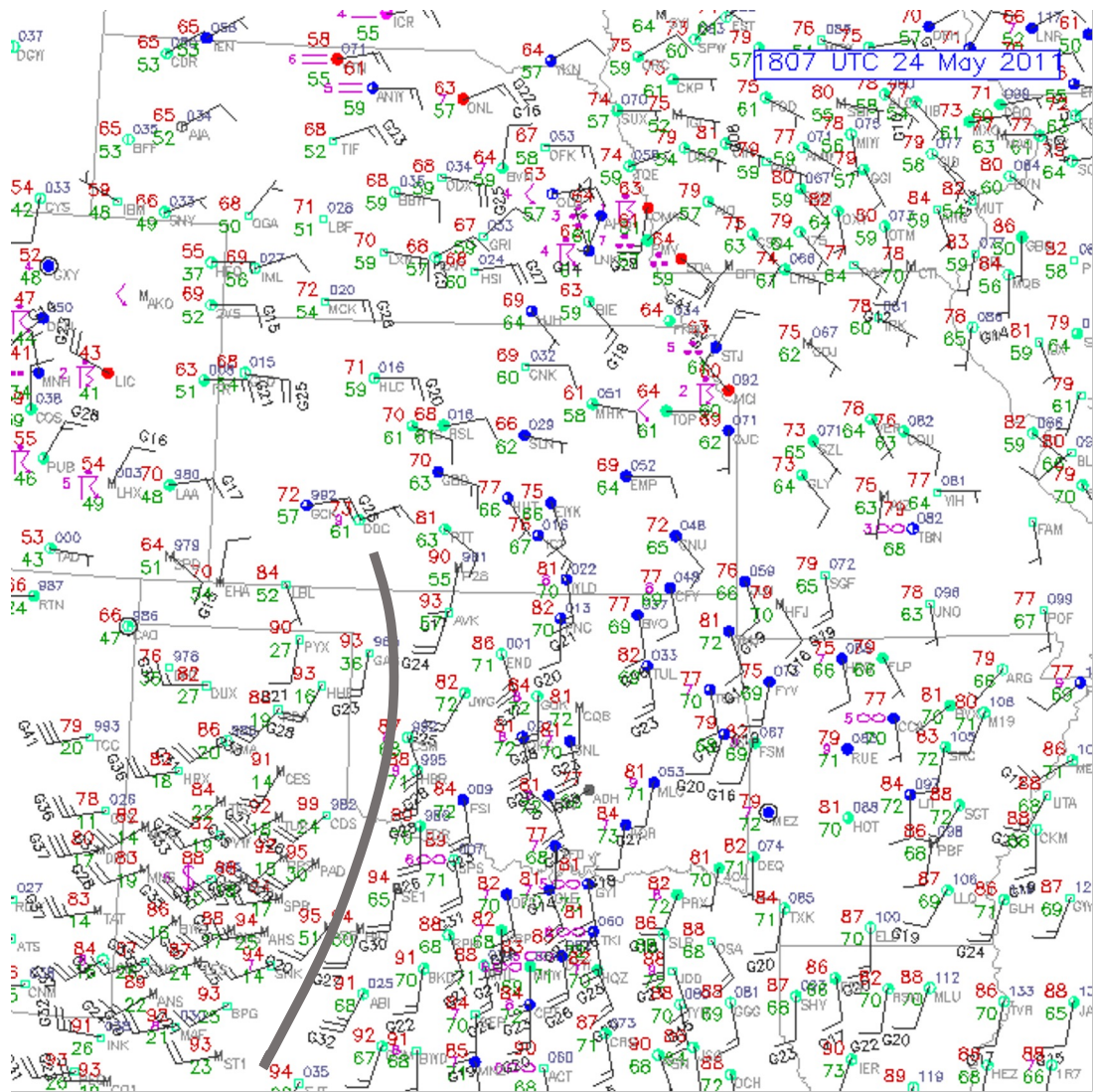
1845 UTC Tue 24 May 2011

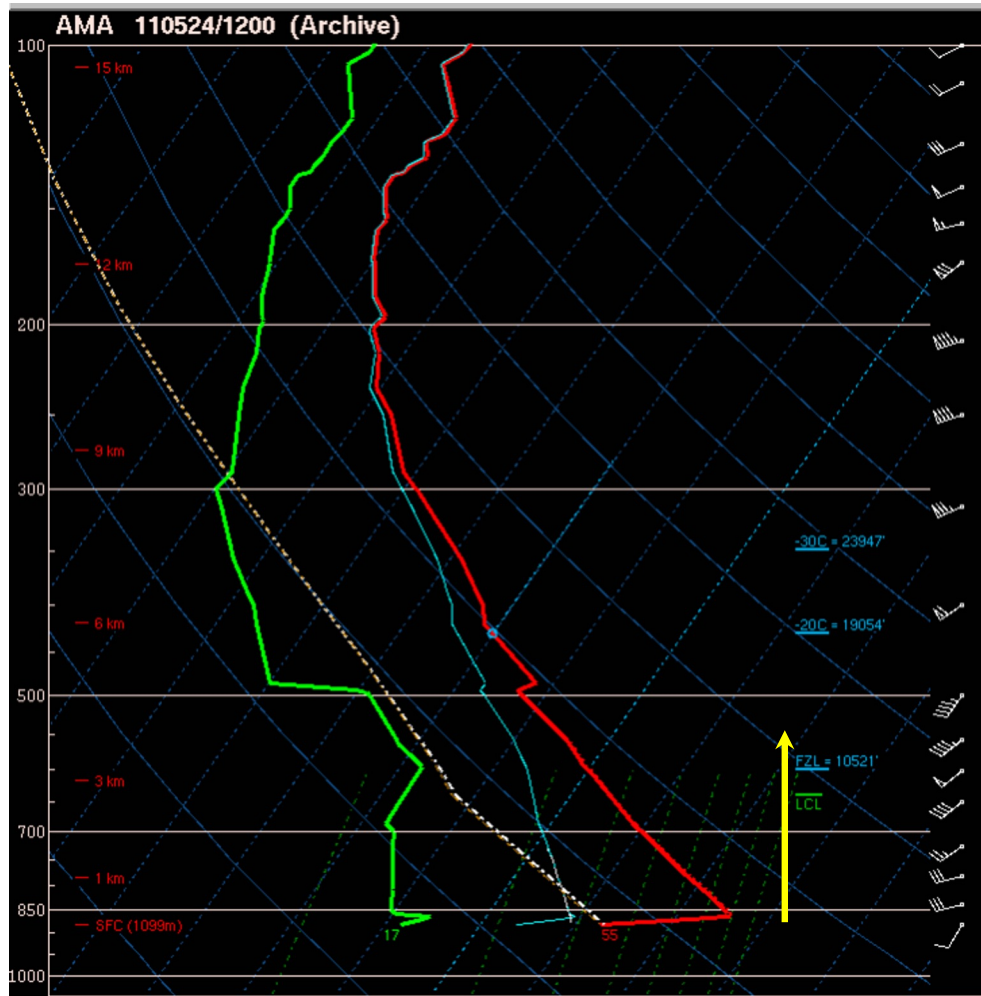
Visible Satellite

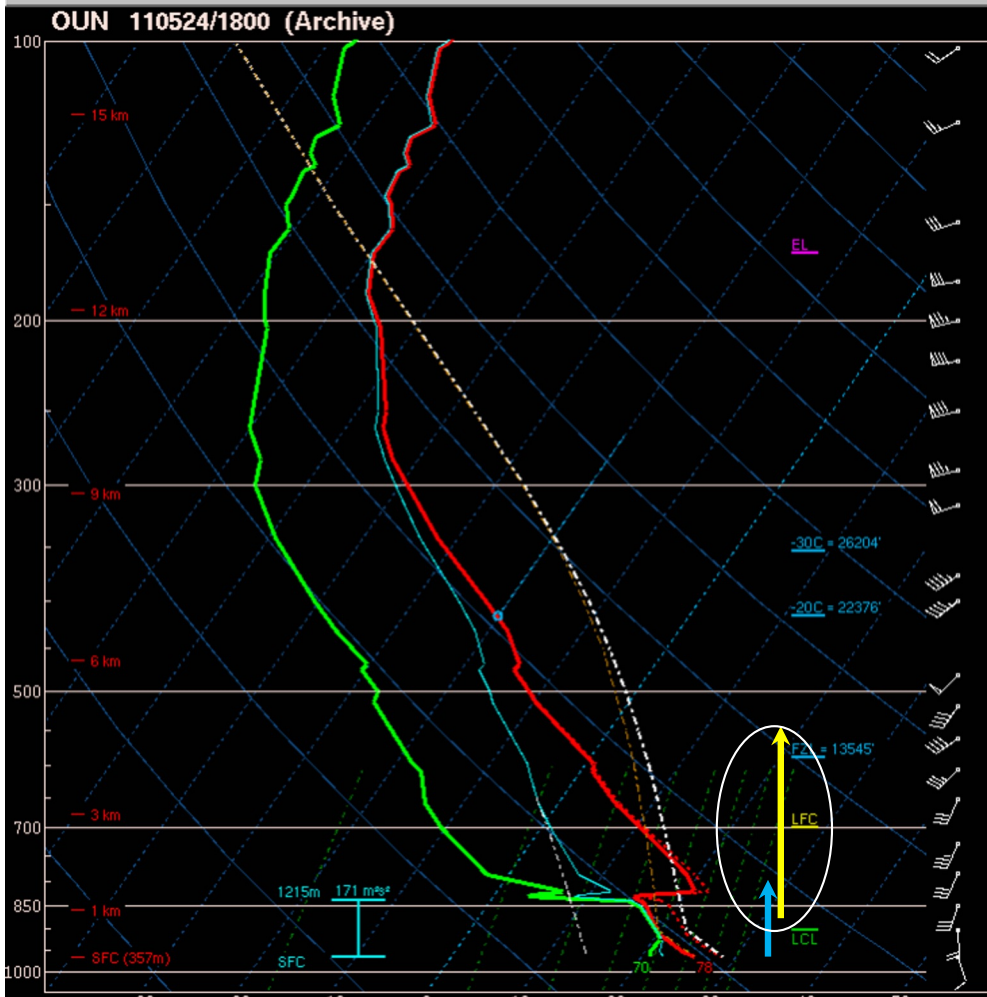
<http://adds.aviationweather.gov>

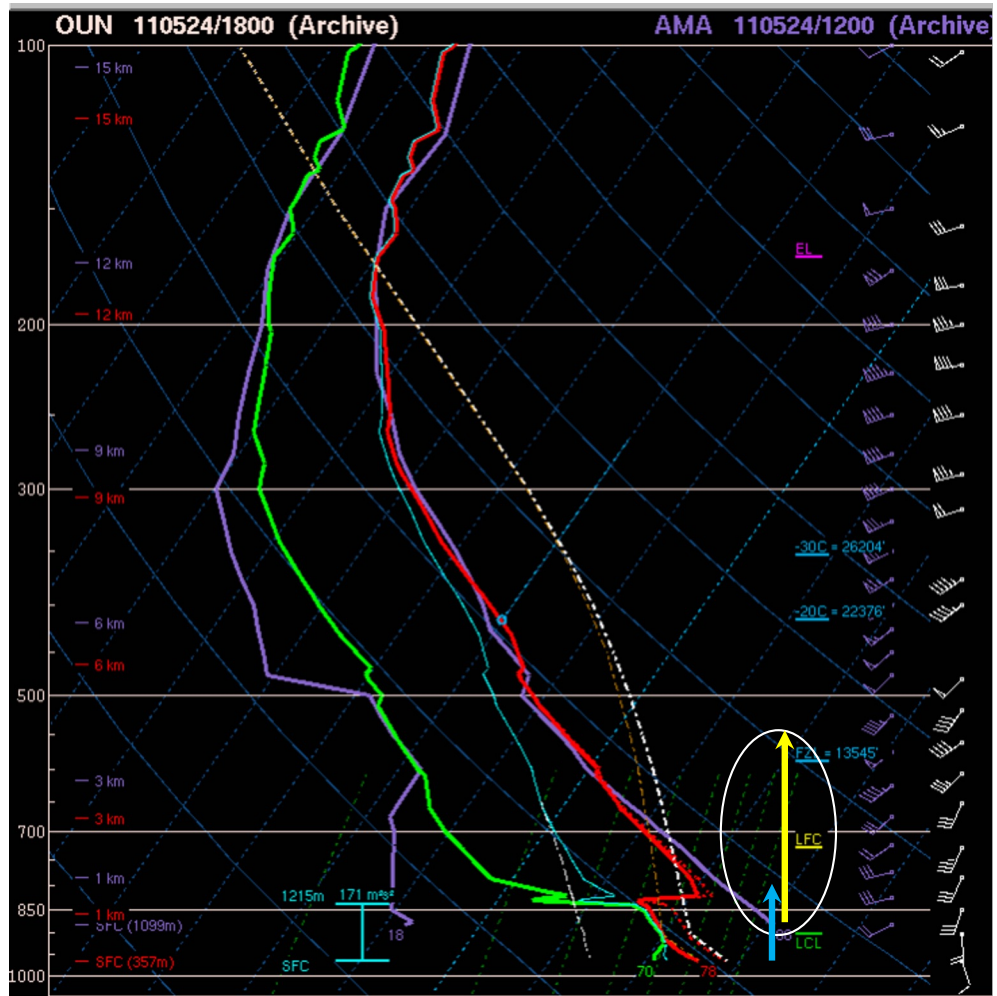
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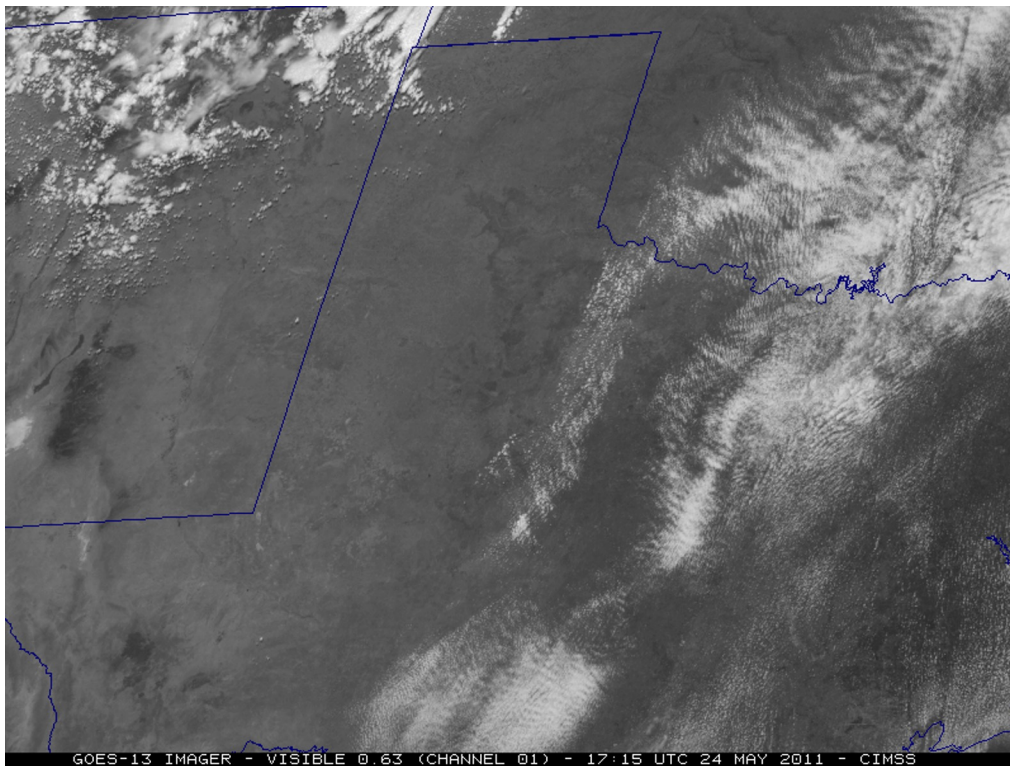












Dryline Summary

- **Focused ascent along dryline can be ~1 m/s (~1 hour to reach LFC)**
 - Lift depends on depth of mixing west of dryline
- **Convection initiation linked to the residence time of air in the zone of ascent**
 - Winds above the surface parallel to dryline keeps air in zone of life longer
- **Convection will initiate in points or bands**
 - Usually the mode of initiation for surface-based supercells in central/southern Plains

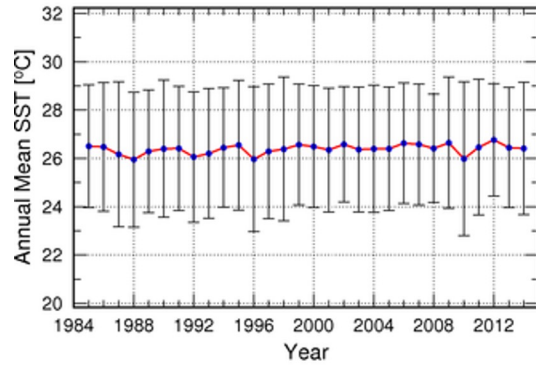
Moisture

- **Moisture in PBL needs to be large and deep enough for convection**
- **Return flow**
- **Inland vertical mixing/diurnal processes**
- **Evapotranspiration**

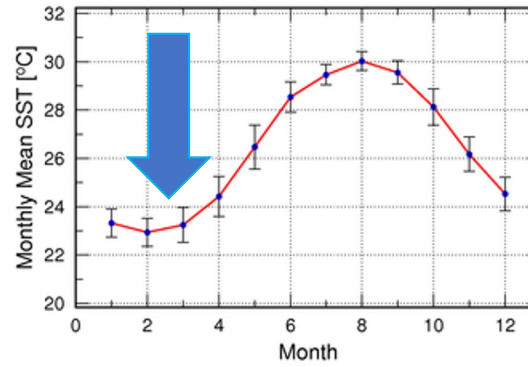
Low-level Moisture Return Flow

- Key elements include air mass modifications over warmer water and return flow trajectories
- Air mass modification is a multi-date process
- Ask yourself the following:
 - Where is the air coming from?
 - What are the underlying ocean characteristics?
 - What is the character of the returning moist layer?

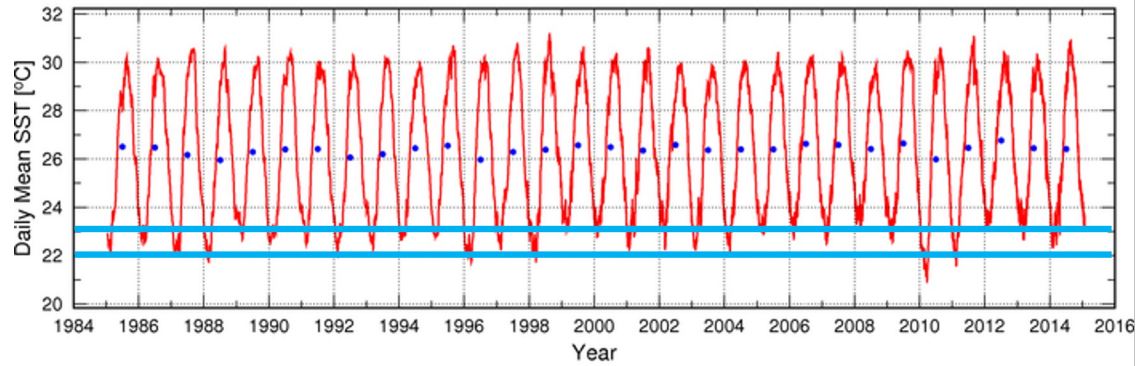
GOM Annual Mean SST



GOM Monthly Mean SST

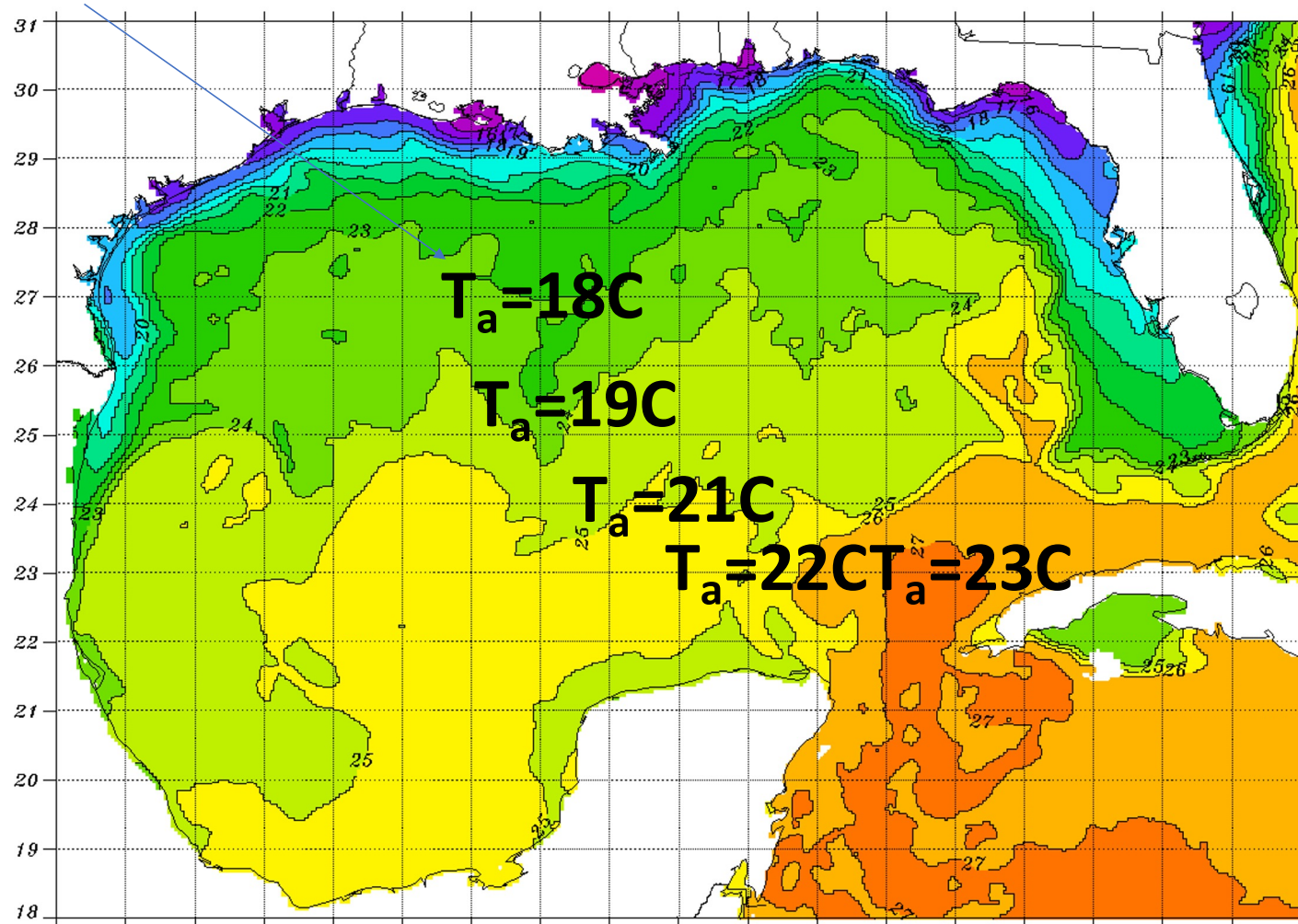


GOM Daily Mean SST

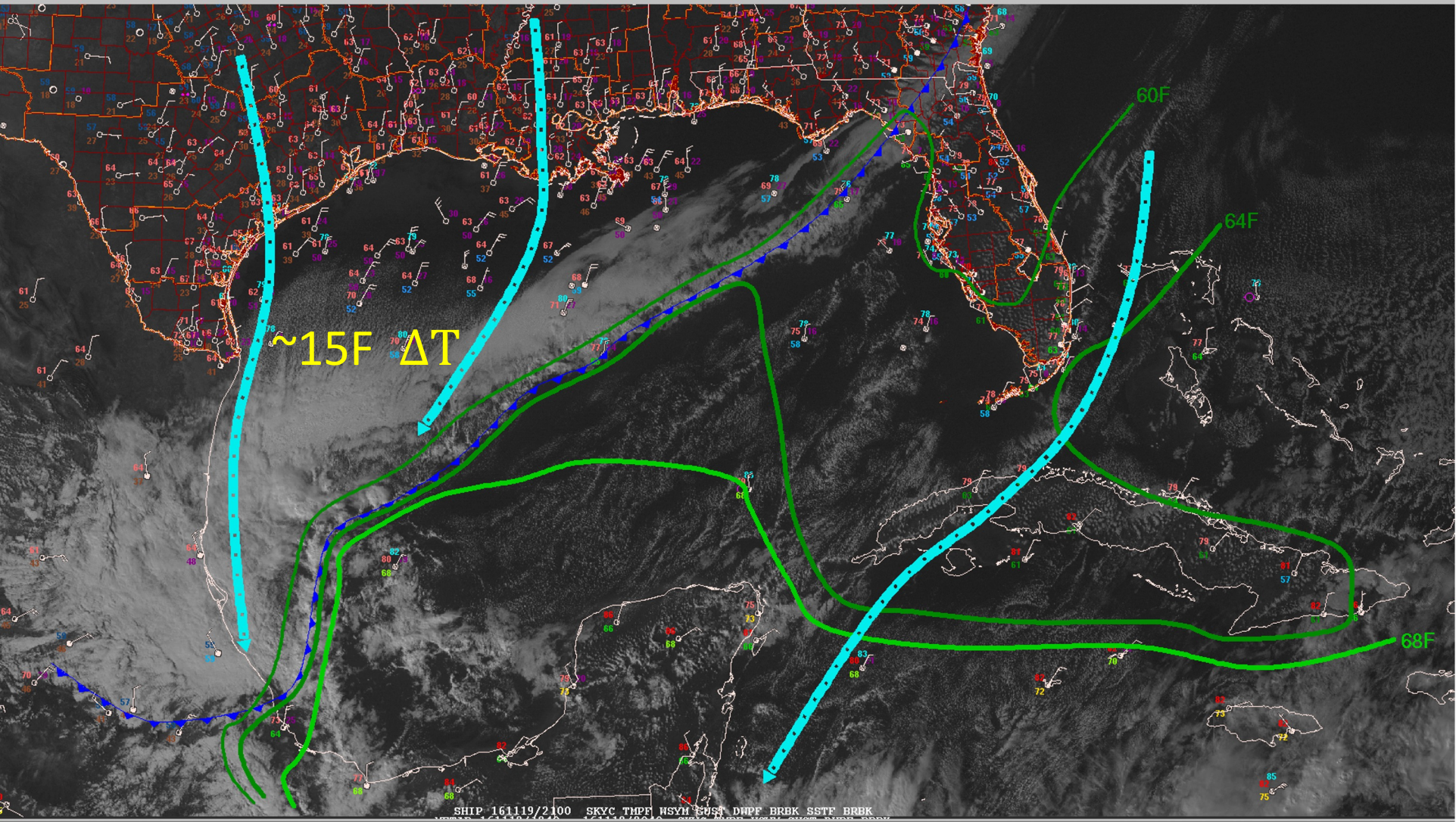


NOAA/NESDIS GEO-POLAR BLENDED 5 km SST ANALYSIS
FOR THE GULF OF MEXICO

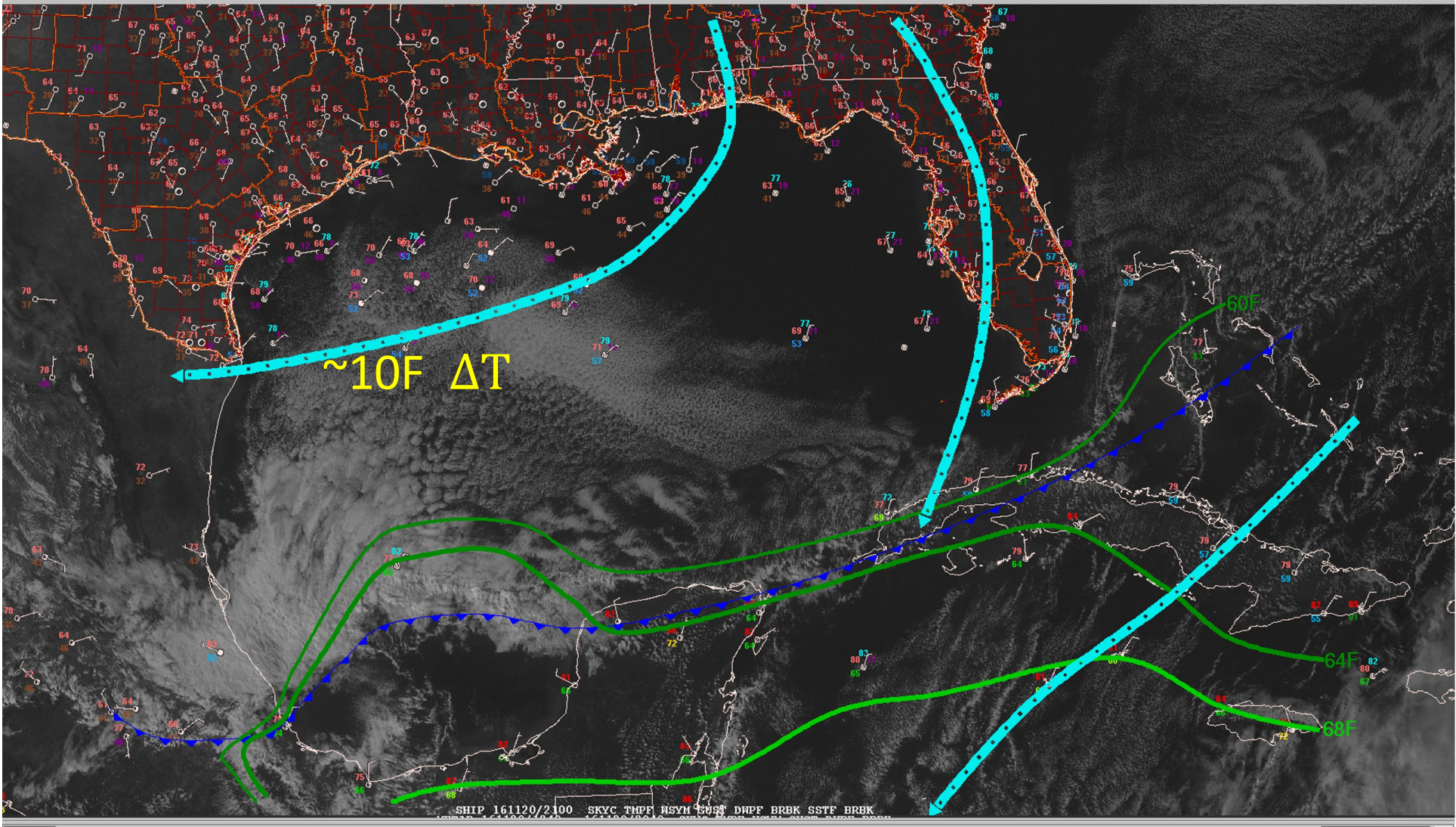
Air temperature after cold front passage



Navigation toolbar with icons for Defeat, AUTO Repeat OFF, MAP, L, PRINT, RELOAD, and a DPT icon. A 'Loop: 2' indicator is present, along with Zoom and Unzoom buttons. A 'Hide Loop' button and a set of navigation arrows are also visible, along with a 'Stop' button.

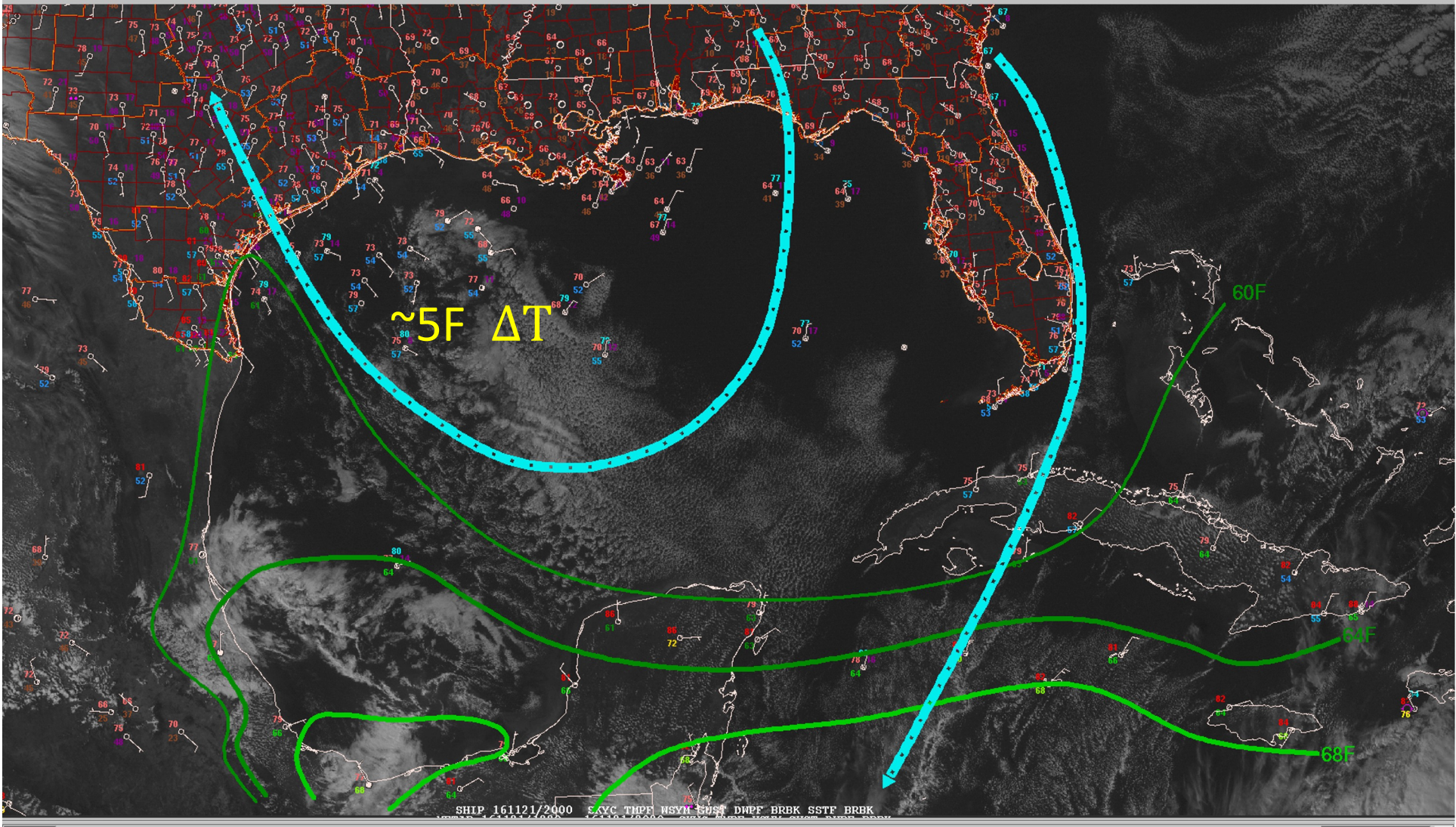


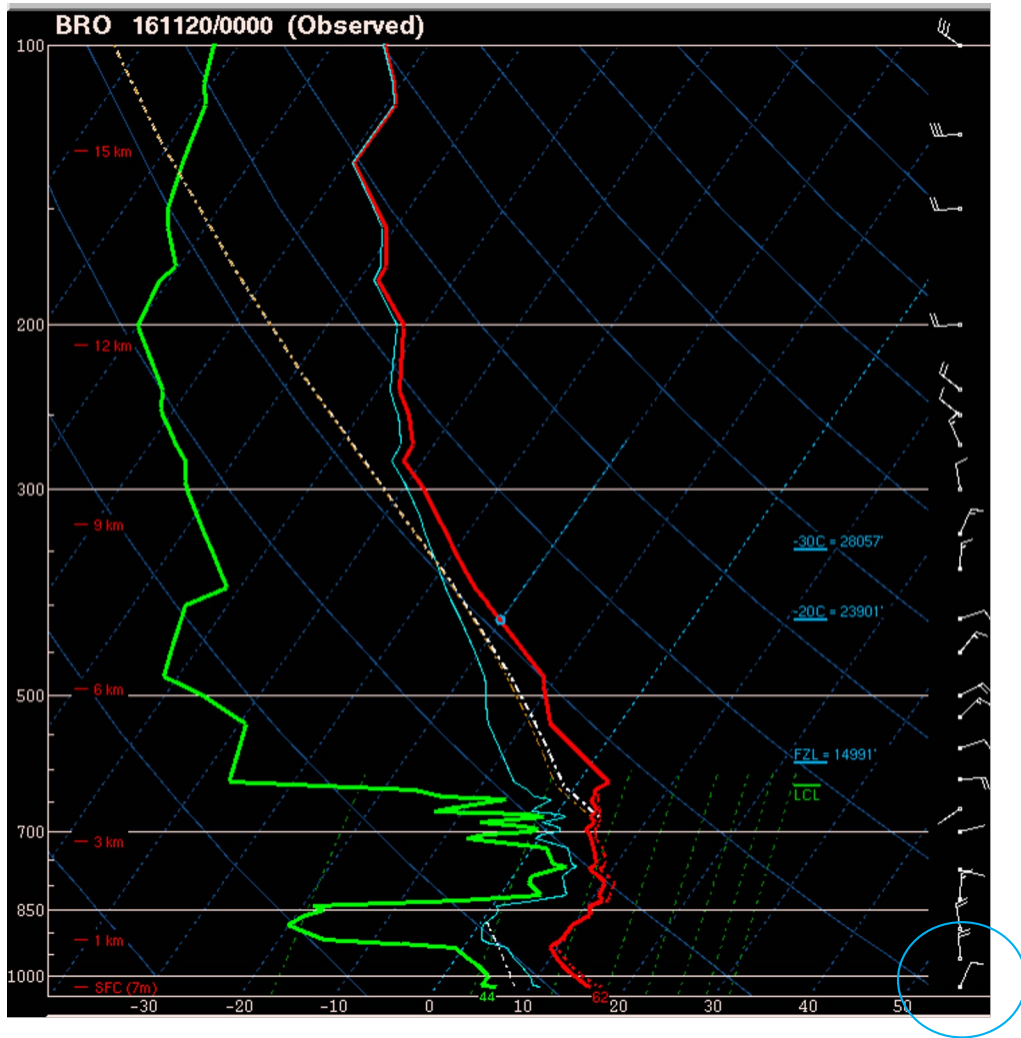
Default AUTO Update OFF MAP L PRINT RELOAD [Navigation icons] Loop: 2 [Zoom Unzoom] Hide Loop [Navigation icons] Stop

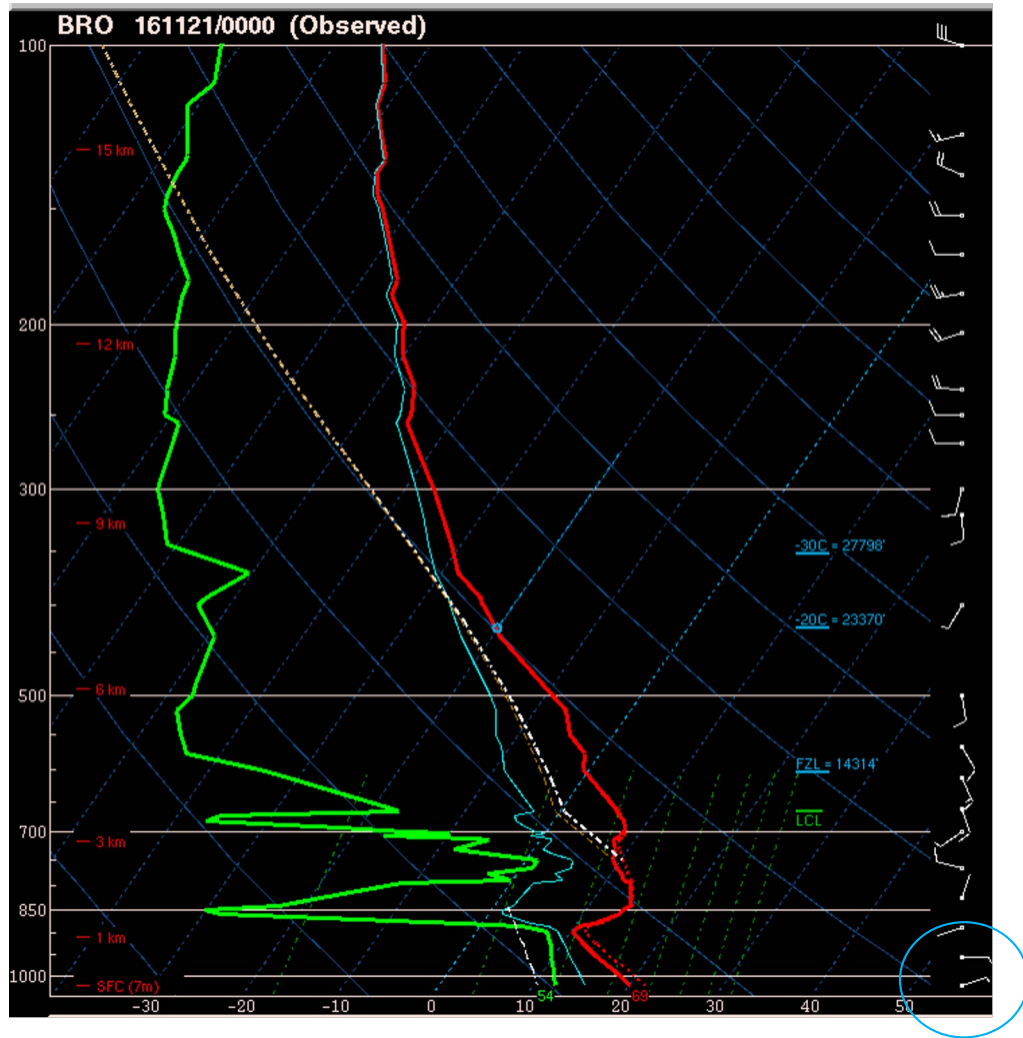


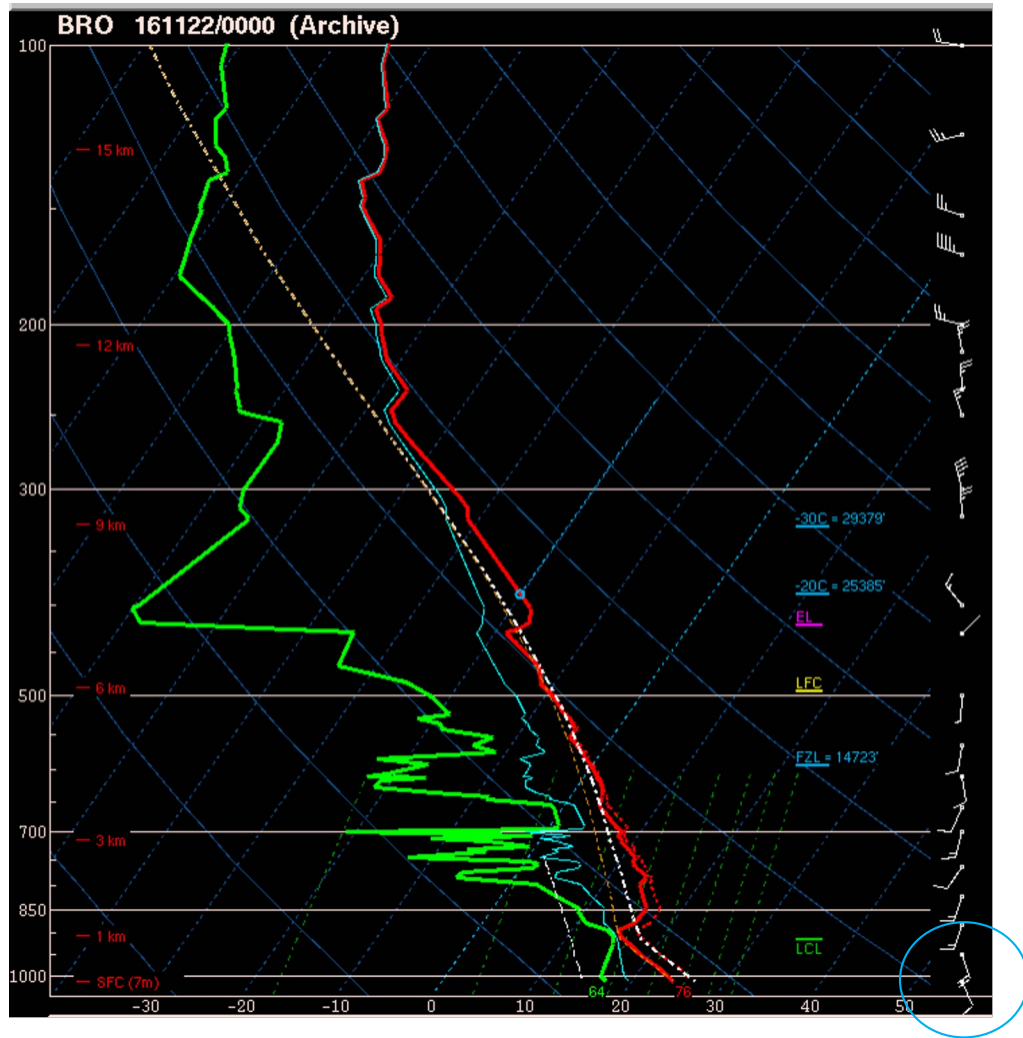
Default AUTO Update OFF MAP L PRINT RELOAD DPT Zoom Unzoom Hide Loop Stop

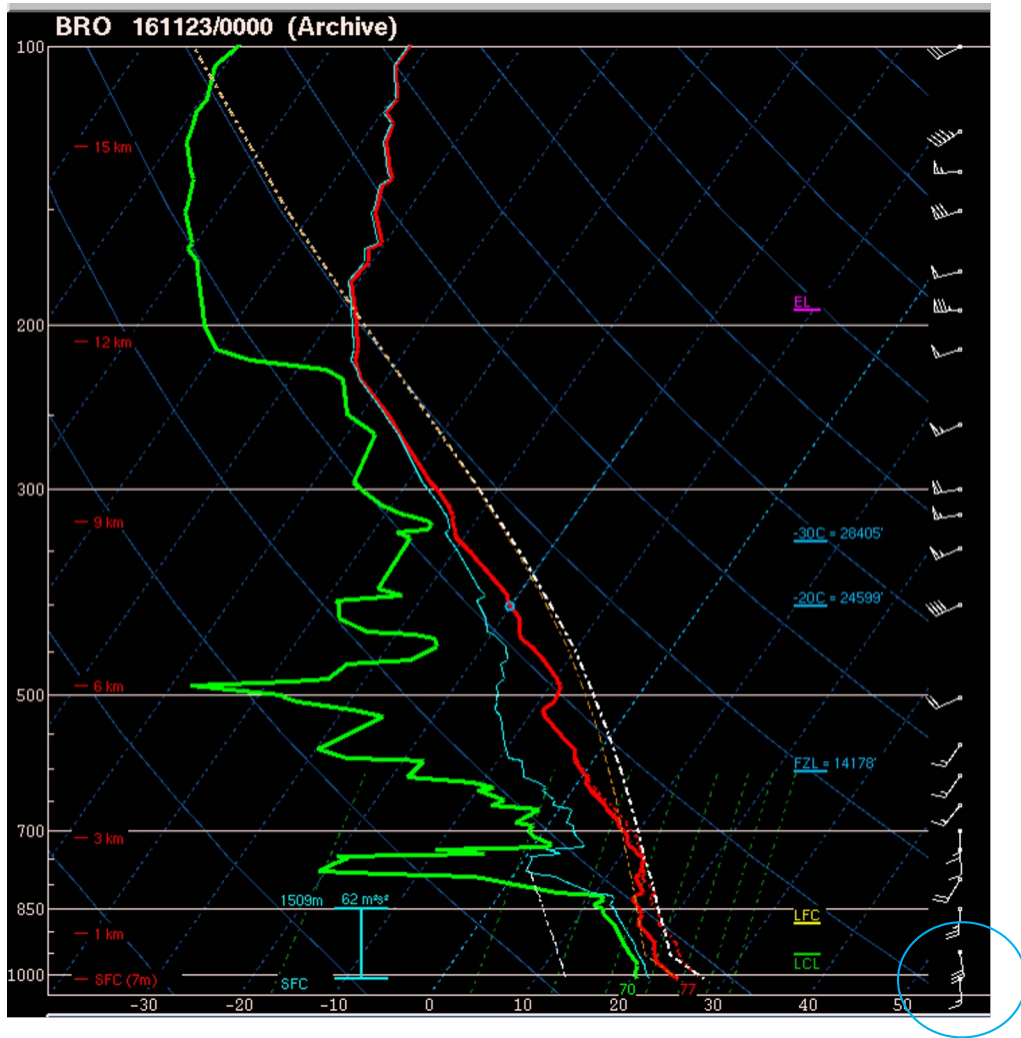
Loop: 2





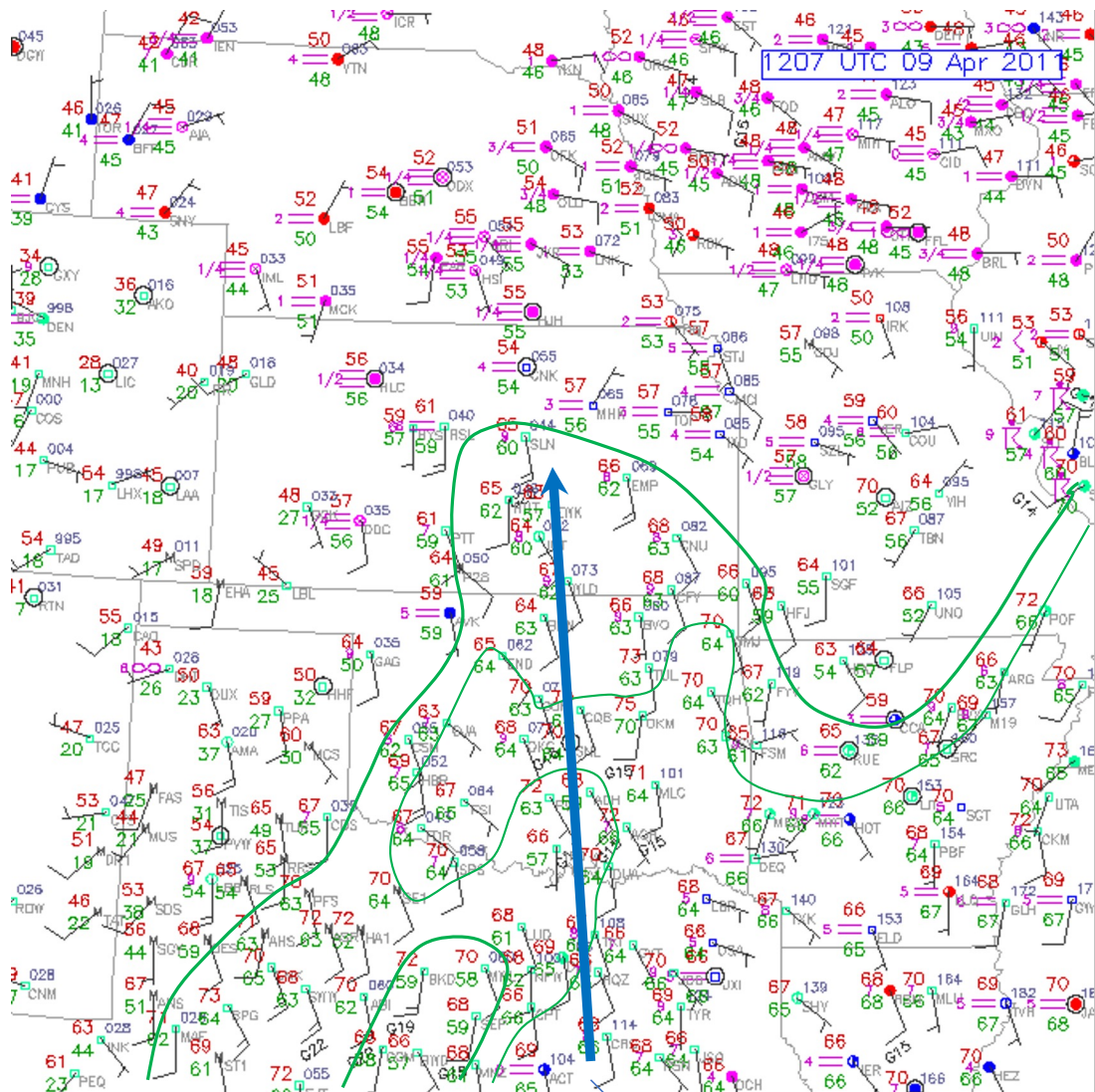


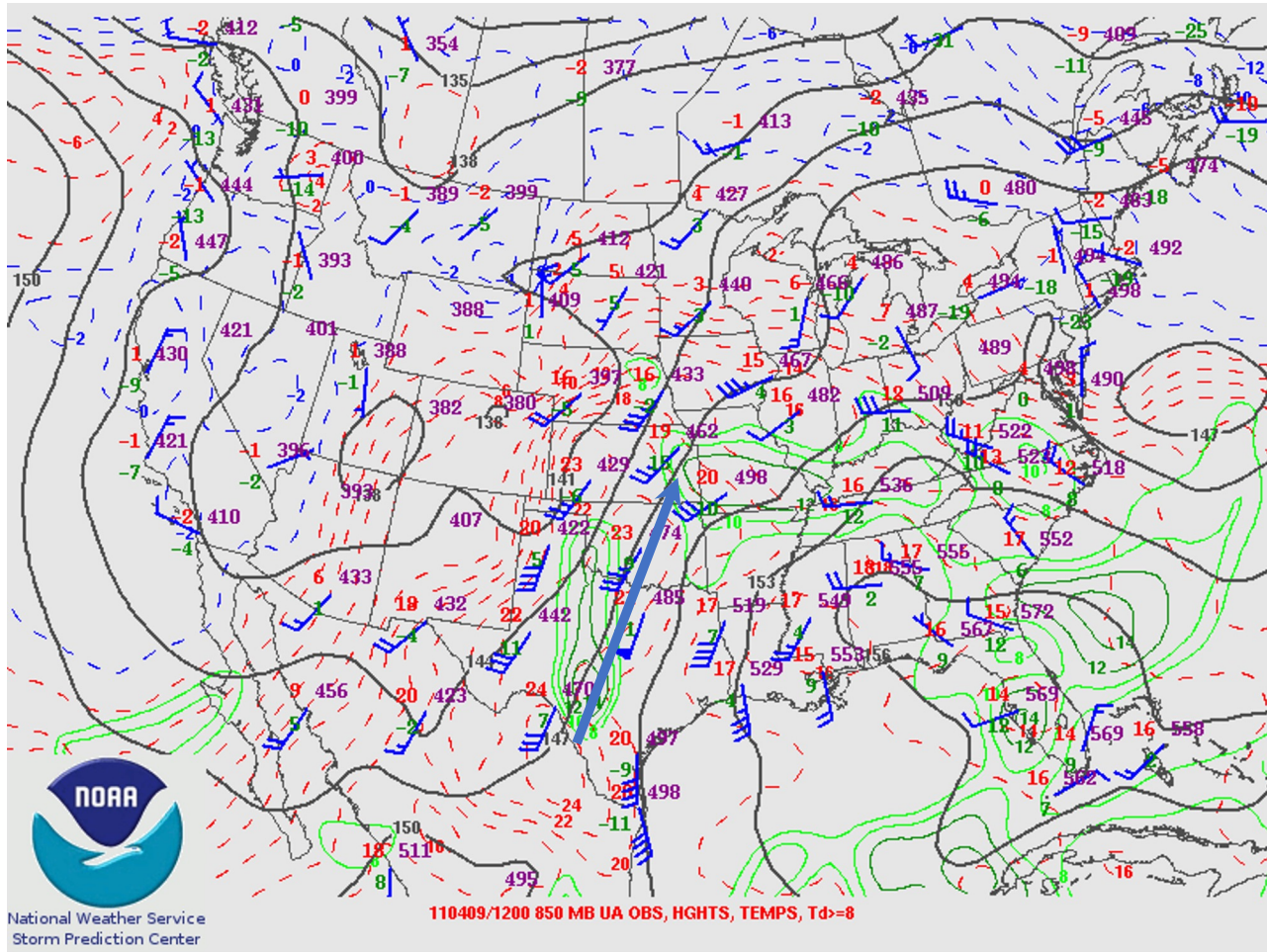




Inland Vertical Mixing

- **Surface moisture can decrease during the daytime when:**
 - Daytime heating/mixing extends deeper than original moist layer depth
 - Moisture advection is not enough to offset mixing
 - Compounded by dry soil, little green vegetation, and ongoing drought
- **Daytime mixing is governed by:**
 - Vertical moisture structure
 - Height and strength of lid/cap
 - Upstream moisture sources
 - Local moisture sources



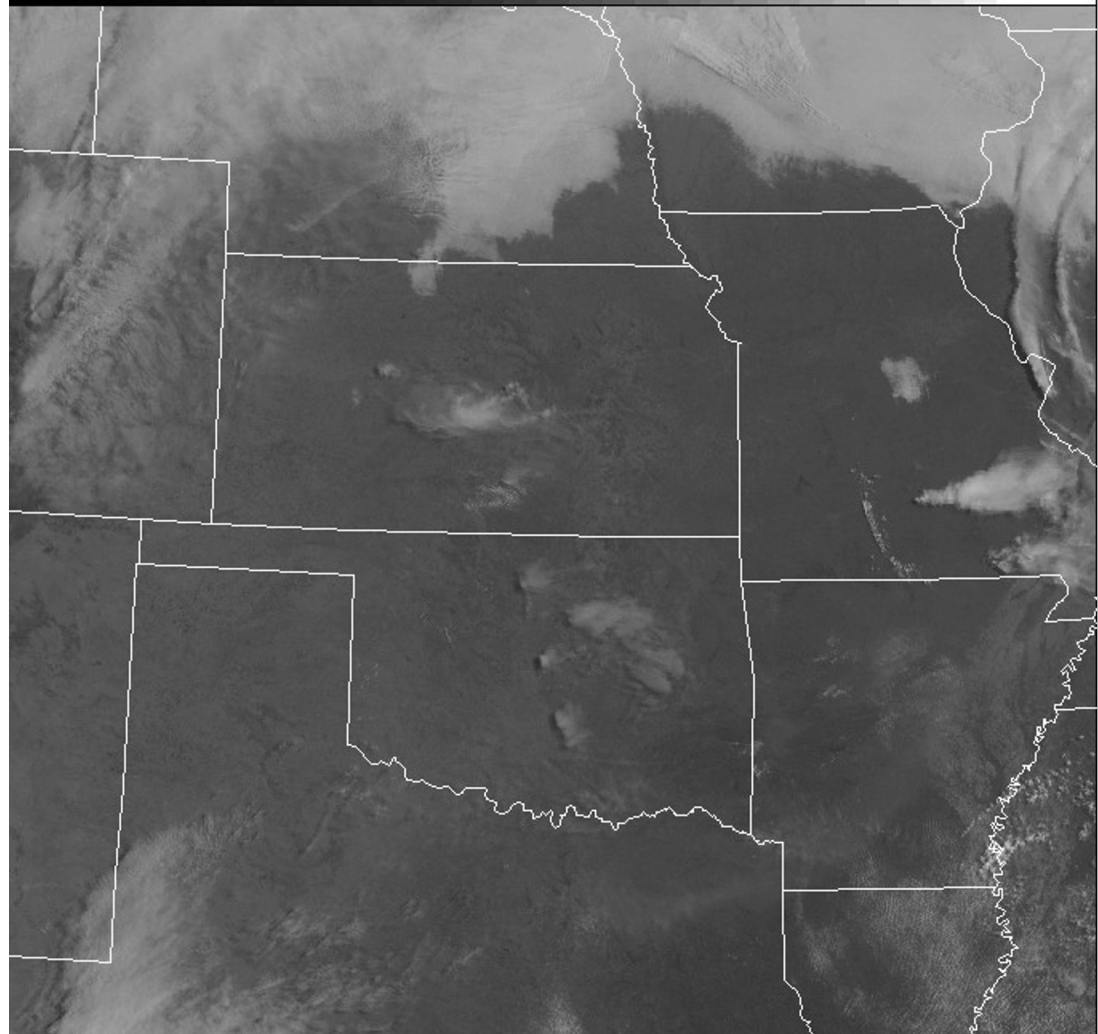


1445 UTC Sat 09 Apr 2011

Visible Satellite

<http://adds.aviationweather.gov>

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U.S. Drought Monitor CONUS

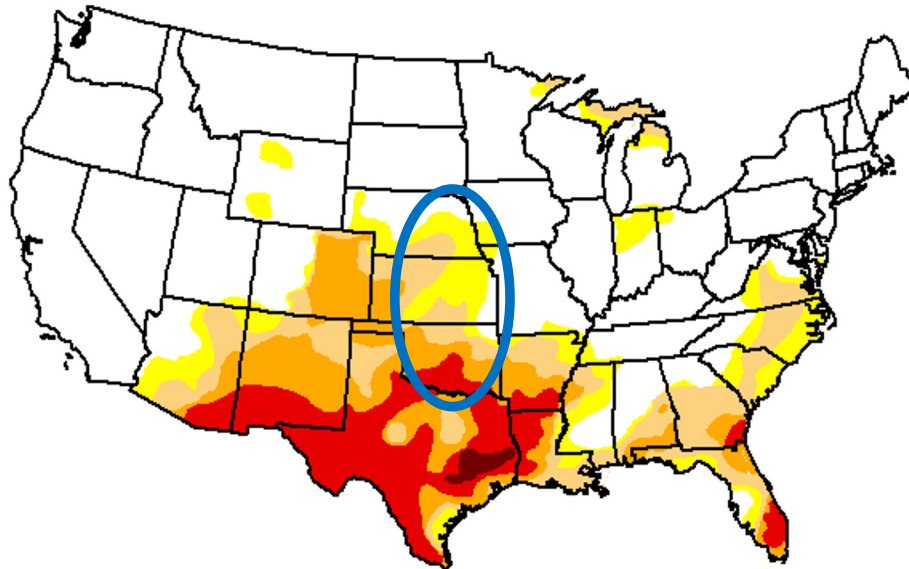
April 5, 2011

(Released Thursday, Apr. 7, 2011)

Valid 7 a.m. EST

Drought Conditions (Percent Area)

	None	D0-D4	D1-D4	D2-D4	D3-D4	D4
Current	62.46	37.54	28.75	18.17	8.75	0.42
Last Week <i>3/29/2011</i>	61.76	38.24	28.14	18.22	6.41	0.00
3 Months Ago <i>1/4/2011</i>	60.50	39.50	21.74	8.50	2.60	0.00
Start of Calendar Year <i>1/4/2011</i>	60.50	39.50	21.74	8.50	2.60	0.00
Start of Water Year <i>9/28/2010</i>	60.05	39.95	13.16	3.09	0.30	0.00
One Year Ago <i>4/6/2010</i>	71.57	28.43	9.10	2.00	0.00	0.00



Intensity:

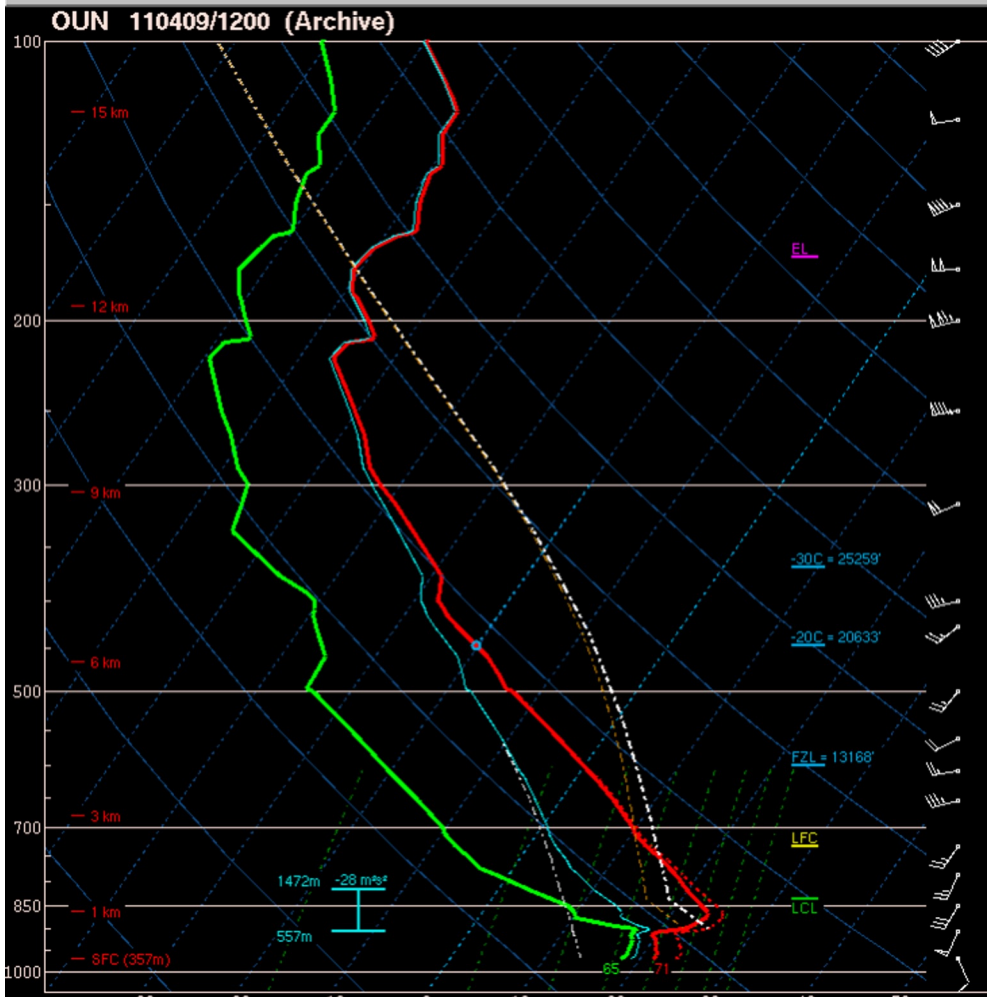
- D0 Abnormally Dry
- D1 Moderate Drought
- D2 Severe Drought
- D3 Extreme Drought
- D4 Exceptional Drought

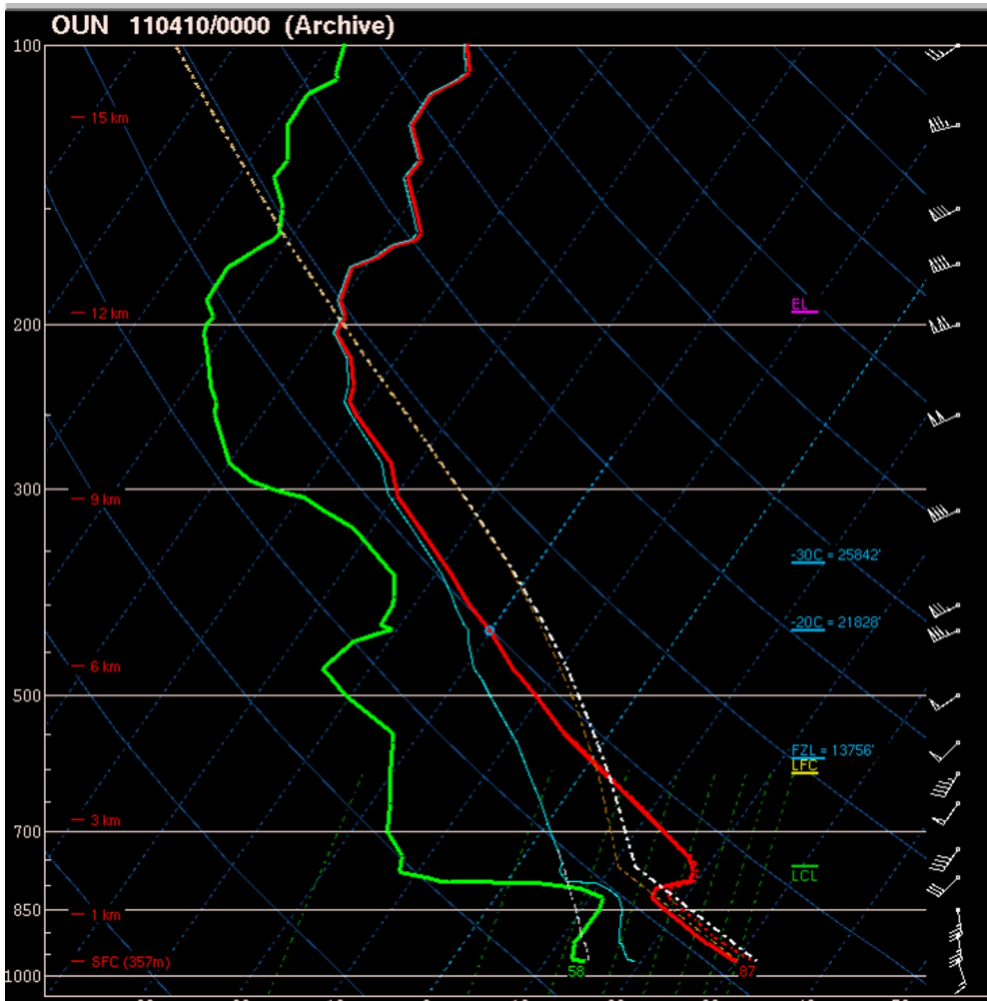
The Drought Monitor focuses on broad-scale conditions. Local conditions may vary. See accompanying text summary for forecast statements.

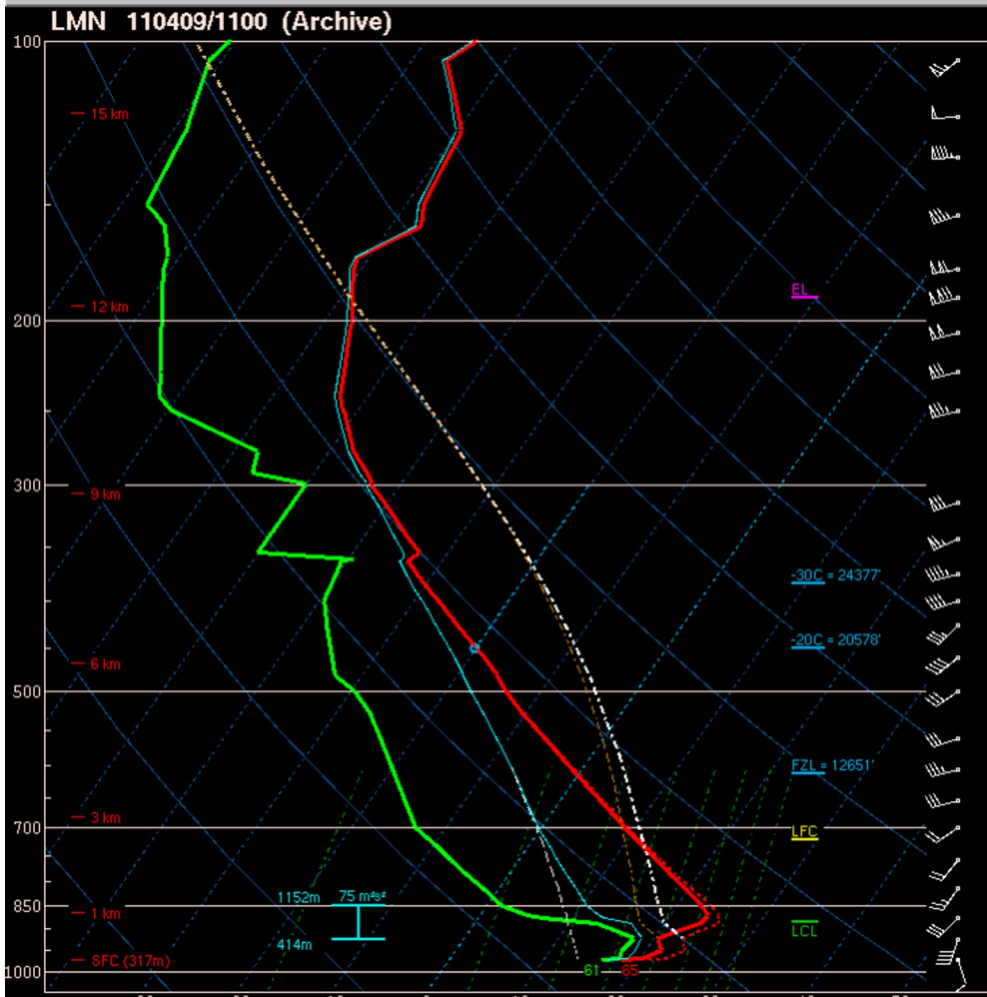
Author(s):

Mark Svoboda
National Drought Mitigation Center

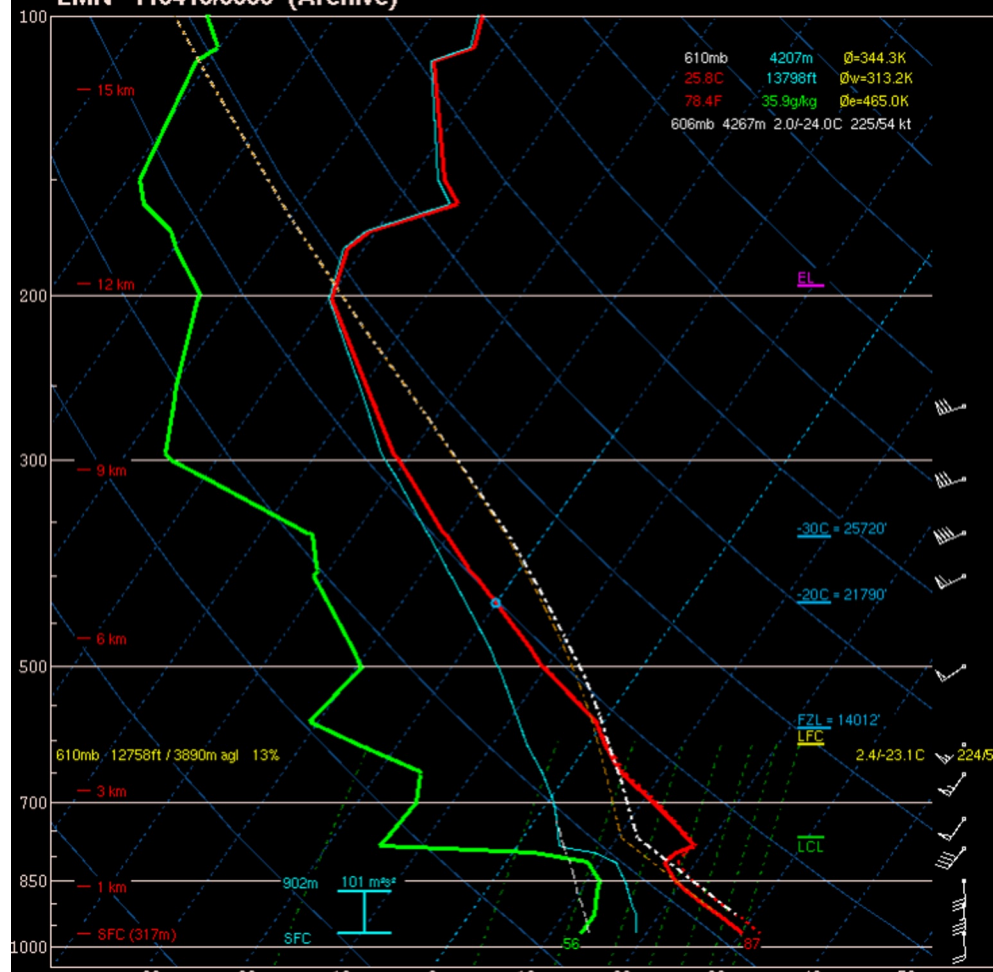


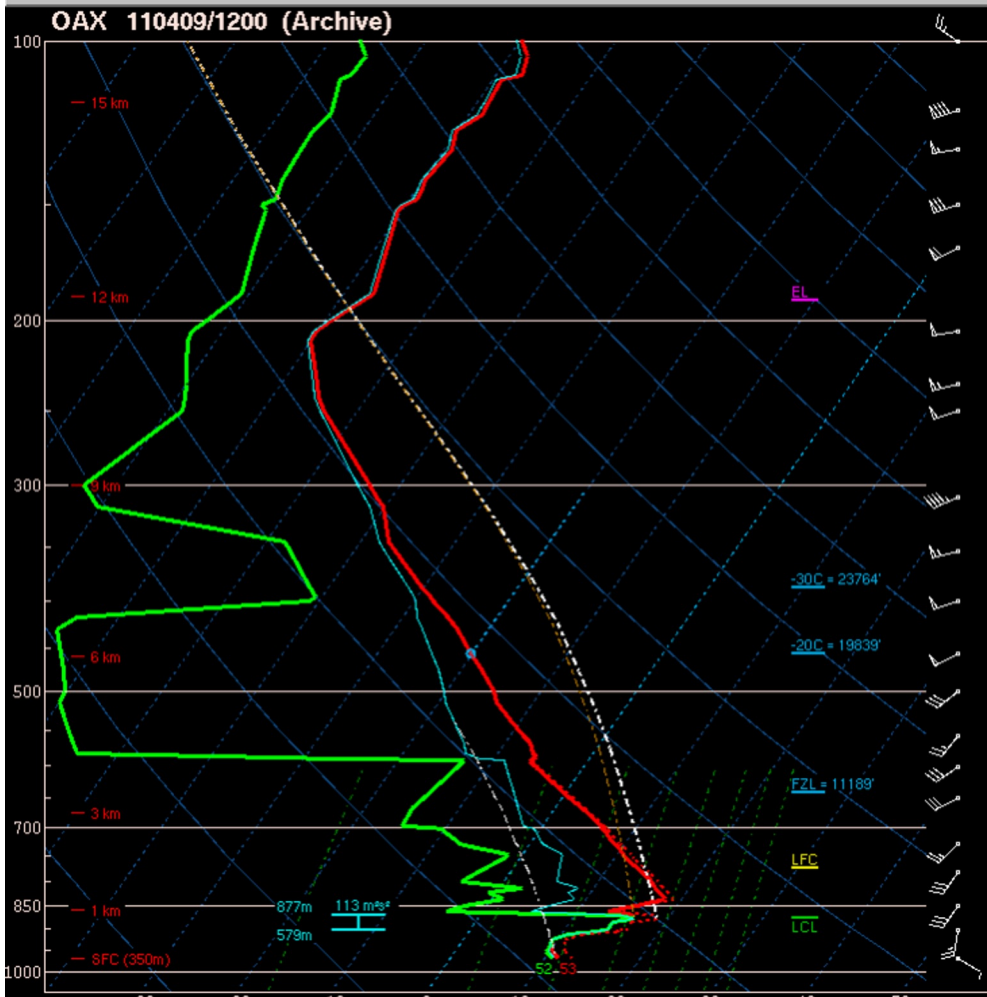


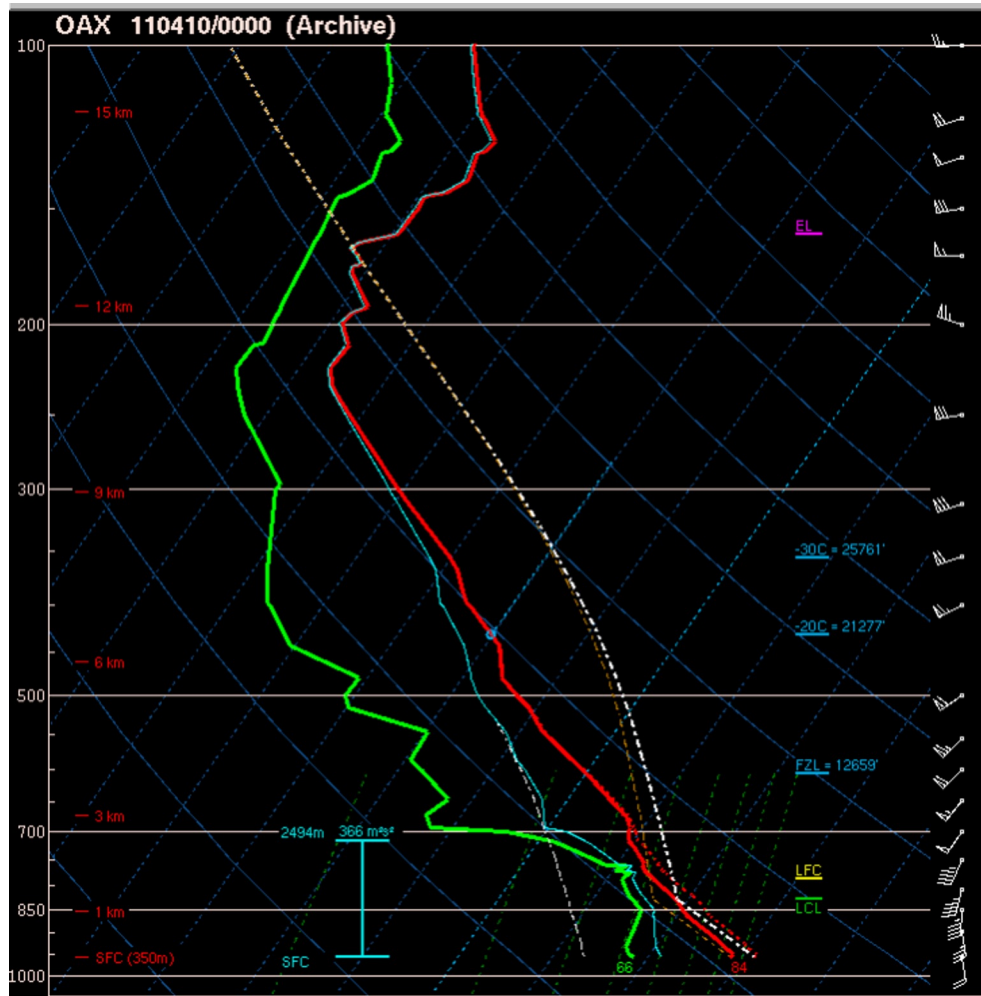




LMN 110410/0000 (Archive)

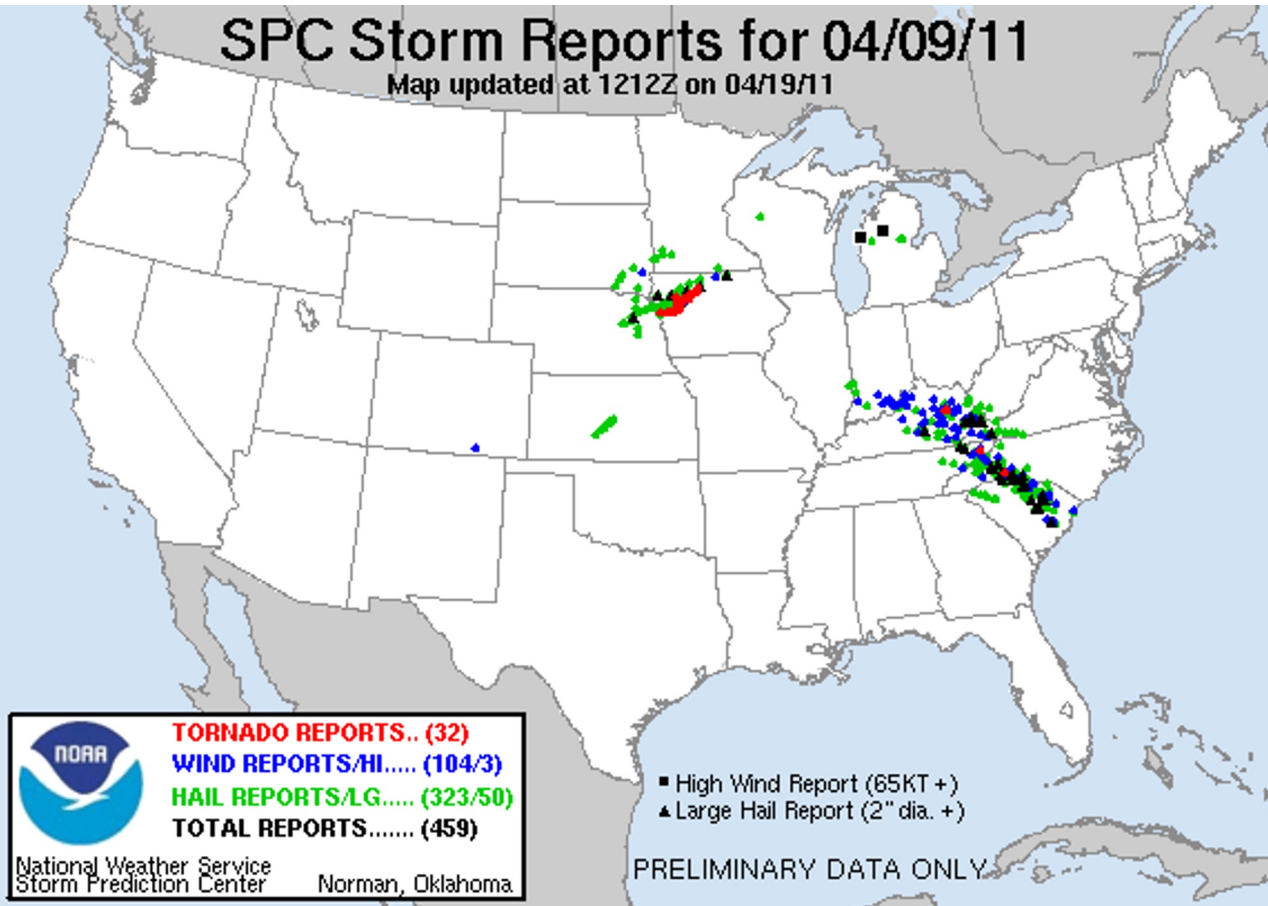






SPC Storm Reports for 04/09/11

Map updated at 1212Z on 04/19/11

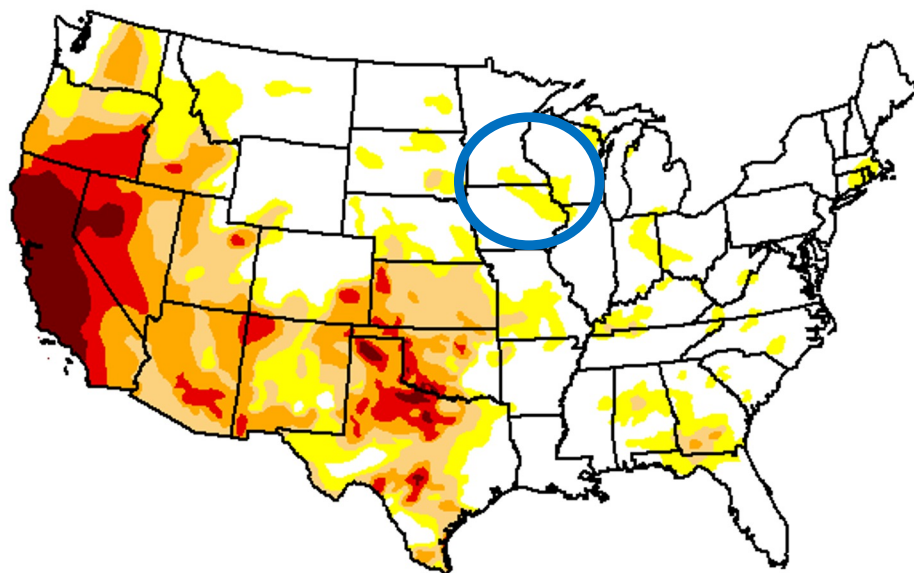


Evapotranspiration

- **Need moist soil and growing vegetation**
- **Plentiful rain previous 1-2 weeks**
- **Maturing crops (wheat, corn, or canola)**
- **Capped boundary layer to trap moisture; relatively weak winds**
- **Almost always a significant return flow contribution (in addition)**

U.S. Drought Monitor CONUS

August 19, 2014
(Released Thursday, Aug. 21, 2014)
Valid 8 a.m. EDT



Drought Conditions (Percent Area)

	None	D0-D4	D1-D4	D2-D4	D3-D4	D4
Current	51.93	48.07	33.56	21.62	10.12	3.80
Last Week <i>8/12/2014</i>	52.81	47.19	33.61	22.03	10.24	3.79
3 Months Ago <i>5/20/2014</i>	52.36	47.64	38.12	28.30	14.47	4.99
Start of Calendar Year <i>12/31/2013</i>	48.24	51.76	30.95	16.67	3.96	0.37
Start of Water Year <i>10/1/2013</i>	39.57	60.43	41.21	20.70	3.06	0.29
One Year Ago <i>8/20/2013</i>	40.02	59.98	45.61	32.23	10.54	1.32

Intensity:

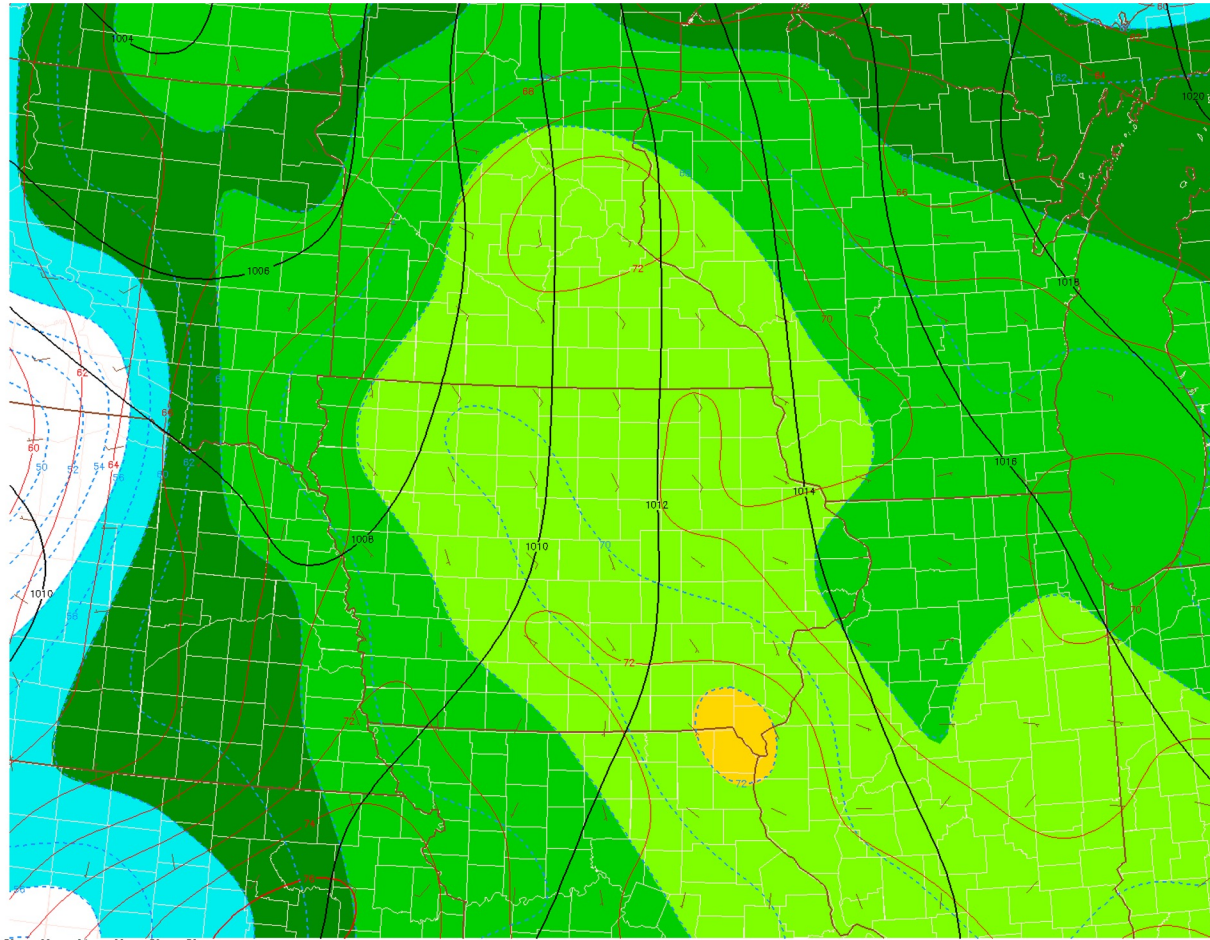
- D0 Abnormally Dry
- D1 Moderate Drought
- D2 Severe Drought
- D3 Extreme Drought
- D4 Exceptional Drought

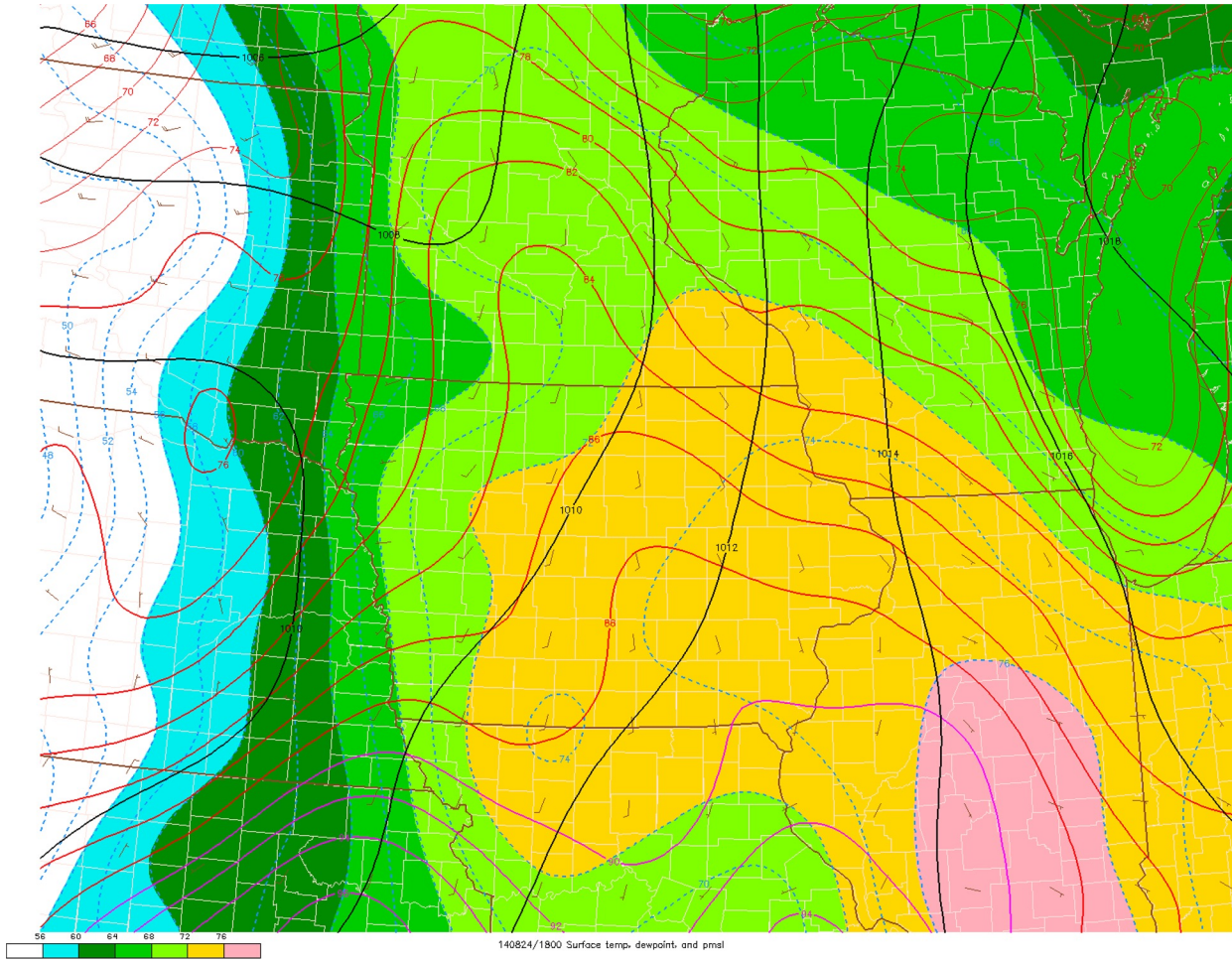
The Drought Monitor focuses on broad-scale conditions. Local conditions may vary. See accompanying text summary for forecast statements.

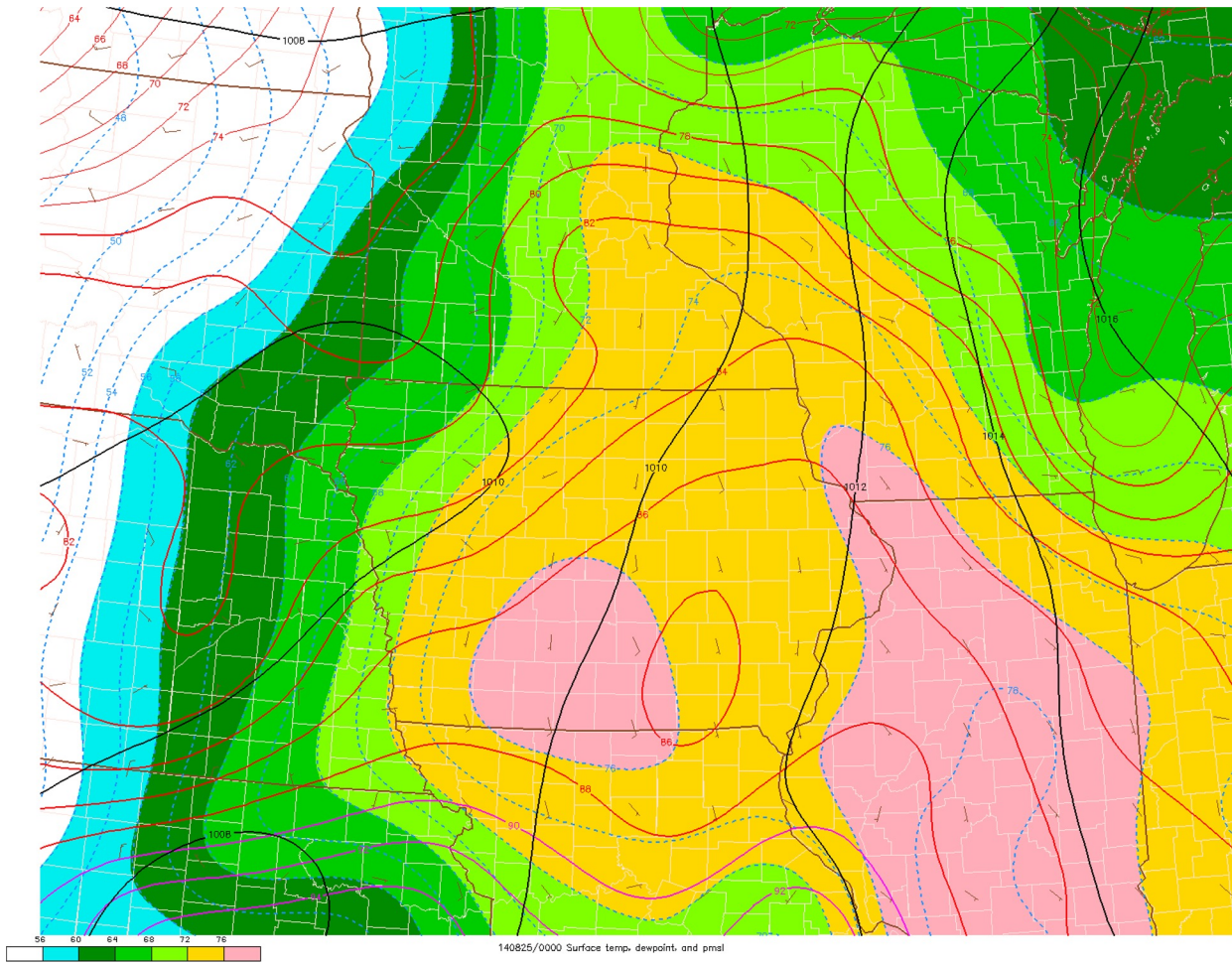
Author(s):

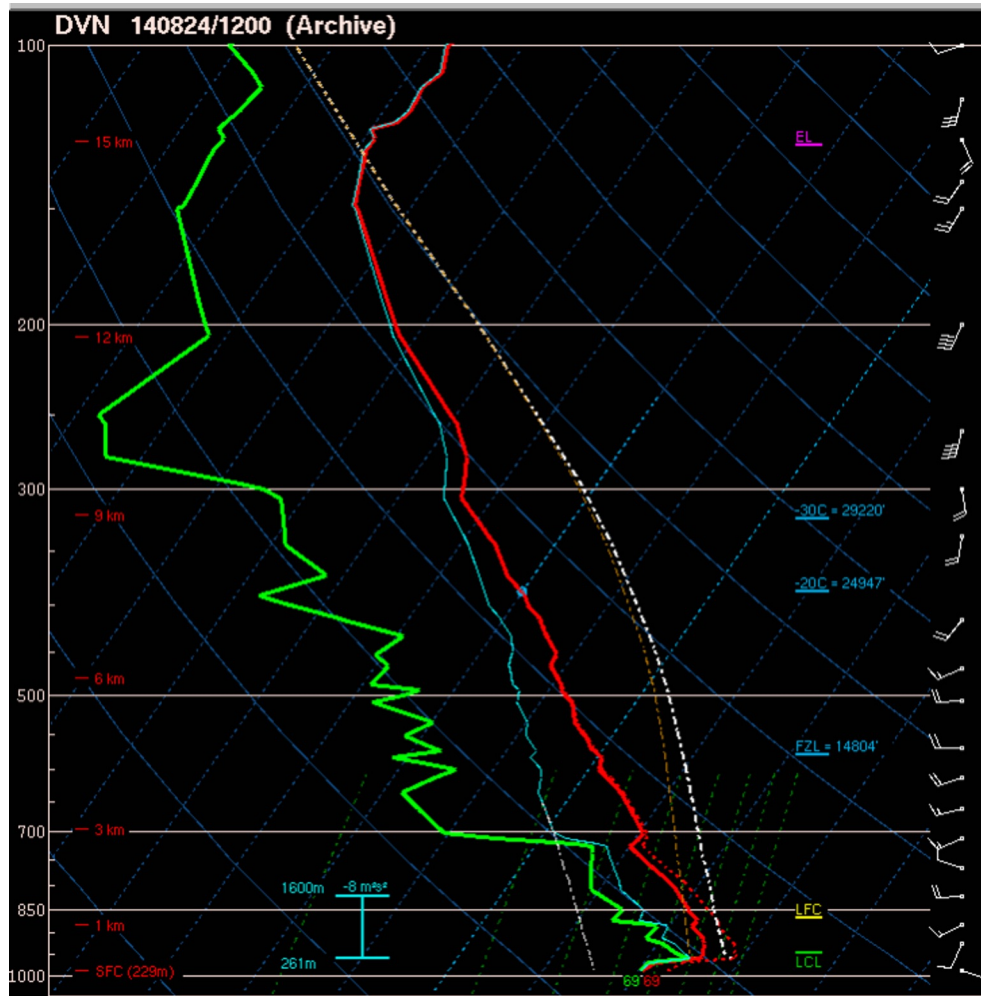
Richard Tinker
CPC/NOAA/NWS/NCEP

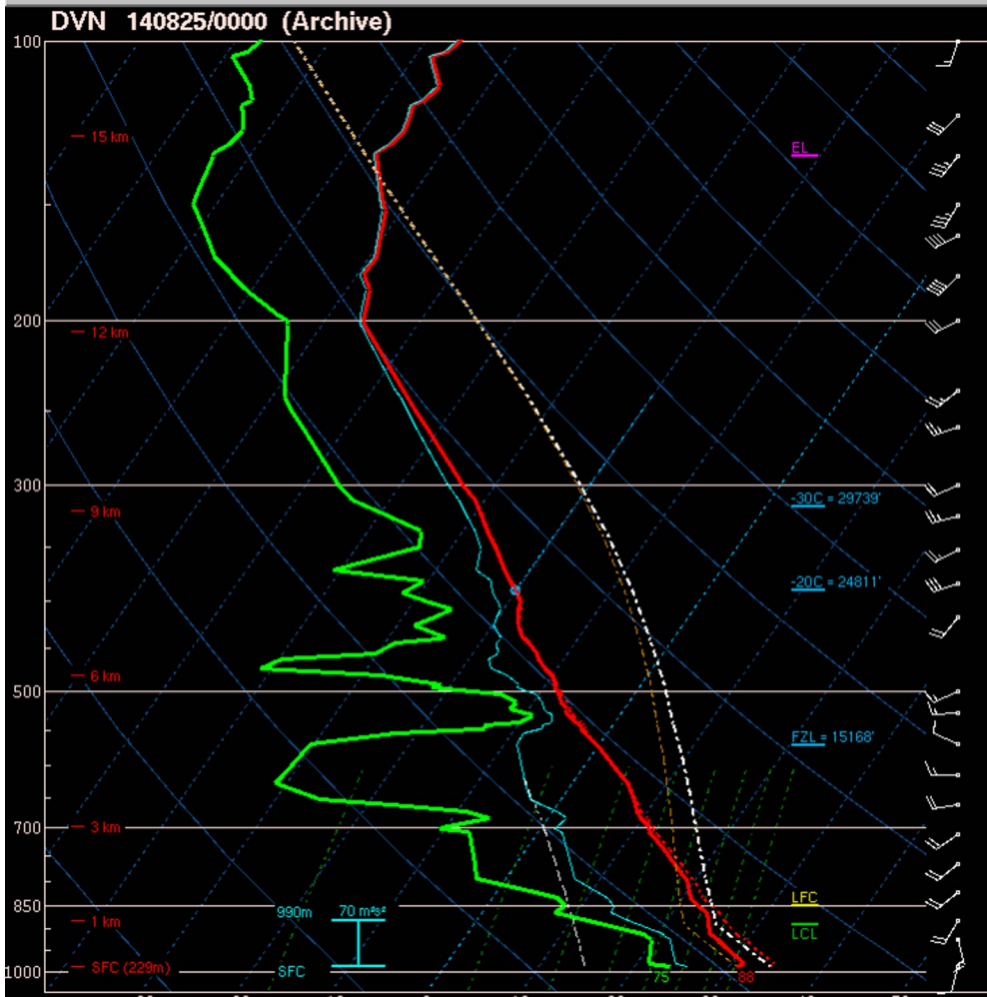


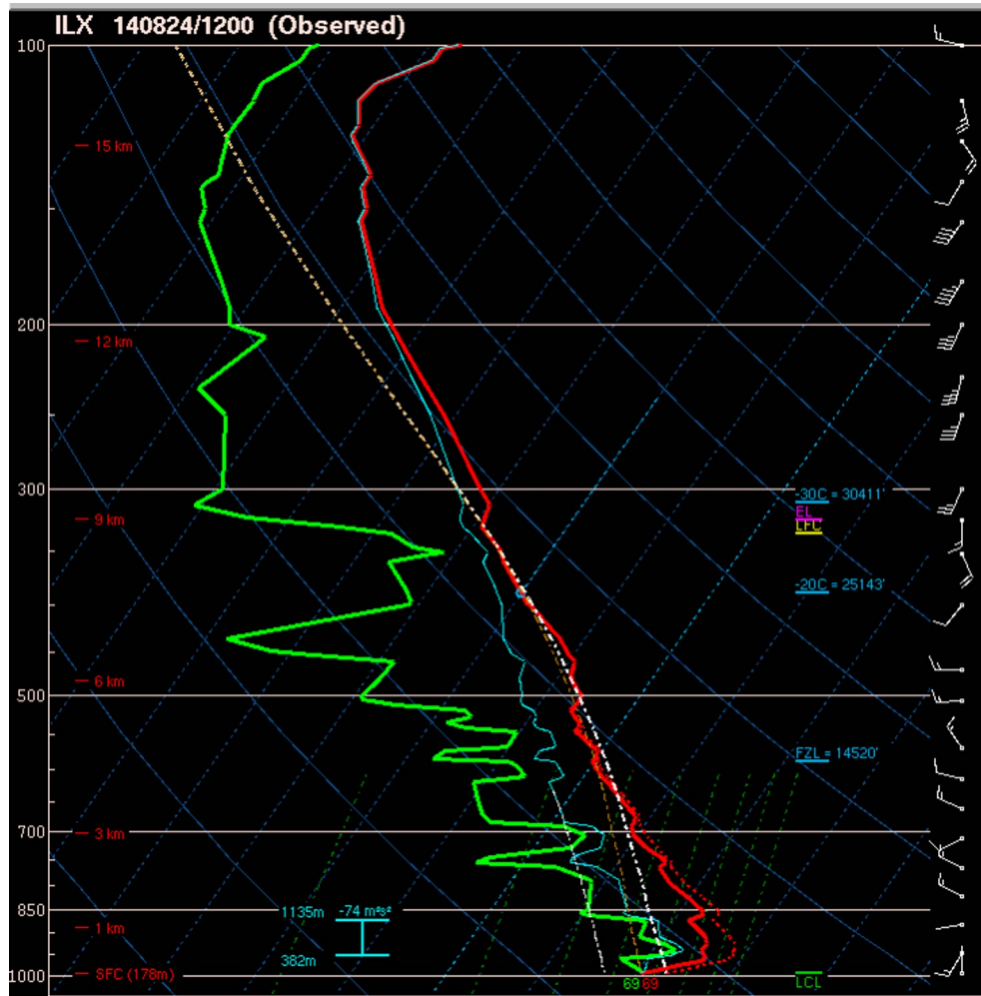


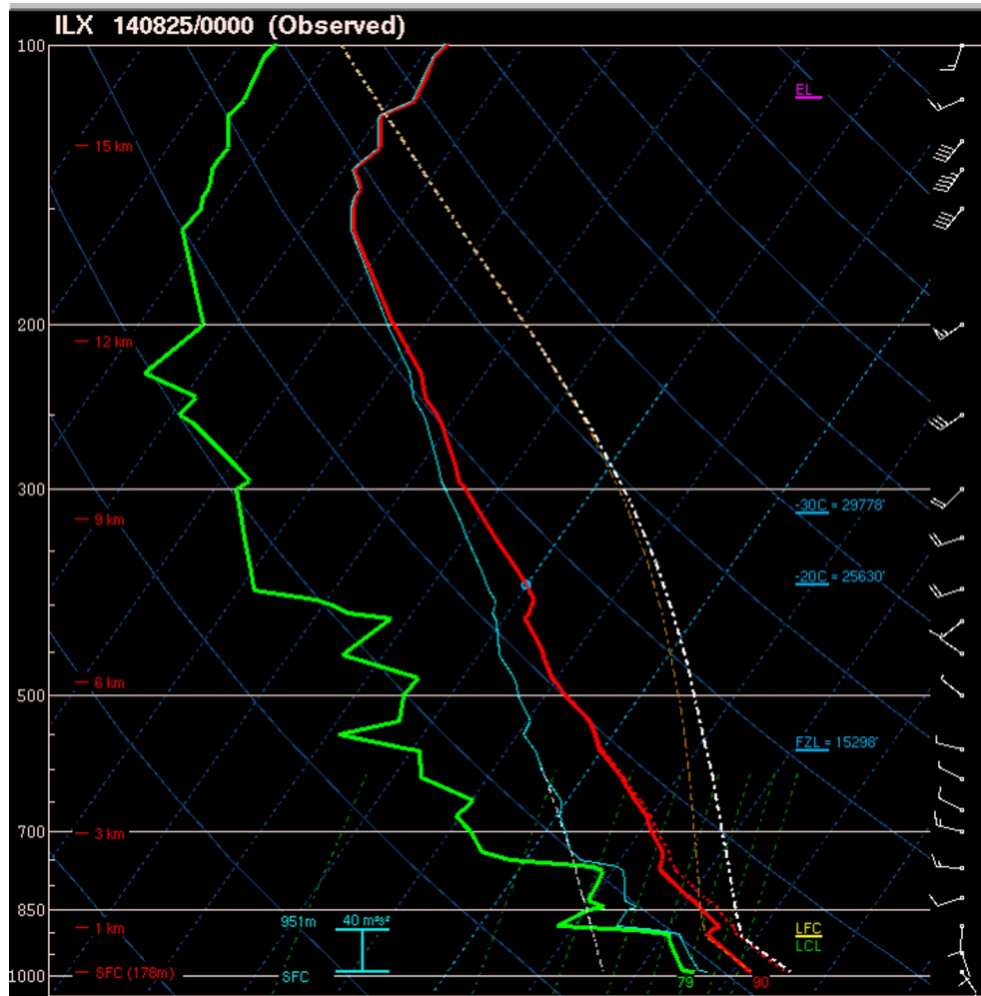






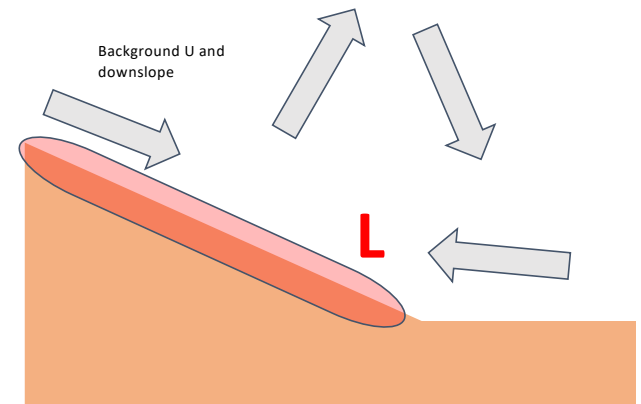
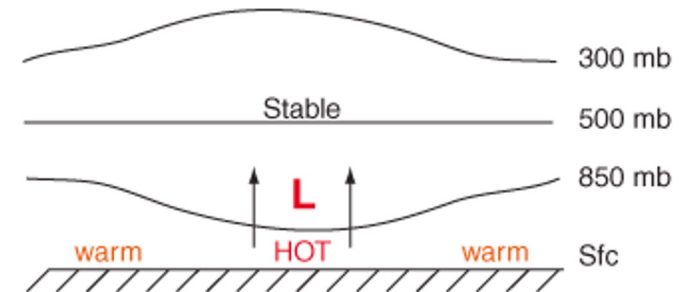


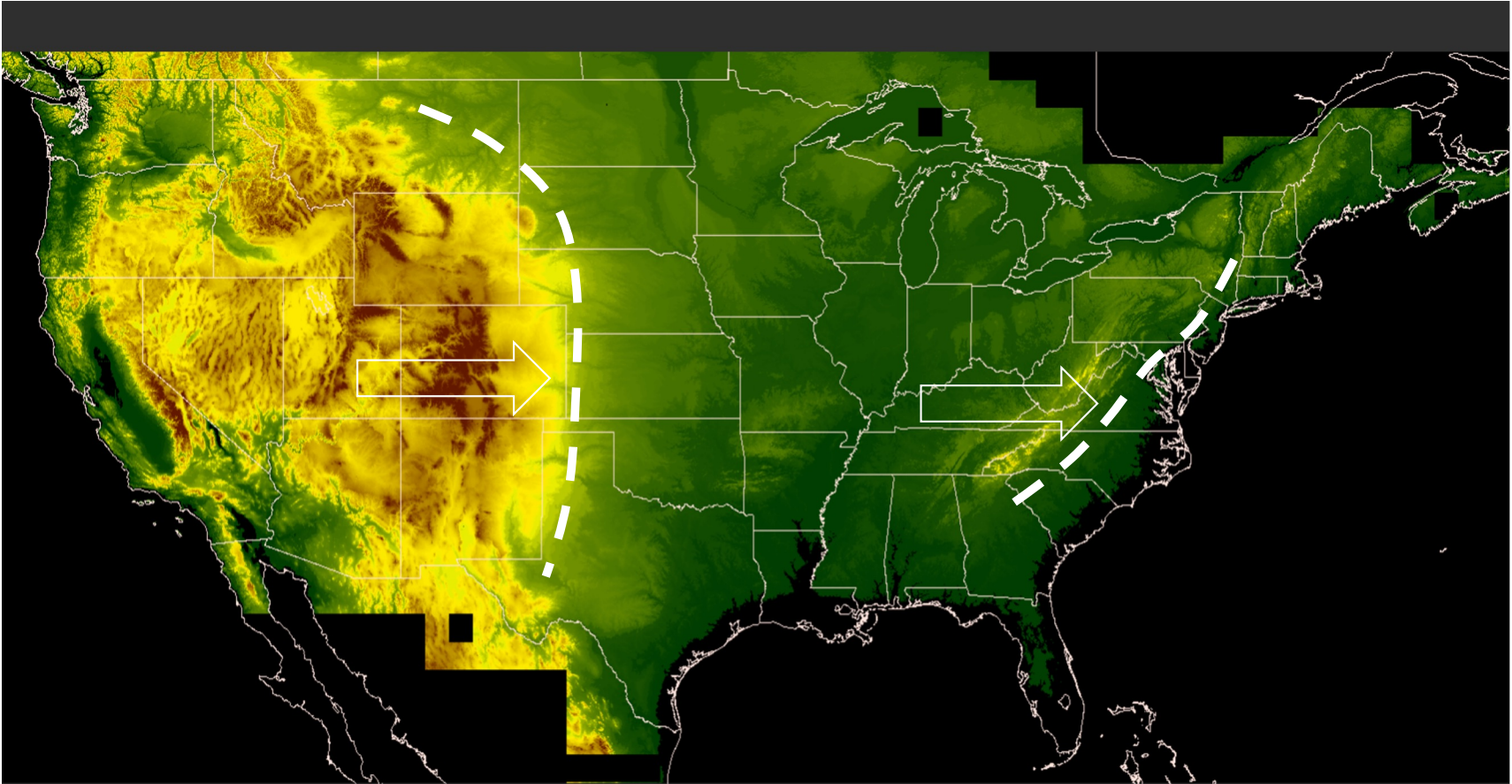




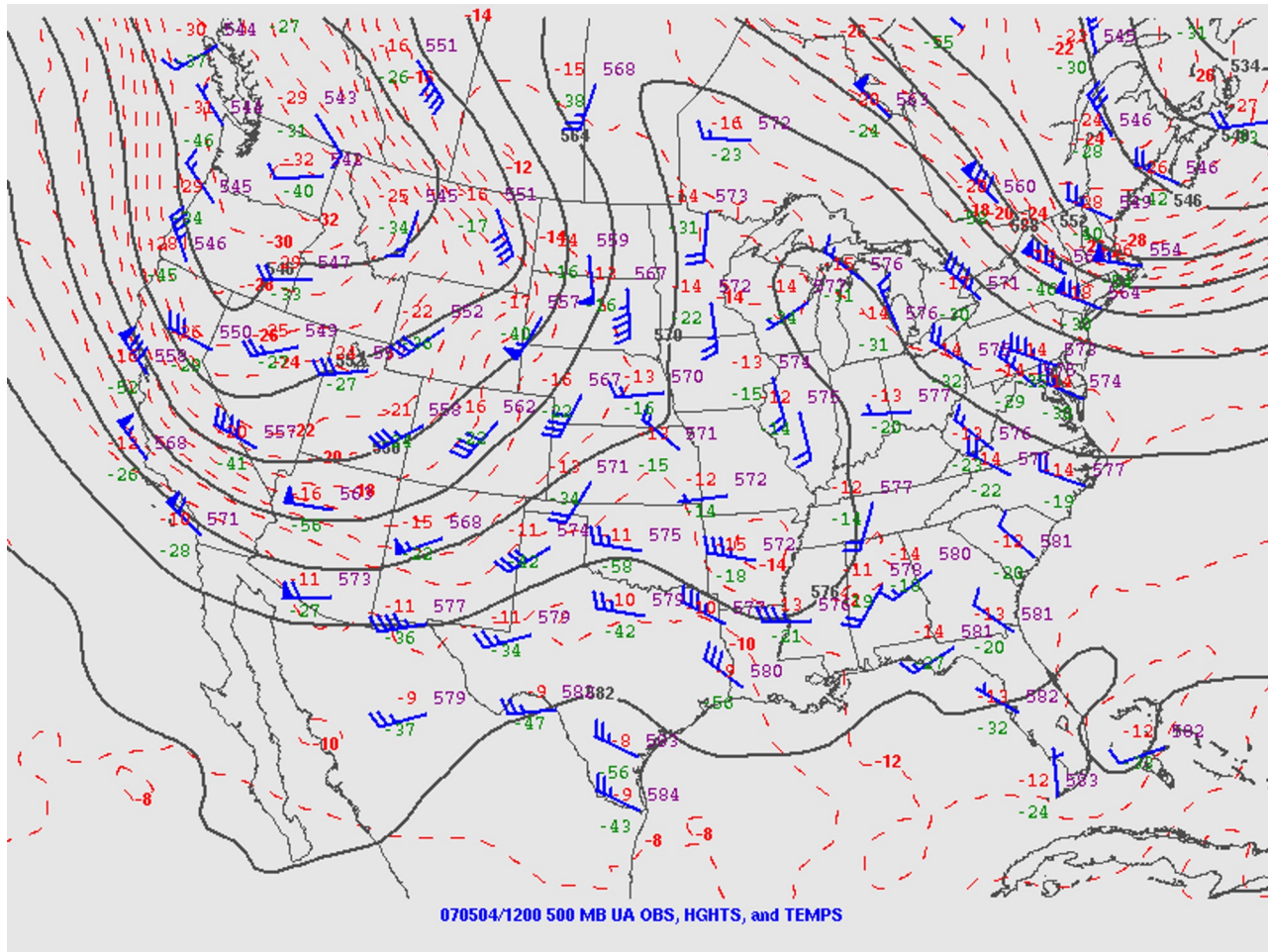
Lee Cyclogenesis

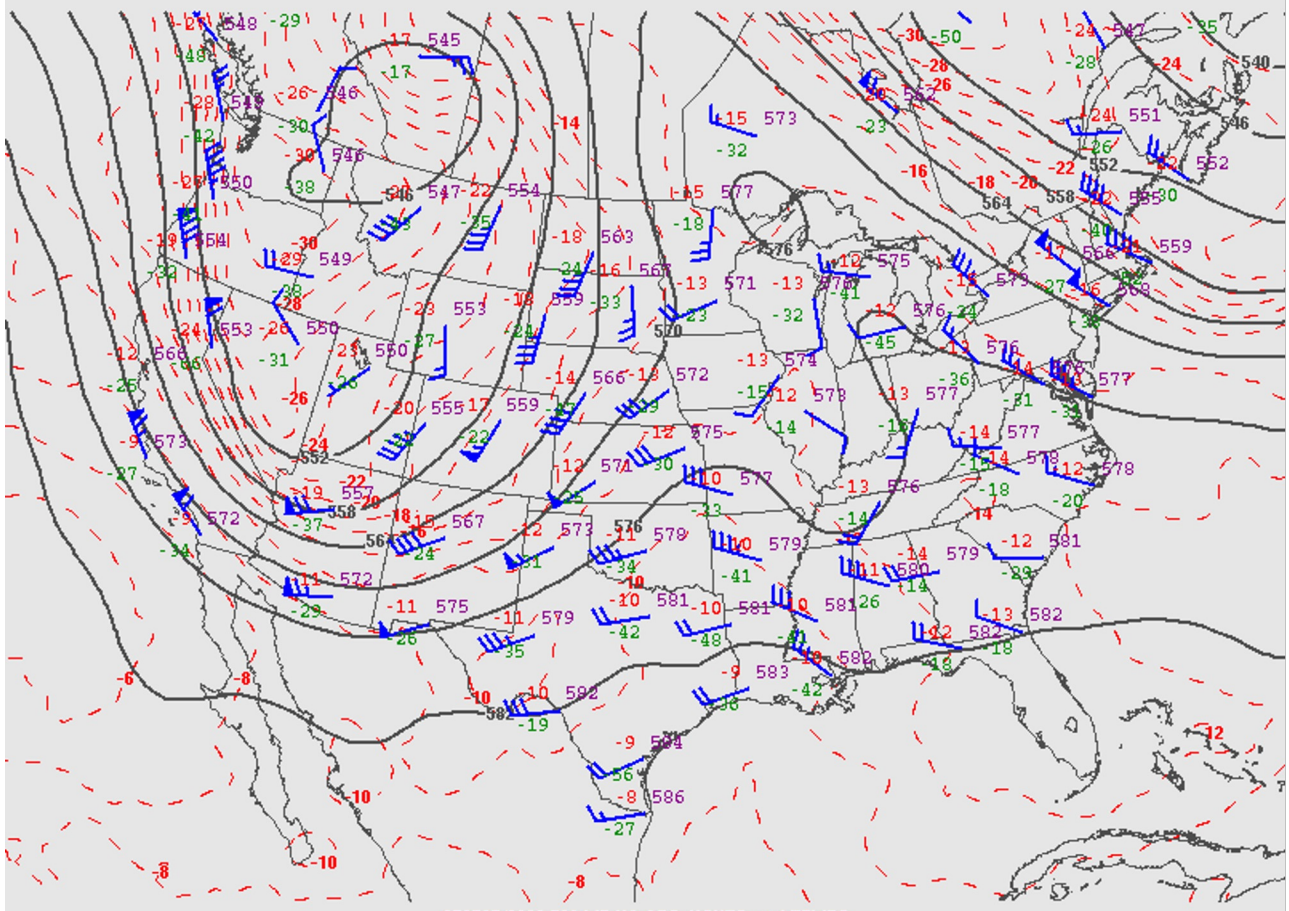
- **Mid-level flow crosses high terrain:**
 - Subsidence and warming E of mountains
 - Stronger diabatic heating across slope results in rising motion from induced katabatic flow. (Frontogenesis)
 - Warming of column and rising motion leads to “warm core” low formation
 - Strongest pressure falls with strongest flow crosses highest terrain
- **The lee trough/cyclone deepens before arrival of strongest Q-G forcing for ascent: (More QG Later)**
 - Head start on differential advection, “loaded gun” sounding, and veering winds with height



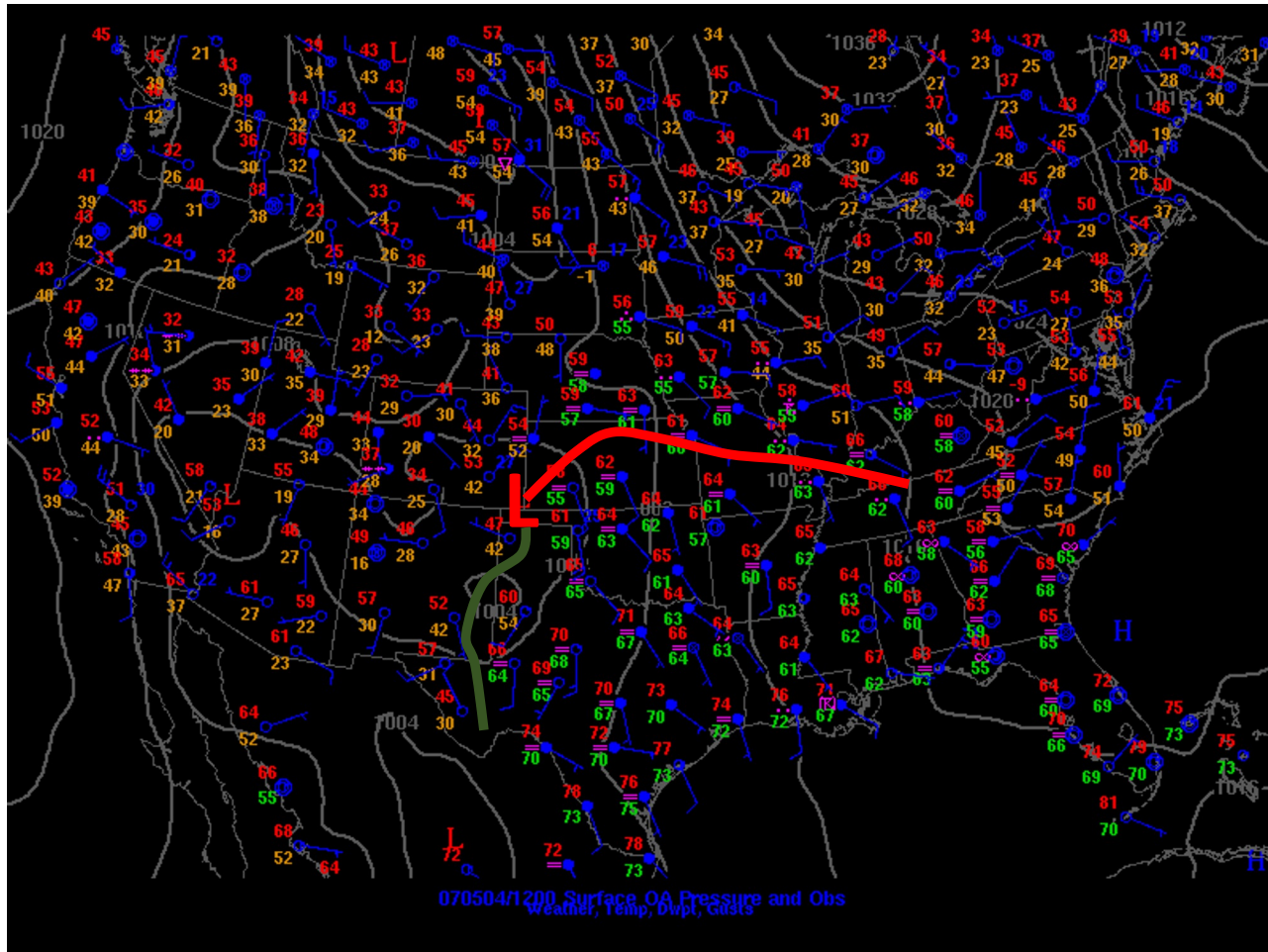


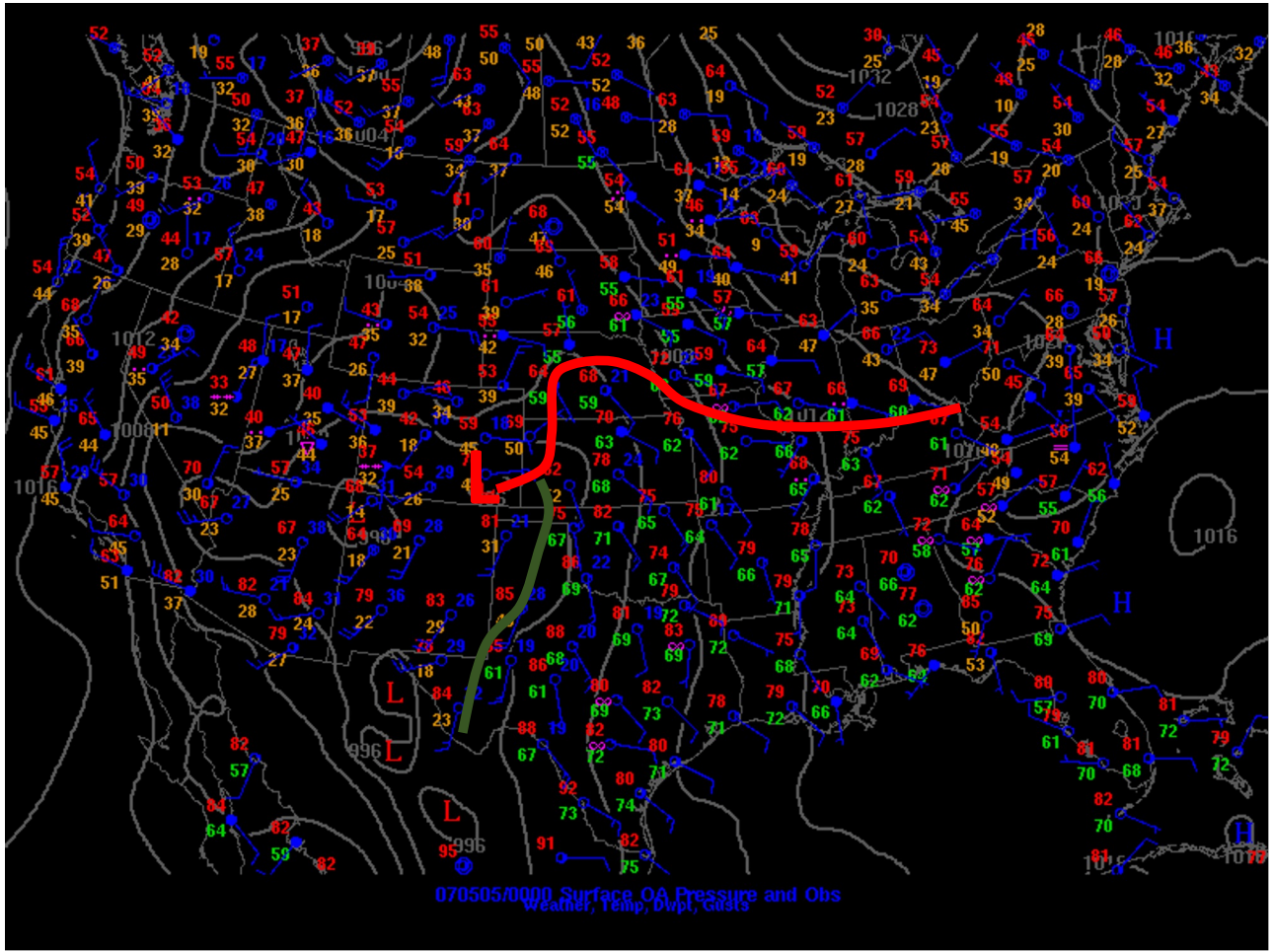
951027/1200 GOES8 SMD





070505/0000 500 MB UA OBS, HGHTS, and TEMPS

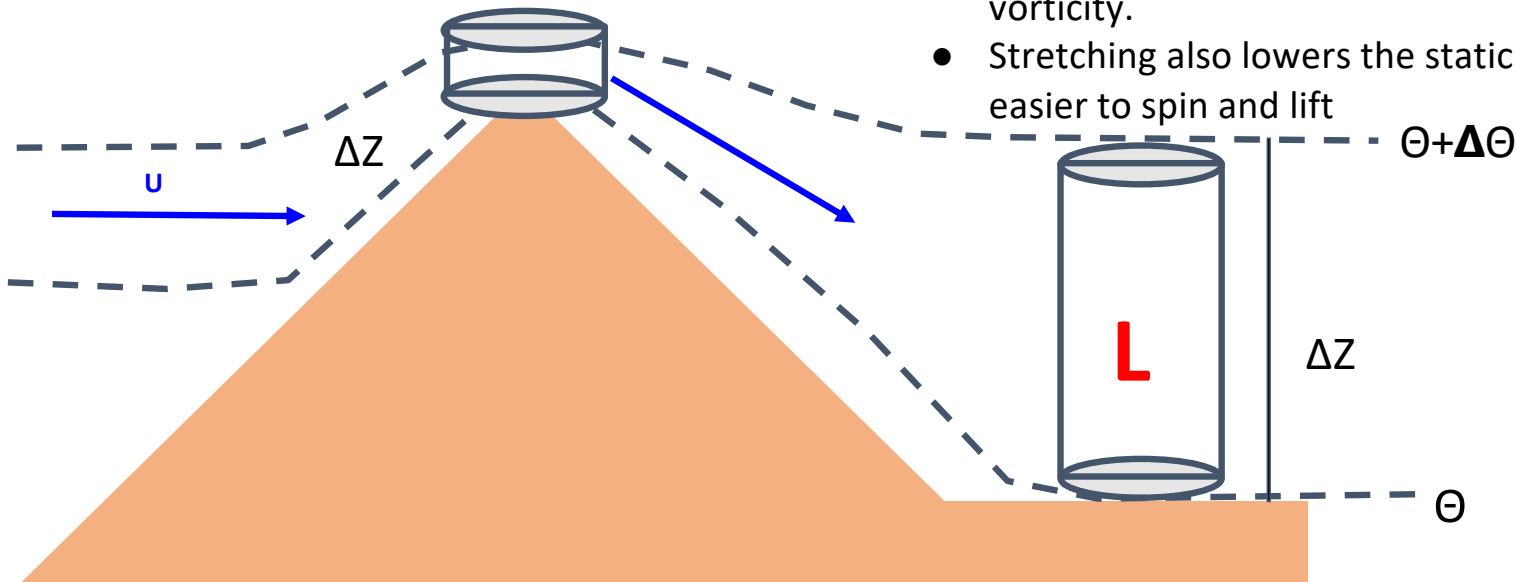




Lee Cyclogenesis PV Framework

$$Pv = \frac{\zeta_{\theta} + f}{\Delta Z} = \text{const}$$

- Conservation of Potential Vorticity
 - Assumptions:
 - Constant density/stable
 - Adiabatic
 - Frictionless
- Stretching of the column vertically requires an increase in relative vorticity between the theta surfaces inducing positive vorticity.
- Stretching also lowers the static stability of the atmosphere. easier to spin and lift

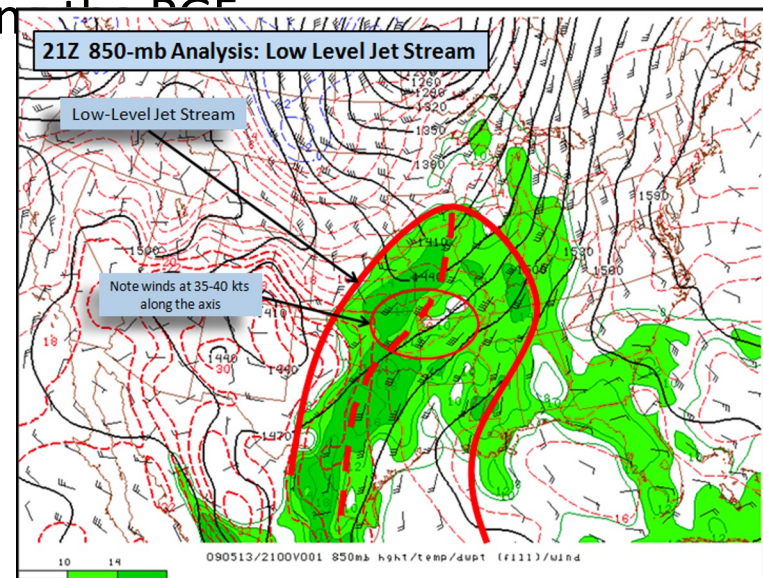
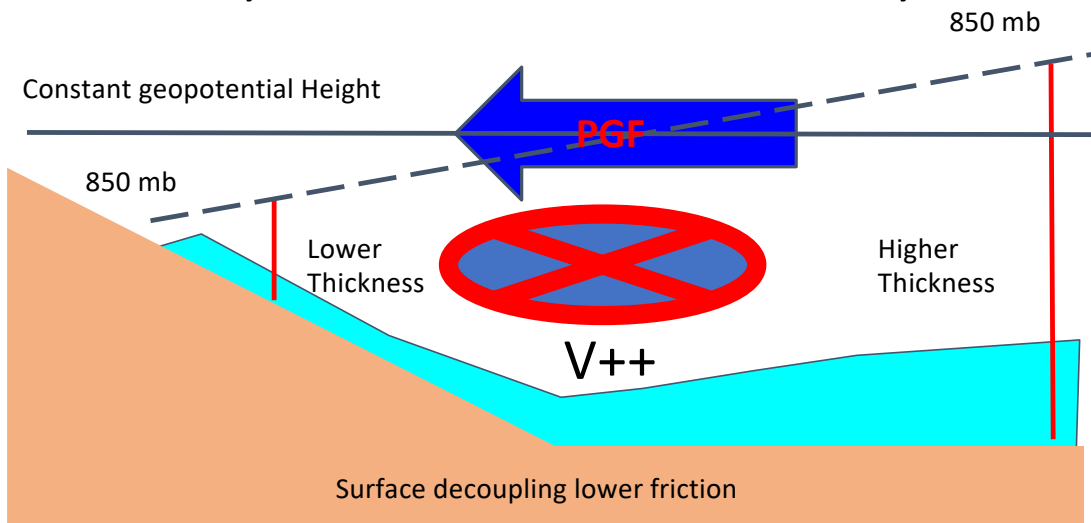


Nocturnal Low-Level Jet (LLJ)

- Plains often see early nocturnal low-level jet (LLJ) ramp up with lee cyclogenesis
- Related to two primary factors:
 - Boundary-layer decoupling and loss of surface friction (“inertial oscillation”)
 - Diurnal temperature variations over sloped terrain (thermal wind)
- Part of the process that can favor late evening/early overnight tornadoes (with favorable moisture/CAPE)

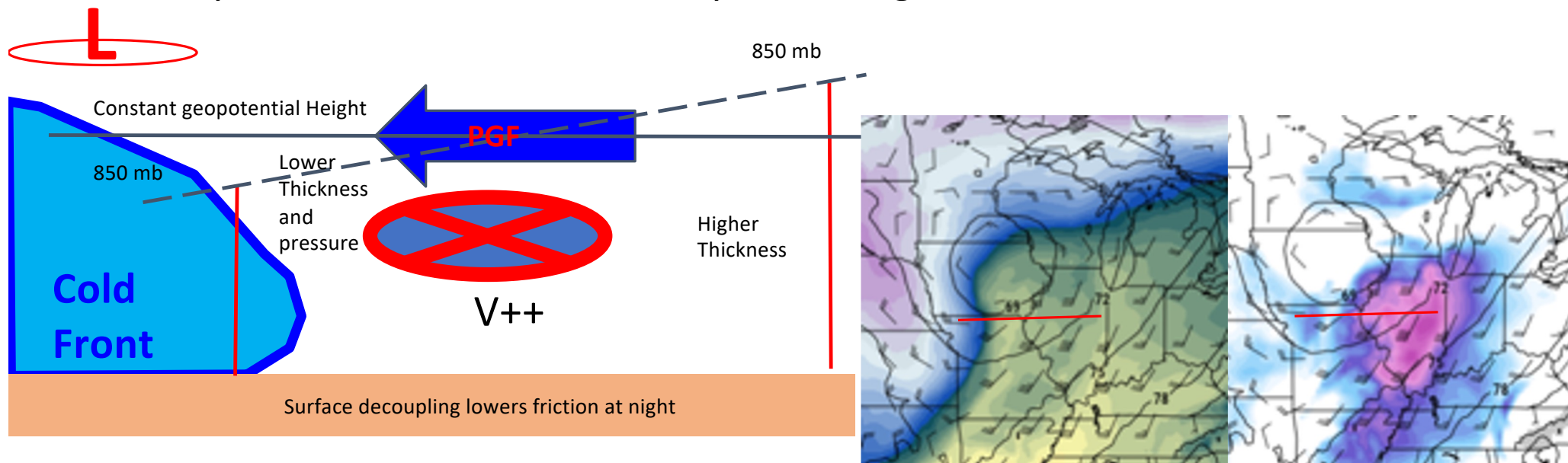
Nocturnal Low-Level Jet (LLJ) Formation

- LLJs form in response to difference in terrain heights, heat fluxes and thickness across pressure surfaces that induces a horizontal Pressure Gradient Force.
- Coriolis forcing turns the momentum northward resulting in an enhanced southerly flow.
- Surface decoupling favors lower friction and acceleration of the jet above the surface
- Lee Cyclones can also enhance it by increasing the PGF

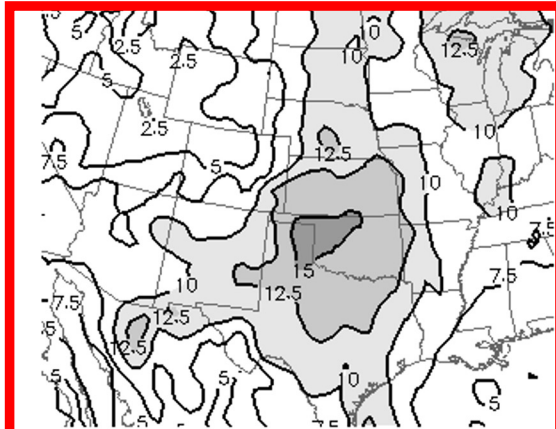


Cyclone induced Low-Level Jet (LLJ) Formation

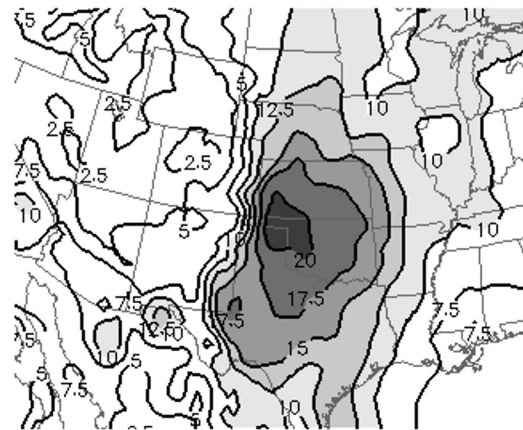
- LLJs form in response to difference in heat fluxes and thickness across pressure surfaces that induces a horizontal Pressure Gradient Force.
- Coriolis forcing turns the momentum northward resulting in an enhanced southerly flow.
- Surface decoupling favors lower friction and acceleration of the jet above the surface
- Lee Cyclones can also enhance it by increasing the PGF.



EF2+ supercell events

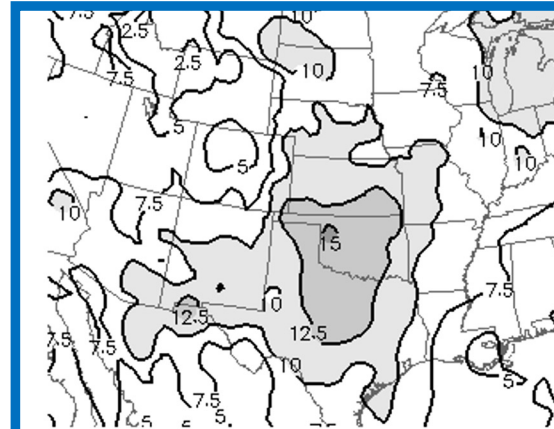


850 hPa isotachs at 0000 UTC (m s⁻¹ SIGTOR)

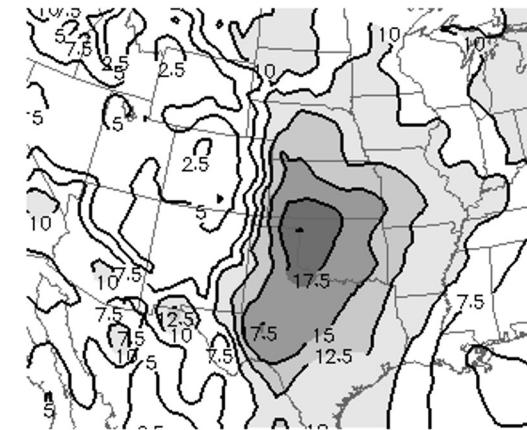


850 hPa isotachs at 0300 UTC (m s⁻¹ SIGTOR)

Nontornadic supercell events



850 hPa isotachs at 0000 UTC (m s⁻¹ NONTOR)



850 hPa isotachs at 0300 UTC (m s⁻¹ NONTOR)

Lee Cyclogenesis Summary

- Lee cyclogenesis occurs where upper-level flow is perpendicular to terrain features (westerly flow for Rockies)
- Lee cyclogenesis more robust with lower static stability
- Lee cyclone helps drive low-level moisture return
- Nocturnal low-level jet associated with lee cyclone is driven by inertial oscillation and thermal wind \Rightarrow increased low-level shear and tornado threat in evening and after dark if adequate CAPE