

2D Boussinesq Equations

The equation of motion for 2D fluid with Boussinesq approximation:

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + w \frac{\partial u}{\partial z} = -\frac{1}{\rho_0} \frac{\partial p}{\partial x}, \quad (1)$$

$$\frac{\partial w}{\partial t} + u \frac{\partial w}{\partial x} + w \frac{\partial w}{\partial z} = -\frac{1}{\rho_0} \frac{\partial p}{\partial z} + g \frac{\theta'}{\theta_0}, \quad (2)$$

where variables, u , w , p , ρ , θ , are wind components in x and z directions, p is pressure, ρ , θ is potential temperature. ρ is constant base state air density, and θ is constant base state potential temperature. $\theta' = \theta - \theta_0$ is the perturbation potential temperature. The second term on the right hand side of (2) is the buoyancy term.

$$\frac{\partial \theta'}{\partial t} + u \frac{\partial \theta'}{\partial x} + w \frac{\partial \theta'}{\partial z} = 0 \quad (3)$$

is the potential temperature conservation equation.

$\frac{\partial \text{Eq.(1)}}{\partial z} - \frac{\partial \text{Eq.(2)}}{\partial x}$ gives the y-vorticity equation:

$$\frac{\partial \eta}{\partial t} + u \frac{\partial \eta}{\partial x} + w \frac{\partial \eta}{\partial z} = -\frac{g}{\theta_0} \frac{\partial \theta'}{\partial x}, \quad (4)$$

where

$$\eta = \frac{\partial u}{\partial z} - \frac{\partial w}{\partial x}.$$

For this 2D non-divergent flow, the velocity components can be written in term of stream function ψ ,

$$u = \frac{\partial \psi}{\partial z}, w = -\frac{\partial \psi}{\partial x}, \quad (5)$$

so that vorticity

$$\eta = \frac{\partial^2 \psi}{\partial x^2} + \frac{\partial^2 \psi}{\partial z^2} = \nabla^2 \psi, \quad (6)$$

or $\nabla^2 \psi = \eta, \quad (7)$

which is an elliptic equation which can be solved for ψ when η is known. In this term project, you are to use equations (3), (4) and (7) to construct a 2D model where prognostic state variables are η and θ' , and ψ is a diagnostic state variable.

Eq. (3) and (4) are first integrated forward to obtain η and θ' at the future time level, then elliptic Eq.(7) is solved to obtain ψ . u and w are then obtained from (5). One time step is then complete.

Term Project Assignment – Part 1: Solve for stream function and velocity given vorticity distribution. Advect the vorticity field with a constant horizontal flow.

Define a computational domain that is H deep, and L wide, and grid intervals dx and dz . Then $L = (nx - 1) dx$ and $H = (nz - 1) dz$ where nx and nz are the number of grid points defined from the left to right boundaries and from the bottom to top boundaries, respectively.

For the implementation lateral boundary conditions, it is convenient to define one extra column of grid points outside the left and right physical boundaries, so that the grid point index goes from -1 to $nx+1$. With Fortran, you can declare arrays like $u(0:nx+1, nz)$ to match such index. With Python, all arrays have to start from 0 index.

For our term project, we will assume periodic lateral boundary conditions. For example, using Fortran, for u ,

$$u(0,:) = u(nx-1,:)$$

$$u(nx+1,:) = u(2,:)$$

The above conditions should be set for all state variables. The top and bottom boundary conditions are rigid walls, therefore $w = 0$ at $z = 0$ and $z = H$.

Problem 1: Assume zero initial environmental flow, therefore initial stream function is zero everywhere. At the top and bottom boundaries, ψ remains zero always, which is the top and bottom boundary conditions for ψ .

For H and $L = 10,000$ m, and $\Delta x = \Delta z = 100$ m,

$$\eta = \eta_0 \cos^2\left(\frac{\pi r}{2R}\right) \text{ for } r \leq R \text{ otherwise } \eta = 0, \quad (8)$$

$$\theta' = \theta'_0 \cos^2\left(\frac{\pi}{2} r_T\right) \text{ for } r_T = \left[(x-x_c)^2/R_x^2 + (z-z_c)^2/R_z^2 \right]^{1/2} \leq 1 \text{ otherwise } \theta' = 0, \quad (9)$$

where (x_c, z_c) is the coordinate of the center of thermal bubble and R_x and R_z are the horizontal and vertical radius of the elliptic thermal bubble.

where r is the radial distance from the domain center, and $R = \frac{1}{4} H$.

For the above vorticity distribution and $\eta_0 = 0.01 \text{ s}^{-1}$ solve for the stream function ψ using successive over-relaxation method. Confirm that your results are correct by plotting (within the physical domain without extra boundary points) the original η and η calculated from (6).

Calculate u and w from (5) and plot the wind vectors in the physical domain.

Problem 2: For constant flow $u = 10 \text{ m/s}$, $w = 0$, and the initial vorticity given by (8), integrate (4) without the right hand side buoyancy gradient term for 2000 s, and plot your η solutions at 500, 1000, 1500, 2000 s.

Use forward-in-time, upstream-in-space scheme, and Courant number 0.25, 0.5, 0.75, and 1.0 to do the integration. Briefly discuss the results.

Term Project Assignment – Part 2. Integrate potential temperature equation (3) and vorticity equation (4) with an initial thermal bubble and zero vorticity, to simulate rising warm bubble and dropping cold bubble.

Problem 1: Rising warm bubble/dry thermal convection

Set initial η to zero (and therefore u and w also to zero).

Define an initial thermal bubble by setting initial perturbation potential temperature as

$$\theta' = \theta'_0 \cos^2\left(\frac{\pi}{2} r_T\right) \text{ for } r_T = \left[(x - x_c)^2 / R_x^2 + (z - z_c)^2 / R_z^2 \right]^{1/2} \leq 1 \text{ otherwise } \theta' = 0, \quad (9)$$

where (x_c, z_c) is the coordinate of the center of thermal bubble and R_x and R_z are the horizontal and vertical radius of the elliptic thermal bubble.

Set the bubble amplitude $\theta'_0 = 5$ K. Set $(x_c, z_c) = (L/2, 2000)$, i.e., center the initial bubble at 2 km above ground and in the center for horizontal domain. Set $R_x = R_z = 2000.0$ meter. Set $H = 12,800$ m and $L = 25,600$ m, and $\Delta x = \Delta z = 100$ m.

Use centered difference for the right hand side term of equation (4) with $g = 9.8 \text{ m s}^{-2}$, $\theta_0 = 300$ K, and integrate equation (3) and equation (4) using forward-in-time, upstream-in-space scheme.

Step 1: Integrate equations (3) and (4) to obtain θ' and η at the next time level (in the interior domain). Assume that the maximum u and w will be 10 m s^{-1} , chose your time step size Δt that so that $\frac{\Delta t U_{\max}}{\Delta x} \leq \frac{1}{\sqrt{2}}$. The $\sqrt{2}$ factor is due to 2D advection.

Set the boundary conditions. For θ' and η , we assume zero gradient normal to the boundary for both. The following is a Python function that sets zero gradient condition on all boundaries.

```
def zero_gradient_bc(var, nx, nz):
    var[0, :] = var [1, :]
    var [nz-1, :] = var [nz-2, :]
    var[:, 0] = var[:, 1]
    var[:, nx-1] = var[:, nx-2]
    return 0
```

For velocity components u and w , we use the free-slip wall boundary conditions. That is, the normal wind component is zero, and the parallel wind component assumes zero normal gradient. For u , a python code would look like:

```
u[t+1, :, 1] = 0
u[t+1, :, nx-1] = 0
u[t+1, 0, :] = u[t+1, 1, :]
u[t+1, nz-1, :] = u[t+1, nz-2, :]
```

Step 2: Solve elliptic equation (7) to obtain ψ then calculate u and w as you did in Task 1, using the vorticity you obtained at the next time level. ψ at the boundaries should be set before the elliptic equation is solved. For the current problem, setting $\psi = 0$ at all boundries is appropriate.

This completes the full time step integration, and you have obtained all state variables at the next time level, from which you intergrate forward to more time steps until end time $T=2000$ s.

Given Δt and T , the total number of time levels involved in the integration will be $N=T/\Delta t+1$. For this 2D problem, memory is not an issue so you can declare arrays for all state variables like $u(0:nx+1, nz, N)$ so that you have values at all time steps for output and plotting purpose.

For 3D NWP problems, we never save states at all time levels in memory (it will require too much memory). For the current problem, we only need to declare potential temperature and vorticity arrays at 2 time levels, like $pt(0:nx+1, nz, 2)$, $vort(0:nx+1, nz, 2)$, and the velocity arrays at one time level, like $u(0:nx+1, nz)$, because within any time level, we don't need to know the states at more than these number of time levels. In that case, state variables need to be

written to disk files periodically, at specified intervals. Current NWP models often write output (forecasts) to disk every hour. Also, at the end of each time step, you need to transfer pt and vort at time level 2 to time level 1, i.e., $pt(:, :, 1) = pt(:, :, 2)$, $vort(:, :, 1) = vort(:, :, 2)$, so that the next time step integration will start again time level 1. You will need a time counter to keep track of the time.

Discuss your results.

Problem 2: Dropping cold bubble/density current along ground

Set $\theta'_0 = -15$ K, $(x_c, z_c) = (L/2, 3000)$ meter, $R_x = 4000$, $R_z = 2000$ meter in equation (9) for the initial cold thermal bubble. Set $H = 6,400$ m and $L = 51,200$ m, and $\Delta x = \Delta z = 100$ m. Simulate the dropping cold bubble/density current in Straka et al. (2003) (see <https://twister.caps.ou.edu/CFD2023/Straka.pdf>). Integrate the equations to 900 s, and output and plot the θ' and u fields using contours every 300s. Choose your time step size so that

$\frac{\Delta t U_{\max}}{\Delta x} \leq \frac{1}{\sqrt{2}}$. For this problem, the same boundary conditions as for the rising bubble can be used.

Discuss your results in comparison with the 100 m grid spacing results in Fig. 3 of the Straka paper at 900 s, and the reference solutions using 25 m grid spacing at additional times in Fig.1 of the Straka paper.

Given the large grid size (~ 512 x 64 grid points), the computation is likely slow, if you use python on a low-end laptop!